



# *Sediment transport dynamics & grain size trends recorded by Oligo-Miocene megafans in the Swiss Molasse basin*

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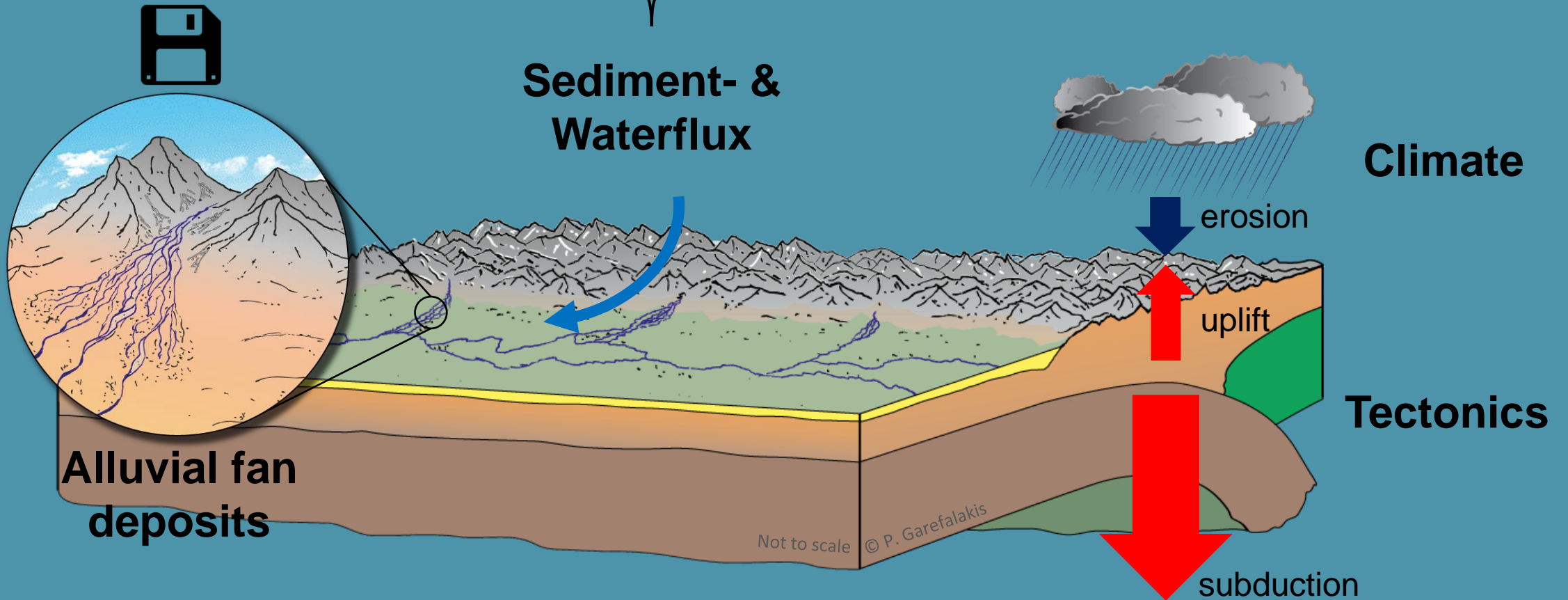
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**Motivation:** Foreland basin → *key to the adjacent mountain belt*  
Foreland basin deposits controlled by *climate* and *tectonics*.  
Information on sediment transport dynamics *stored in fan deposits*.

