

# **Growth, replacement and element diffusion in garnet**

## **Part II: OXYGEN DIFFUSION**

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Uni Lausanne, ANU, Columbia University

- 1. University of Bern, Institute of Geological Sciences, Switzerland*
- 2. Institut de Sciences de la Terre, University of Lausanne, Switzerland*

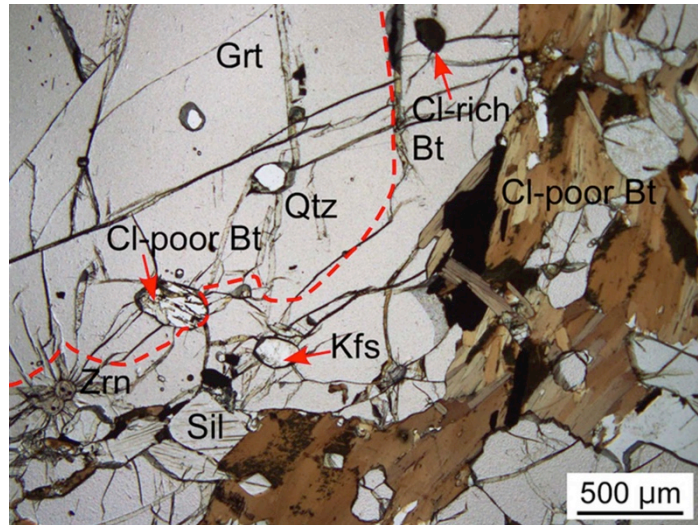
# Element diffusion in garnet

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## Granulites, Sør Rondane, East Antarctica

Pelitic gneiss: Grt-Sill-Bt

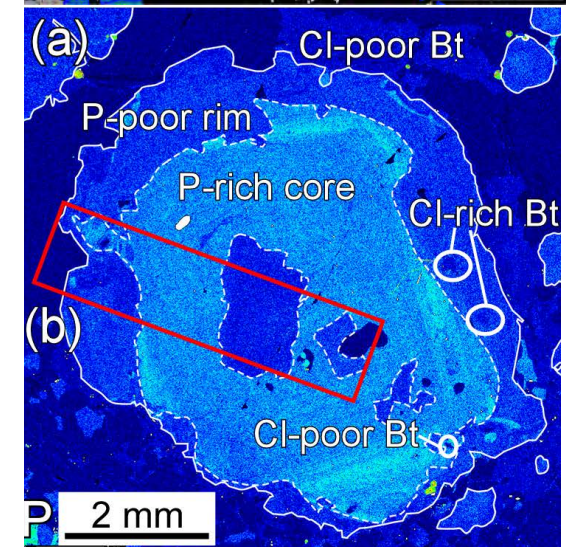
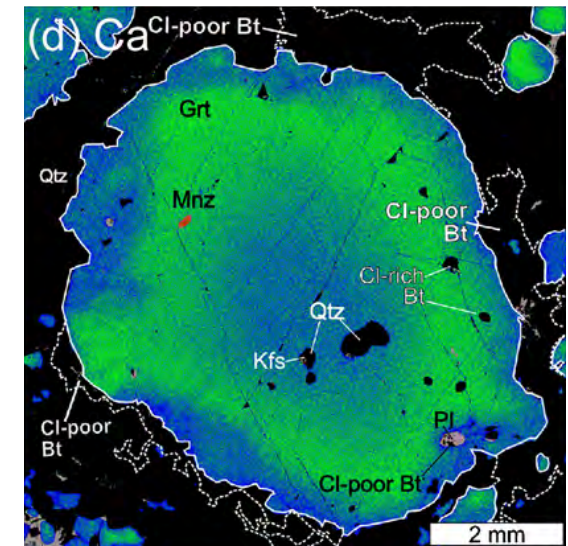


- ✓ Relaxed major element zoning
- ✓ Step-wise P zoning

Peak conditions 850°C, 1.1 GPa

Garnet rim growth 800°C, 0.8 GPa *Higashino et al. 2013*

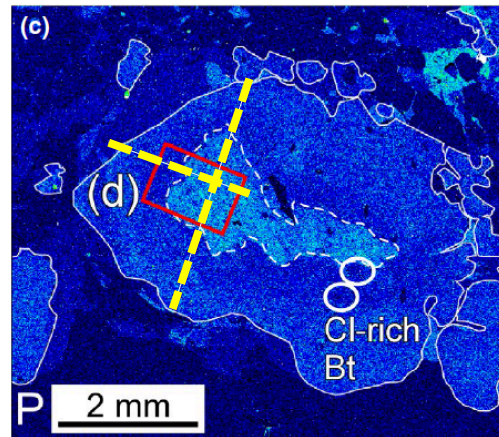
- ✓ Veins indicating penetration of Cl-rich brines
- ✓ Fluid infiltration at 800°C



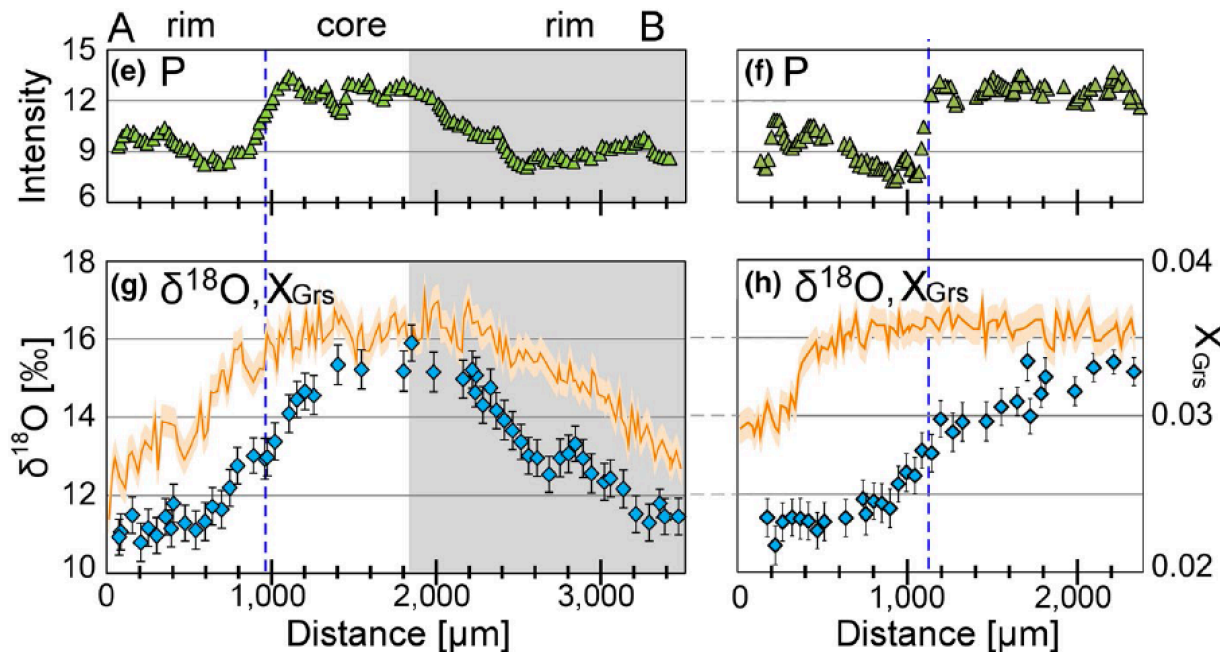
# Oxygen diffusion in garnet

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- ✓ Relaxed major element zoning
- ✓ Step-wise P zoning
- ✓ Symmetric  $^{18}\text{O}/^{16}\text{O}$  zoning



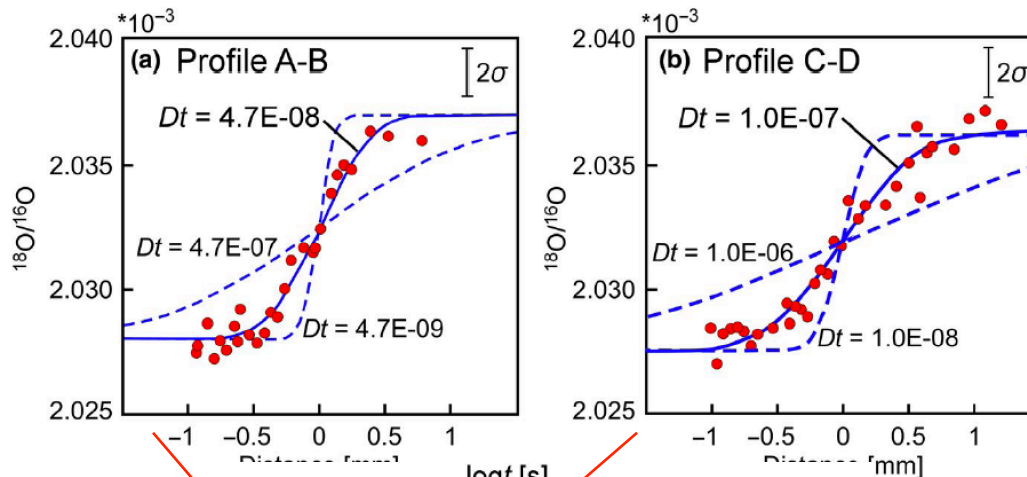
## Oxygen isotopes

- Preserved  $\delta^{18}\text{O}$  zoning at 800°C!
- Core - rim zoning  $\delta^{18}\text{O} = 15.5$  to  $11.0$  ‰
- Input of external fluids

# Oxygen speedometry

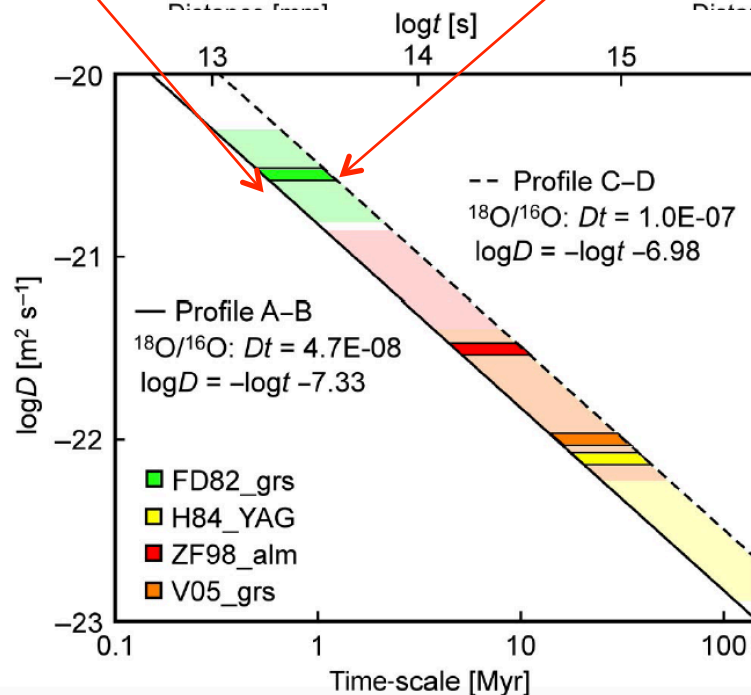
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Data fitted to Fick's second law  
(Crank, 1975)

$$C = C_{\min} + 0.5 (C_{\max} - C_{\min}) \operatorname{erfc} [x / (4Dt)^{0.5}]$$



Duration of high T  
metamorphism (800°C)  
between 0.5-40 Ma  
(maximum)

Depends on diffusion coefficient D

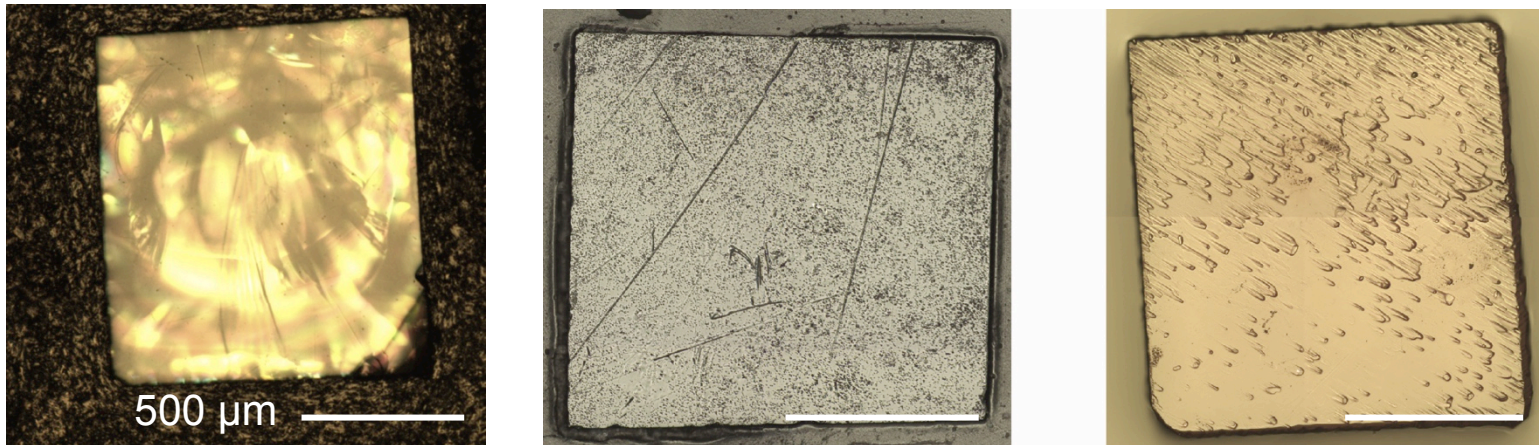


## Oxygen diffusion in garnet

Experimental study using YAG ( $\text{Y}_3\text{Al}_5\text{O}_{12}$ ) and pyrope ( $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ ) garnet in a  $^{18}\text{O}$  enriched matrix

