

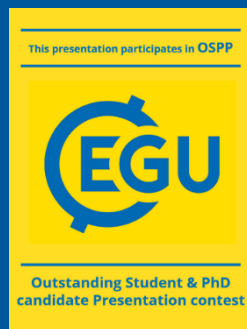
Insights into the deep geothermal potential from heat in place calculations of the Berlin/Brandenburg region

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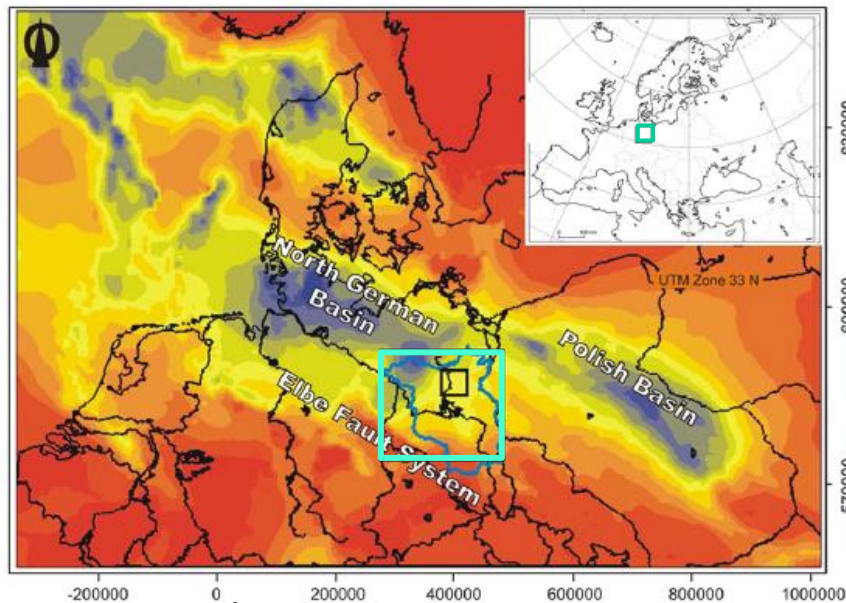
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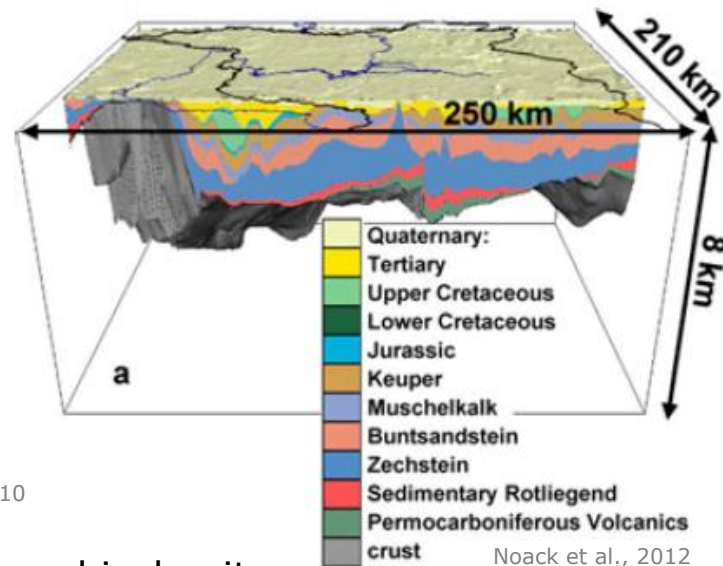
Go to Abstract
and judging option