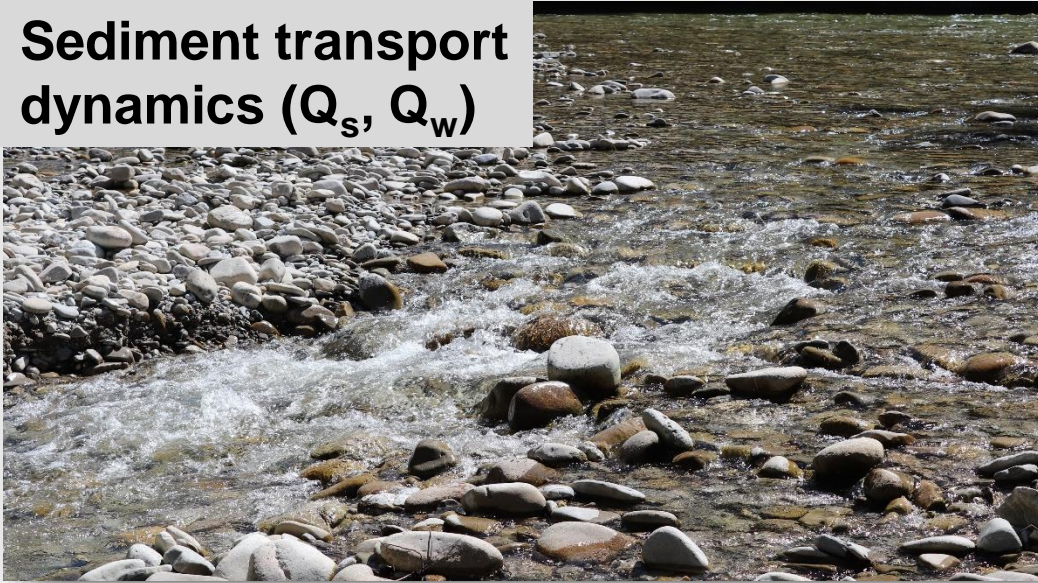




# Focus: Extraction of sediment fluxes from conglomerates

- i. Estimates of actual fan sediment fluxes  $Q_s$  by applying a self-similar model  
(e.g. Fedele & Paola, 2007; D'Arcy et al., 2017)

**Sediment transport  
dynamics ( $Q_s$ ,  $Q_w$ )**



We can extract sediment fluxes ( $Q_s$ ) from the stratigraphic record by applying a self-similar grain size fining model (e.g. Fedele & Paola, 2007; D'Arcy et al., 2017).

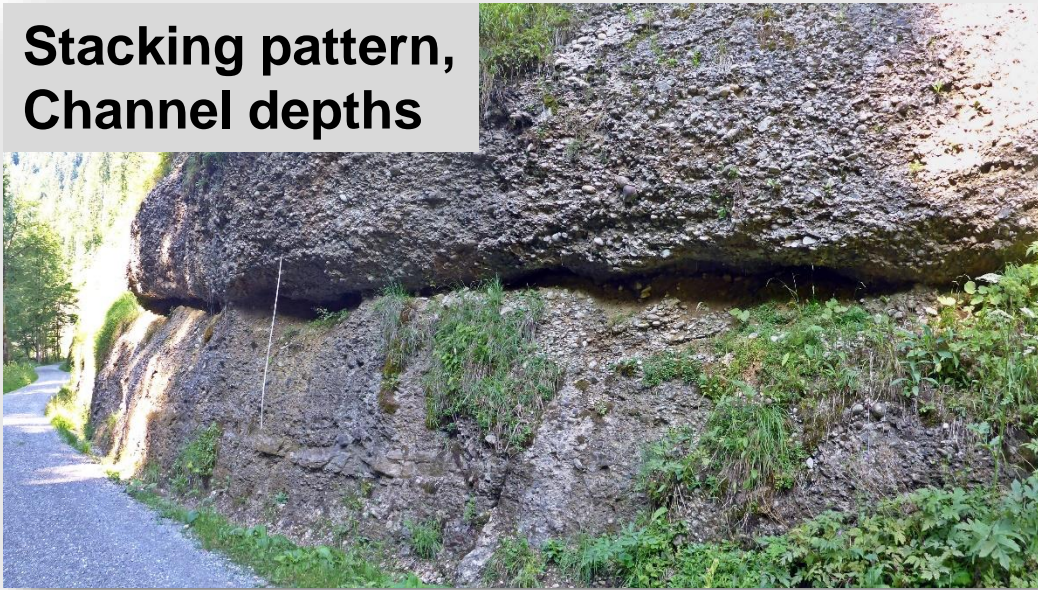
The model is based on:

- sediment accumulation rates / subsidence
- grain size (input grain size, fining rate)
- grain size distribution follows self-similarity

# Focus: Extraction of sediment fluxes from conglomerates

- i. Estimates of actual fan sediment fluxes  $Q_s$  by applying a self-similar model  
(e.g. Fedele & Paola, 2007; D'Arcy et al., 2017)

**Stacking pattern,  
Channel depths**



First we need to have information on the fan (or channel) morphometry by analysing the stacking pattern (architecture of the fan) and measuring channel depths preserved during bankfull flow-depth conditions.

This helps to estimate fan surface slopes, in combination with measurements on grain size...



