

The Regional Fractional Energy Balance Equation: Projections to 2100

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Agenda

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Bayesian Parameter Estimation and Error Models

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Parameter Estimates

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Regional Projections

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Energy Balance Models

- The classical energy balance equation (EBE) assumes the Earth is a uniform box or a set of N-boxes, we will consider N=1.

$$C \frac{dT}{dt} + AT = F(t)$$

- Solving with Green's functions we have:

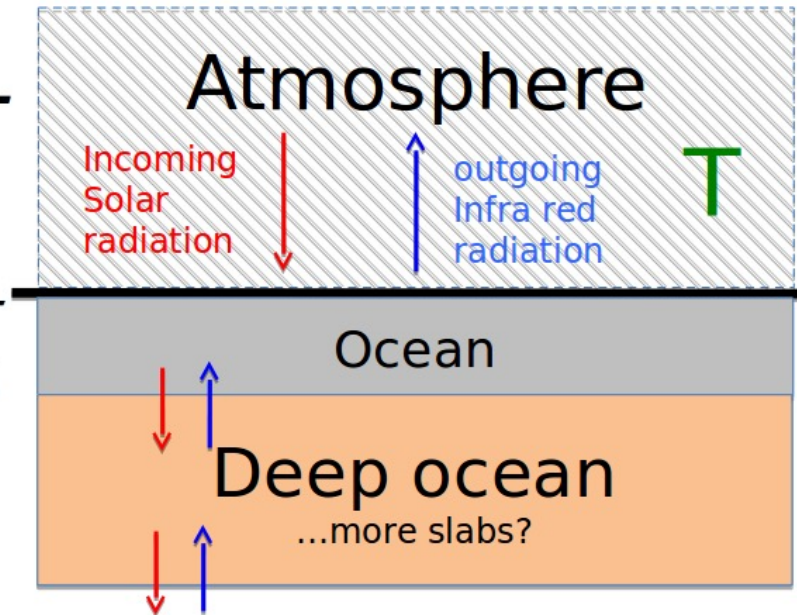
$$T(t) = s \int_{-\infty}^t G_0(t-u) F(u) du$$

Linearized heat gained by radiative exchange:

$$AT$$

Heat absorbed by box:

$$C \frac{dT}{dt}$$



Where the Green's function is:

$$s = \frac{1}{A} \quad \tau = Cs$$

$$G_0(t) = \frac{1}{\tau} e^{-\frac{t}{\tau}}$$

Fractional Energy Balance Equation

- We now generalize the derivative operator to a Weyl fractional derivative:

$$\tau^h {}_{-\infty}D_t^h + T(t) = sF(t)$$

Where the solution again can be represented in terms of a Green's function:

$$T(t) = s \int_{-\infty}^t G_{0,h}(t-u)F(u)du$$

The Dirac response Green's function is given as:

$$G_{0,h}(t) = \tau^{-1} \left(\frac{t}{\tau}\right)^{h-1} E_{h,h} \left(-\left(\frac{t}{\tau}\right)^h\right)$$

Using the 2-Parameter Mittag-Leffler function:

$$E_{\alpha,\beta} = \sum_k^{\infty} \frac{z^k}{\Gamma(\alpha k + \beta)}$$

Model and Forcing Parameters

$$T(t) = s \int_{-\infty}^t G_{0,h}(t-u)F(u)du$$

$$F(t) = F_{GHG} + \alpha F_{Aer} + F_{Vol} + \frac{\langle F_{Vol} \rangle}{\langle F_{Vol}^v \rangle} F_{Vol}^v + \gamma$$

Model Parameters: h and τ

- h characterizes the memory of the system.
- τ characterizes the approach of the system to equilibrium.

Forcing Parameters: α and v

- α is the aerosol linear scaling factor that effectively re-calibrates the aerosol forcing.
- N is the volcanic intermittency correction exponent that allows high intermittent volcanic forcing to be used in a linear framework.

Climate Sensitivity Parameter: s

- s is the climate sensitivity parameter s , refers to the equilibrium change in the annual GMST following a unit change in radiative forcing.

Bayesian Parameter Estimation and Error Models

The error model was not ad hoc, rather predicted by the model itself: the internal variability response to white noise internal forcing, a fraction Relaxation noise (fRn).

Fractional Gaussian Noise Correlation Function:

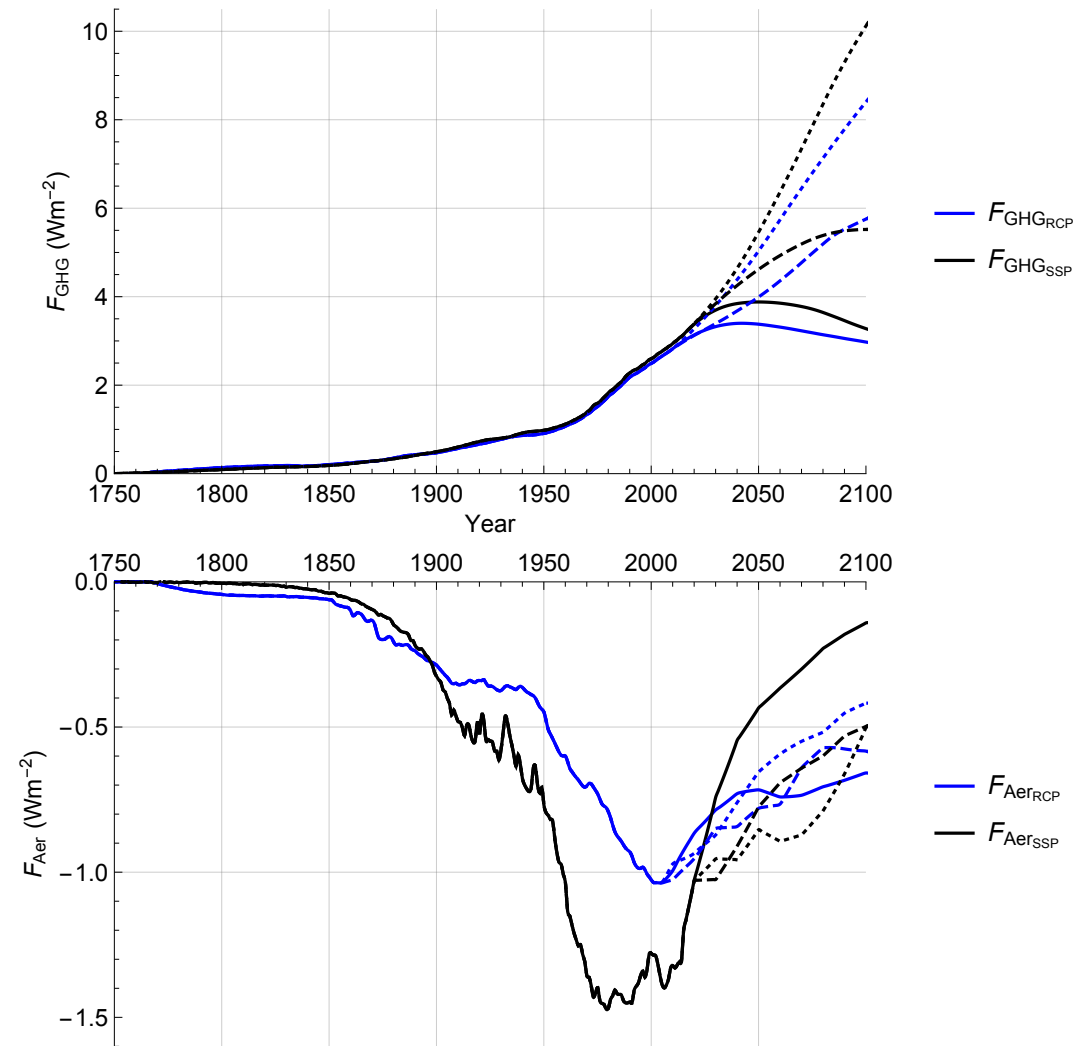
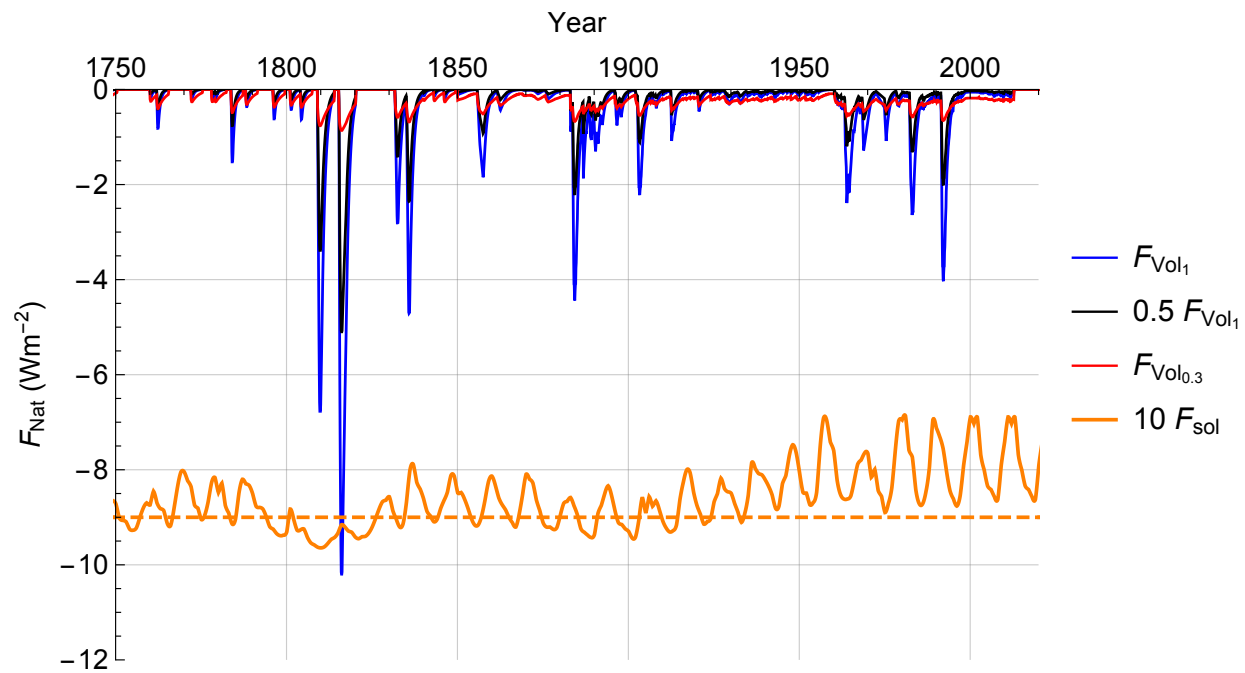
$$R_h^{(fGn)}(\lambda) = \frac{1}{2} \left((\lambda + 1)^{2h+1} + (\lambda - 1)^{2h+1} - 2\lambda^{2h+1} \right)$$

The fGn is only the high frequency approximation of the more general fRn.

Fractional Relaxation Noise Correlation Function:

$$R_h^{(fRn)}(\lambda) = \frac{\tau^2}{2} \left(V_h \left(\frac{\lambda - 1}{\tau} \right) - V_h \left(\frac{\lambda + 1}{\tau} \right) - 2V_h \left(\frac{\lambda}{\tau} \right) \right)$$