1. Gas furnace with Ar flux 1500–1600°C
2. Piston cylinder 900–1600°C, 1.0–2.5 GPa, dry and wet

### *Recovered garnet cubes after experiment*



*Scicchitano, Jollands et al. 2021*

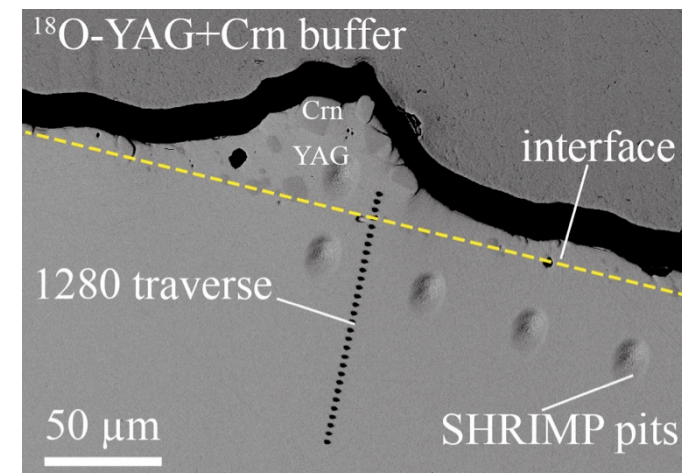
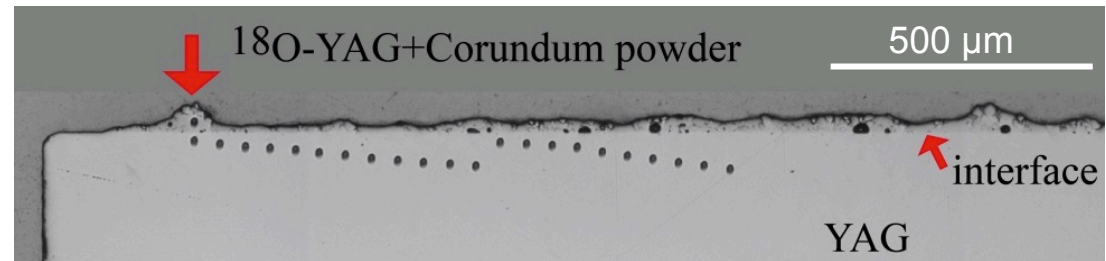
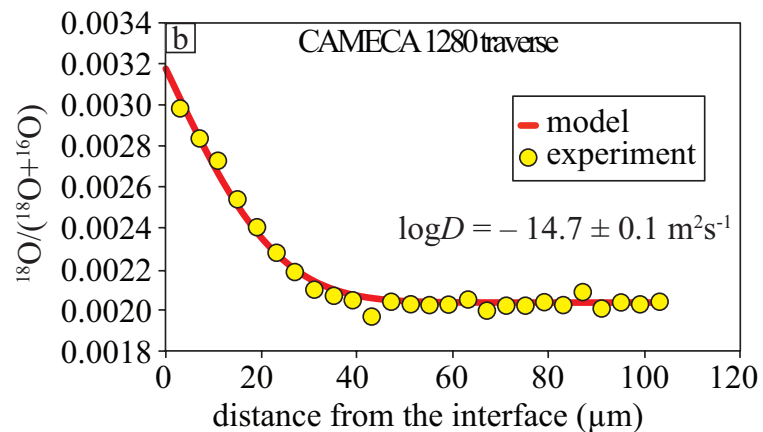
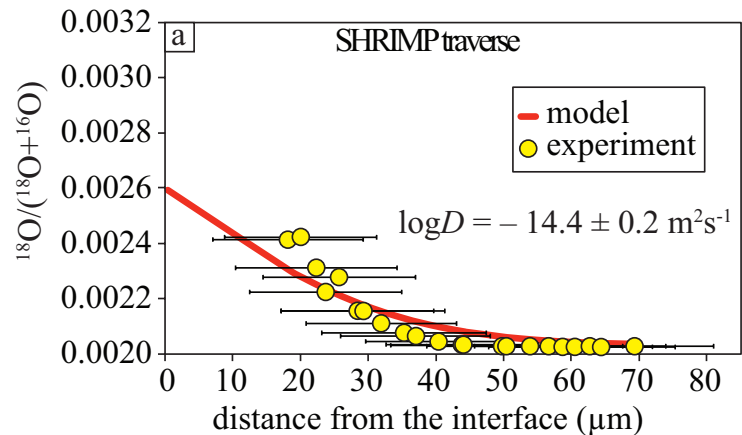
# Oxygen diffusion in garnet

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## Measurement of $^{18}\text{O}/^{16}\text{O}$ with three methods

### 1. SIMS (SHRIMP and Cameca 1280) traverse – resolution 15–3 microns



*Scicchitano, Jollands et al. 2021*

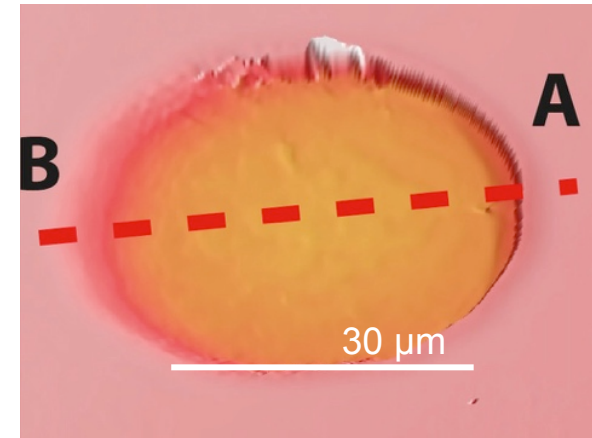
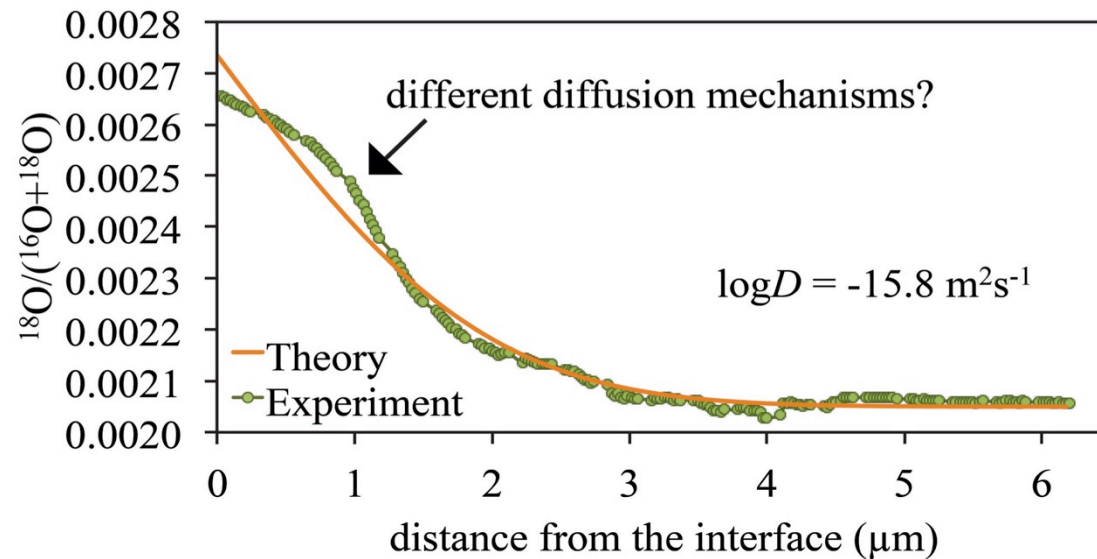
# Oxygen diffusion in garnet

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## Measurement of $^{18}\text{O}/^{16}\text{O}$ with three methods

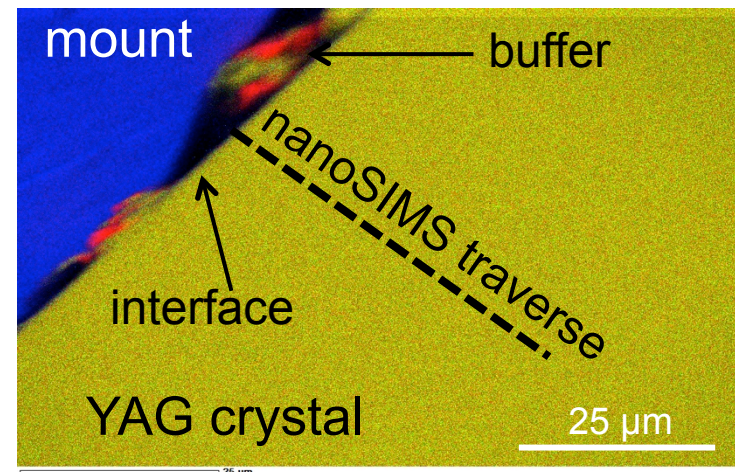
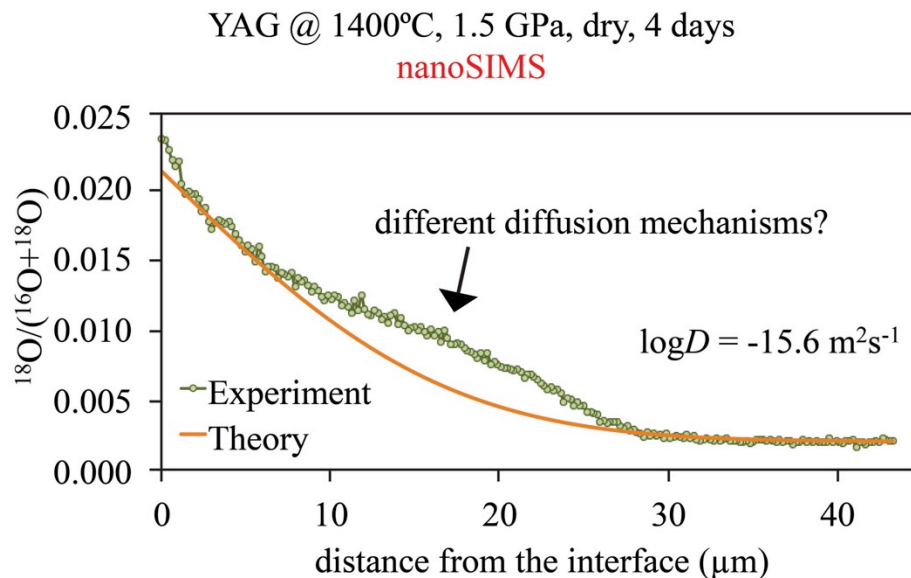
### 2. SIMS (SHRIMP and Cameca 1280) depth profiling – resolution $\sim 20$ nm



