

SANS: Publicly Available Daily Multi-Scale Seismic Ambient Noise Source Maps



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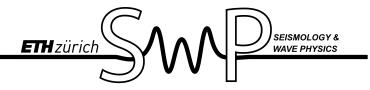
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Outline

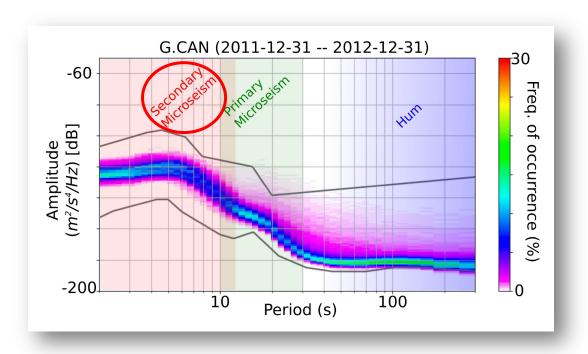


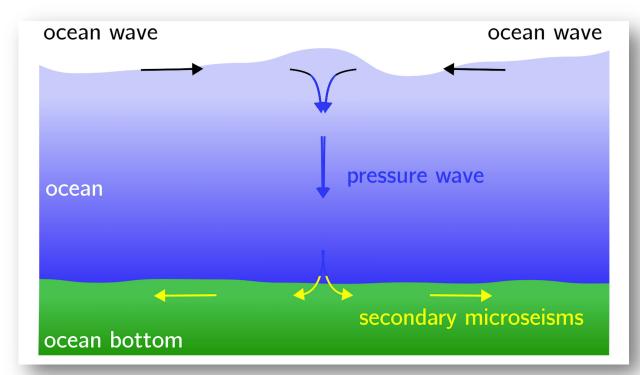
- Secondary microseisms
- Matched Field Processing (MFP)
- Finite-frequency Inversion
- **SANS**: Publicly Available Daily Seismic Ambient Noise Source Maps
- Conclusions and Outlook

Secondary microseisms



1. We focus on the secondary microseisms which are in a frequency range of 0.1-0.2 Hz.



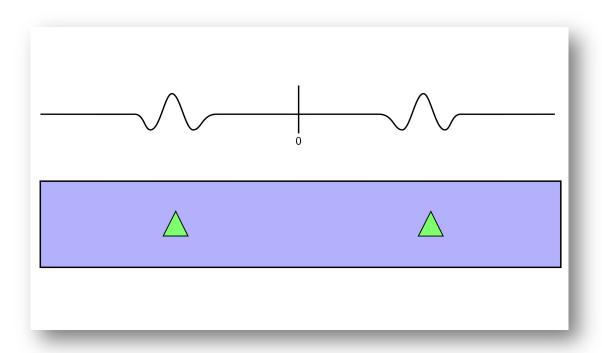


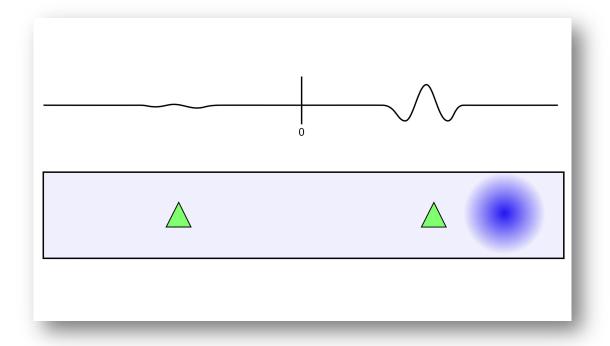
2. Generation mechanism: Two overlapping ocean waves travelling in opposite directions create vertical pressure waves which leads to a tiny displacement at the ocean bottom.