# Focus: Extraction of sediment fluxes from conglomerates

- i. Estimates of actual fan sediment fluxes  $Q_s$  by applying a self-similar model  
(e.g. Fedele & Paola, 2007; D'Arcy et al., 2017)

**Grain size distribution**  
**Grain size fining trends**



...which we measured on digital photographs using the Wolman (1954) approach.

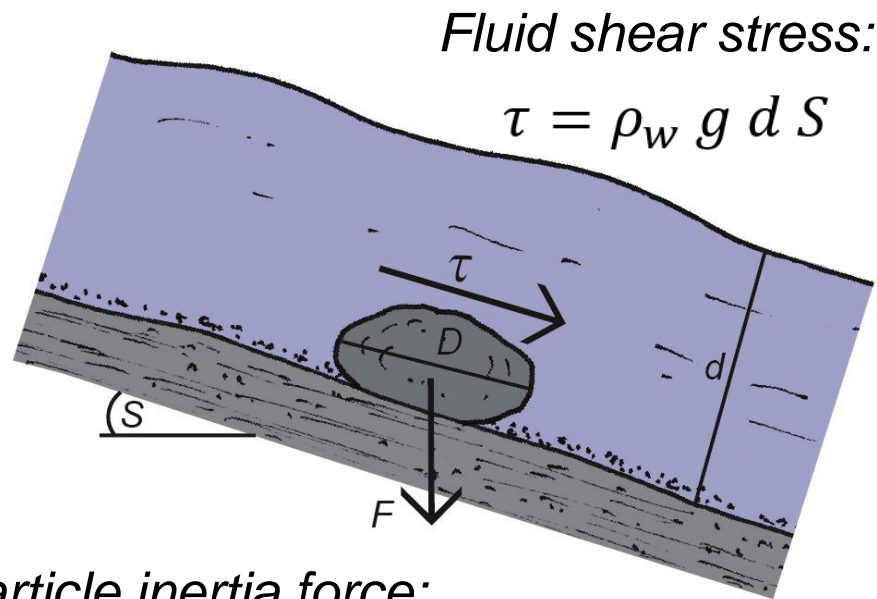
For each outcrop we measure 100 grains and calculate the percentiles of interest (e.g.  $D_{50}$  for the 50<sup>th</sup> percentile, or  $D_{84}$  for the 84<sup>th</sup> percentile).

In combination with the channel depth and hydrological formulas based on initial sediment transport, we can estimate the fan surface slopes.

Furthermore, each section of interest discloses proximal-distal geometries from the Apex towards distal positions, which is needed to calculate the grain size fining trends (decrease of grain size along distance).

# Focus: Extraction of sediment fluxes from conglomerates

- ii. Estimates of fan capacities (or bedload gravel-fluxes  $Q_b$ ) by paleo-hydrological calculations to get fan intermittencies (e.g. Meyer-Peter & Müller, 1948; Wong & Parker, 2006)