*Scicchitano, Jollands et al. 2021*

## Measurement of $^{18}\text{O}/^{16}\text{O}$ with three methods

### 3. nanoSIMS traverses – spot resolution 200 nm



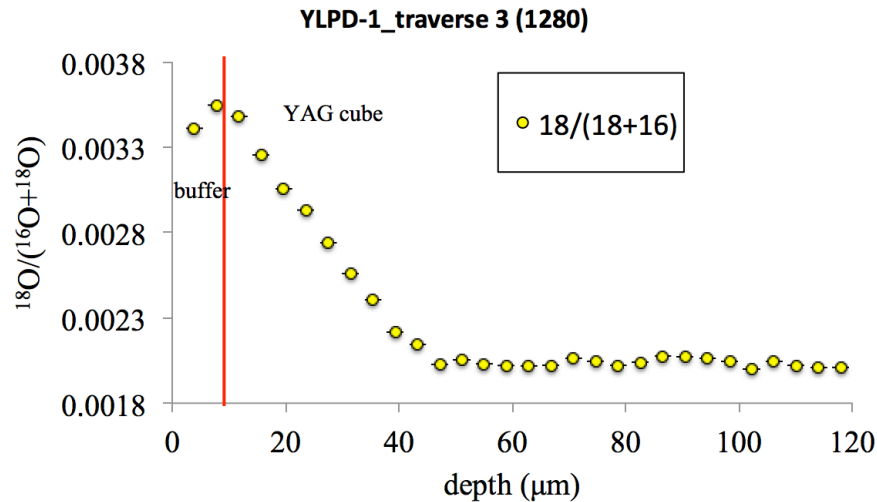
Scicchitano, Jollands et al. 2021



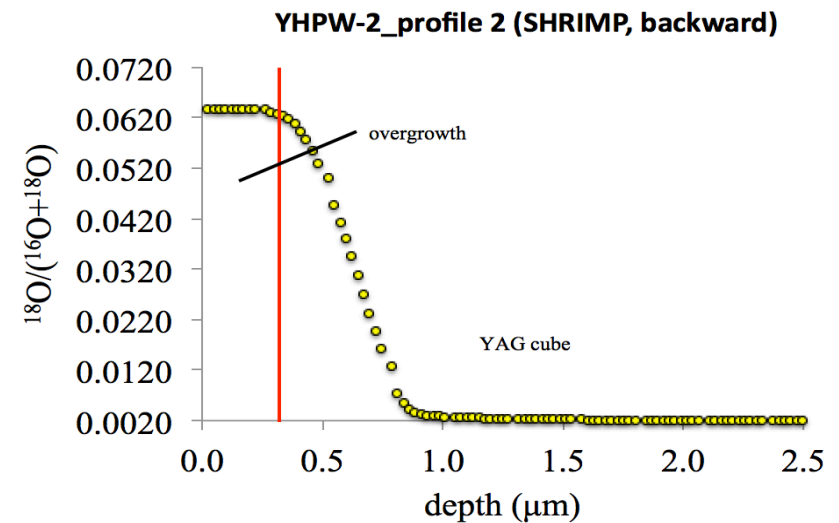
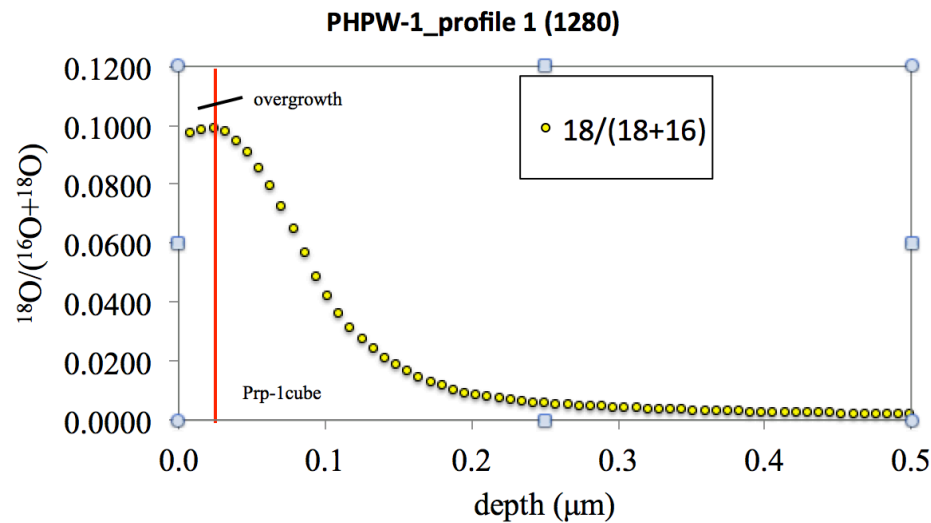
# Different types of profile

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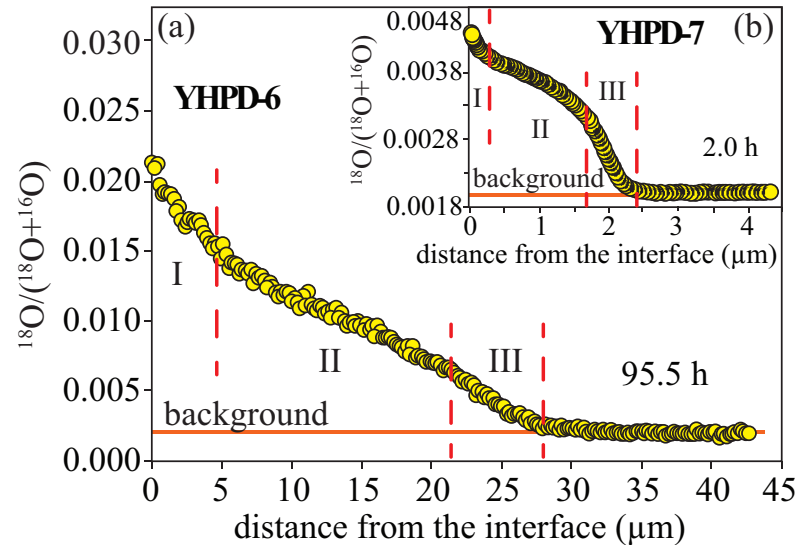
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## 1. Error function profiles In LP-HT runs, in HP-LT “wet” runs

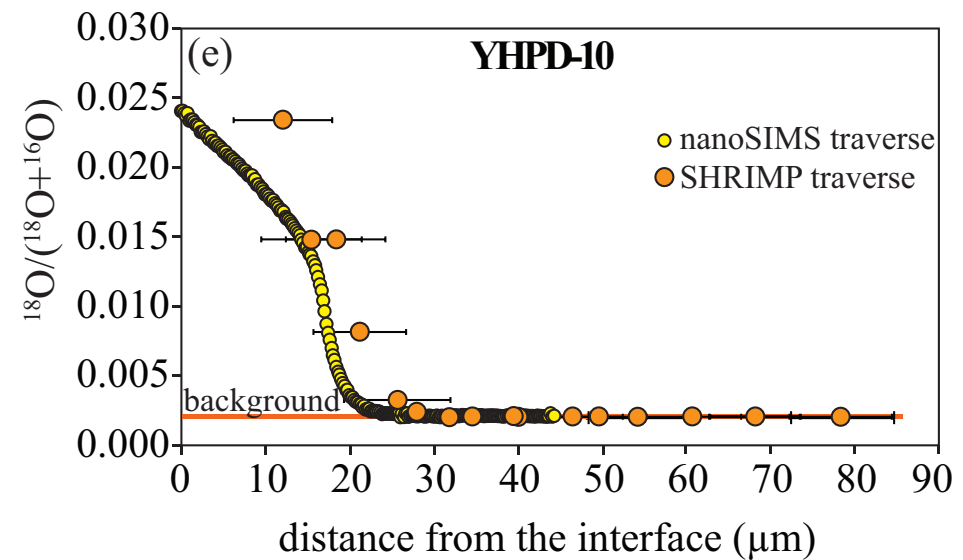
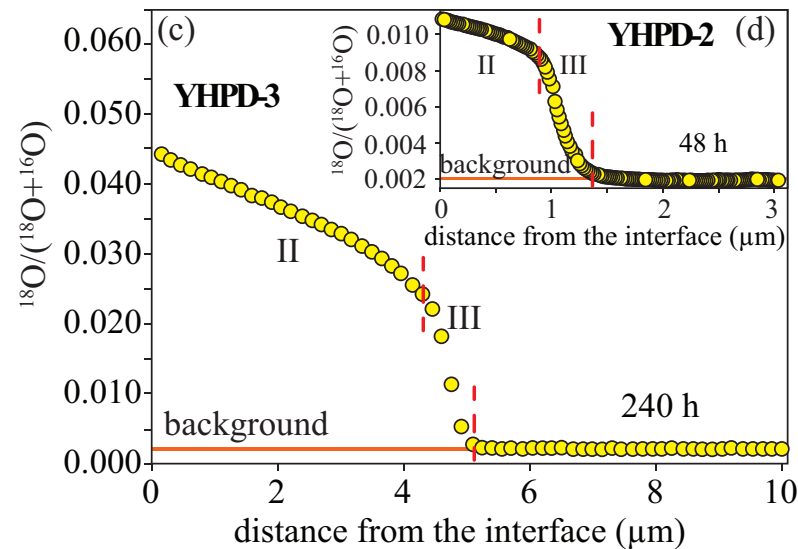


# Different types of profile



## 2. Stepped profiles

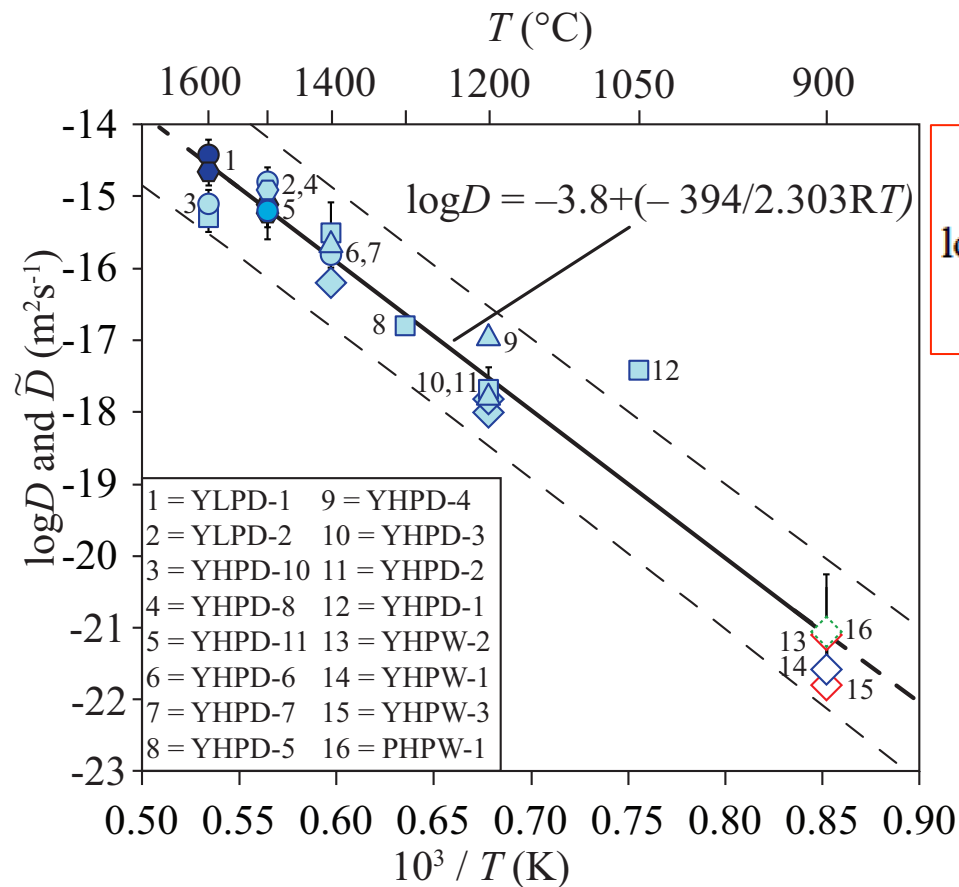
In all HP “dry” experiments, 1 excl.