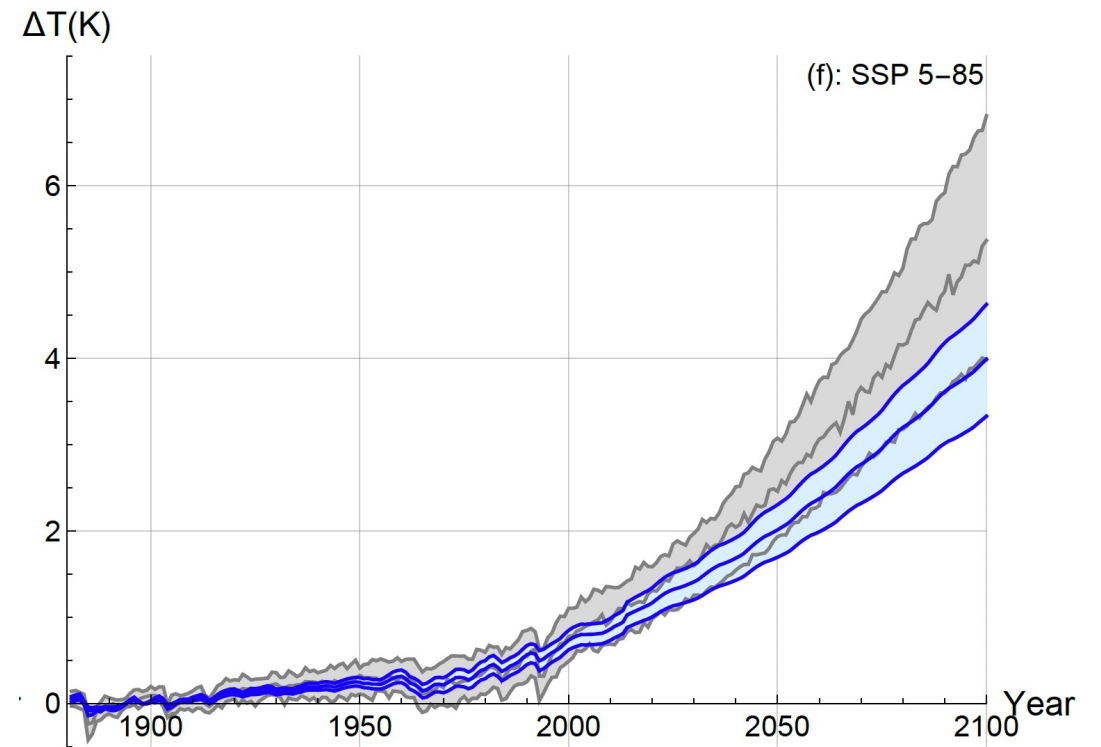
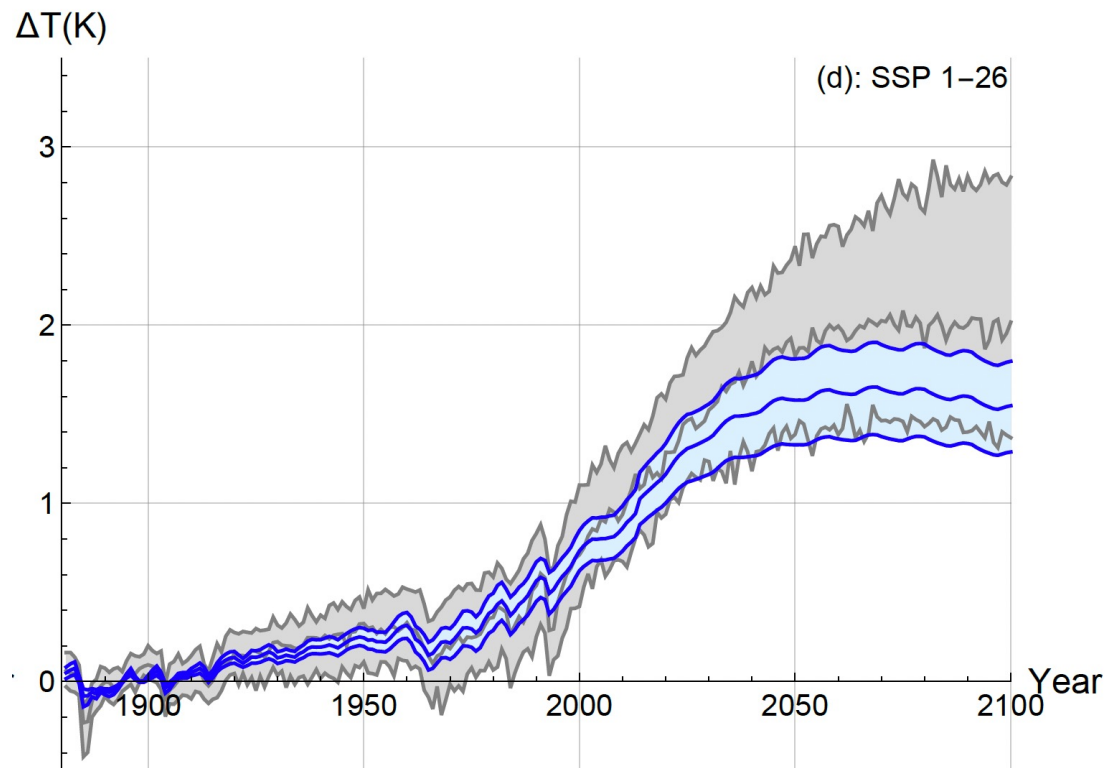
Forcings