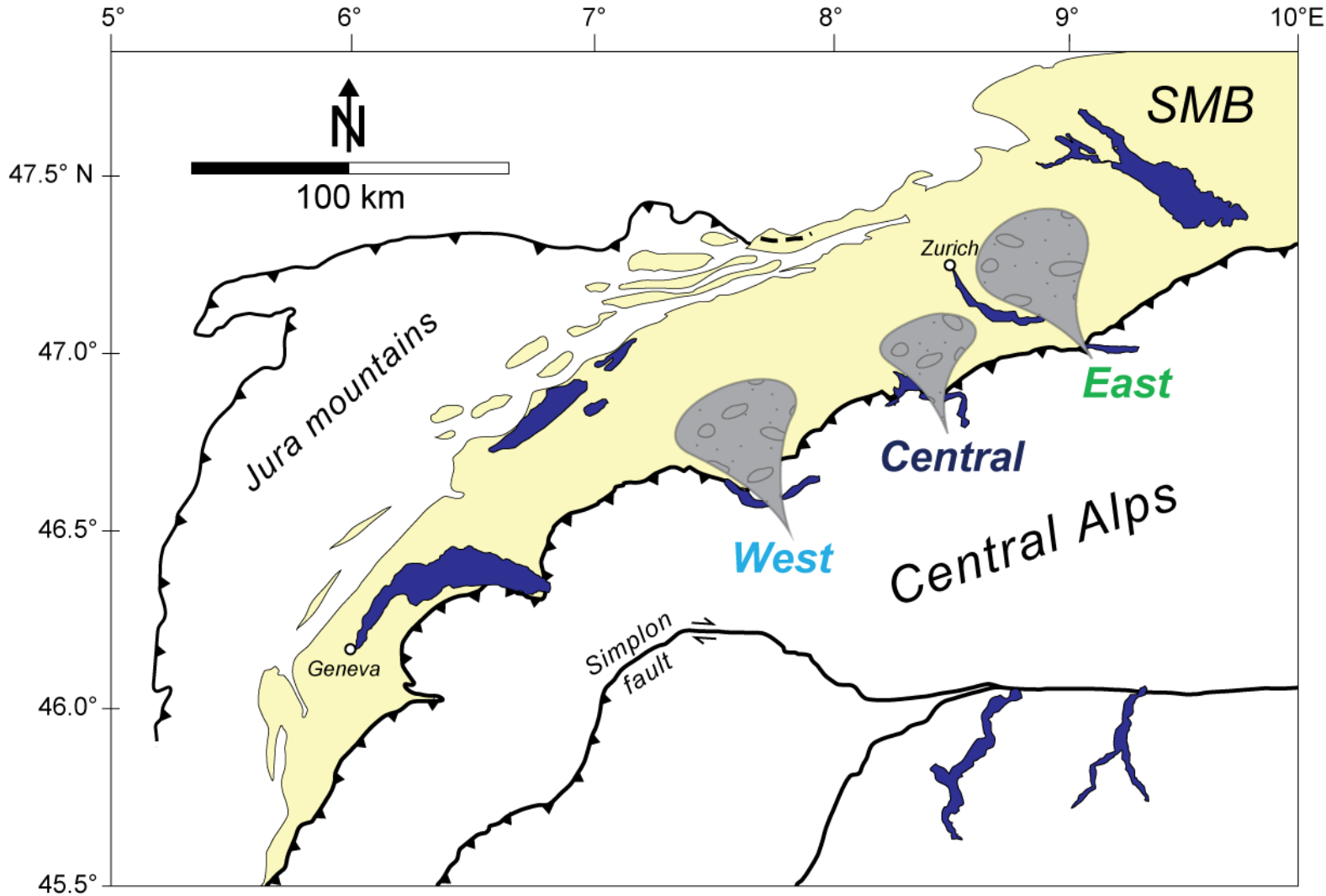


We can calculate the bedload transport rate in dry weight per unit channel width  $q_b$  (or alternatively the bedload gravel-flux or the sediment flux capacities) of the fans by combining:

- The dimensionless bedload transport capacity after Meyer-Peter & Müller (1948) and
- The dimensionless bedload transport rate after Einstein, H.A. (1950)
- We used the recalibrated formula based on the reanalysis of the original Meyer-Peter & Müller (1948) work by Wong & Parker (2006)