# Results, first approximation

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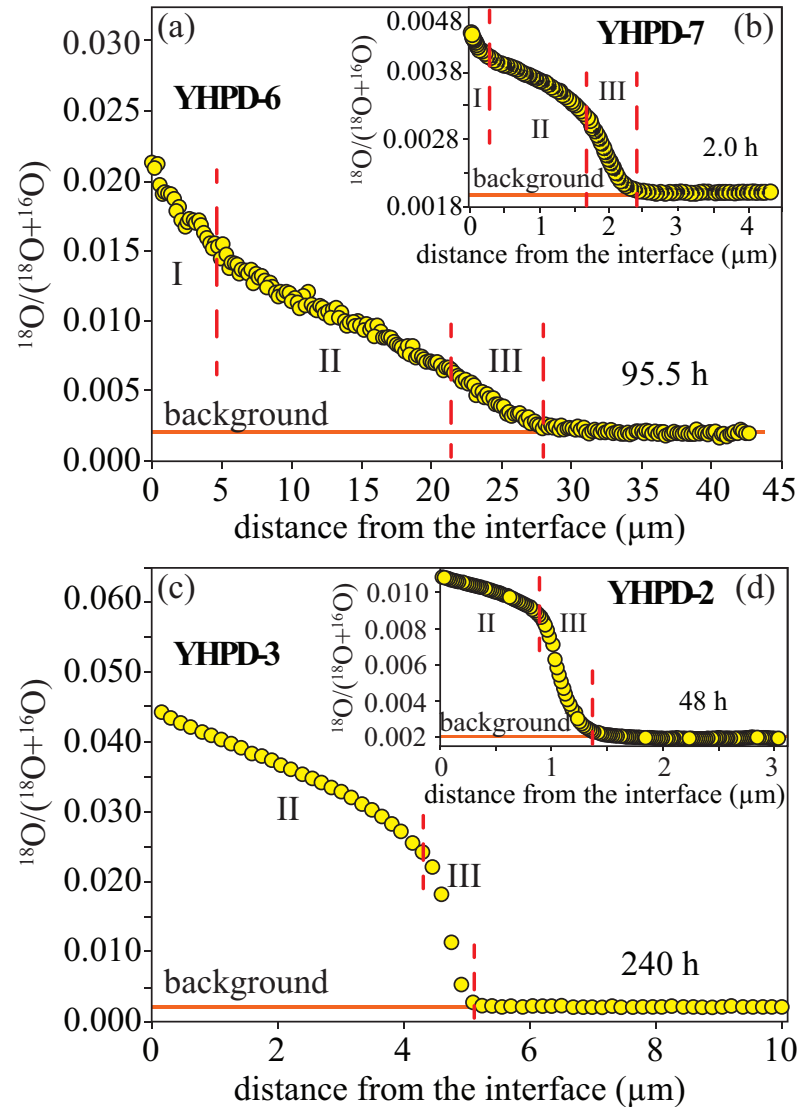
$$\log \tilde{D} \text{ m}^2\text{s}^{-1} = -3.8 (\pm 0.7) + \left( \frac{-394 (\pm 19) \text{ kJmol}^{-1}}{2.303RT} \right)$$

**Full dataset (partial fitting)**

No composition effect  
See no  $\text{H}_2\text{O}$  effect  
Cannot estimate P effect

- |                          |                   |                   |
|--------------------------|-------------------|-------------------|
| ○ SHRIMP traverse        | ● YAG dry 1.5 GPa | ◇ YAG wet 1.0 GPa |
| □ NanoSIMS               | ● YAG dry 2.5 GPa | ◇ YAG wet 1.5 GPa |
| △ SHRIMP depth profiling | ● YAG dry 1 atm   | ◇ Prp wet 1.0 GPa |
| ◇ 1280 traverse          |                   |                   |
| ◇ 1280 depth profiling   |                   |                   |

# Results: stepped profiles



Overgrowth + diffusion is unlikely

Not concentration-independent diffusion on a single site

Two diffusion mechanisms likely  
fast pathway + slow pathway

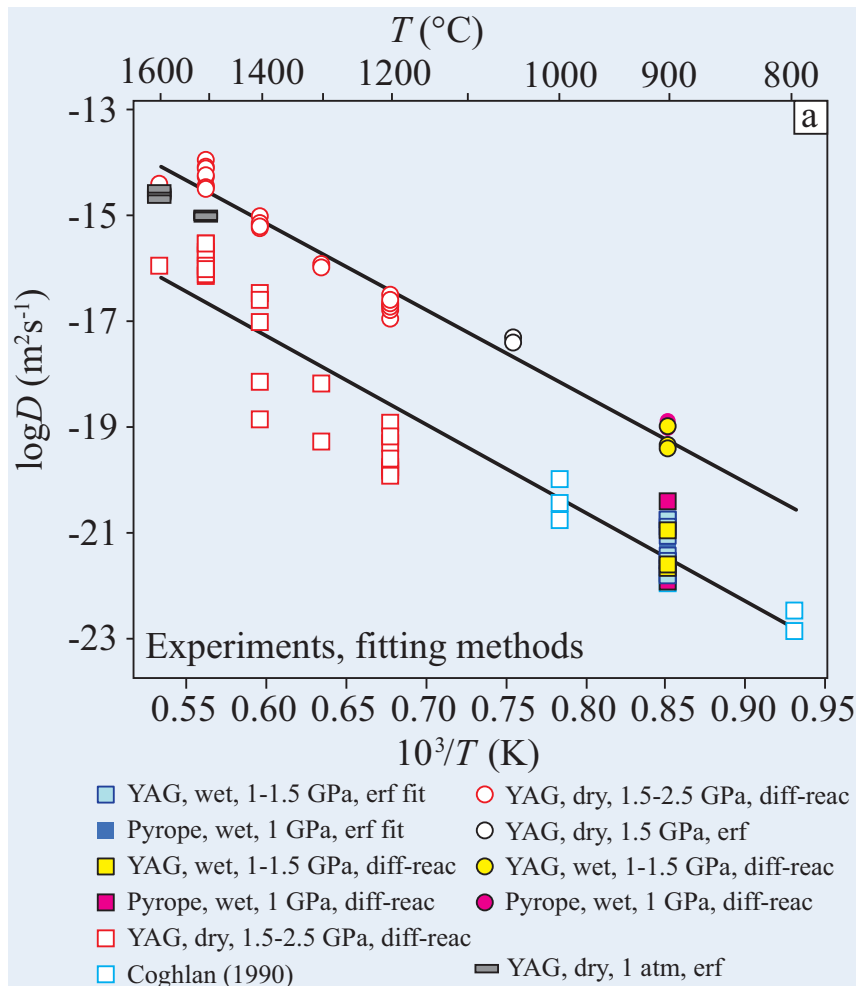
Diffusion on vacancy + interstitial



# Results: two mechanisms

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Two diffusion mechanisms likely  
fast pathway + slow pathway

Diffusion on vacancy + interstitial

slow (most profiles)

$$\log D \text{ m}^2 \text{ s}^{-1} = -7.2 (\pm 1.3) + \left( \frac{-321 (\pm 32) \text{ kJ mol}^{-1}}{2.303RT} \right)$$

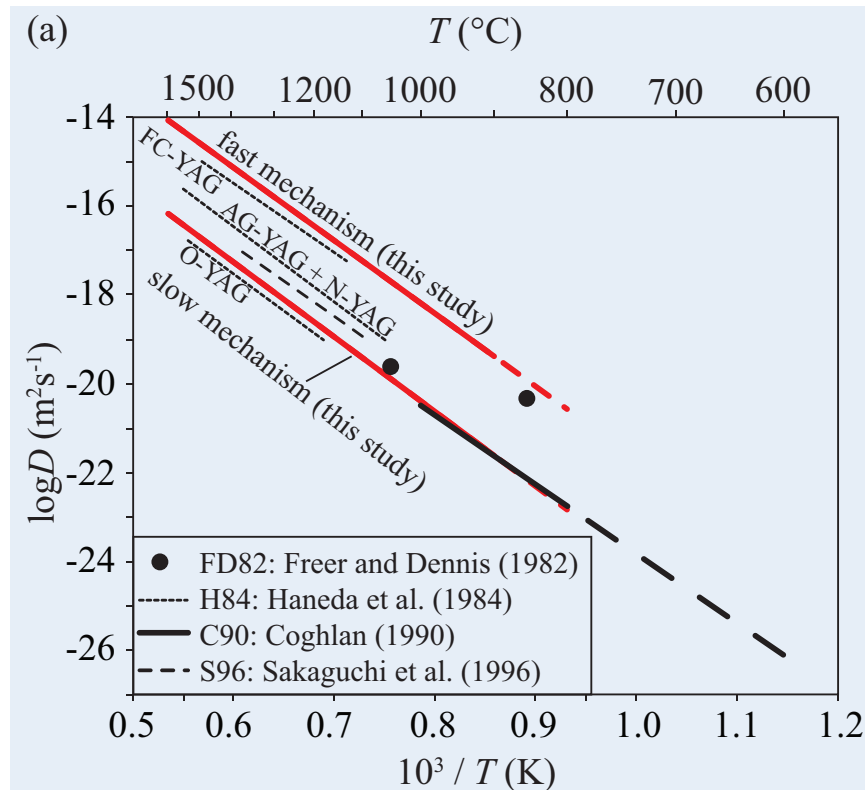
fast

$$\log D \text{ m}^2 \text{ s}^{-1} = -5.4 (\pm 0.7) + \left( \frac{-312 (\pm 20) \text{ kJ mol}^{-1}}{2.303RT} \right)$$

# Comparison to previous studies

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**All previous experimental/  
natural/computational results  
fall within the range**

slow (most profiles)

$$\log D \text{ m}^2 \text{ s}^{-1} = -7.2 (\pm 1.3) + \left( \frac{-321 (\pm 32) \text{ kJ mol}^{-1}}{2.303RT} \right)$$

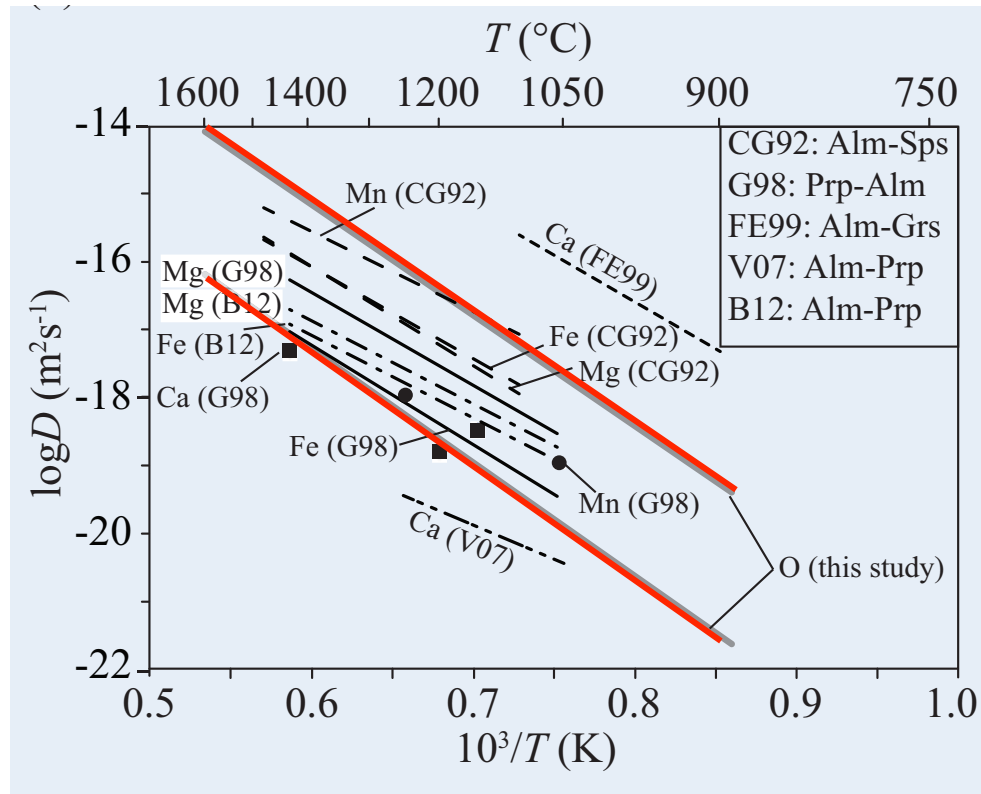
fast

$$\log D \text{ m}^2 \text{ s}^{-1} = -5.4 (\pm 0.7) + \left( \frac{-312 (\pm 20) \text{ kJ mol}^{-1}}{2.303RT} \right)$$

# Comparison to cations

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**Activation energy O > major divalent cations**

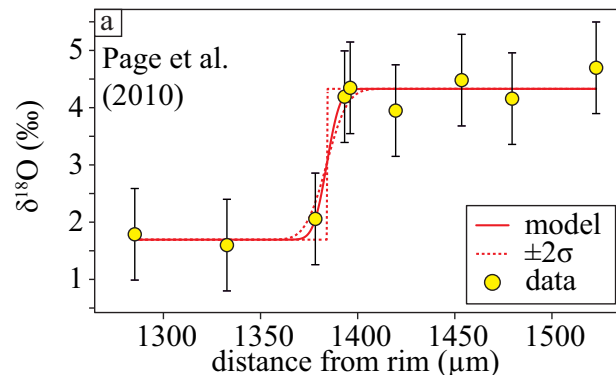
⇒ Extrapolation to  $T < 850^\circ\text{C}$  would result in slower O diffusivity relative to major cations

Data are not normalized to a fixed oxygen fugacity.  
CG92: Chakraborty and Ganguly (1992); G98: Ganguly et al. (1998); FE99: Freer and Edwards (1999); V07: Vielzeuf et al. (2007); B12: Borinski et al. (2012).