Seismic Noise Cross-Correlations



1. To study the seismic noise source distribution we use cross-correlations, specifically their (a)symmetry.

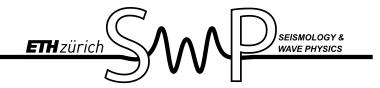




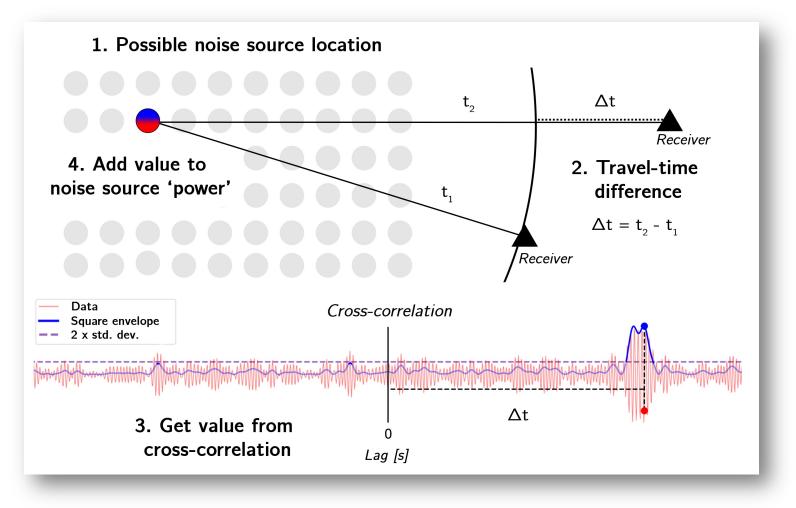
2. If the noise source distribution within the domain is homogeneous we would get a symmetric cross-correlation.

3. A heterogeneous distribution would result in an asymmetric cross-correlation.

Matched Field Processing

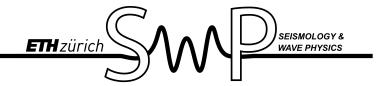


1. One method to locate noise sources is Matched Field Processing (MFP) [1]. It is an efficient data-driven method that allows us to get a rough idea of where the dominant noise sources are.

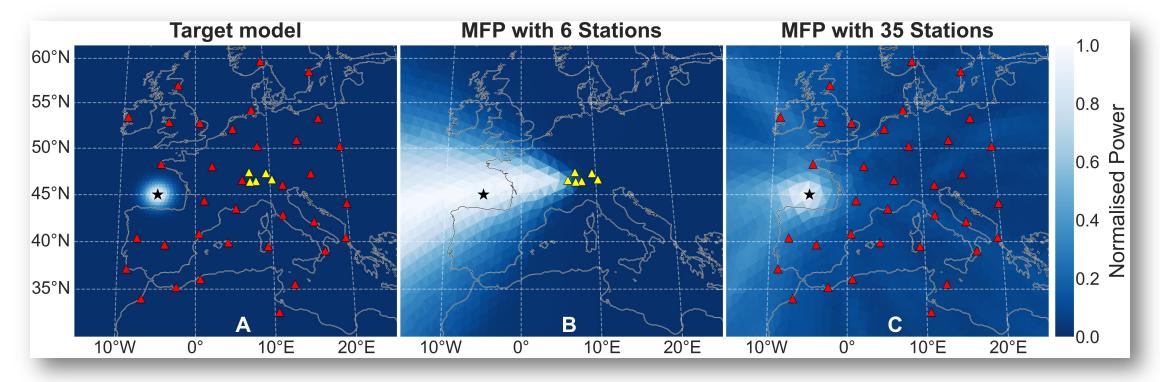


- 2. MFP works by iterating over a grid of possible noise sources and calculating the expected travel times to each station.
- The expected travel time difference indicates at what lag in the crosscorrelation we would expect a source for this location given a constant surface wave speed.
- 4. The value of the cross-correlation (or in our case the square envelope of the cross-correlation) is added to the 'power' of that noise source location.
- 5. This is repated for all cross-correlations.

Matched Field Processing



1. Here we show a synthetic example of the MFP algorithm. We forward model cross-correlations using the target model (A) and run MFP for two different station distributions (B) and (C).



- 2. When the stations (△) do not surround the source (B) MFP is not able to spatially restrict the noise sources but mainly provides a direction.
- 3. When the stations (▲) surround the source more (C) MFP can also restrict the spatial extent of the noise source.