Mod. after Garefalakis & Schlunegger, 2018

# Study area: Swiss Molasse Basin (SMB)



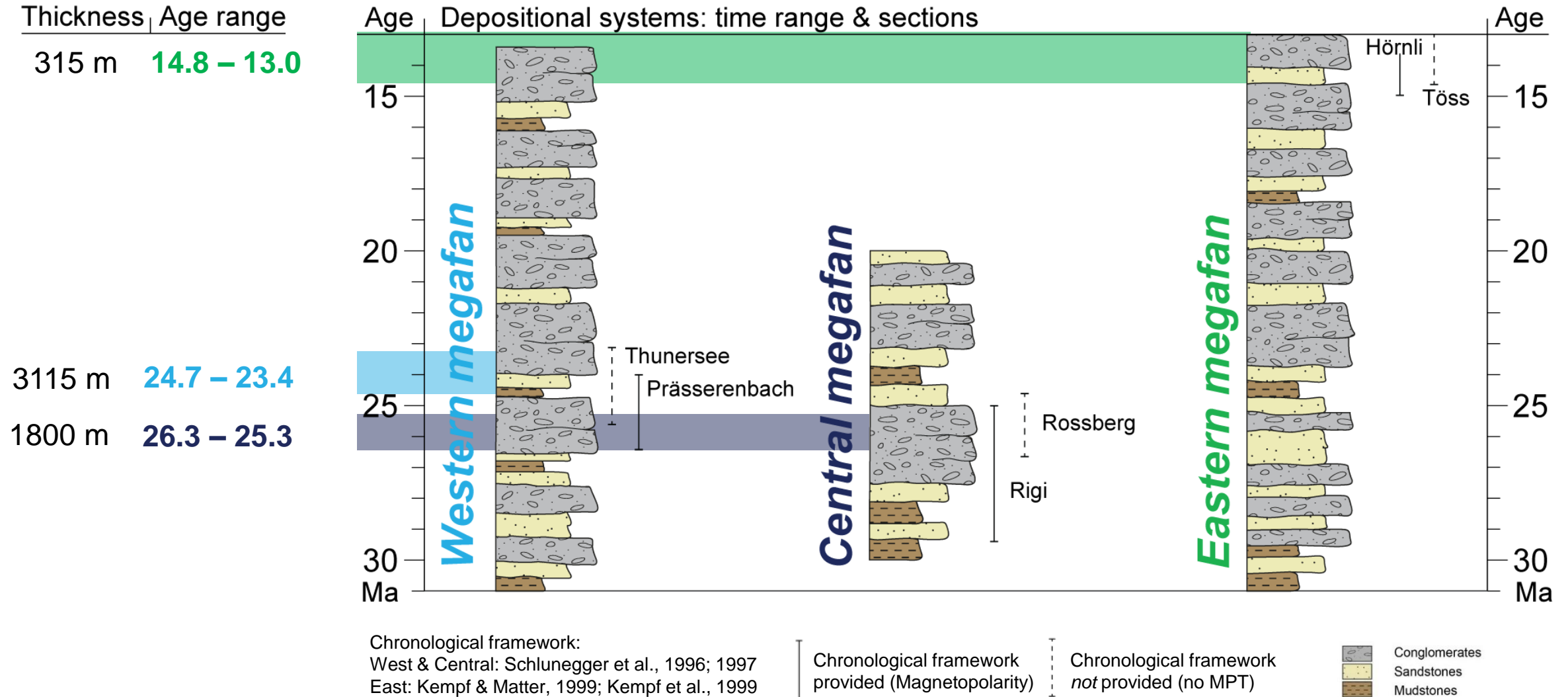
Depositional centres (megafans)  
active during last c. 35 – 10 Ma

Recorders of the erosional history of  
the Alps and analysed for:

- temporal & spatial relationships  
to Alpine formation
- paleogeography, source areas  
and clast composition

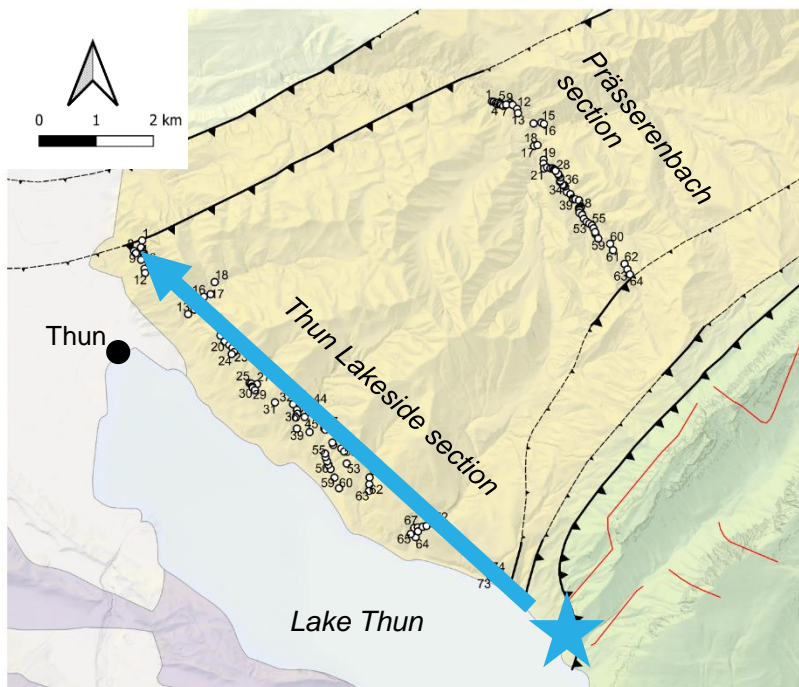
Here, we focus on three individual  
sections with proximal-distal  
geometries and extract grain size  
fining trends and sediment fluxes at  
the section scale.