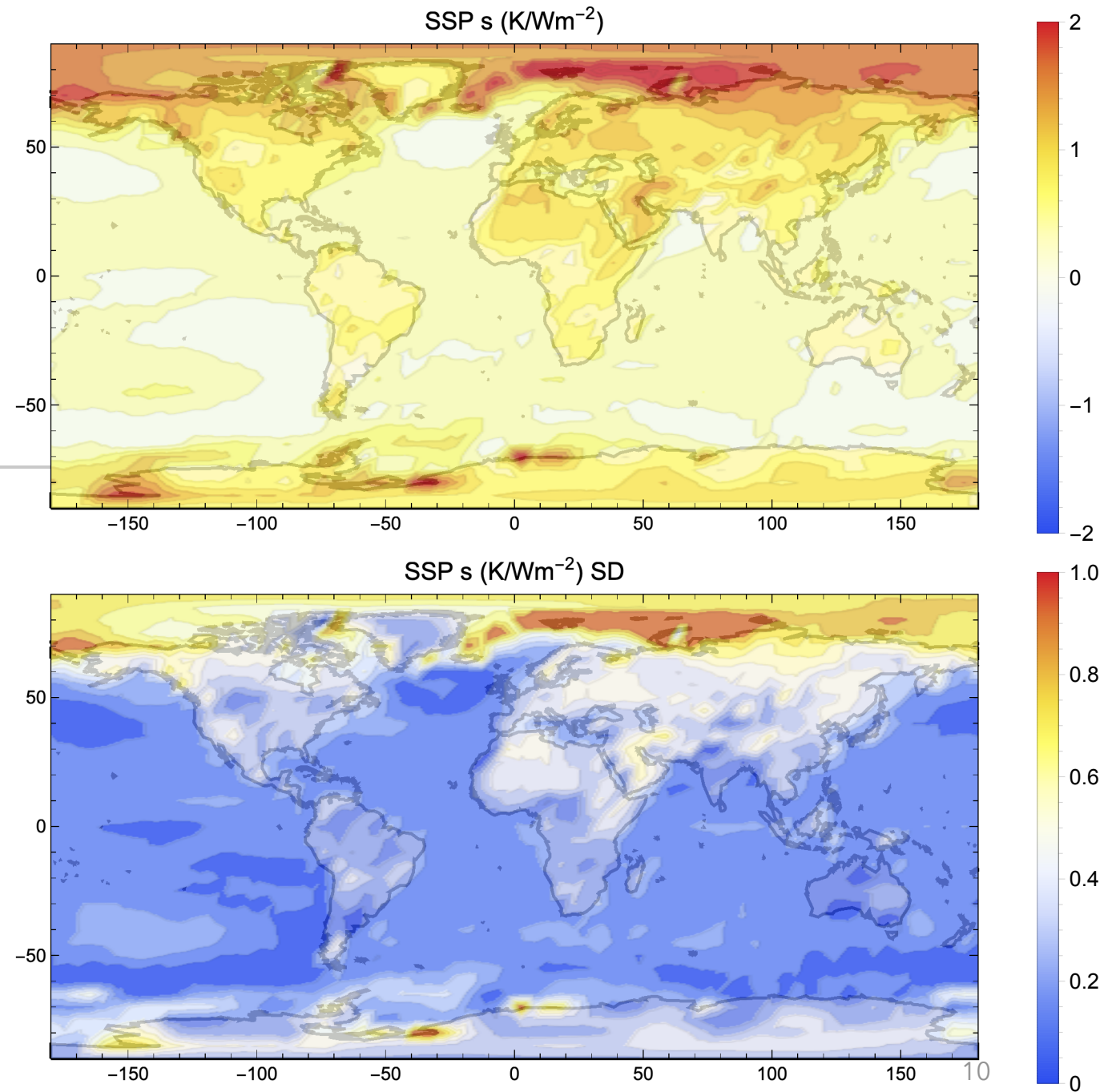
Global Parameter Estimates

	h	τ (years)	α	ν	ECS
Median	0.38	4.7	0.33	0.28	1.8
90% CI	[0.32, 0.44]	[2.4, 7.0]	[0.05, 0.61]	[0.16, 0.40]	[1.5, 2.2]