Berlin/Brandenburg as target area



Noack et al., 2010

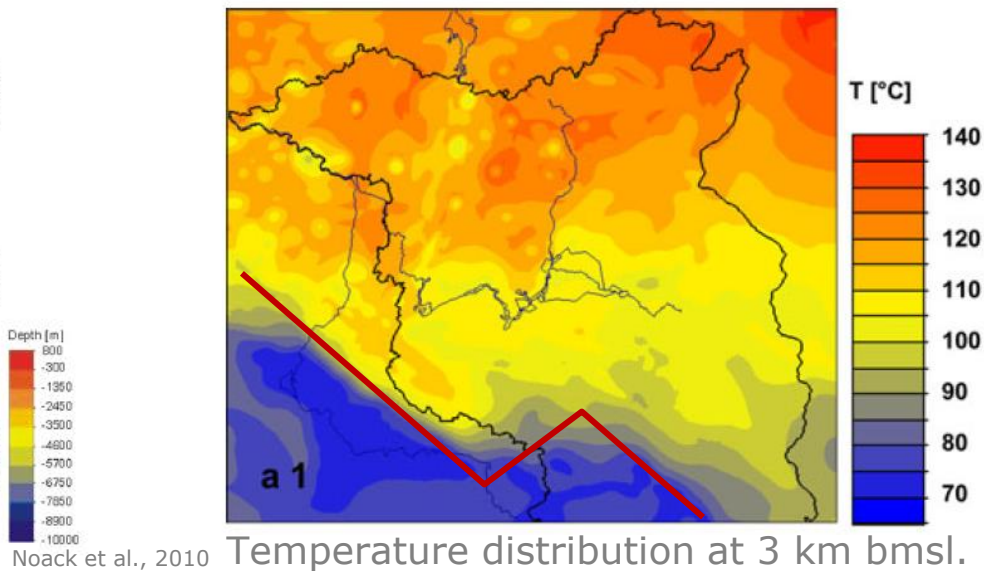
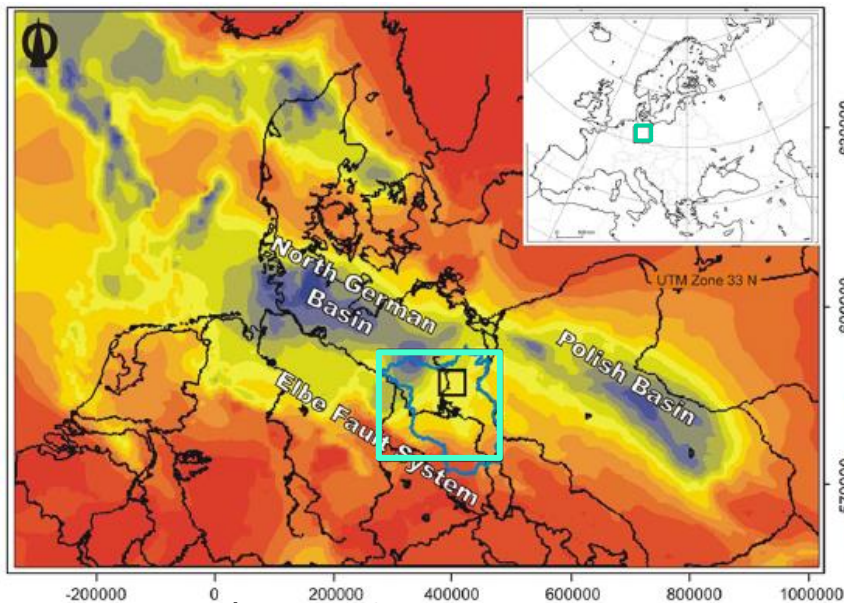


Noack et al., 2012

- at North East German Basin rim
- Heat in place (HIP) calculated for each litho-stratigraphical unit
- reference model includes 3D structural geological model



Berlin/Brandenburg as target area



- at North East German Basin rim
- Heat in place (HIP) calculated for each litho-stratigraphical unit
- reference model includes 3D structural geological model and thermal conductive model



Berlin/Brandenburg as target area



How does the value of models change with HIP additional to temperature?



