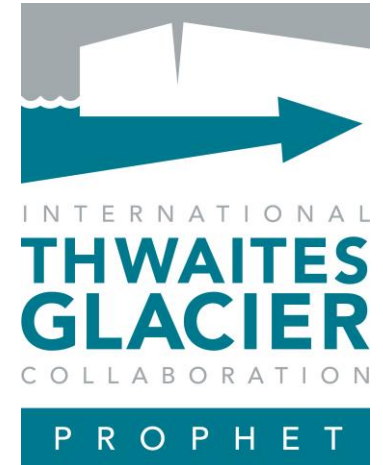


Treatment of Density Variations in Ice-Flow Models using the Shallow Ice Stream Approximation

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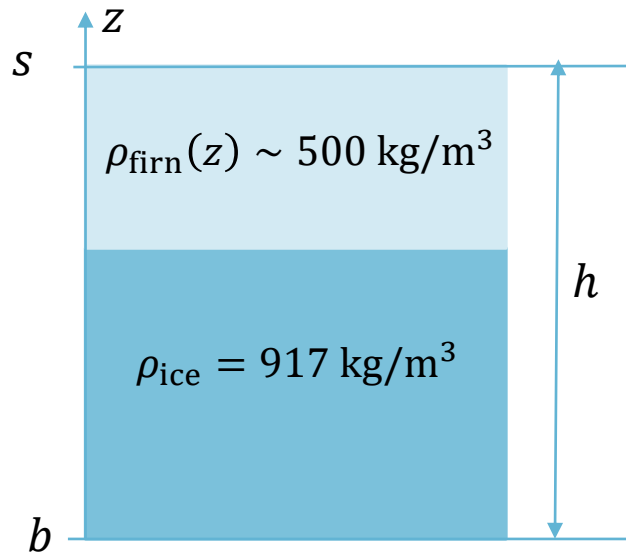
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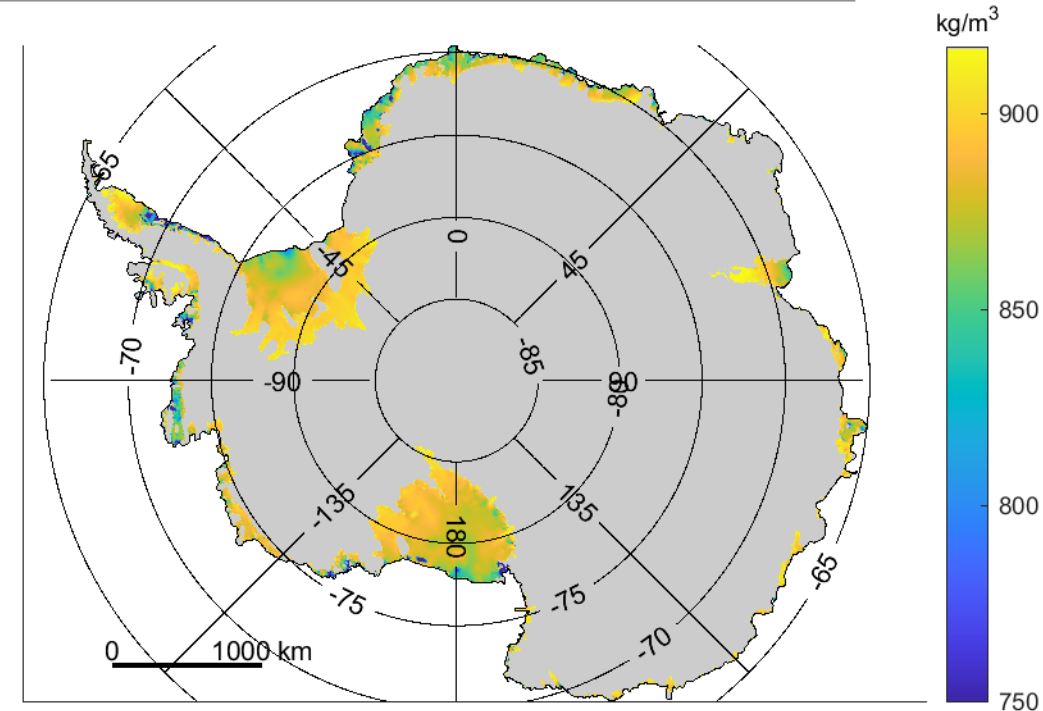
Density Variations in Ice Sheets