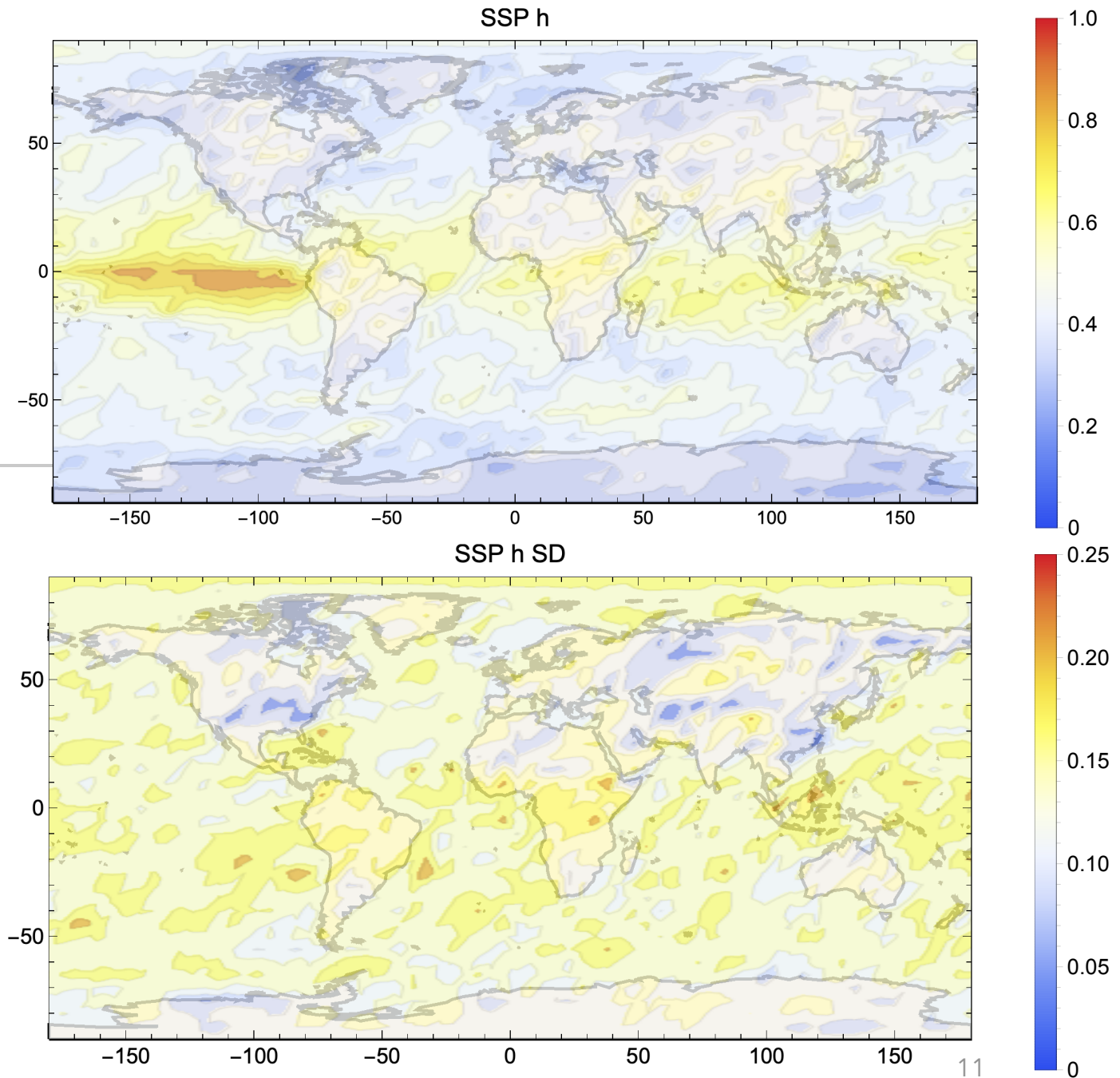
Global Projections



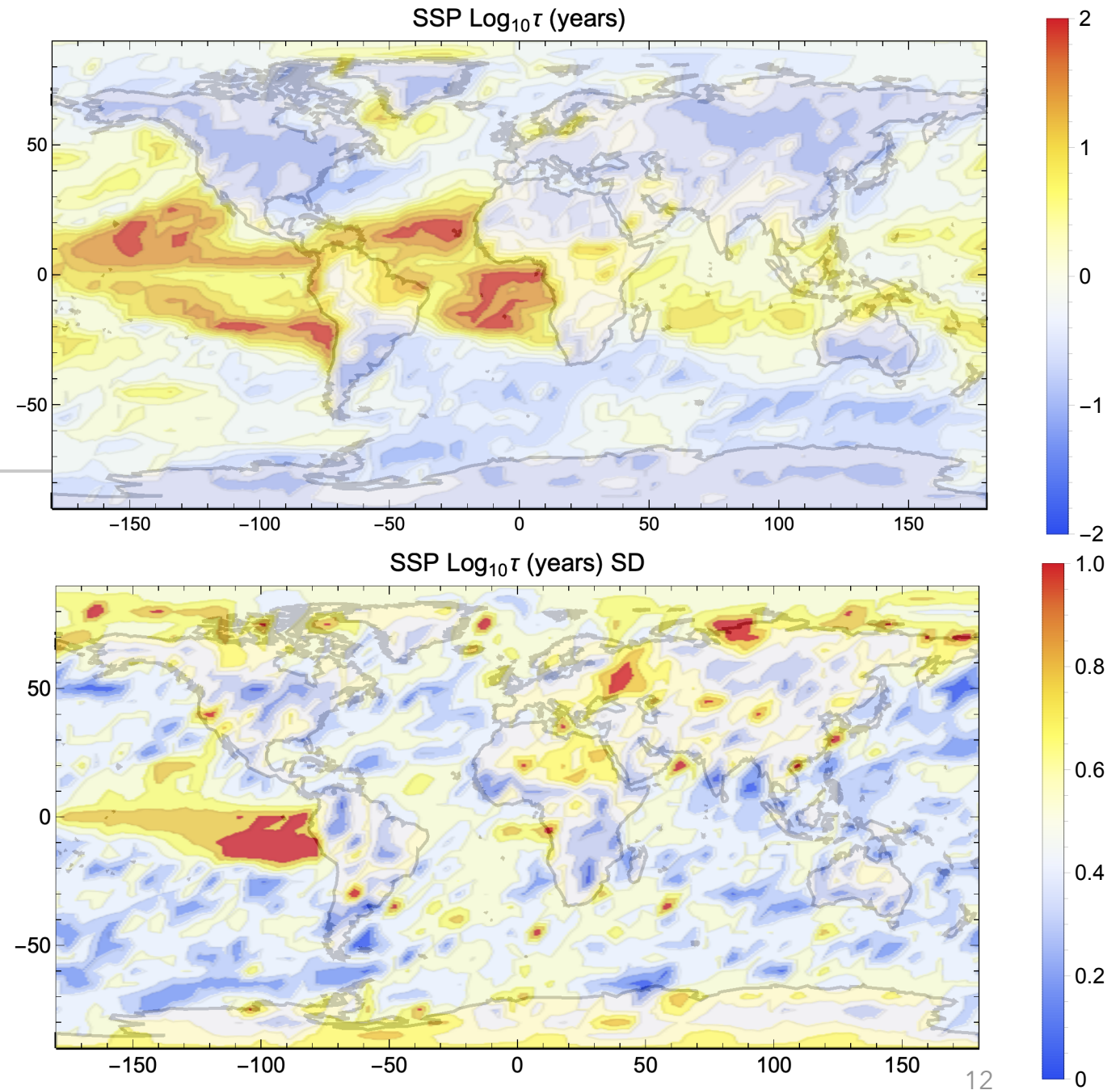
Regional Parameters: Climate Sensitivity