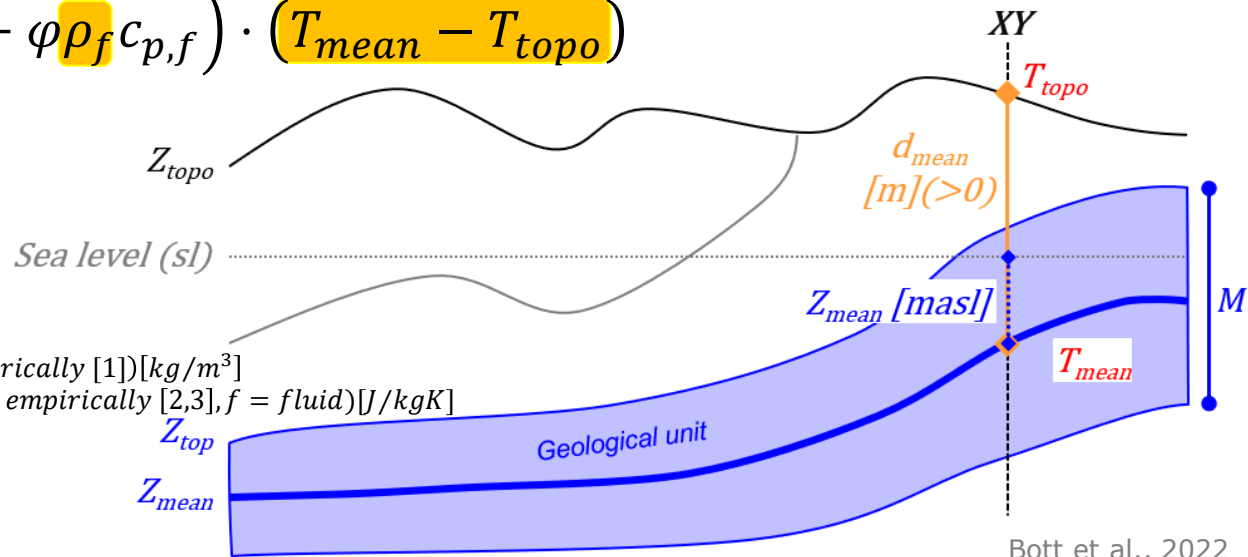
HIP method for geothermal potential



$$H = M \cdot \left((1 - \varphi) \rho_s c_{p,s} + \varphi \rho_f c_{p,f} \right) \cdot (T_{mean} - T_{topo})$$

Vary in every point

H = Heat in place [J/m^2]
 M = Reservoir thickness [m]
 φ = porosity [-]
 ρ = density (s = solid, f = fluid calculated empirically [1]) [kg/m^3]
 c_p = specific heat capacity (s = solid calculated empirically [2,3], f = fluid) [J/kgK]
 T_{mean} = mean temperature [$^{\circ}\text{C}$]
 T_{topo} = surface temperature [$^{\circ}\text{C}$]



Bott et al., 2022



HIP contributing and empirical formulas

For hydrostatic pressure:

$$p = \rho_{brine\ 0} \cdot g \cdot d_{mean} + 101325\ Pa$$

For mean reservoir temperature:

$$T\ at\ Z_{mean,XY} = Z_{min,XY} - \frac{M_{XY}}{2}$$

[1] Batzle and Wang, 1992

$$\rho_{fw} = 1 + 10^{-6} \cdot (-80T - 3.3T^2 + 1.75 \cdot 10^{-3}T^3 + 489p - 2Tp + 0.016T^2p - 1.3 \cdot 10^{-5}T^3p - 0.333p^2 - 0.002Tp^2)$$

$$\rho_{brine} = \rho_{fw} + S \cdot (0.688 - 0.44S + 10^{-6} \cdot (300p - 2400pS + T \cdot (80 + 3T - 3300S - 13p + 47pS)))$$

[2] Bär, 2012; [3] Vosteen and Schellschmidt, 2003 for sedimentary layers (1), and for magmatic rocks (2)