Ice Sheet = Ice layer + Firn layer



Vertically-averaged density,

$$\langle \rho \rangle(x, y) = \frac{1}{h} \int_b^s \rho(z) dz$$



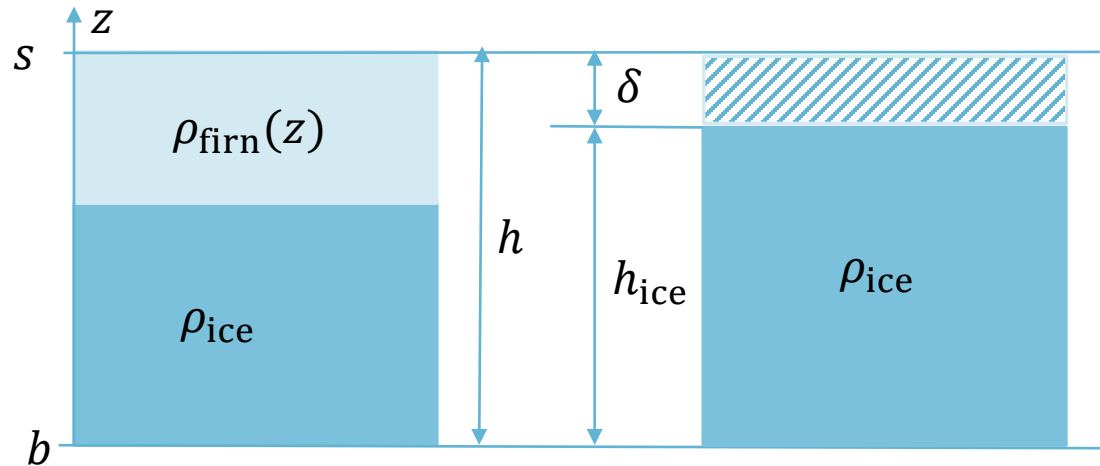
Vertically-averaged density of Antarctic Ice Shelves (BedMachine Antarctica ©Morlighem *et al.* 2020):

$\langle \rho \rangle \sim 917 - 790 \text{ kg/m}^3$,
more than 10% spatial variation in density.

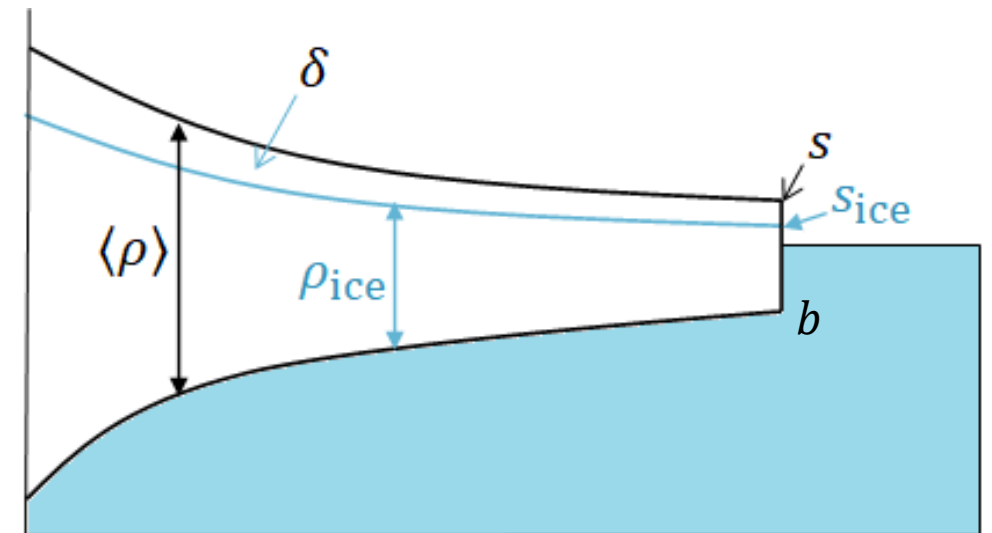
Density-to-Thickness Adjustment [D2T]

$$\text{Firn air-content: } \delta \equiv h \left(1 - \frac{\langle \rho \rangle}{\rho_{\text{ice}}} \right)$$

D2T adjustment: set $\rho = \rho_{\text{ice}}$ and adjust the thickness of the ice sheet: $h_{\text{ice}} = h - \delta$.