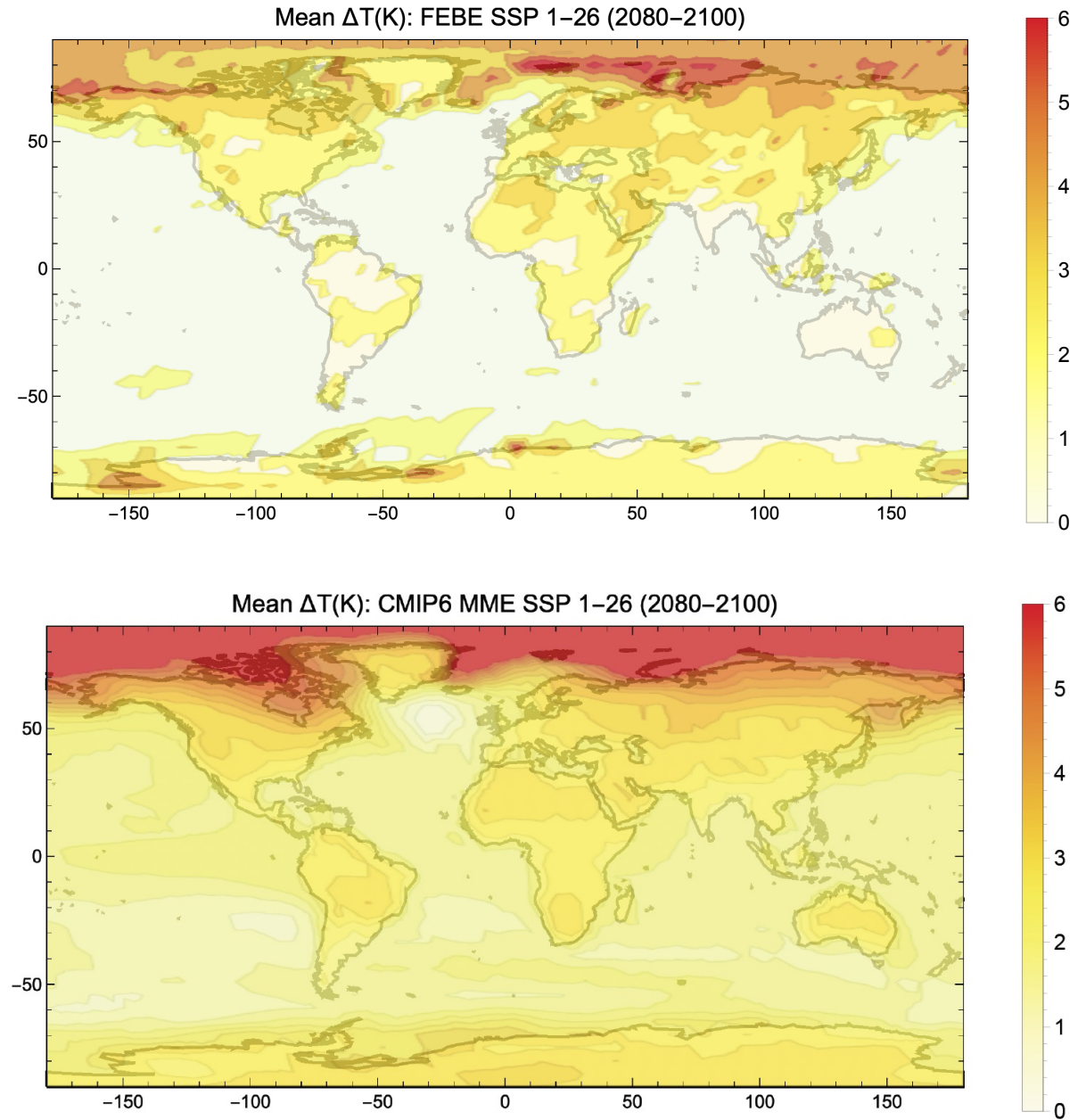
Regional Parameters: h Parameter



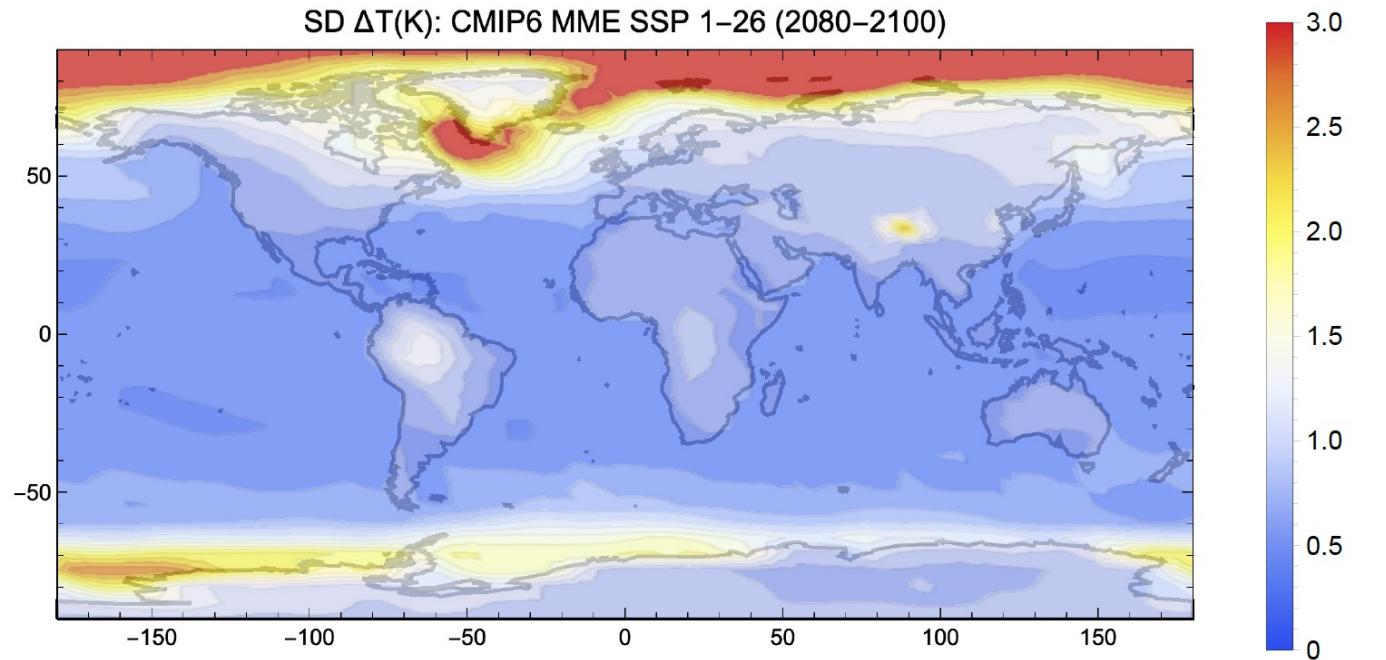
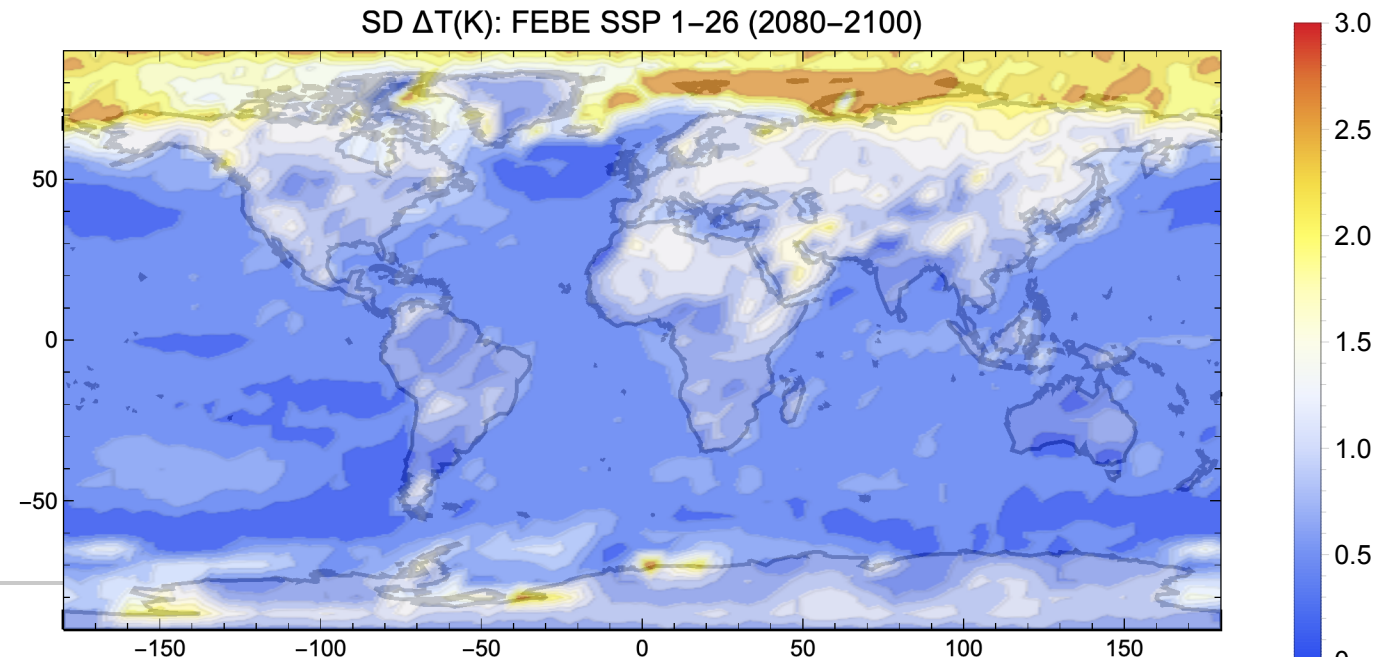
Regional Parameters: Relaxation Time



