



Observational study of time-varying climate feedback parameter

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Equilibrium climate sensitivity

Energy budget equation

$$\left[\begin{array}{rcl} N & = & F + \lambda T \quad (W \cdot m^{-2}) \\ \text{incoming rad.} - \text{outgoing rad.} & = & \text{rad. forcing} + \text{rad. response} \end{array} \right]_{\text{TOA}}$$

[CHARNEY *et al.*, 1979; RAMANATHAN, 1987]

TOA : *Top of atmosphere*

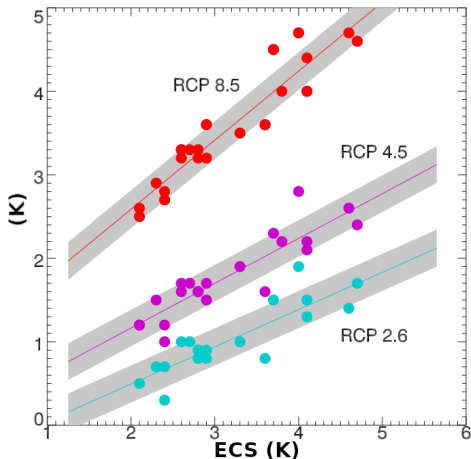
- N : Earth energy imbalance
- F : radiative forcing
- T : surface temperature
- λ : **climate feedback parameter**

Equilibrium climate sensitivity (*ECS*)

[ARRHENIUS, 1896; MANABE & WETHERALD, 1967; CHARNEY *et al.*, 1979]

$$ECS = -\frac{F_{2x}}{\lambda}$$

Fundamental metric of climate change amplitude and projections

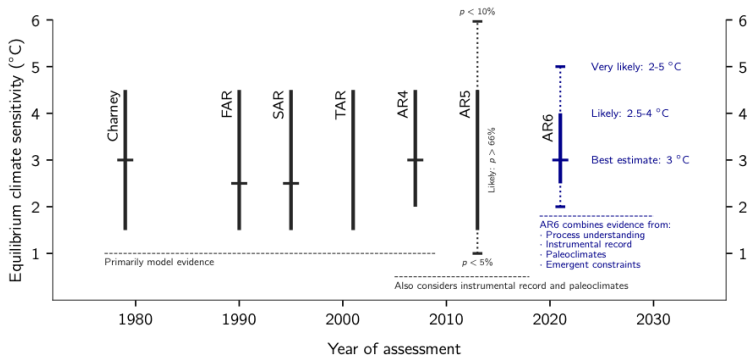


$T(2100) \propto ECS$ for three IPCC socio-economic scenarios

Adapted from [SHERWOOD *et al.*, 2020]

Problem : ECS is still very uncertain !

a) Evolution of equilibrium climate sensitivity assessments from Charney to AR6



1979-2013 :

$1.5 \leq ECS \leq 4.5$ K (likely)

[CHARNEY *et al.*, 1979; IPCC, 2013]

Recently :

$2.3 \leq ECS \leq 4.5$ K (likely) [SHERWOOD *et al.*, 2020]

$2.5 \leq ECS \leq 4.0$ K (likely) [ARIAS *et al.*, 2021]

$2.0 \leq ECS \leq 5.0$ K (very likely) [ARIAS *et al.*, 2021]

Inconsistencies between methods despite recent attempts of reconciliation between methods

[ANDREWS *et al.*, 2018; SHERWOOD *et al.*, 2020]

- Observational estimates : low values
- Climate models estimates : high values

Estimate of ECS over 1971-2017 with observations of :

- F from [SHERWOOD *et al.*, 2020] (non aerosol) and [BELLOUIN *et al.*, 2020] (aerosols)
- T from [COWTAN & WAY, 2014] scaled by [RICHARDSON *et al.*, 2016]

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- $N \propto dOHC/dt$ where OHC is the ocean heat content from in situ T/S measurements ;
we use an ensemble of 5 solutions : [GOURETSKI & KOLTERMANN, 2007 ; LEVITUS *et al.*, 2009 ; LEVITUS *et al.*, 2012 ; GOOD *et al.*, 2013 ; CHENG *et al.*, 2017 ; ISHII *et al.*, 2017]

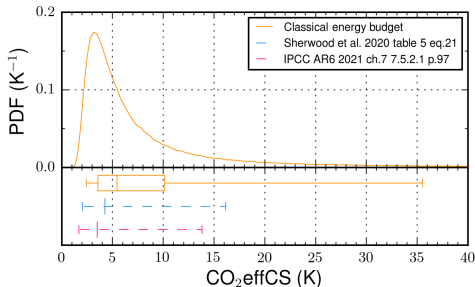
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Method : **regression of $N - F$ over T**
with rigorous handling of uncertainties

Results published in *Journal of climate*
[CHENAL *et al.*, 2022]

ECS = 5.5 [2.4 ; 35.6] K (very likely)



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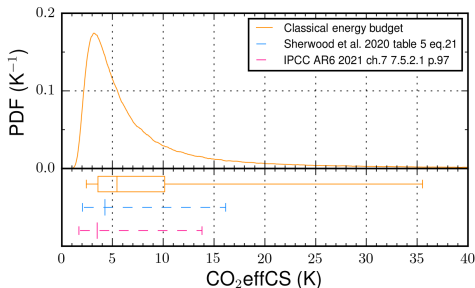
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With comparable approach :