# Study area: Proximal-distal sections and temporal framework

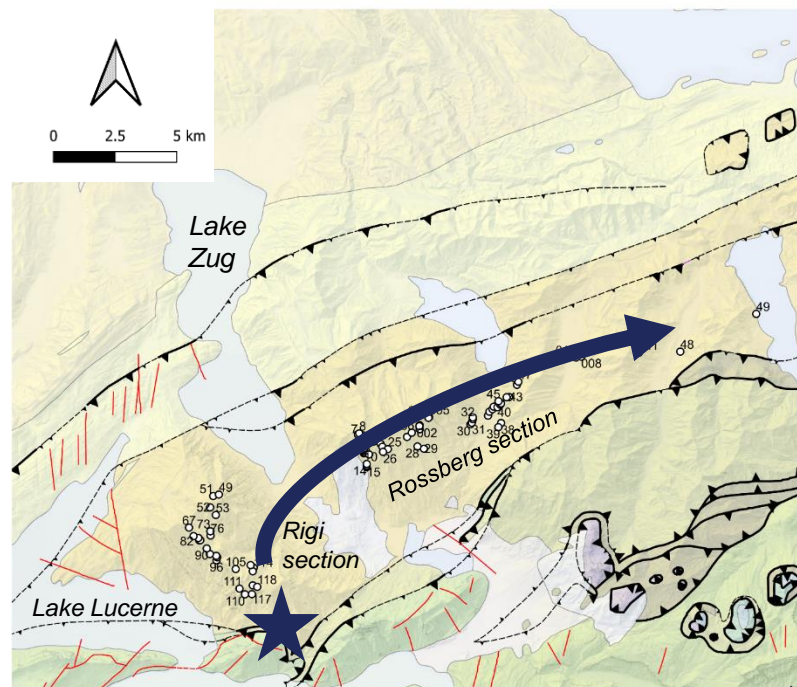




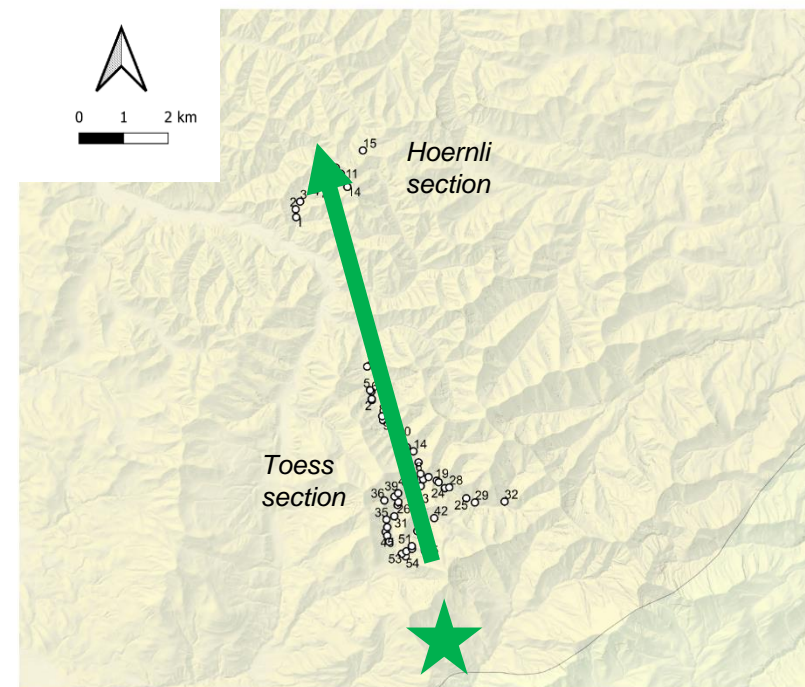
# Study area: Proximal-distal sections in the Swiss Molasse Basin