Finite-frequency Inversion: Correlations



1. A more complex method to locate noise sources is to perform a finite-frequency inversion [2,3]. This requires more computational power but models the wave propagation more accurately.

$$\mathbf{C}(x_1, x_2, \omega) = \int \mathbf{G}(x_1, \xi, \omega) [\mathbf{G}^*(x_2, \xi, \omega) \mathbf{S}(\xi, \omega)] d\xi$$

C = Cross-correlation wavefield

G = Green's Function

S = Power Spectral Density of Source Distribution

2. This equation allows us to forward model a cross-correlation C for any noise source distribution S if we have the Green's functions G for the stations at locations x_1 and x_2 .



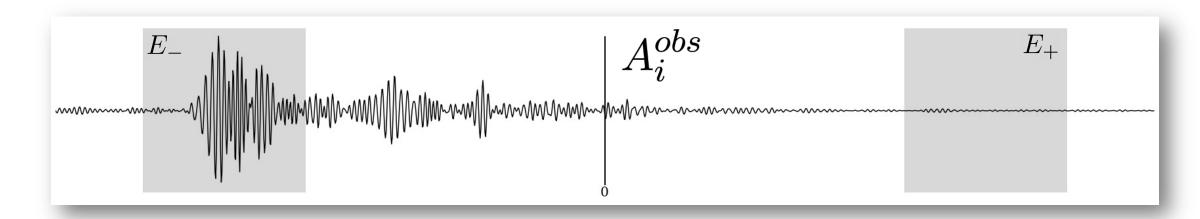
Finite-frequency Inversion: Measurement



1. For the inversion we use a certain measurement in the cross-correlation: the logarithmic energy ratio. This is chosen due to its relative insensitivity to unknown 3-D Earth structure.

Logarithmic
$$A$$
 : Energy Ratio

$$A = \ln\left(\frac{E_+}{E_-}\right)$$



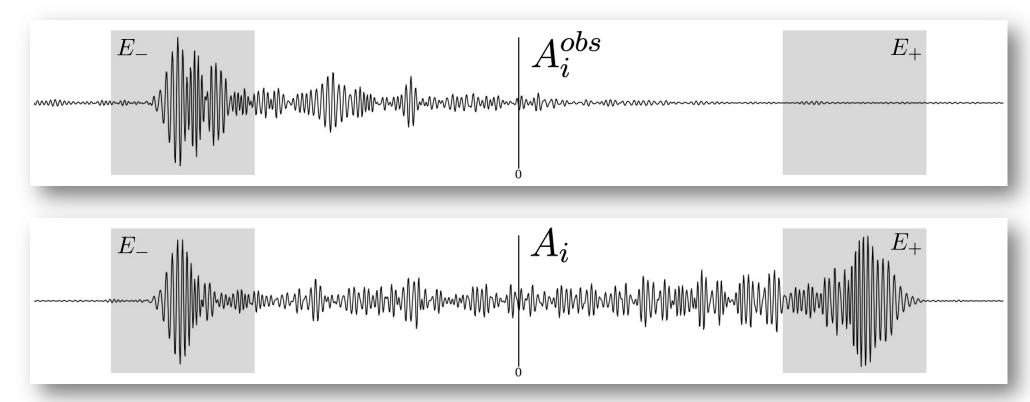
2. It measures the asymmetry in the cross-correlation by taking the natural logarithm of the ratio of energies in the expected surface wave arrival windows in the causal (E_{+}) and acausal (E_{-}) part of the correlation.