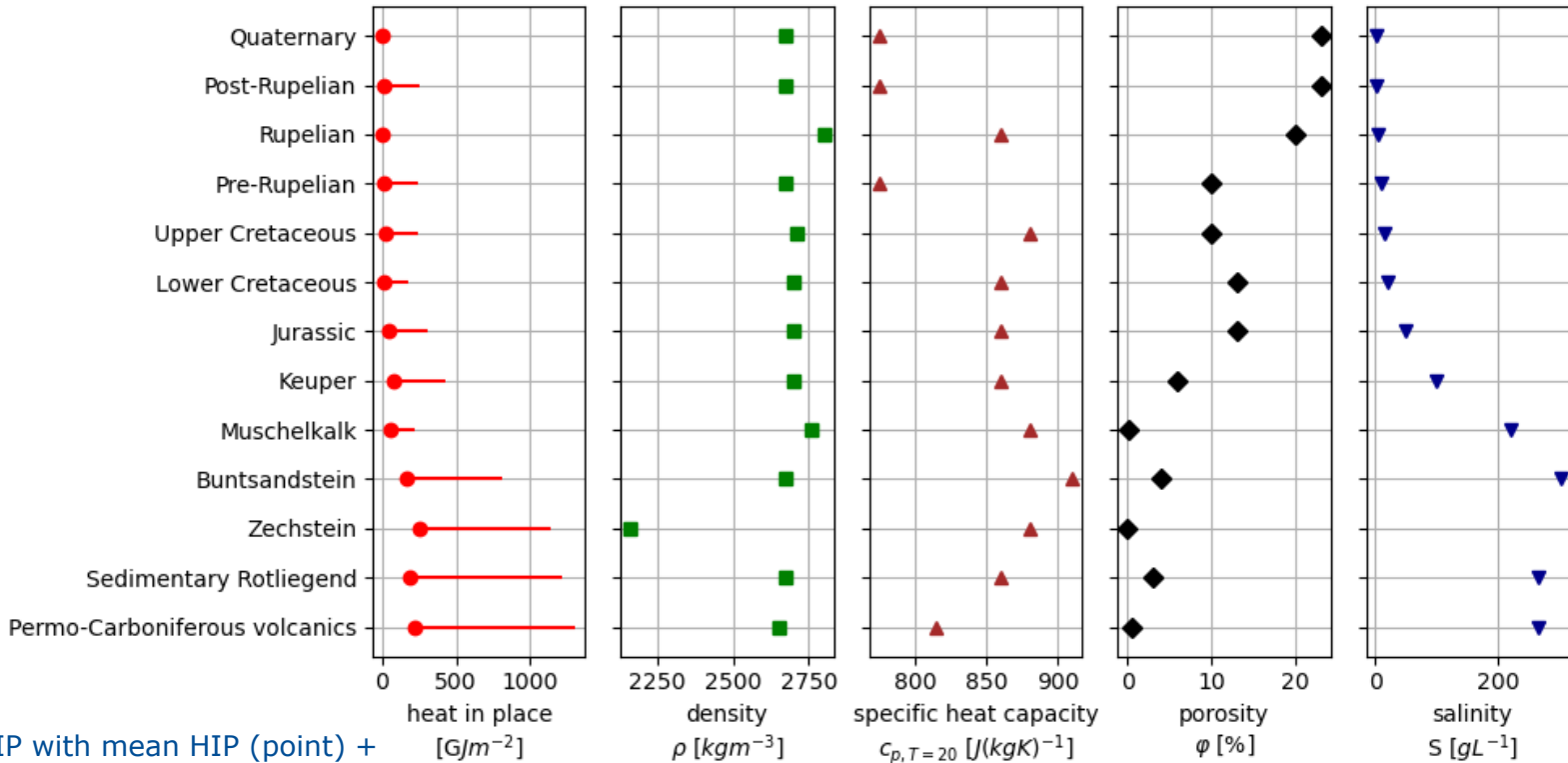
$$(1) \ c_{p,m}(T) = 6.295 \cdot 10^{-6} \cdot T^3 - 4.99 \cdot 10^{-3} \cdot T^2 + 1.71 \cdot T + c_{p,m,(T=20^\circ C)}$$

$$(2) \ c_{p,m}(T) = 4.619 \cdot 10^{-6} \cdot T^3 - 3.945 \cdot 10^{-3} \cdot T^2 + 1.46 \cdot T + c_{p,m,(T=20^\circ C)}$$

p = hydrostatic pressure [Pa], for [1] in [MPa]
ρ_{brine 0} = water density at surface [kg/m³]
g = gravitational acceleration [m/s²]
d_{mean} = distance of water column [m]
T = mean temperature [°C]
Z_{mean,XY}, Z_{min,XY} = Z coordinate of point (X,Y)[m]
M_{XY} = thickness at point (X,Y)[m]
ρ_{fw}, ρ_{brine} = density (fw = freshwater)[g/cm³]
S = salinity [weight fraction = ppm/1000000]
c_{p,m} = specific heat capacity [J/kgK]



Lithology-dependent HIP & parameters



HIP with mean HIP (point) +
error bar from min to max HIP



All lithology-dependent HIP parameters

Lithology (all values for presented model)	φ porosity [%]	ρ density _(s) [kg/m^3]	c_p specific heat capacity ($T=20^\circ C$) [J/kgK]	Salinity [g/L] If $\varphi = 0$ then no value
Quaternary	23 [1]	2670 [4,5]	775 [6]	2 [7]
Tertiary Post-Rupelian	23 [1]	2670 [4,5]	775 [6]	2 [7]
Tertiary Rupelian	20 [1]	2800 [4,5]	860 [6]	5 [7]
Tertiary Pre-Rupelian	10 [1]	2670 [6]	775 [6]	10 [7]
Upper Cretaceous	10 [1]	2710 [10]	880 [6]	15 [7]
Lower Cretaceous	13 [1]	2700 [4,5]	860 [6]	20 [7]
Jurassic	13 [1]	2700 [4,5]	860 [6]	50 [7]
Keuper	6 [1]	2700 [4,5]	860 [6]	100 [7]
Muschelkalk	0.1 [1]	2760 [6]	880 [6]	220 [7]
Buntsandstein	4 [1,2]	2670 [4,5]	910 [6]	300 [7]
Zechstein	0 [1]	2160 [4,5]	880 [6]	-
Sedimentary Rotliegend	3 [9]	2670 [1,5]	860 [6]	265 [8]
Permo-Carboniferous volcanics	0.5 [1,3]	2650 [1,5]	815 [6]	265 [8]
Pre-Permian	0 [1]	2700 [1,5]	850 [6]	-
Upper Crust	0 [1]	2800 [5,11]	850 [6]	-
Lower Crust	0 [1]	2800 [5,11]	850 [6]	-
Lithospheric mantle	0 [1]	3300 [5,11]	1005 [6]	-