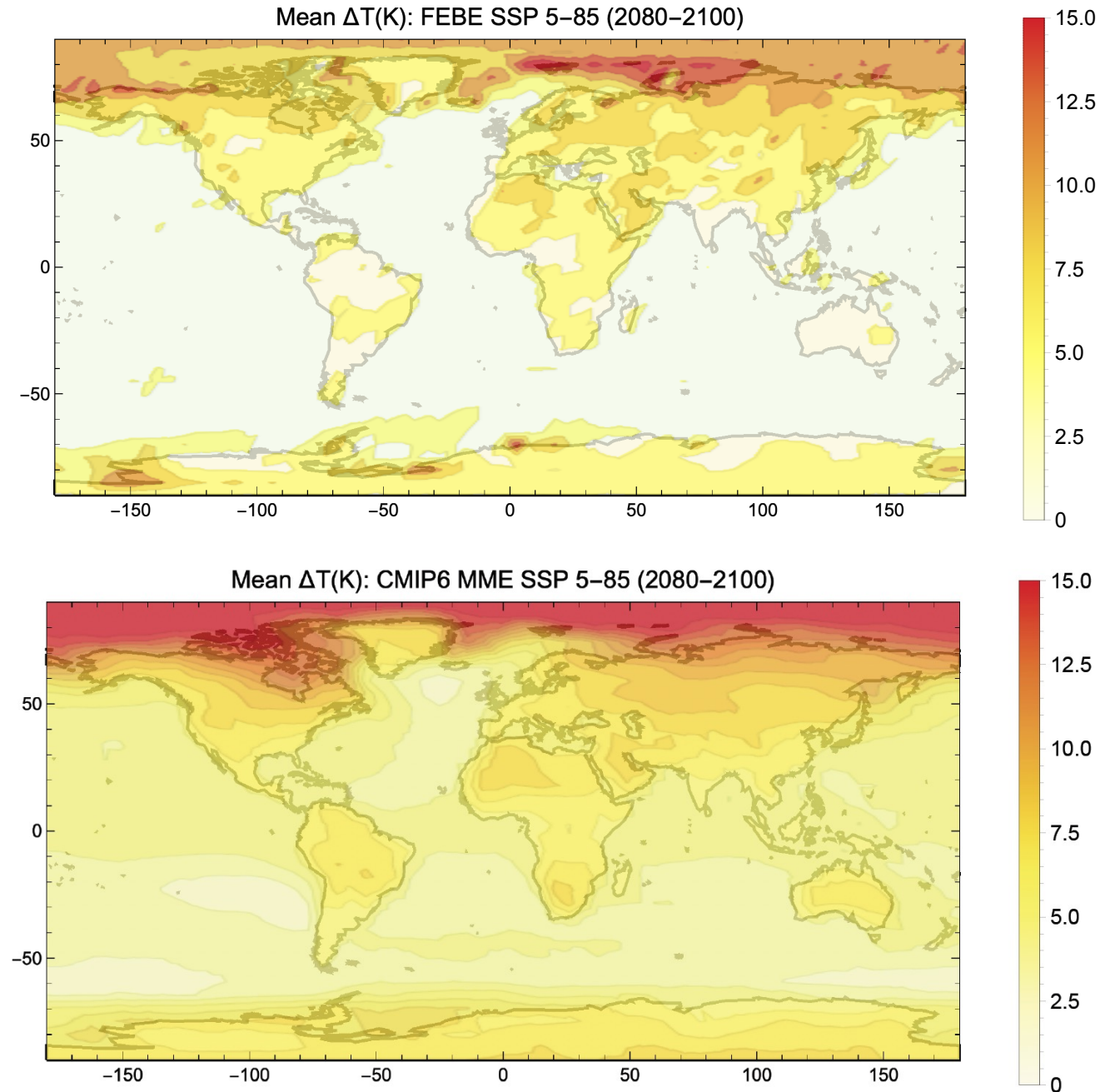
Climate Projections: SSP 1-26



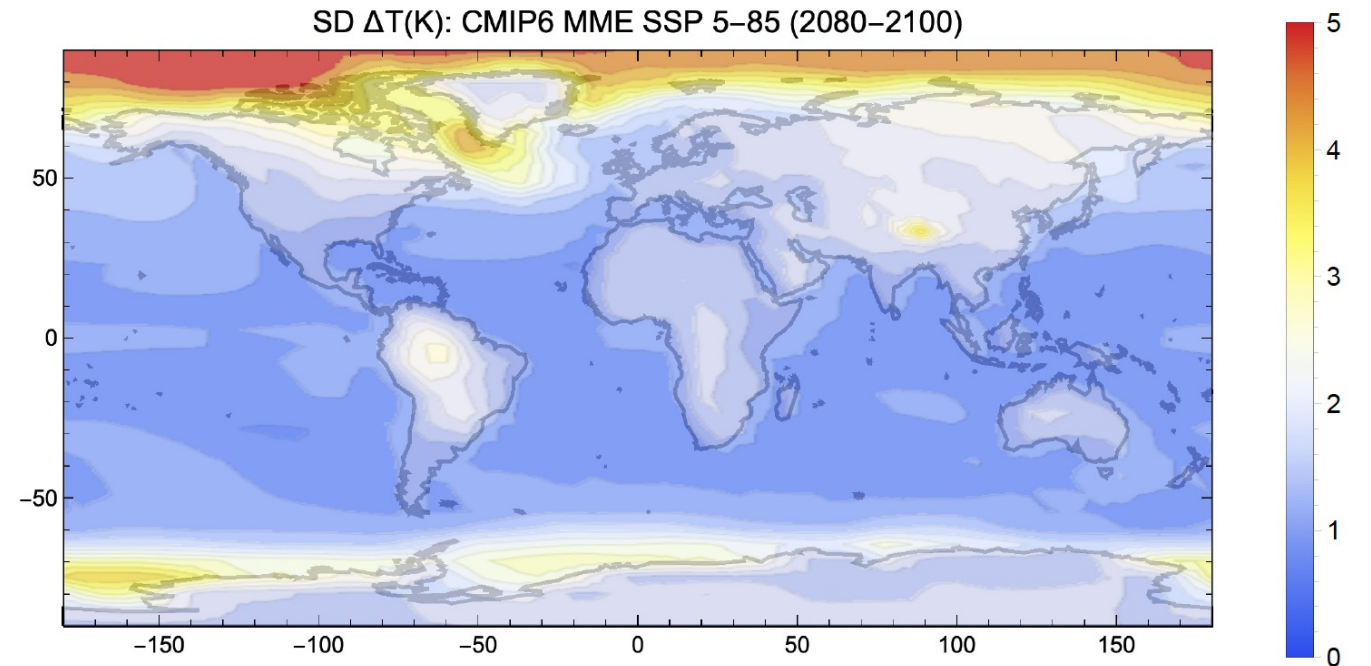
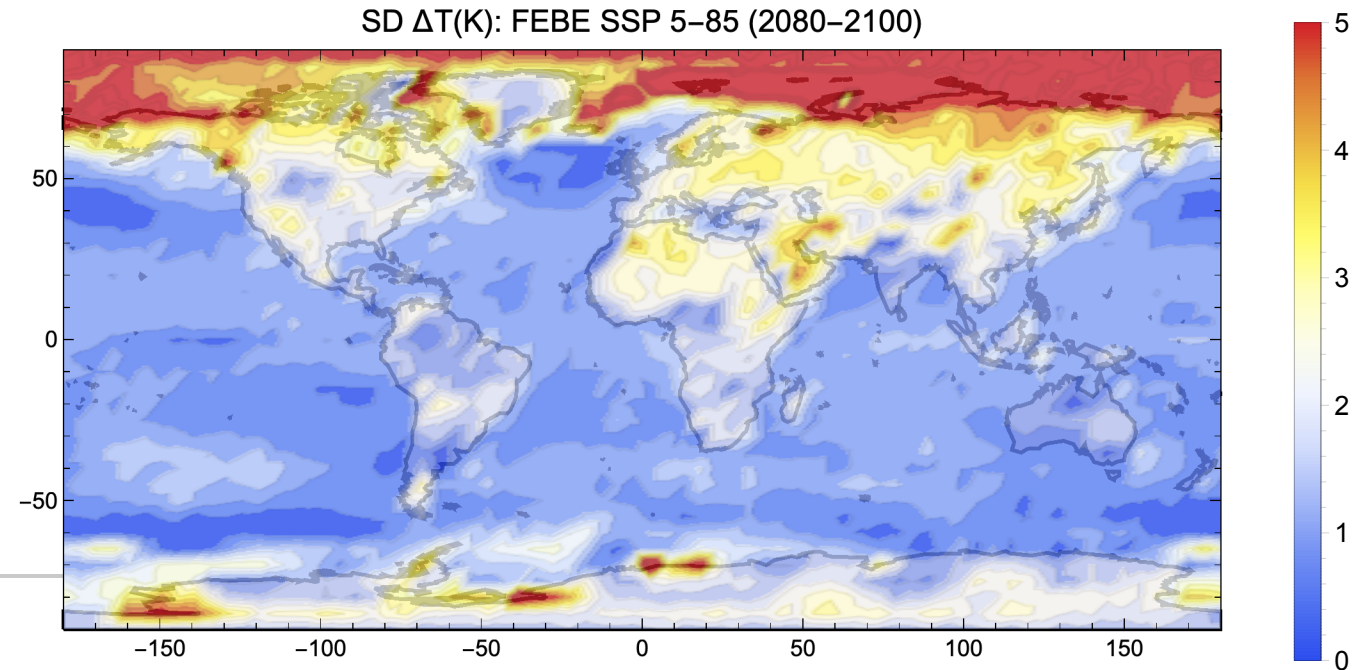
Climate Projections SD: SSP 1-26



Climate Projections: SSP 5-85



Climate Projections SD: SSP 5-85





Future Work

- Investigate generational differences between CMIP models by calibrating using the historical runs of the CMIP models in order to perform a feedback analysis.
- Updating our parameter estimates from calibrations on GCMs allows for GCM-FEBE hybrid projections.
- The regional FEBE can be expanded to allow for pixel-to-pixel interaction through energy transfers.
- Investigate non-linear feedbacks such as the ice-albedo feedbacks.
- Extensions to precipitation at global and regional scales since the FEBE model is consistent with space-time scaling processes in historical precipitation data.

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Thank you



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