**West – c. 12 km**



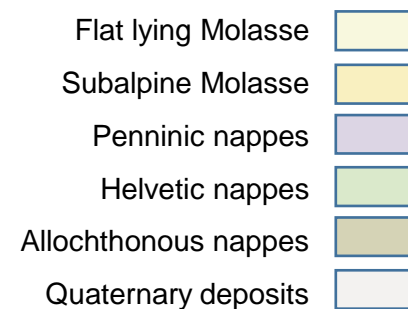
**Central – c. 30 km**



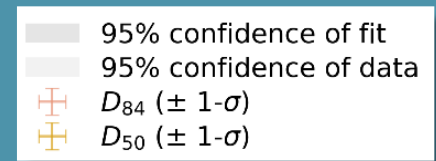
**East – c. 12 km**

## Discharge direction based on paleo-flow measurements

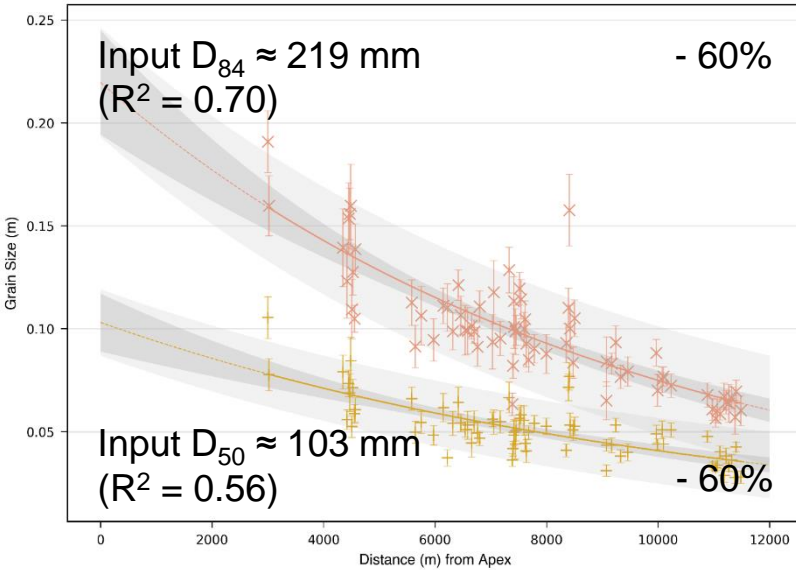
- preserved channel geometries
- sets of imbricated clasts, gutter casts or cross-beds
- identification of paleo-apex from stacking pattern and clast morphometry ★



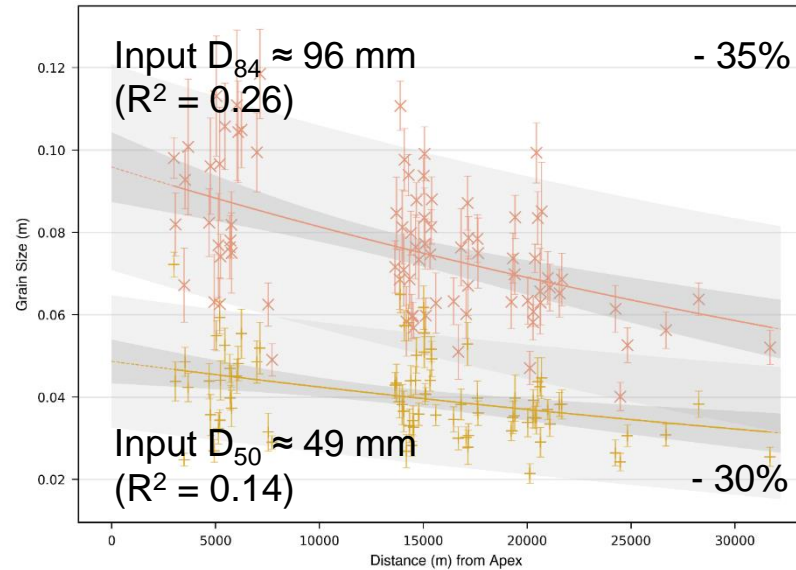
# Results: Grain size trends along distance & characteristics



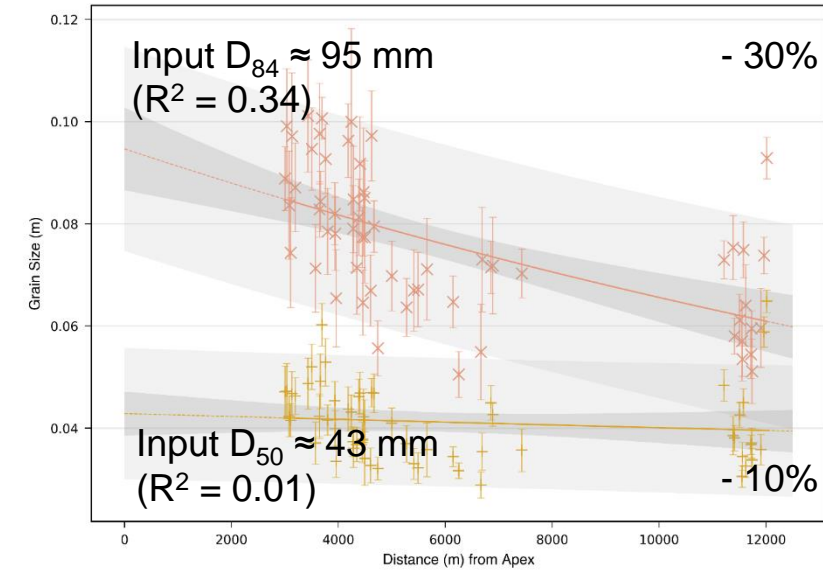
## West (c. 12 km)



## Central (c. 30 km)



## East (c. 12 km)



$D_{84}$  avg.

