

Dual-domain modeling of discharge dynamics in a laboratory-scale fractured porous matrix system

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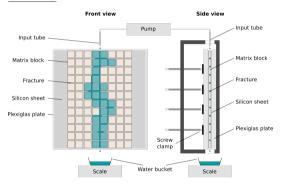


 \rightarrow Fractured porous media contribute to \sim 75% of global aquifers (Dietrich et al., 2005)

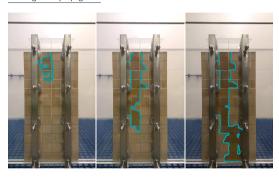
- Vadose zone, above groundwater table (variably saturated)
 - · Fractures and fracture networks
 - Matrix
 - → High contrasts in hydraulic conductivities
- → Water migrates on a wide range of time and spatial scales due to their homogeneous natures (Pruess, 1998)
- This study
 - Simplified analog infiltration experiments at lab-scale
 - · Modeling recharge dynamics with a dual-domain approach
 - Rüdiger et al. (2022)
- Our aims
 - Process understanding (Appendix A)
 - Fracture flow- and intersection dynamics
 - Fracture-matrix interactions
 - Model validation



General setup

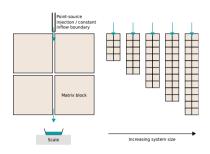


Wetting front propagation



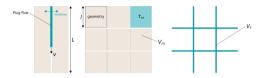


Approach: Simplified laboratory setup and dual-domain modeling



- $m \times 2$ porous blocks network systems
 - Varying total vertical length (m=2, 4, 6, 8, 10, 12)
 - Constant aperture of 1 mm
 - \bullet Constant inflow rate of 1.5 ml \cdot min $^{-1}$
 - Seeberger sandstone
 - Homogenous matrix (18.6 % effective porosity)
- \blacksquare Infiltration experiments \rightarrow measuring discharge
- $\blacksquare \ \ \mathsf{Qualitative} \ \mathsf{observations} \ \to \ \mathsf{Appendix} \ \mathsf{B}$

Dual-porosity model after Neuweiler et al. (2012)

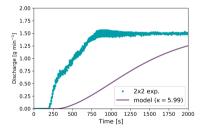


- Memory function ϕ (Equation 3)
 - describes fracture-matrix mass exchange
- Parameters
 - Matrix-fracture volume ratio $\kappa = V_m/V_f$
 - ullet Characteristic imbibition time au_m
 - Fracture flow velocity $v = Q_0/aW$
- Model assumptions
 - Plug-flow in fractures
 - Perfectly coupled domains
- Appendix C and D



Original parameterization

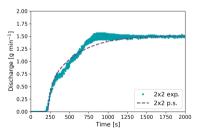
- $au_m = 600 \text{ s}$
- lacksquare v o experimentally determined L/t_1
- lacksquare κ derived from setup geometry (= 5.99)



■ Strong deviations of first arrival and dispersion of outflow signal

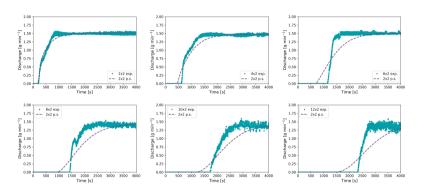
Calibration

- $\tau_m = 600 \text{ s}$
- $\mathbf{v} \to \text{calibrated } (\mathbf{v}^*)$
- $\kappa \to \text{calibrated } (\kappa^*)$



■ Calibrated model recovers discharge dynamics very well

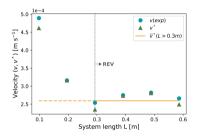




■ Deviations increase with increasing system length (L)

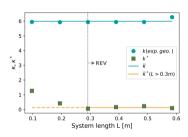


Velocity



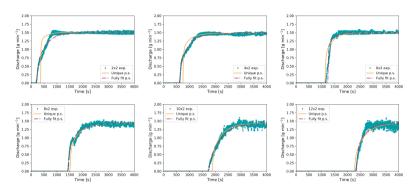
- Observed v and calibrated v^* almost the same
- lacksquare Both converge toward limit for $L \geq 0.3 \ \mathrm{m}$
- Averaged velocity $\overline{v^*} = 2.6e^{-4} \text{ m s}^{-1}$

Domain-coupling parameter



- \blacksquare Calibrated κ^* < geometrical determined κ
- Calibrated converges toward limit for $L \ge 0.3$ m
- Averaged transfer coefficient $\overline{\kappa^*} = 0.1356$
- \blacksquare Scaling factor $\alpha \ = \ \overline{\kappa^*}/\kappa = \text{0.0226}$

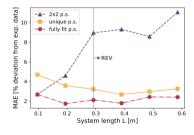




- Fully calibrated fits recover observed discharge dynamics
- Unique parameter set performs better for larger systems



Parameter sets



Unique

- Performs less well for very small systems below REV
- \blacksquare Describes discharge dynamics in sufficiently large systems with high accuracy (MAE <4%)
- \blacksquare Performs almost as good as the model fully calibrated for each L individually

2×2

- Deviation increases with increasing system length *L*
- Strong impact of partitioning dynamics at fracture intersection on discharge dynamics



Thank you for your attention!

If you have any questions feel free to contact me (fruediger@gwdg.de)



References



Dahan, O., Nativ, R., Adar, E. M., Berkowitz, B., & WeisbrodB, N. (2000). On fracture structure and preferential flow in unsaturated chalk. Groundwater, 38(3), 444–451.