Finite-frequency Inversion: Misfit



1. We take the logarithmic energy ratio measurement on the synthetic (A_i) and observed cross-correlation (A_i ^{obs}) and compare the two using the L_2 norm. We aim to minimise this misfit during the inversion.

$$\mathcal{L}_2$$
 norm $\chi = \frac{1}{2} \sum_{i=1}^{N} \left[A_i - A_i^{obs} \right]^2$



Finite-frequency Inversion: Sensitivity Kernel



1. Adjoint techniques allow us to use this measurement to calculate the adjoint sources f and thus the sensitivity kernels. They indicate where an increase or decrease in noise source strength would decrease the misfit.

$$K(\xi,\omega) = 2f(\omega) \cdot (G^*(x_1,\xi,\omega) \cdot G(x_2,\xi,\omega))$$

K = Sensitivity Kernel

G = Green's Function

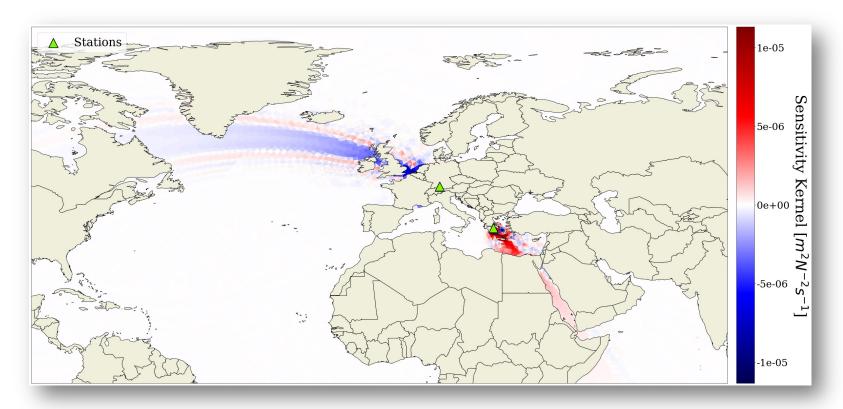
f = Measurement-dependent Adjoint Source

- 2. Similar to the correlation equation, the sensitivity kernel equation uses the Green's functions G for the stations at location x_1 and x_2 .
- 3. We reduce the computational cost of the forward model by (i) pre-computing the Green's functions G using AxiSEM with a simple PREM 1-D velocity model and (ii) using a spatially variable grid to reduce the number of source locations that need to be modelled.

Finite-frequency Inversion: Gradient

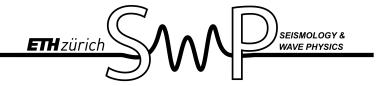


1. Here we see an example of a synthetic sensitivity kernel. Blue indicates that an increase in noise source strength would decrease the misfit.

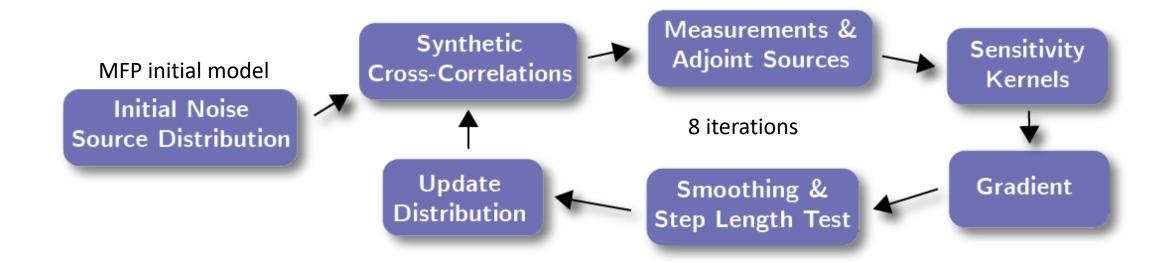


2. For multiple stations we sum up all the sensitivity kernels to obtain the gradient. This is then used to update the noise source distribution for the next iteration of the inversion.