98 mm

$D_{50}$  avg.

51 mm

$d_{bf}$  avg.

1.40 m

$S_{(D84)}$  avg.

0.60% or 0.36°

$S_{(D50)}$  avg.

0.30% or 0.18°

76 mm

40 mm

1.45 m

0.50% or 0.30°

0.25% or 0.15°

76 mm

41 mm

1.50 m

0.45% or 0.26°

0.24% or 0.14°

# Results: Sediment fluxes & activities of target sections

## → Actual sediment fluxes per unit width:

- sediment accumulation rates / subsidence
- grain size (input grain size, fining rate)
- self-similarity
- e.g., Fedele & Paola, 2007; D'Arcy et al., 2017

Average **actual** sediment fluxes ( $D_{84}$  &  $D_{50}$ )

**West** =  $18.8 \pm 1.45 \text{ km}^2 \cdot \text{Myr}^{-1}$

**Central** =  $39.8 \pm 3.74 \text{ km}^2 \cdot \text{Myr}^{-1}$

**East** =  $6.6 \pm 1.6 \text{ km}^2 \cdot \text{Myr}^{-1}$

## → Gravel bedload capacity per unit width:

- based on sediment transport equations
- grain size, channel depth and slope
- critical shear stress (Shields stress)
- e.g., Wong & Parker, 2006

Average **gravel bedload** capacities ( $D_{84}$  &  $D_{50}$ )

**West** =  $1200 \text{ km}^2 \cdot \text{Myr}^{-1}$

**Central** =  $1300 \text{ km}^2 \cdot \text{Myr}^{-1}$

**East** =  $1200 \text{ km}^2 \cdot \text{Myr}^{-1}$

## → Fan activity (intermittency):

- ratio between actual sediment flux and gravel-bedload capacity

**West** =  $1 - 4 \%$  (or  $3 - 14 \text{ days / yr}$ )

**Central** =  $2 - 6 \%$  (or  $8 - 21 \text{ days / yr}$ )

**East** =  $0.2 - 2 \%$  (or  $1 - 7 \text{ days / yr}$ )



# Discussion: Sediment fluxes & activities of target sections

## → Actual sediment fluxes per unit width:

- Presumably controlled by differences in mountain building processes (tectonics / erosion)
- Paleo-climate: climate was probably similar (e.g., Zachos, 2001; Mosbrugger et al., 2005)

Average **actual** sediment fluxes ( $D_{84}$  &  $D_{50}$ )

**West** =  $18.8 \pm 1.45 \text{ km}^2 \cdot \text{Myr}^{-1}$

**Central** =  $39.8 \pm 3.74 \text{ km}^2 \cdot \text{Myr}^{-1}$

**East** =  $6.6 \pm 1.6 \text{ km}^2 \cdot \text{Myr}^{-1}$

## → Gravel bedload capacity per unit width:

- Similar fan characteristics (average grain size, channel depths, slopes, stacking pattern / architecture and appearance of conglomerates are prone for a similar fan (or channel) morphometry)

Average **gravel bedload** capacities ( $D_{84}$  &  $D_{50}$ )

**West** =  $1200 \text{ km}^2 \cdot \text{Myr}^{-1}$

**Central** =  $1300 \text{ km}^2 \cdot \text{Myr}^{-1}$

**East** =  $1200 \text{ km}^2 \cdot \text{Myr}^{-1}$

## → Fan activity (intermittency):

- Highest sediment concentration in Central fan, probably driven by tectonic processes (Garefalakis & Schlunegger, 2018)

**West** = 1 – 4 % (or 3 – 14 days / yr)

**Central** = 2 – 6 % (or 8 – 21 days / yr)

**East** = 0.2 – 2 % (or 1 – 7 days / yr)

# Questions and comments? Happy to help!

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