Dietrich, P., Helmig, R., Hötzl, H., Sauter, M., Köngeter, J., & Teutsch, G. (2005). Flow and transport in fractured porous media. Springer Science & Business Media.



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Nimmo, J. R. (2012). Preferential flow occurs in unsaturated conditions. Hydrological Processes, 26(5), 786-789.



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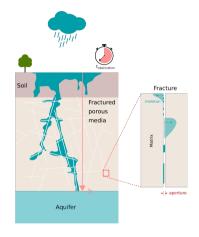
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Shigorina, E., Rüdiger, F., Tartakovsky, A. M., Sauter, M., & Kordilla, J. (2021). Multiscale smoothed particle hydrodynamics model development for simulating preferential flow dynamics in fractured porous media. Water Resources Research, 57(3), e2020WR027323.



- Preferential flow and recharge dynamics as function of
 - Input conditions (infiltration)
 - Magnitude
 - Concentrated vs. diffuse
 - Temporal vs. steady
 - Fracture network properties
 - Density
 - Aperture
 - Topology (intersections, connectivity)
 - Fracture-matrix interactions
 - Imbibition
 - Fracture flow regimes
 - Discrete slugs
 - Films
- Flow path formation (volume or interface area)
- → Observed arrival times contradict common notion that flow in fractures only occur under equilibrium conditions (Dahan et al., 2000)
- → No ubiquitous classification for the onset of preferential flow exists (Nimmo, 2012)

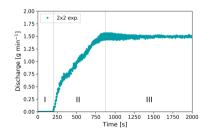




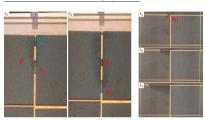
Appendix B: Qualitative observations

Example: 2 x 2 blocks network

- Three main phases (I, II, III)
 - I Redistribution water through fracture network and matrix imbibition
 - II Onset of discharge, ongoing matrix imbibition
 - III Quasi-steady state, matrix reached maximum saturation



Flow modes and partitioning dynamics

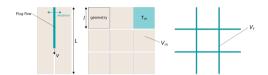


- Two dominant flow modes
 - Slugs (S₁,S₂)
 - Films (F)
- Matrix imbibition (im) slows down fracture flow progression
- Slugs (S_2) and films (F) above the wetting front move extremely fast
- Intersections (partitioning dynamics)
 - Horizontal infiltration (hi)
 - By-passing



Appendix C: Dual-domain modeling

Dual-porosity model after Neuweiler et al. (2012)



- Equation 1 analytically solved
 - Initial boundary condition
 - Solution in Equation 5
- Memory function ϕ (Equation 3)
 - describes fracture-matrix mass exchange
- Parameters
 - ullet Matrix-fracture volume ratio $\kappa = V_m/V_f$
 - \bullet Characteristic imbibition time τ_m (Equation 4)
 - Fracture flow velocity $v = Q_0/aW$
- Model assumptions
 - Plug-flow in fractures
 - Perfectly coupled domains

Main equations

$$\frac{\delta S_f(z,t)}{\delta t} + \kappa \frac{\delta S_m(z,t)}{\delta t} - v \frac{\delta S_f(z,t)}{\delta z} = 0$$
 (1)

$$S_m = \int_0^t dt' \phi(t - t') S_f(t')$$
 (2)

$$\phi^*(\lambda) = \frac{1}{\sqrt{\lambda \tau_m}} \tanh\left(\sqrt{\lambda \tau_m}\right) \tag{3}$$

$$\tau_m = \frac{l^2}{D} \tag{4}$$

$$j^*(\lambda) = Q_0 \cdot \lambda^{-1} \cdot \exp[-La\lambda(1 + \kappa \phi^*)/v]$$
 (5)

 S_f, S_m Saturation fracture, -matrix

 V_f, V_m Volume fracture, -matrix

D Diffusion coefficient

/ Length scale characteristic geometry

a Fracture aperture

W Fracture width

 $j^*(\lambda)$ Volumetric outflow rate (Laplace space)

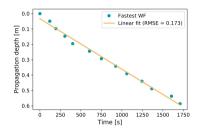
Q0 Volumetric inflow rate

L Vertical system length



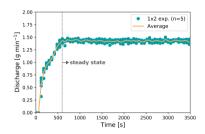
Appendix D: Preliminary investigations - Flow progression and time-scale of fracture-matrix interaction

12 x 6 blocks network



- Infiltration front into initially air-dry system progresses with constant velocity
- Supports plug-flow assumption
- $v \approx 3.28e^{-4} \text{ m s}^{-1}$

1 x 2 blocks network



- Characteristic imbibition time $\tau_m(exp.)$ ≈ 600 s
- $\tau_m(ana.) = l^2/D \approx 147 \text{ s}$
- / Block length = 0.04785 m
- D Diffusion coefficient = $1.56e^{-5}$ m s⁻¹
 - Based on parameter estimation (Shigorina et al., 2021)
 - Calculated with equation 20 (Neuweiler et al., 2012)
 - Taken for water content = 0.9



Outlook: Field work plans

- Difficult to set up experiments
 - Stability criteria
 - Study suitability of fracture network
- New potential sites: Ossenfeld and Vogelbeck
- Study analog to lab experiments



Ossenfeld



Vogelbeck