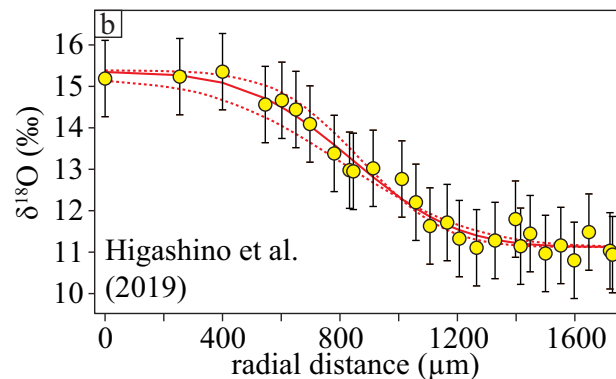
**ADVANTAGE for O**

- is not a multicomponent system
- is not dependent on composition

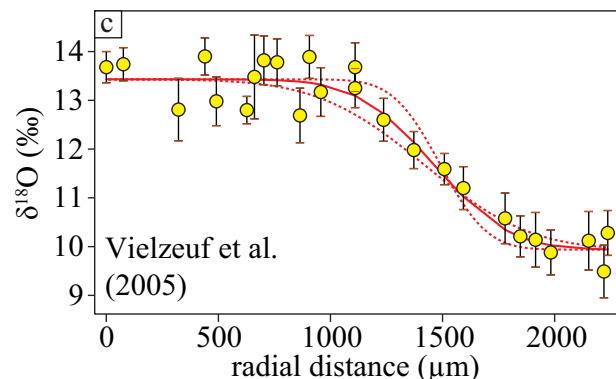
# Natural data



Regional metamorphism:  
Best fit =  $\log Dt$  ( $\text{m}^2$ ) of  $-10.8$   
 $750^\circ\text{C}$ ,  $\log D$  ( $\text{m}^2\text{s}^{-1}$ ) of  $-24.2 \pm 0.5$  = max 1.6 M.y.



Regional metamorphism:  
Fit to spherical geometry =  $\log Dt$  ( $\text{m}^2$ ) =  $-7.4 \pm 0.2$   
 $800^\circ\text{C}$  regression for the slow diffusivity gives  
 $\log D$  ( $\text{m}^2\text{s}^{-1}$ ) of  $-22.9 \pm 0.5$  ( $2\sigma$ ) = times of 30-320 M.y.



Crustal melting:  
Fit to spherical geometry =  $\log Dt$  ( $\text{m}^2$ ) of  $-7.5 \pm 0.3$   
At  $800^\circ\text{C}$  regression for the slow diffusivity gives  
 $\log D$  ( $\text{m}^2\text{s}^{-1}$ ) of  $-22.2 \pm 0.4$ , = 4 to 50 M.y.



- > Garnet O zoning can be retained at high T metamorphic conditions
- > Diffusion of O in garnet is regulated by two mechanisms with the slow one being the most relevant to natural settings
- > Diffusion of O is likely to be slower than that of major cations
- > The diffusivity of O may be a “goldilocks” case in between major and trace elements

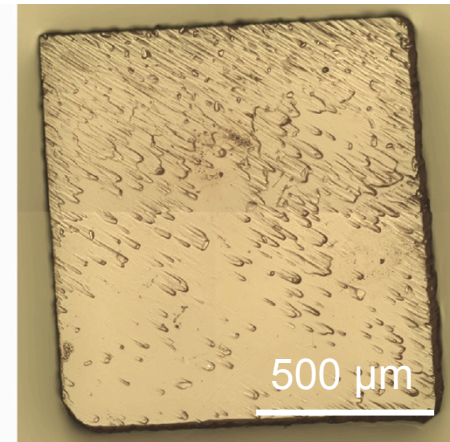
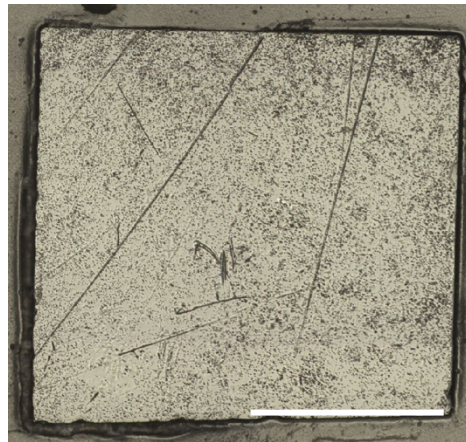
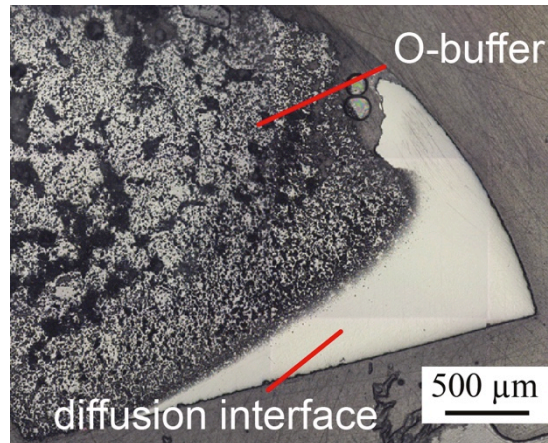
# How robust is oxygen in garnet?

## Oxygen diffusion in garnet

Experimental study using YAG ( $\text{Y}_3\text{Al}_5\text{O}_{12}$ ) and pyrope ( $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ ) garnet in a  $^{18}\text{O}$  enriched matrix

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### *Recovered garnet cubes after experiment*



*M R Scicchitano, PhD thesis 2017*

# Pervasive fluid flow in subducted crust

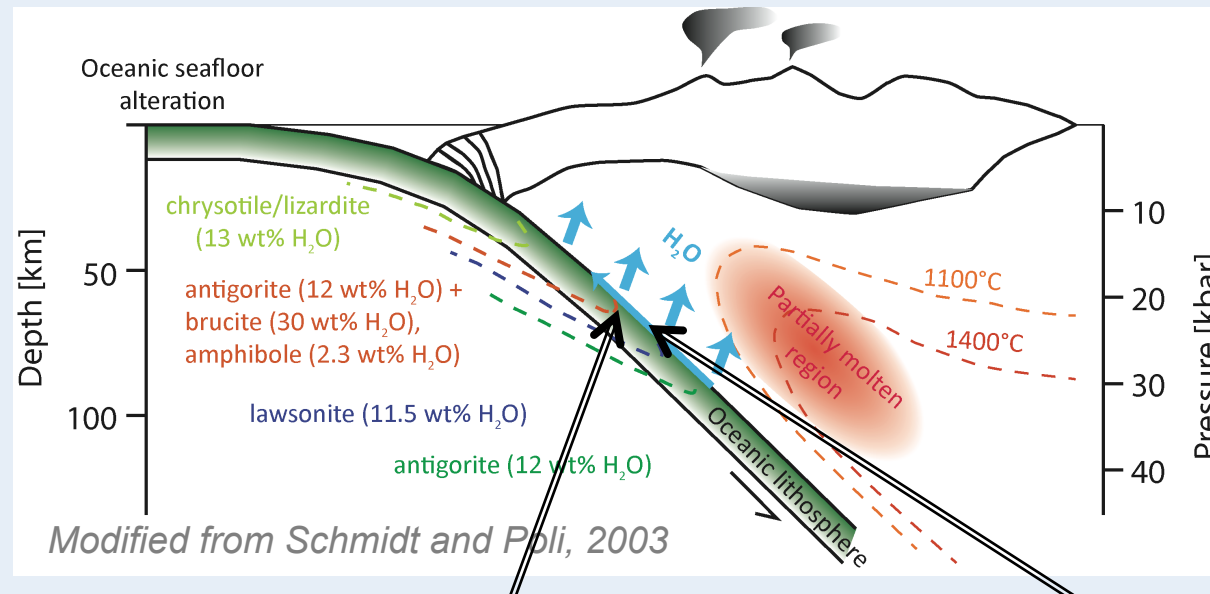
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Daniela RUBATTO, Thomas BOVAY, Pierre LANARI  
University of Bern, Institute of Geological Sciences, Switzerland

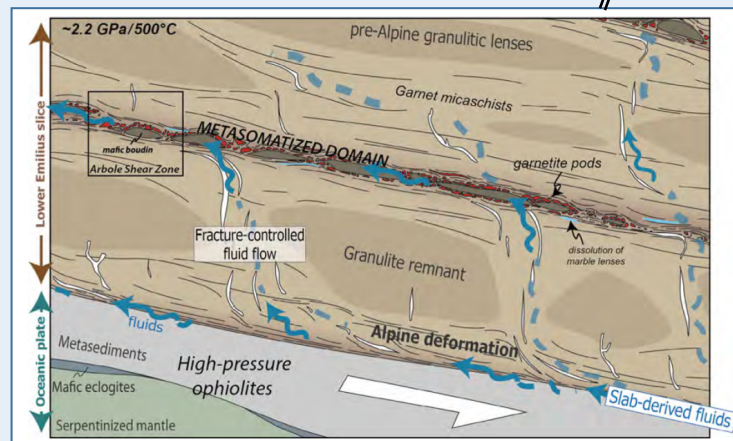
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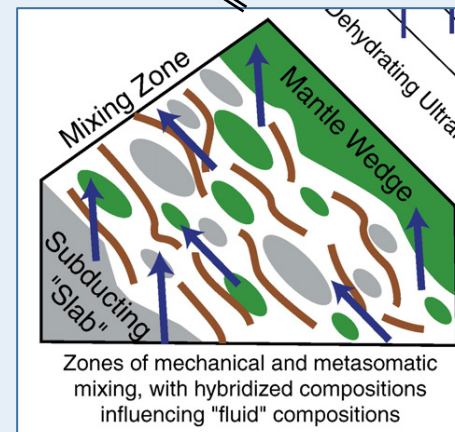
SWISS NATIONAL SCIENCE FOUNDATION



Fluid  
production  
is known



Angiboust et al., 2016



Bebout, 2007

Fluid flow  
dynamics??

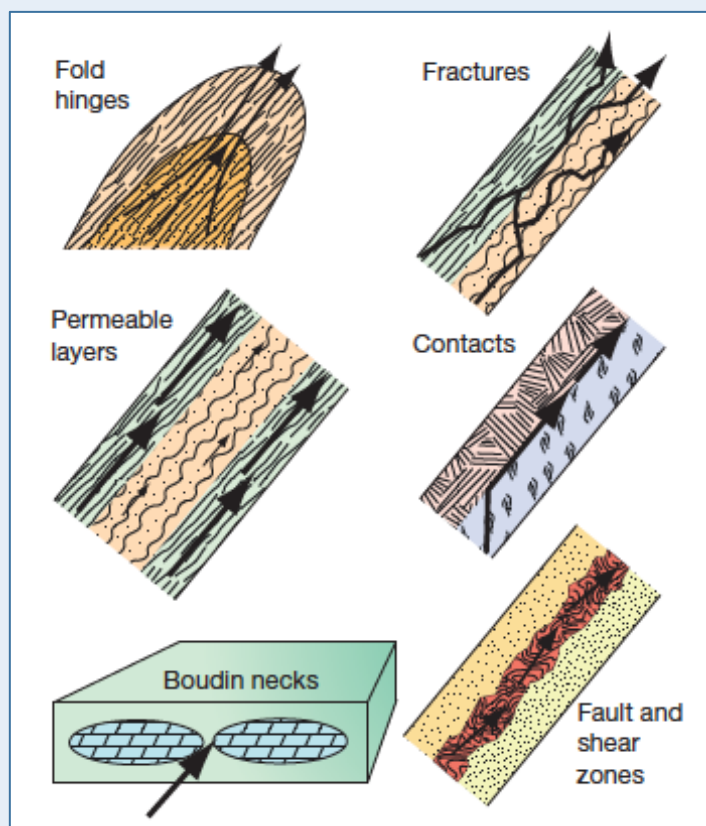
# Pervasive fluid flow in subducted crust

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Daniela RUBATTO, Thomas BOVAY, Pierre LANARI  
*University of Bern, Institute of Geological Sciences, Switzerland*

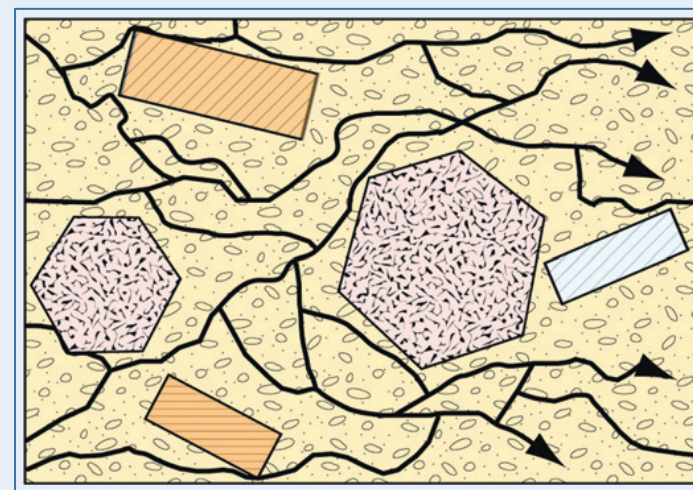
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- **channelized**  
(veins, shear zones, contacts)