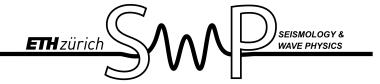
Finite-frequency Inversion



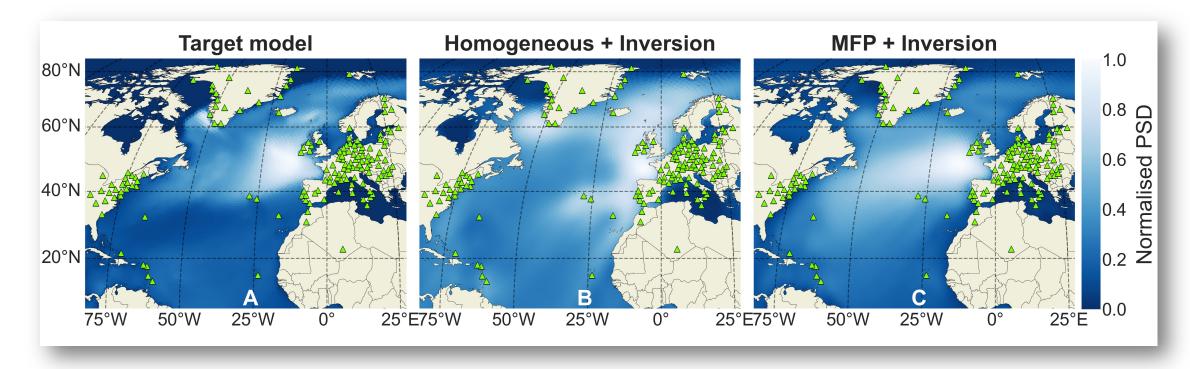
- 1. We use a gradient-based iterative method to decrease the misfit in our inversion. The initial noise source distribution is a normalised smoothed MFP model.
- 2. Synthetic and real data tests have shown that the noise source distribution does not change much after roughly 5 iterations. We run 8 iterations to ensure that the model explains the data well.



Finite-frequency Inversion



1. The MFP initial model is able to steer the inversion in the right direction and helps to ensure that we are not in a local minimum in the inversion.



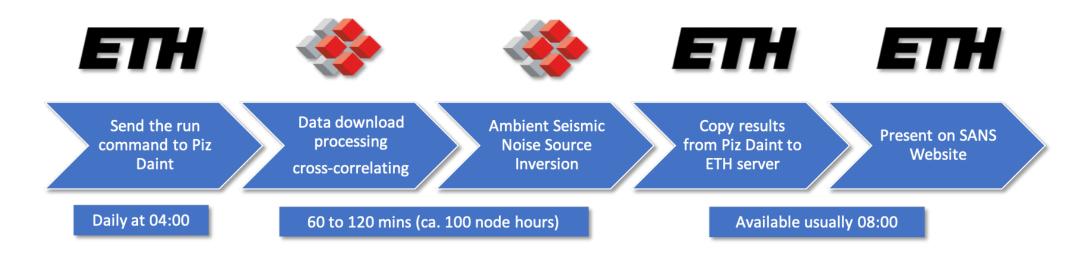
2. Here we forward model synthetic correlations with added noise for a target model (A) and stations (△) surrounding the North Atlantic.

3. The MFP starting model (C) leads to a more accurate noise source distribution compared to a homogeneous starting model (B) after 8 iterations.

SANS: Daily Seismic Ambient Noise Sources



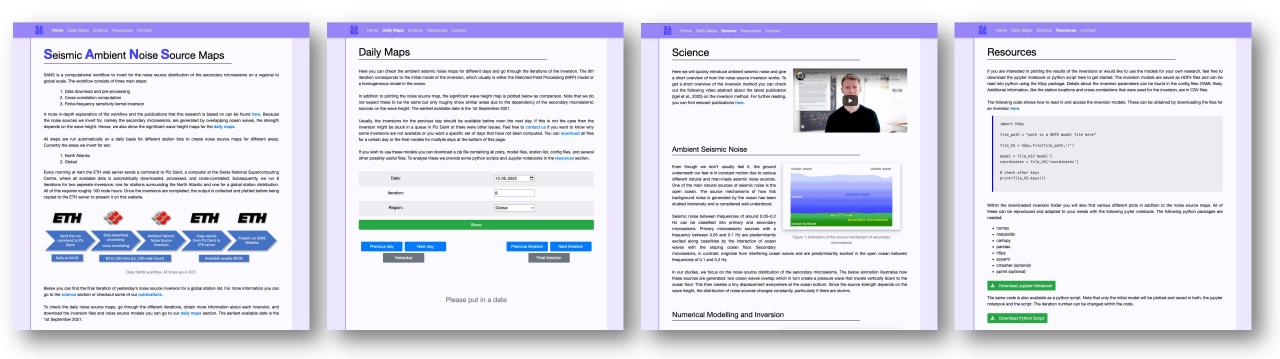
1. SANS is a framework where the previously mentioned inversion method is run on a daily basis on a regional to global scale. The daily Seismic Ambient Noise Source maps are made available to the public on the website: sans.ethz.ch. The workflow of this framework is illustrated here where Piz Daint is a supercomputer at the Swiss National Supercomputing Centre CSCS:



2. Users can study the noise source maps online or download them for implementation in their own studies. We provide example code on how to open and plot the files.

SANS: Daily Seismic Ambient Noise Sources



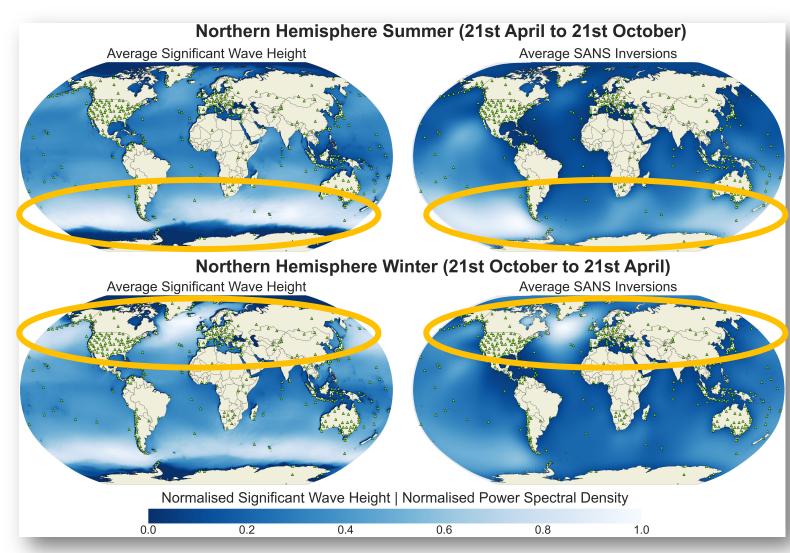


Check it out here: sans.ethz.ch

SANS: Seasonal analysis



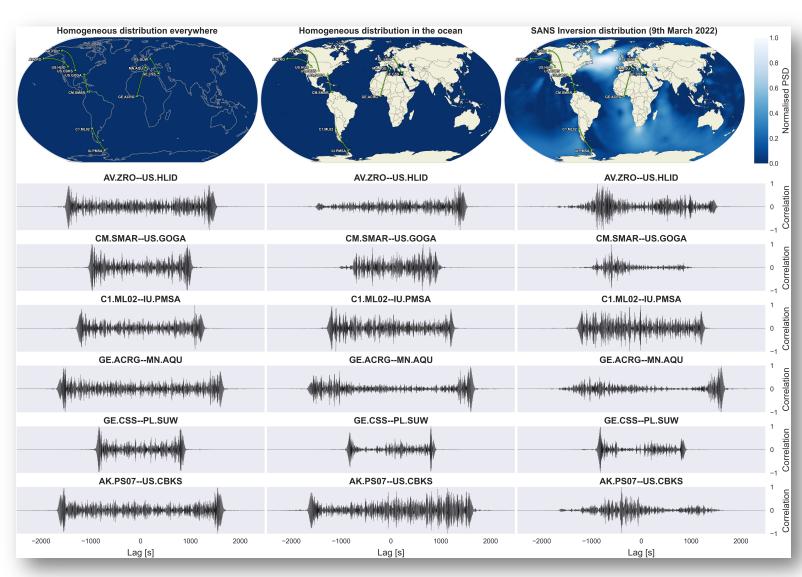
- 1. We analyse nearly a full year of daily global inversions from the 21st May 2021 to the 2nd May 2022.
- 2. The significant wave height maps from the WaveWatch III model [4] should only be taken as a rough reference and not a direct comparison as the generation mechanism is more complicated.
- 3. Averaging all inversions and wave heights for northern hemisphere summer and winter show the clear variations due to the seasons. NH summer is dominated by sources in the southern hemisphere and NH winter is dominated by sources in the North Atlantic. The wave height maps show similar features.