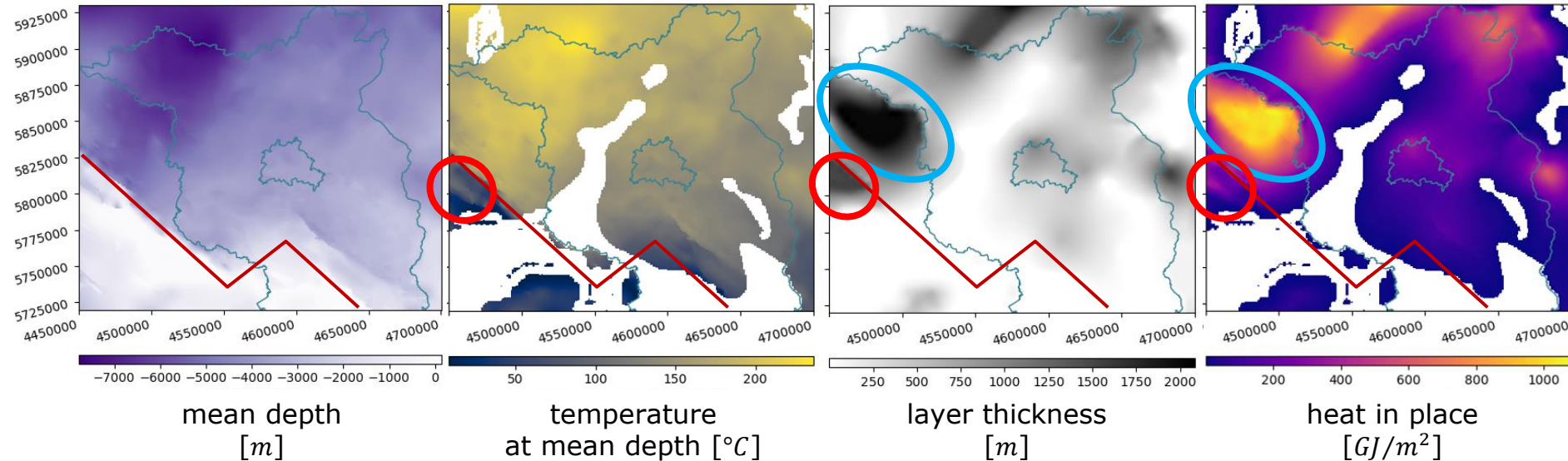


HIP f(temperature, thickness)