e.g. Floating Ice Shelf



$$\langle \rho \rangle \times h = \rho_{\text{ice}} \times h_{\text{ice}}$$

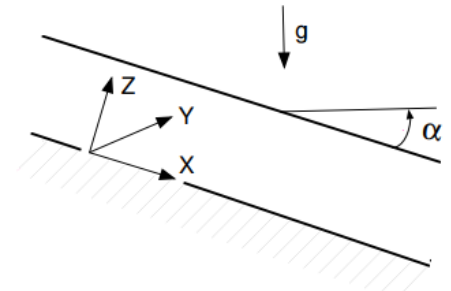
The SSA Equations with Horizontal Density Variations

Vertically-integrated SSA momentum-conservation:

$$\blacksquare \partial_x \left(4h\eta \partial_x u + 2h\eta \partial_y v + \frac{2h\eta}{\rho} \frac{D\rho}{Dt} \right) + \partial_y \left(h\eta (\partial_x v + \partial_y u) \right) - t_{bx} = \rho gh (\partial_x s \cos \alpha - \sin \alpha) + \frac{1}{2} g h^2 \partial_x \rho \cos \alpha$$

$$\blacksquare \partial_y \left(4h\eta \partial_y v + 2h\eta \partial_x u + \frac{2h\eta}{\rho} \frac{D\rho}{Dt} \right) + \partial_x \left(h\eta (\partial_x v + \partial_y u) \right) - t_{by} = \rho gh \partial_y s \cos \alpha + \frac{1}{2} g h^2 \partial_y \rho \cos \alpha$$

- On the **LHS**: additional momentum transfer between regions of low and high density.
- On the **RHS**: additional gravitational driving term due to horizontal density gradient.
- Both terms contribute equally $\sim \Delta\rho/\rho \lesssim 10\%$. However the RHS correction may dominate for grounded ice-caps, and also ice-streams/ice-shelves where $\partial_x s \approx 0$.



Mass-conservation:

$$\frac{D\rho}{Dt} + \rho \nabla \cdot \mathbf{v} = 0$$

Horizontal Density Variation Approaches

Typical
ice-flow
models

Density-to-Thickness adjustment [D2T]:

Set $\rho = \rho_{\text{ice}}$ and adjust the thickness of the ice sheet by the firm air-content: $h_{\text{ice}} = h - \delta$.

- + no modification to SSA eqns required
- modifies all terms involving h in the SSA eqns, not just density gradients

Density Variations - Body Force Only [DV-BF]:

Model a static density distribution $\rho(x, y)$ but neglect correction term on LHS of mom. eqn.

- + more realistic model of density variations
- may not be accurate in scenarios with significant longitudinal stretching and steep gradients

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Density Variations Advected [DVA]:

Allow density to evolve, e.g. set $D\rho/Dt = 0$ such that $\partial_t \rho = -\mathbf{v} \cdot \nabla \rho$.

- + more flexible formulation
- requires snow accumulation and compaction model to be realistic

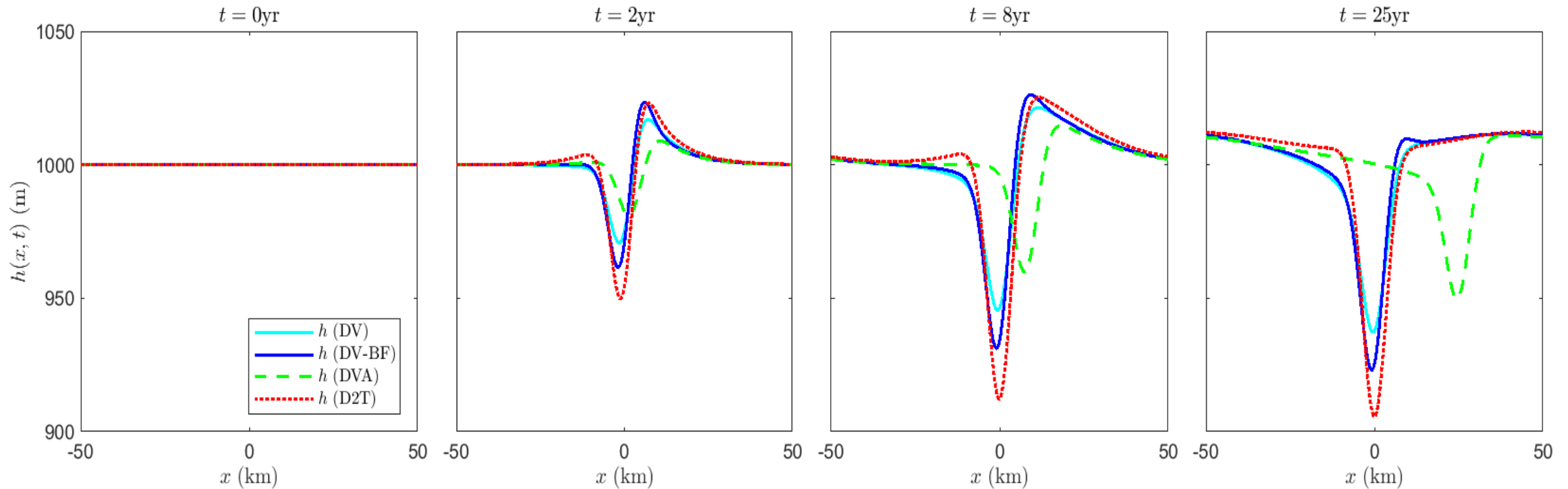
Density Variations [DV]:

Model a static density distribution $\rho(x, y)$ including both correction terms, i.e. $D\rho/Dt = \mathbf{v} \cdot \nabla \rho$.

- + more realistic model of density variations
- momentum eqn is no longer frame-invariant and so difficult to implement in large-scale flow models

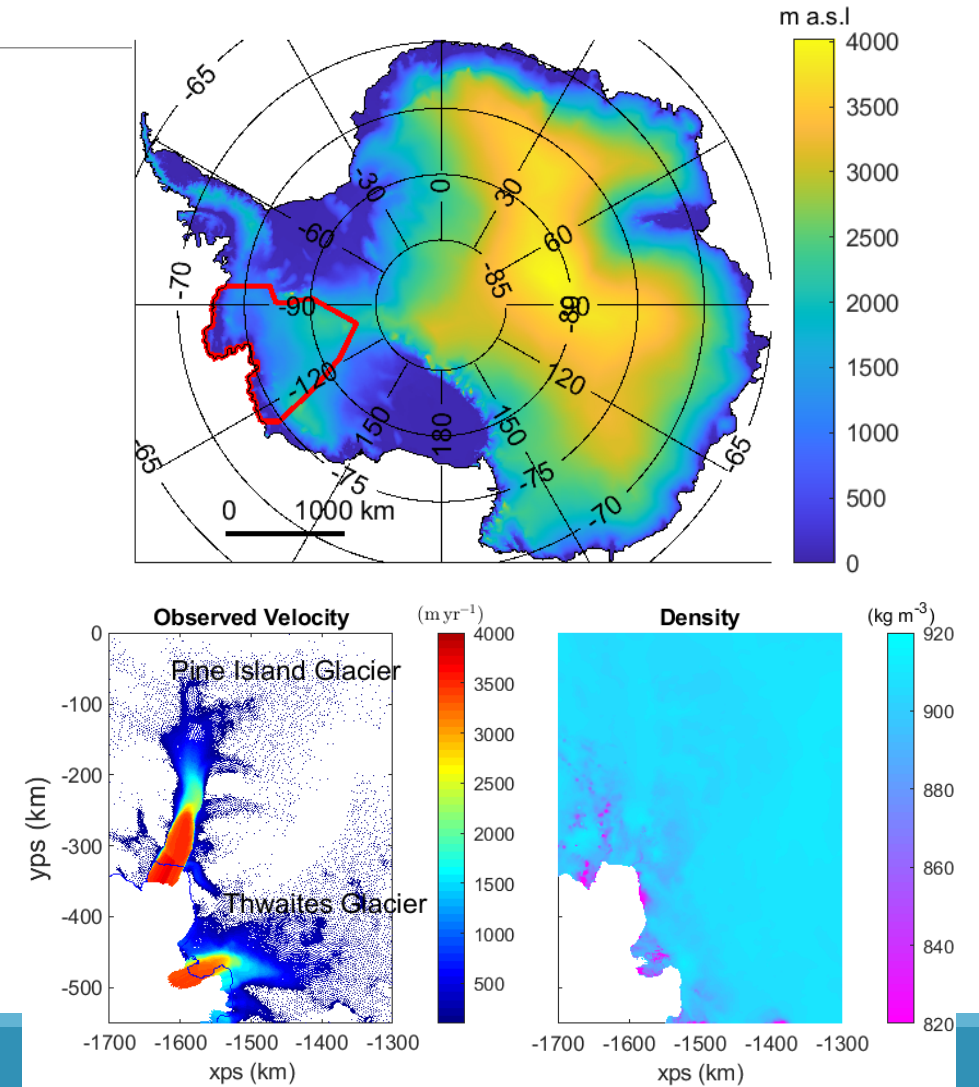
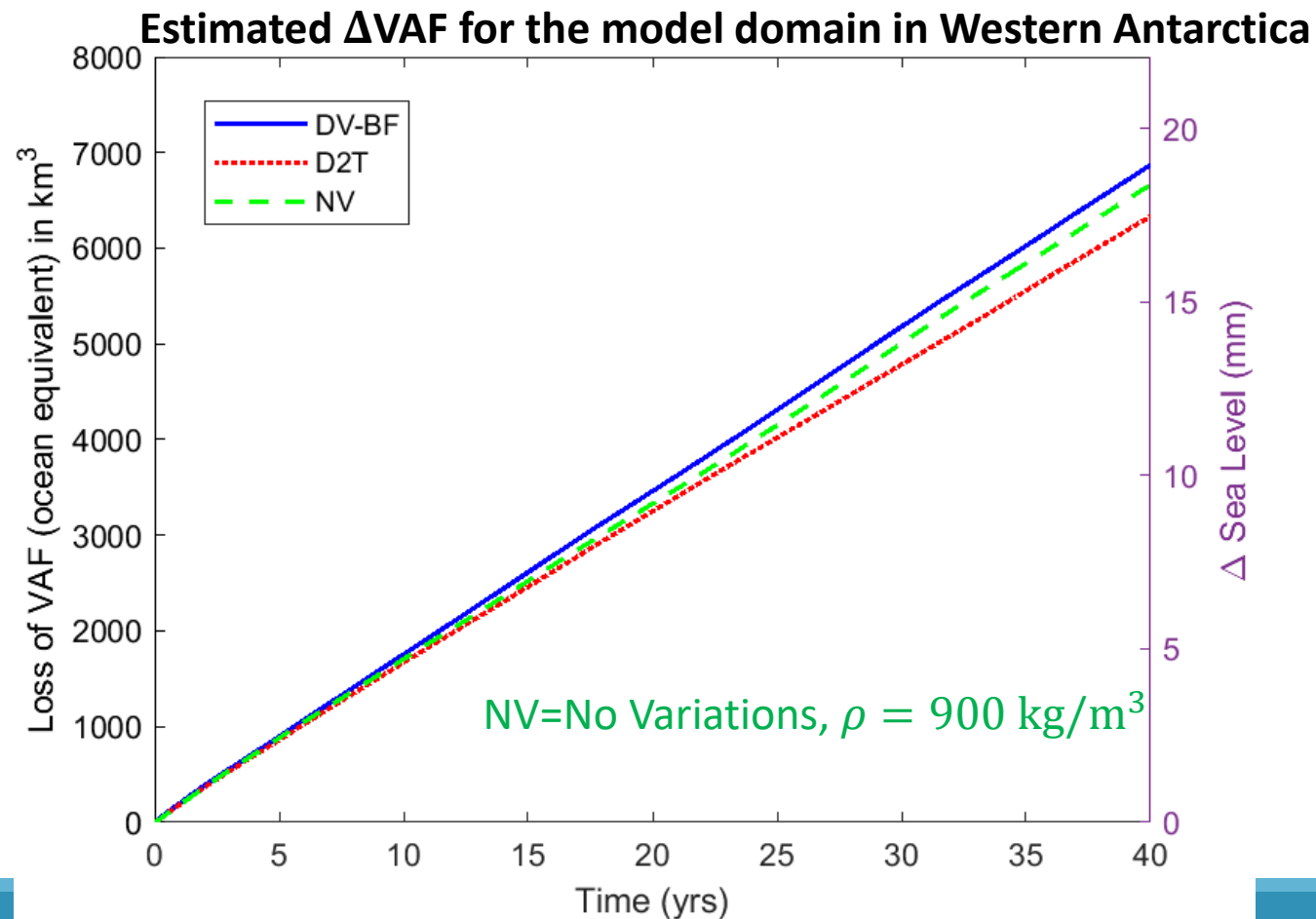
Analytical Simulation: Grounded Ice Sheet Example

infinite horizontal span, at an inclined slope α



Surface perturbation in response to a small Gaussian density perturbation applied to the ice sheet

Numerical Simulation of Western Antarctica



Conclusions

- We have derived a new theoretical framework for the inclusion of Horizontal Density Variations in large-scale ice-flow models.
- Different formulations are proposed for handling Horizontal Density Variations: D2T, DV-BF, DV, DVA.
- A ball-park estimate expects approximately a 10% correction due to observed density variations.
- Perturbation Analysis (to first order in the density variations) reveals a number of qualitative and quantitative differences between different density formulations.
- Transient simulations of the Western Antarctic Ice Sheet suggest that the impact on predicted sea level change could be approximately 10% between different formulations.