*Ague, 2014*

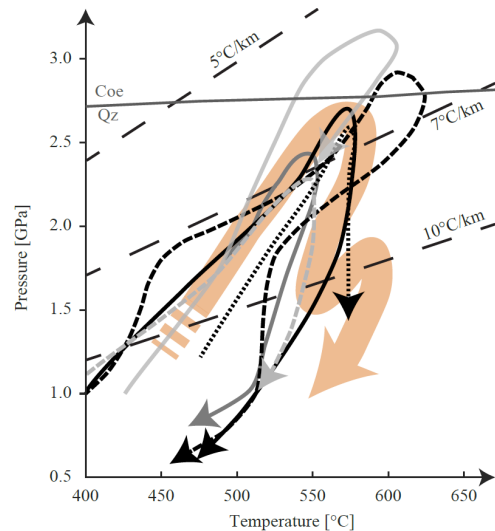
- **pervasive**  
(intergranular porosity)



*Ague, 2014*

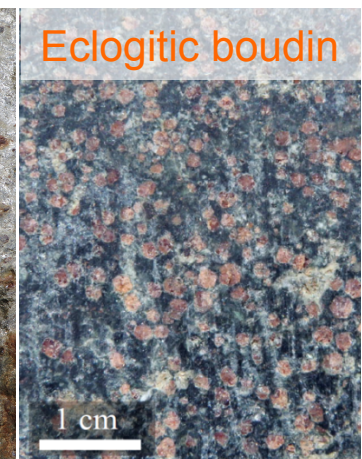
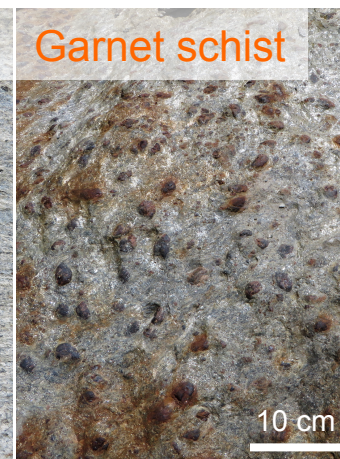
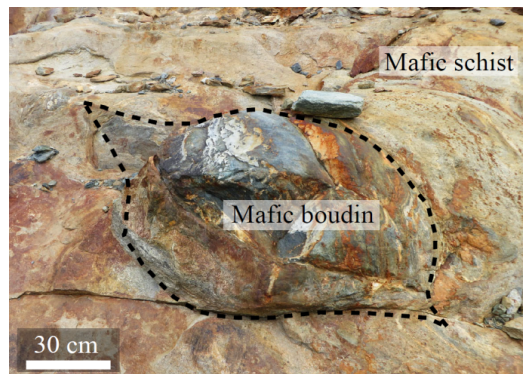
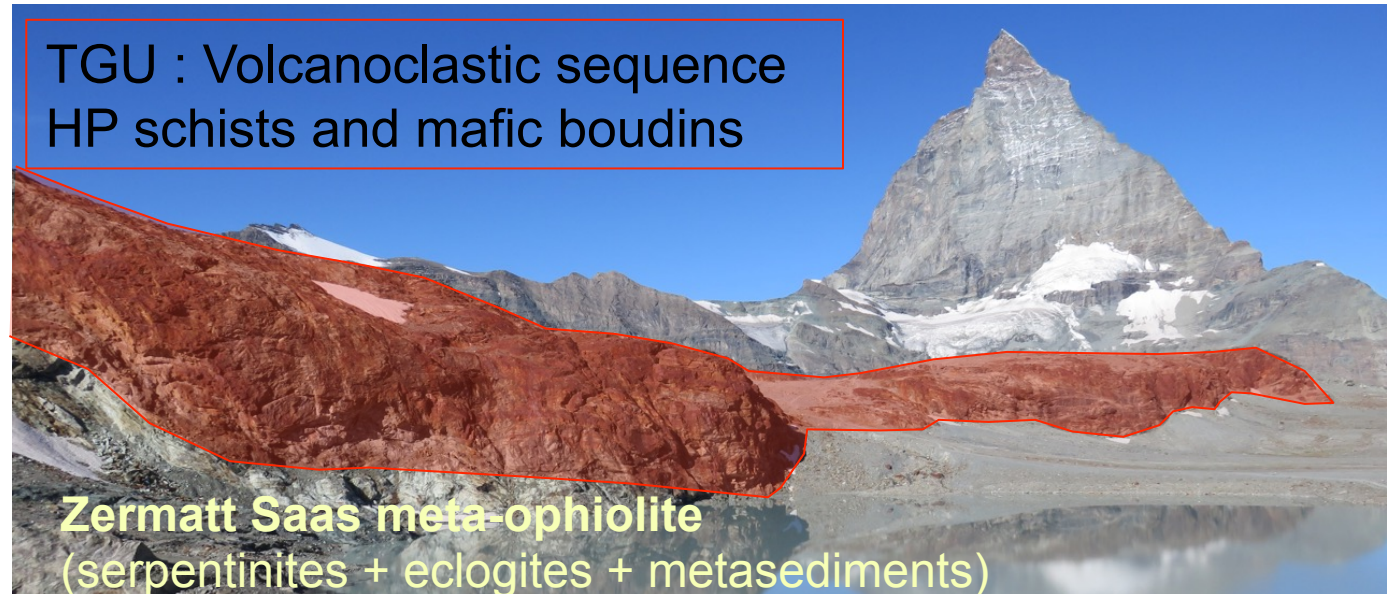


# Theodul Glacier Unit (TGU)



Zermatt Saas Zone	Continental outliers
— Bucher et al. 2005	— Angiboust et al. 2017
— Angiboust et al. 2009	— Fassmer et al. 2016
— Groppo et al. 2009	<b>TGU</b>
— Reinecke et al. 1998	— this study

TGU : Volcanoclastic sequence  
HP schists and mafic boudins





# Garnet textures

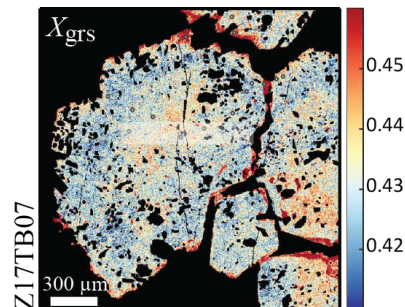
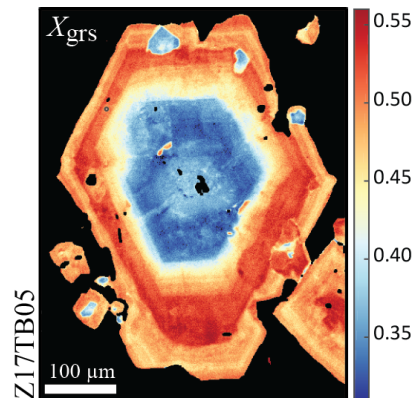
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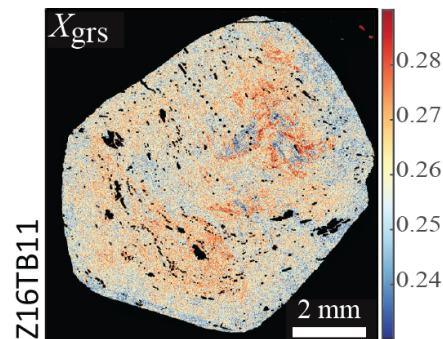
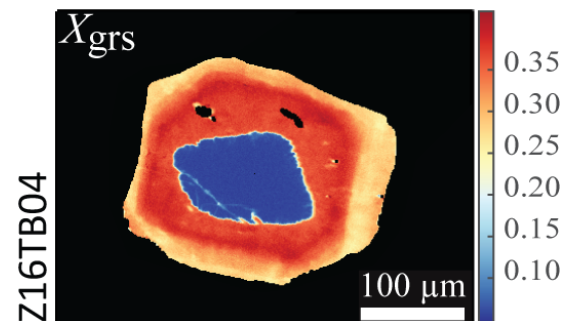
Continuous zoning = continuous growth in a fractionating reactive bulk

Discontinuous zoning = resorption and regrowth induced by fluids

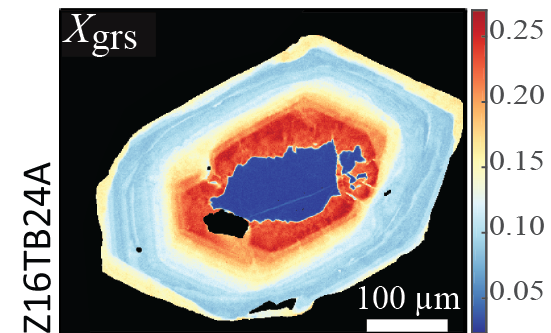
Mafic boudins



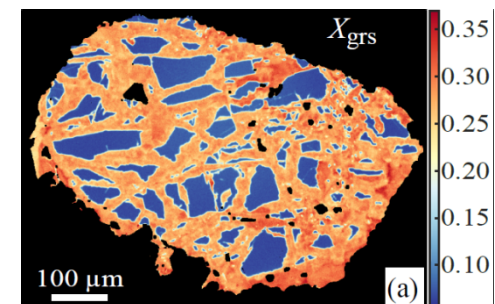
Mafic schist



Chloritoid schist



Garnet schist



# Garnet age

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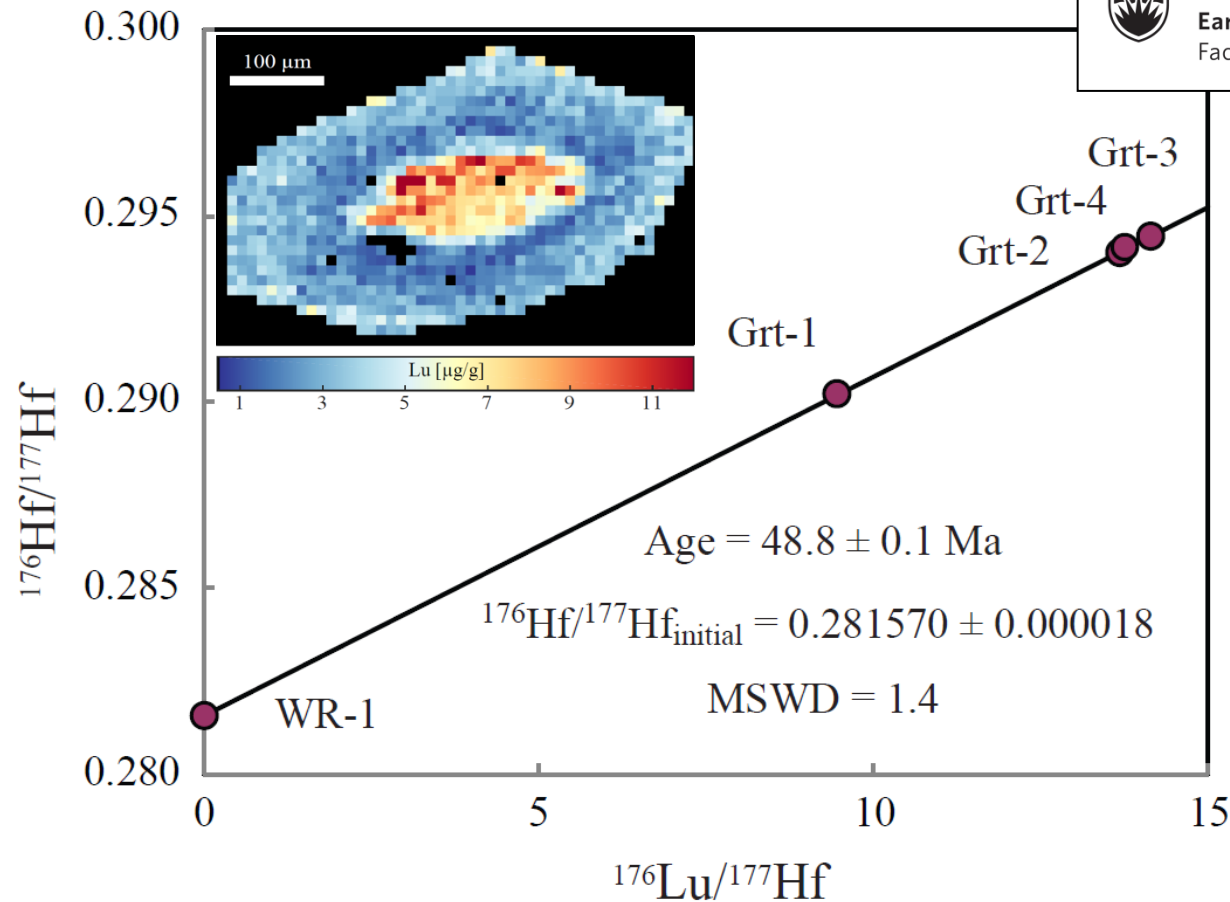