Permo-Carboniferous volcanics

reference temperature model

HIP



- temperature and thickness dependency shows very well in HIP maps

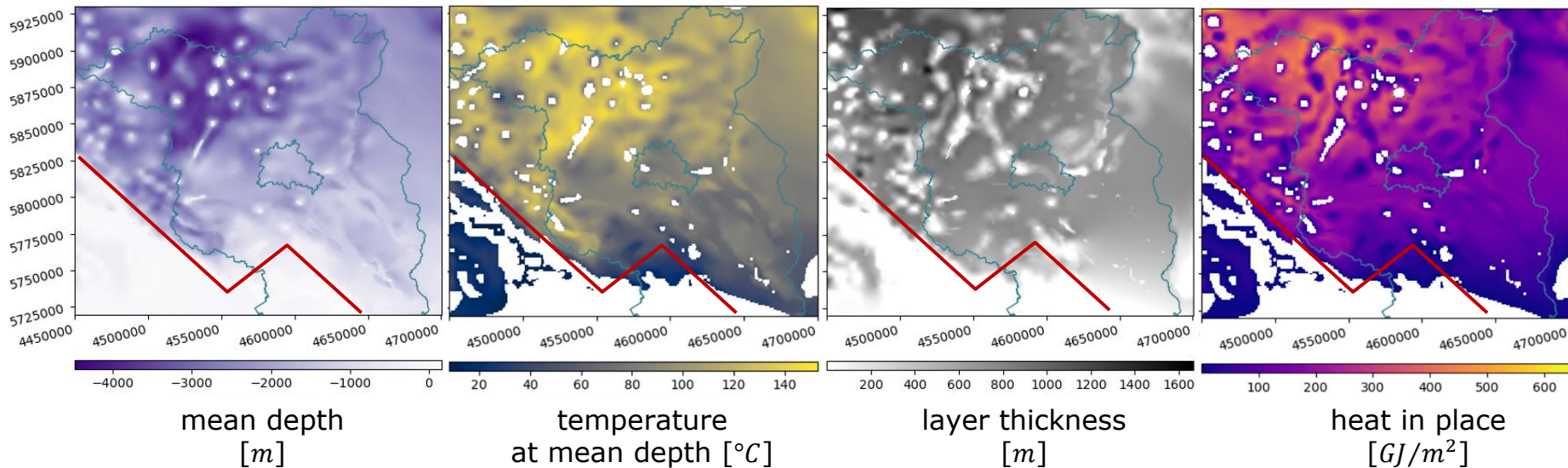


HIP variations due to salt

Buntsandstein

reference temperature model

HIP



- fluctuating pattern due to salt

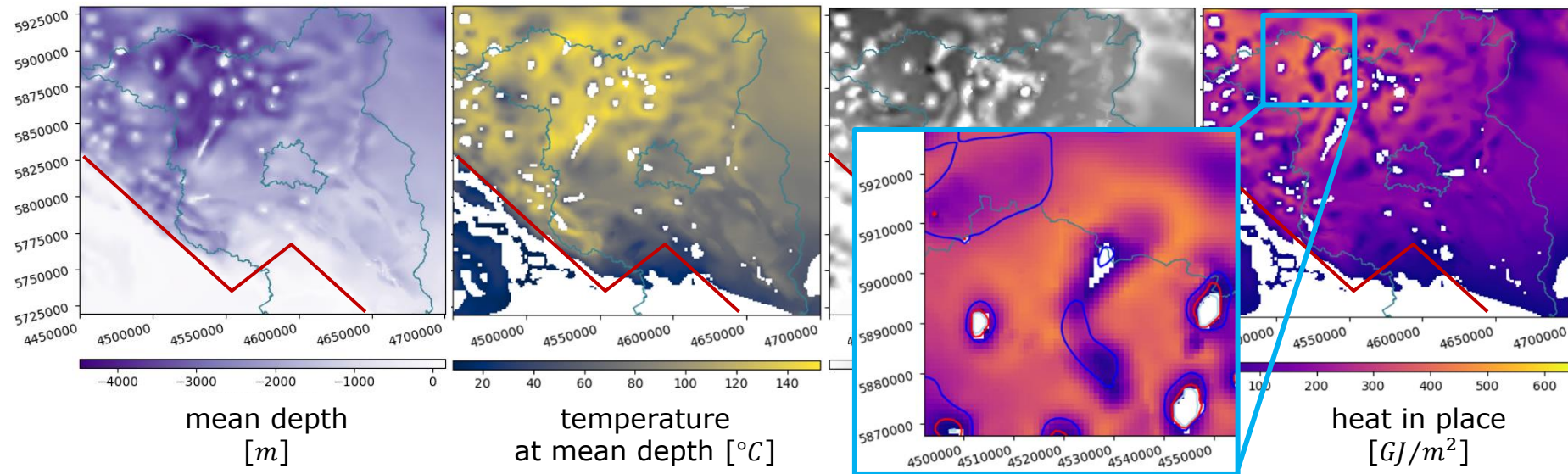


HIP variations due to salt

Buntsandstein

reference temperature model

HIP



- fluctuating pattern due to salt

HIP with salt dome isolines (1800, 3000, 4000 m thickness)



Summary & outlook on HIP

- ✓ Temperature models not expressive enough! Use HIP!
- ✓ Temp & thickness dependency most influential
- ✓ Above-salt lithologies show higher inconsistency
- ✓ For drilling:
 - ✓ data-consistent geological structure, high-resolution models
- ✓ Master-thesis outlook:
 - ✓ parameter analysis
 - ✓ conductive vs. coupled thermal-hydraulic HIP



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and judging option

GitLab

Basin Modelling / Python / Py4HIP

Py4HIP: Python script for heat-in-place calculations



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Programming languages/tools:

Python: <https://www.python.org/>

Jupyter Notebook: <https://jupyter.org/>