SANS: Correlation comparison



- 1. We compare cross-correlations for three different models: homogeneous everywhere, homogeneous in the ocean, and the final SANS map for the 9th March 2022. Note that blue for the homogeneous distributions indicates 1 everywhere.
- 2. As we introduce more realistic noise source distributions the correlations become more and more asymmetric.
- 3. This can have a strong effect on ambient noise tomography studies that assume a homogeneous distribution, especially full waveform methods.

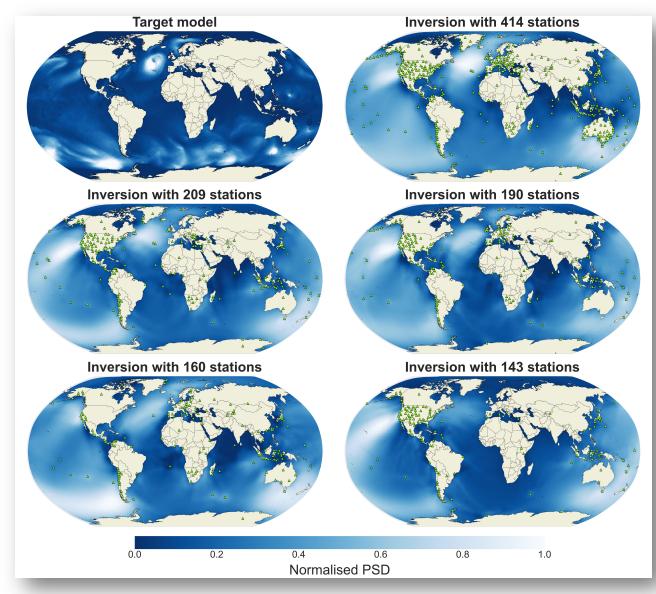


SANS: Resolution



- Due to lack of data or earthquakes that have to be removed from the waveforms the data availability and thus station distribution changes every day.
- We study the effect of different station distributions by forward modelling correlations and inverting for that data using several different station lists.

3. Despite the lack of data for fewer stations the dominant noise source are included in the final model for most station lists. Generally, the resolution is lower in the southern hemisphere due to the lack of stations.



Conclusions & Outlook



- 1. We present a combination of two methods, MFP and finite-frequency inversion, to invert for the noise source distribution of the secondary microseisms on a regional to global scale on a daily basis.
- 2. The SANS framework makes these daily Seismic Ambient Noise Source maps available to the public: sans.ethz.ch
- 3. Analysis of one year of daily inversions shows the expected spatio-temporal variations due to the seasons, i.e. northern hemisphere winter has stronger sources in the southern hemisphere and northern hemisphere summer has the most dominant noise sources in the North Atlantic.
- 4. Implementation of noise source distribution knowledge into other ambient noise studies is strongly encouraged as it has a significant effect on the cross-correlations and possible travel time picks.

Feel free to contact us if you have any questions or want to collaborate: jonas.igel@erdw.ethz.ch



References



- 1. Bowden, D. C., Sager, K., Fichtner, A. & Chmiel, M. Connecting beamforming and kernel-based noise source inversion. *Geophysical Journal International* **224**, 1607–1620 (2021).
- 2. Igel, J. K. H., Ermert, L. A. & Fichtner, A. Rapid finite-frequency microseismic noise source inversion at regional to global scales. *Geophysical Journal International* **227**, 169–183 (2021).
- 3. Ermert, L. *et al.* noisi: A Python tool for ambient noise cross-correlation modeling and noise source inversion. *Solid Earth* 1597–1615 (2020).
- 4. Tolman, H. L. & Chalikov, D. Source terms in a third-generation wind wave model. *Journal of Physical Oceanography* **26**, 2497–2518 (1996).