THE UNIVERSITY OF BRITISH COLUMBIA

Earth, Ocean and Atmospheric Sciences

Faculty of Science

with Matthijs Smit

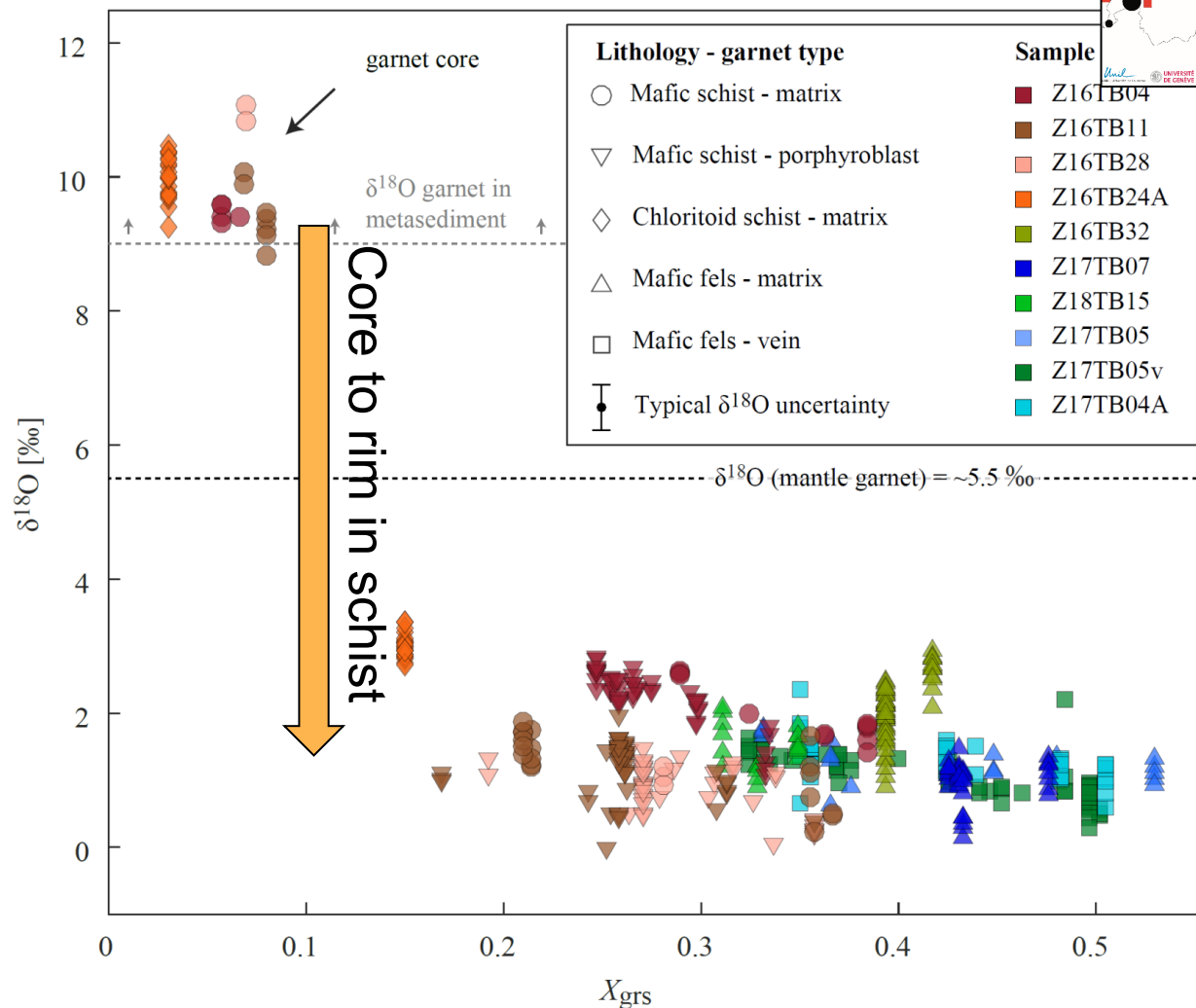


Monometamorphic  
Alpine evolution

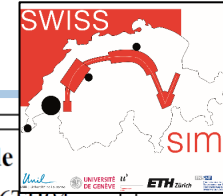
# Garnet oxygen composition

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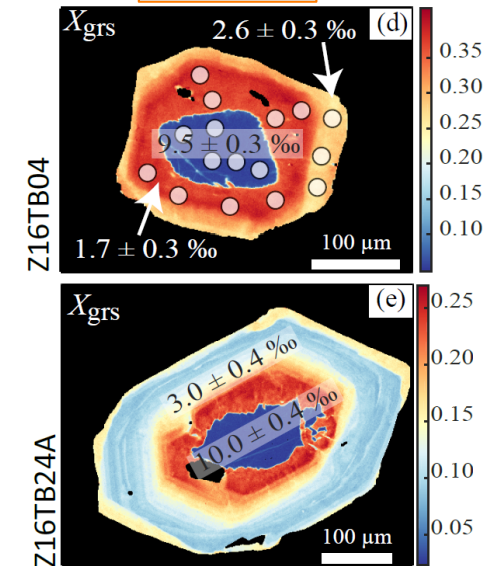
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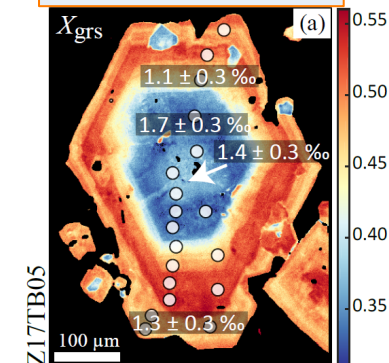
=> Open system behavior with incoming external fluids at high pressure over the entire unit



## Schists



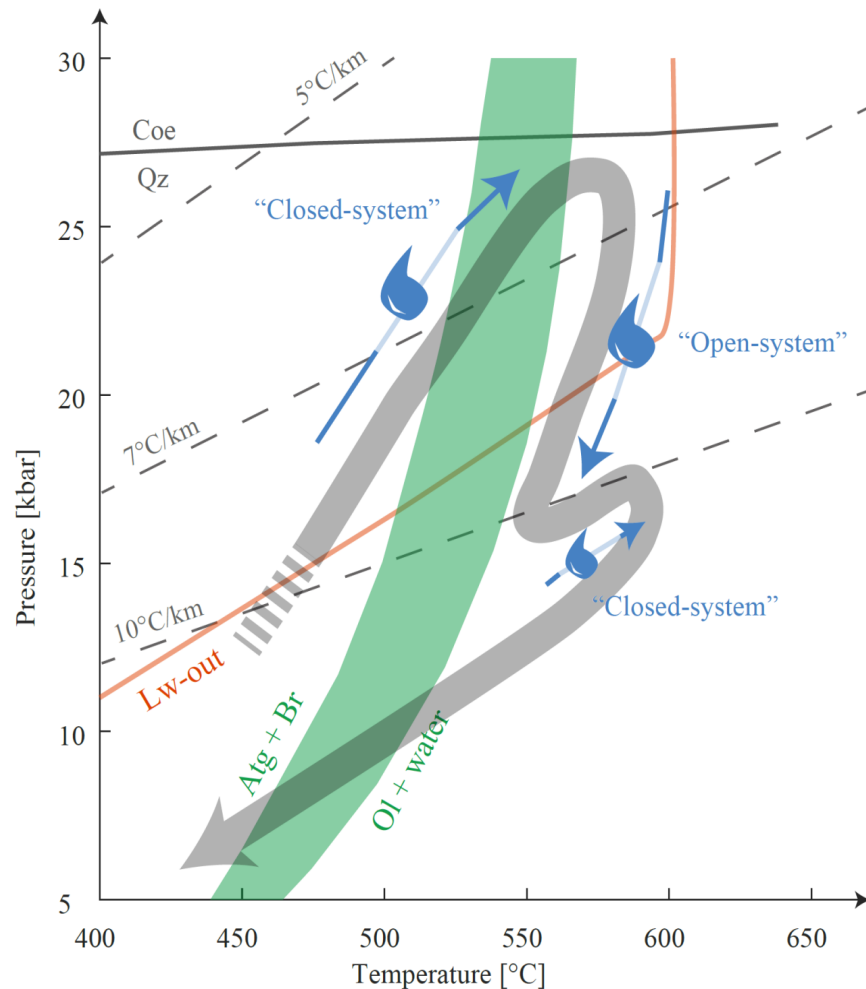
## Mafic boudins



# Fluid modelling

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Garnet  $\delta^{18}\text{O}$  1 ‰ at 580 °C  
 $\Rightarrow \text{H}_2\text{O} \leq 3.5 \text{ ‰}$

Lawsonite breakdown in  
mafic rocks  $\sim 2 \text{ wt\% H}_2\text{O}$

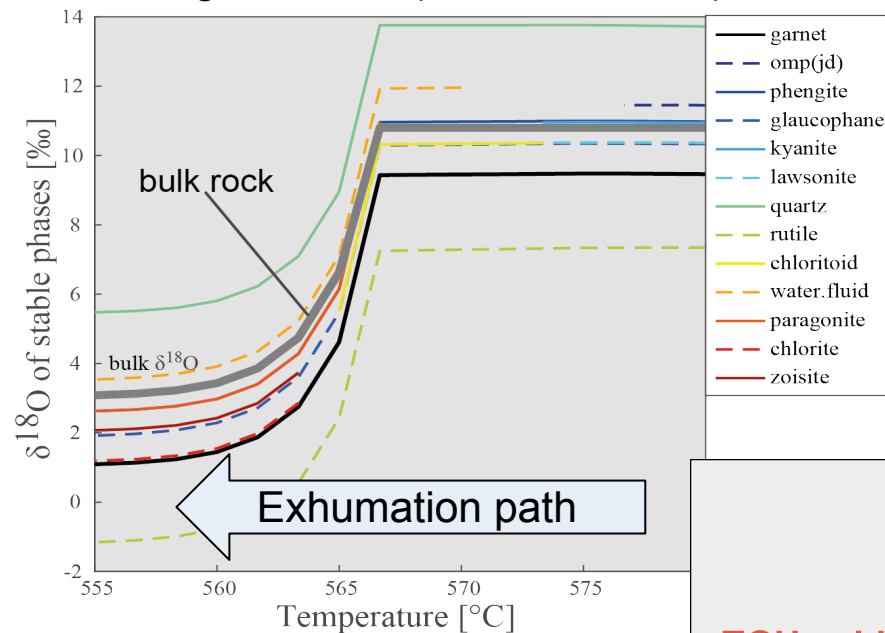
Brucite + antigorite reaction in  
serpentinites = 3–7 wt%  $\text{H}_2\text{O}$

# Fluid modelling

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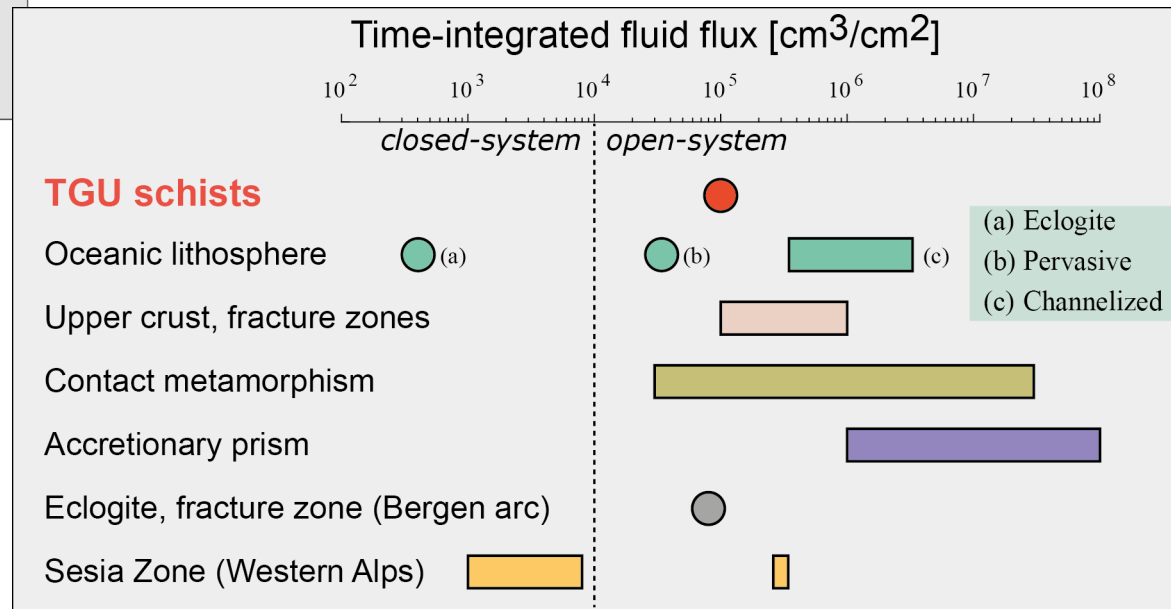
Using P-TLOOP (Vho et al. 2020)