- [SHERWOOD *et al.*, 2020] :
 $ECS = 4.26 [2.04; 16.13] K$
- [FORSTER *et al.*, 2021] :
 $ECS = 3.5 [1.7; 13.8] K$



Time-varying climate feedback parameter

Influence of the mean epoch and duration of observations on the estimate of parameter λ

We extend our EEI solution on 1957-2017 from [MEYSSIGNAC *et al.*, in prep.] :

- **thermsteric component of [FREDERIKSE *et al.*, 2020] sea level reconstruction by GMSL - GMBSL (low-pass filter, 15yr)**
- **in situ solutions (5-solutions ensemble + ARANN [BAGNELL & DE VRIES, 2021]) (low-pass filter, 10yr)**

Radiative forcing and temperature : low-pass filter, 10yr

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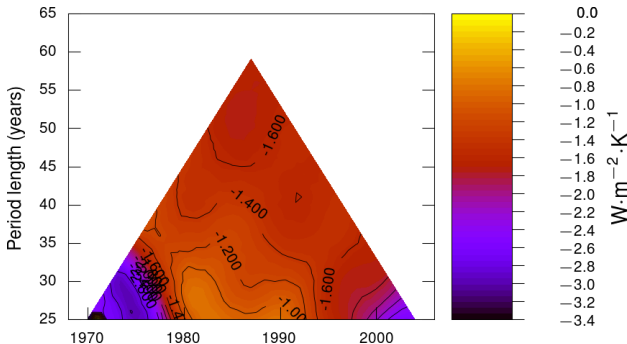
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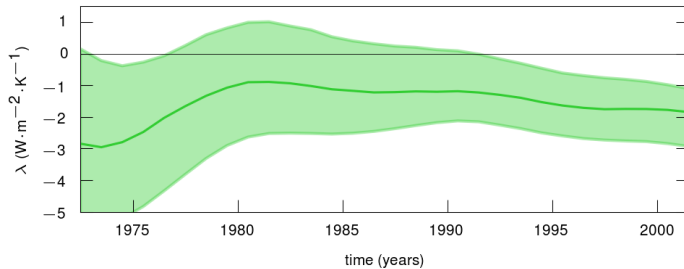
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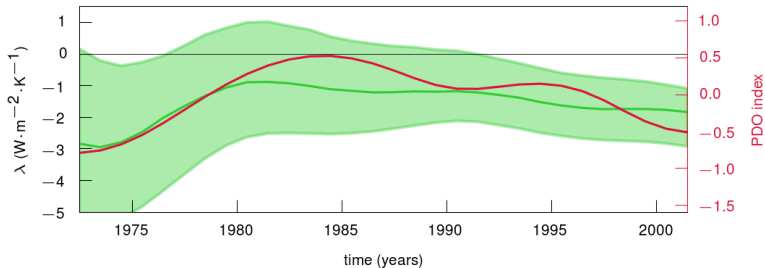
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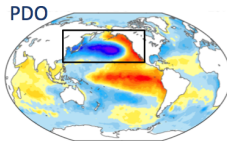
30-yr duration λ estimates [MEYSSIGNAC *et al.*, in prep.]



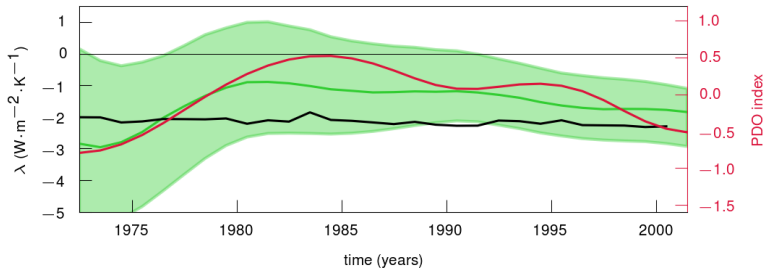
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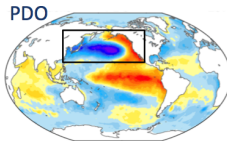
Role of the **Pacific Decadal Oscillation** on parameter λ ?
(here NOAA ERSST v5 PDO indexed, low-pass filtered)



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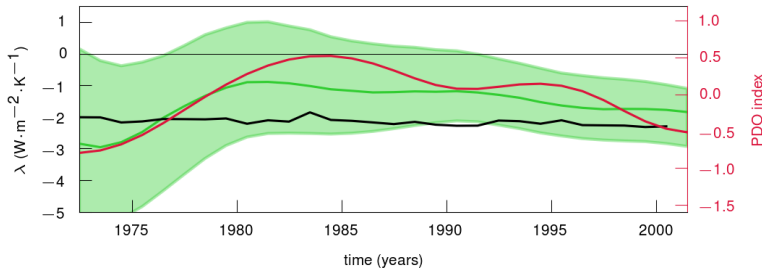


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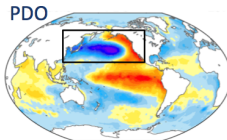


However **climate models** poorly reproduce $\lambda(t)$...
(here ensemble-mean from 8 climate models amip-piForcing simulations)

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But very negative λ can be a possible explanation of other observational ECS low estimates

Conclusions

Conclusions

- Rigorous handling of uncertainties due to climate variability in climate sensitivity estimate
 - ▶ low ECS (≤ 2.4 K) are very unlikely
 - ▶ Median ECS $\gtrsim 3.0$ K
 - ▶ Reconciliation of observational and models estimates
- **Observational constraint on $\lambda(t)$ for climate models and ECS observational studies**

Thanks for your attention

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