Garnet  $\delta^{18}\text{O}$  1 ‰ at 580 °C  
 $\Rightarrow \text{H}_2\text{O} \leq 3.5 \text{ ‰}$

time integrated

Bulk rock  $\delta^{18}\text{O}$  shift from 11 to 3 ‰  
 $\Rightarrow \text{H}_2\text{O flux } 1.1 \times 10^5 \text{ cm}^3/\text{cm}^2$



Modified from Zack and John, 2007

Data from Zack and John 2007, Ferry and Gerdes 1998, Ague 2003, Philippot and Rumble 2000, Konrad-Schmolke et al. 2011, Vho et al. 2020

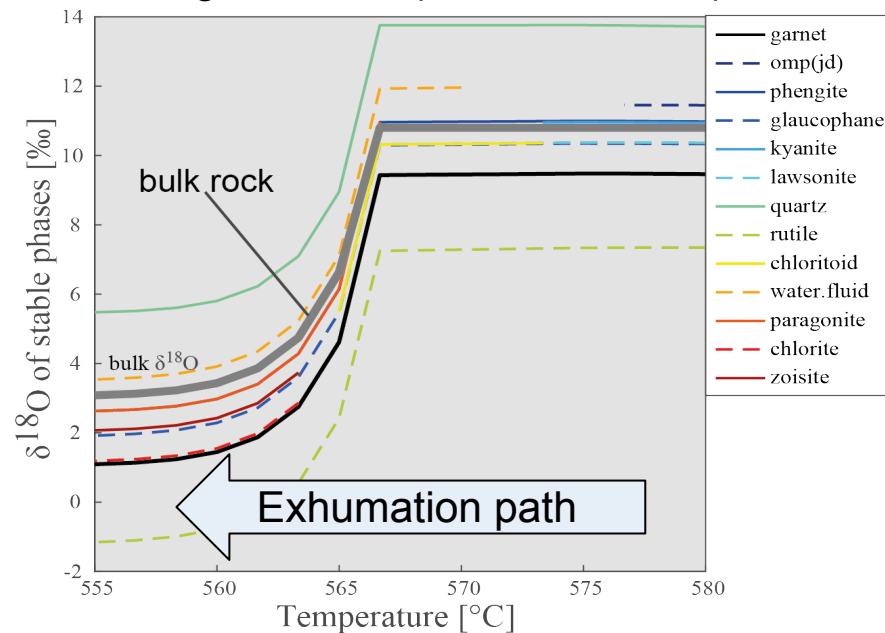


# Fluid modelling

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Using P-TLOOP (Vho et al. 2020)



Garnet  $\delta^{18}\text{O}$  1 ‰ at 580 °C  
 $\Rightarrow \text{H}_2\text{O} \leq 3.5 \text{ ‰}$

time integrated

Bulk rock  $\delta^{18}\text{O}$  shift from 11 to 3 ‰  
 $\Rightarrow \text{H}_2\text{O flux } 1.1 \times 10^5 \text{ cm}^3/\text{cm}^2$

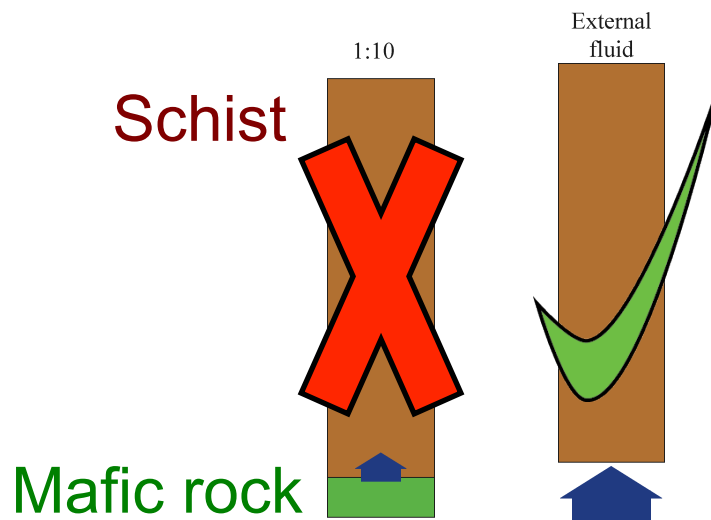


Lawsonite breakdown in mafic rocks  
**cannot** produce

- water with low  $\delta^{18}\text{O} \leq 3.5 \text{ ‰}$
- such large amount of water



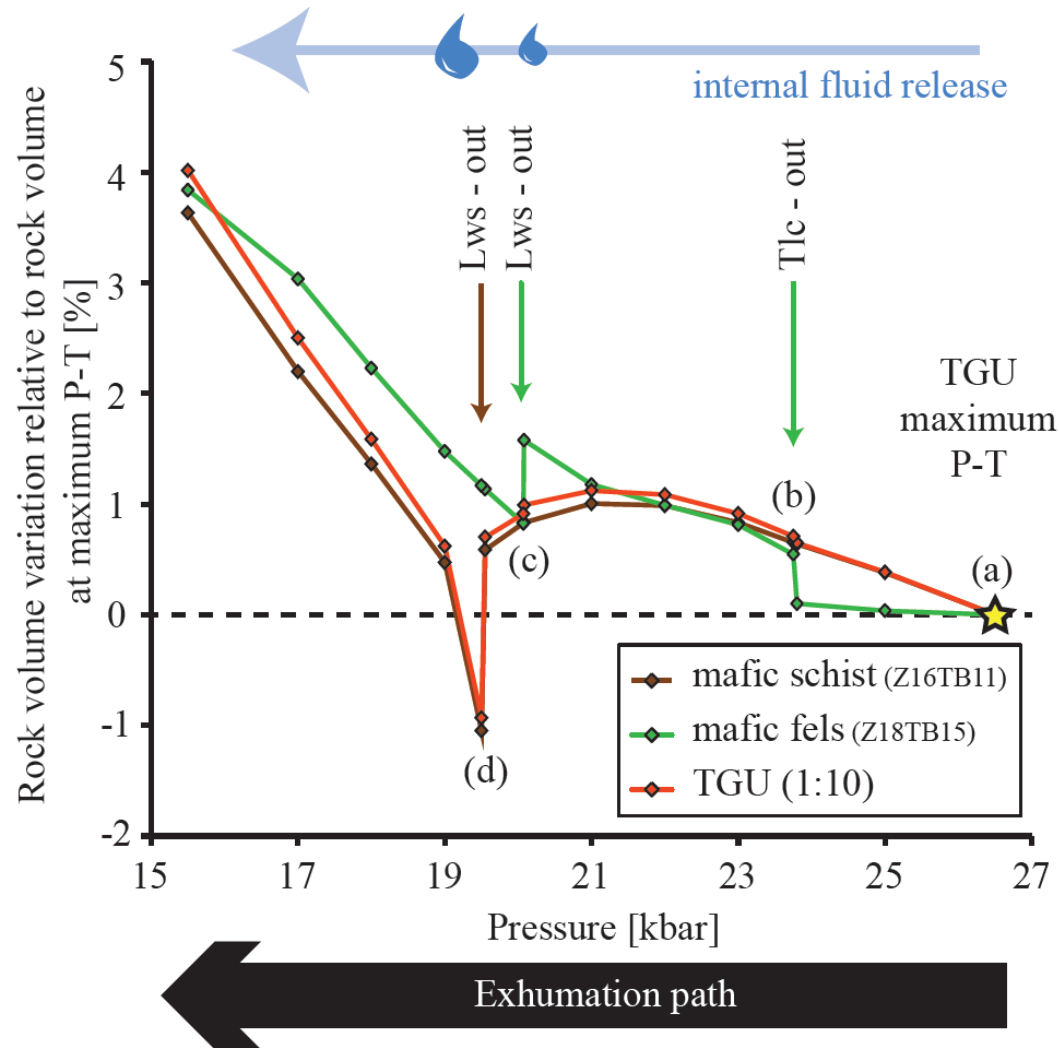
Brucite+ antigorite reaction in  
 serpentinites = 3–7 wt% H<sub>2</sub>O



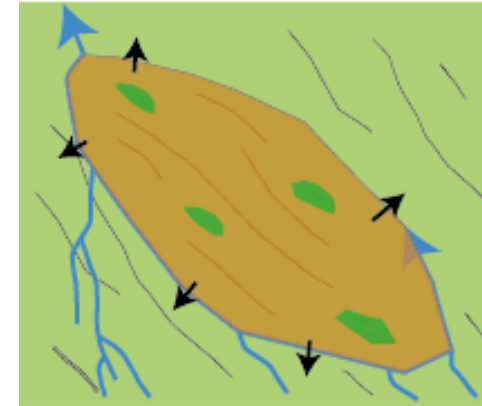
# Volume variation modelling

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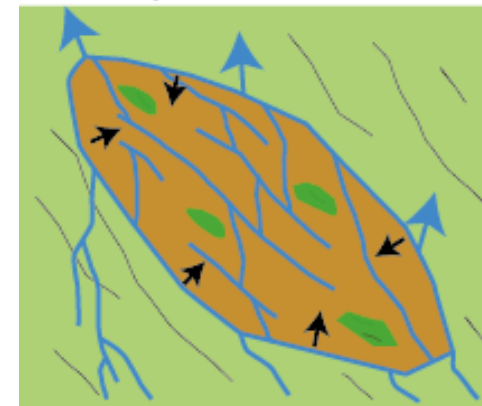
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Schist closed-system



Schist pervasive fluid flow

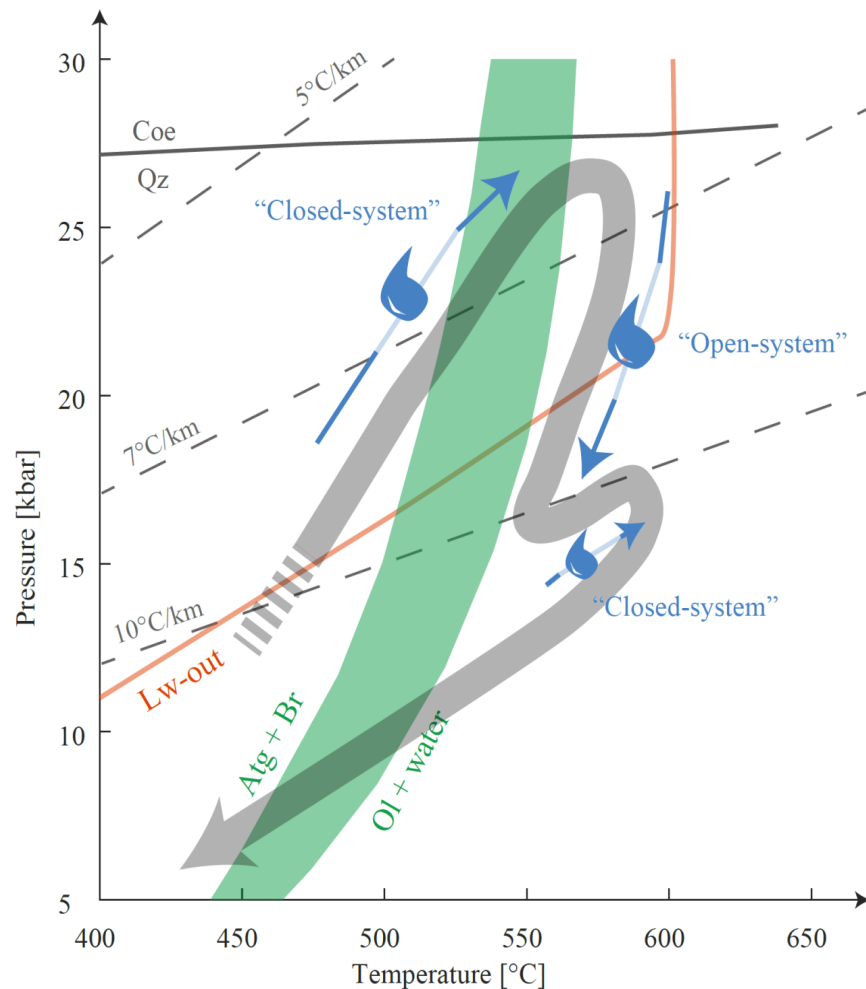


# Pervasive fluid flow in subducted crust

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Modelling  $\Rightarrow$  H<sub>2</sub>O flux at HP  $1.1 \times 10^5 \text{ cm}^3/\text{cm}^2$



- Elevated time-integrated fluid fluxes over the entire unit **requires pervasive fluid flow**

**Trigger:** transient rock volume variations caused by lawsonite breakdown

**Fluid source:** surrounding serpentinite

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