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# Are winds and moisture necessary to cause Indian summer monsoon extremes?

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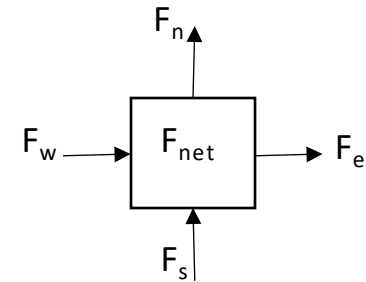
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# Moisture budget theory

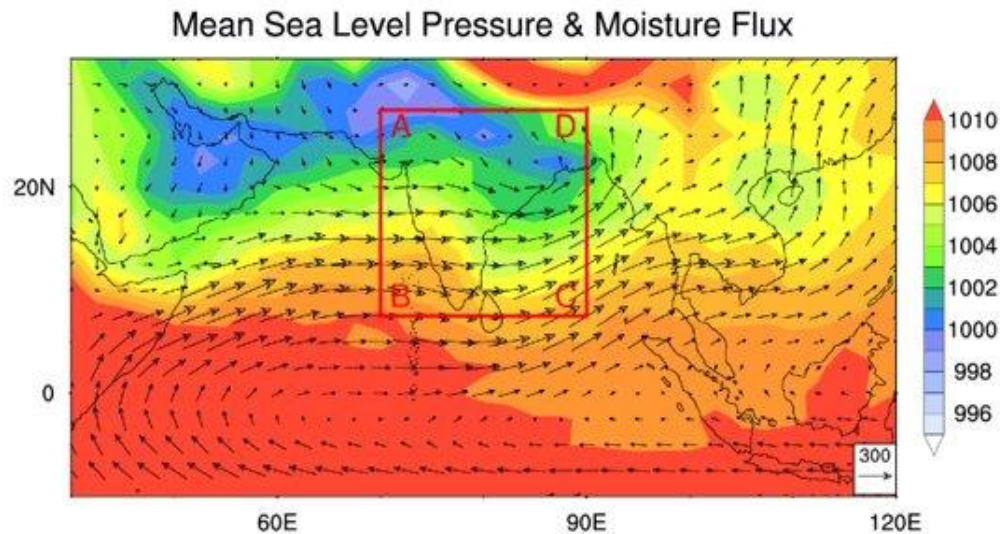
$$\frac{1}{A}(F_w - F_e + F_s - F_n) = \bar{P} - \bar{E} + \frac{\partial P_{wat}}{\partial t},$$

$$\frac{1}{A}(F_{net}) = \bar{P} - \bar{E} + \frac{\partial P_{wat}}{\partial t}$$

Vertically integrated moisture fluxes (F)



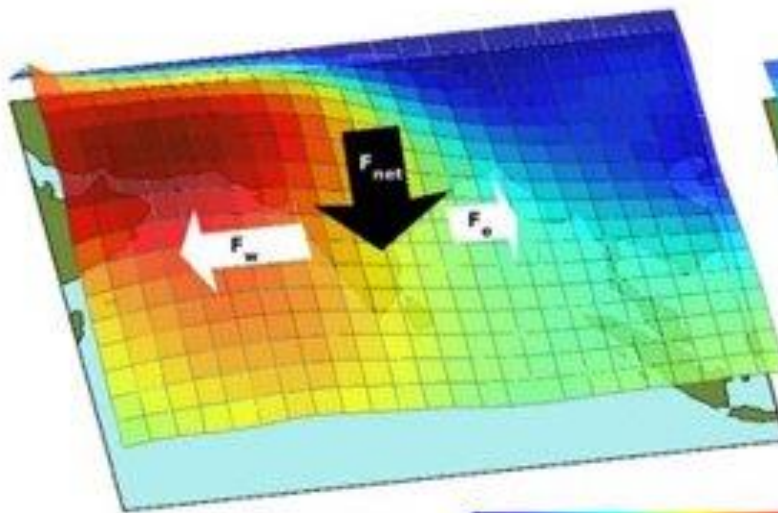
ISMR (P)  $\propto$  Net moisture convergence ( $F_{net}$ )



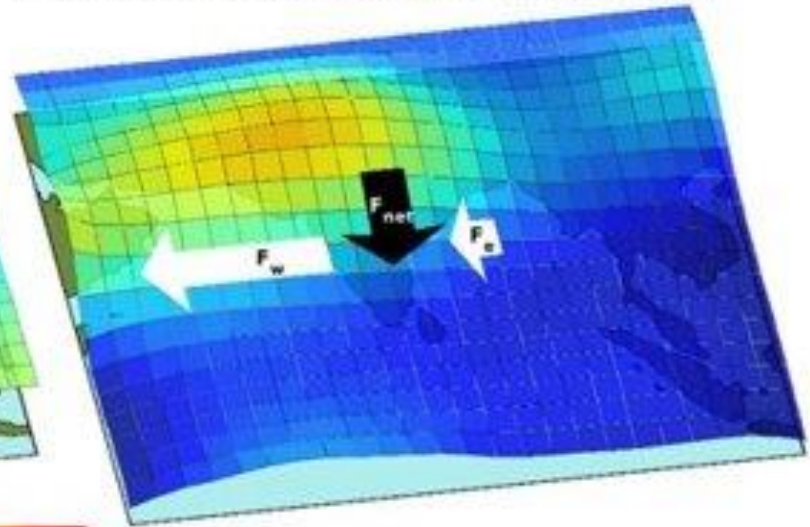
June-September climatological mean sea level pressure along with vertically integrated moisture flux

Extremes:

(a) ENSO : Drought *minus* Flood



(b) Non-ENSO : Drought *minus* Flood



Chakraborty & Singhai, 2021

Schematic showing variations in surface pressure surrounding the Indian monsoon region due to (a) ENSO and (b) Non-ENSO forcing leading to Indian summer monsoon rainfall extremes (drought and flood).

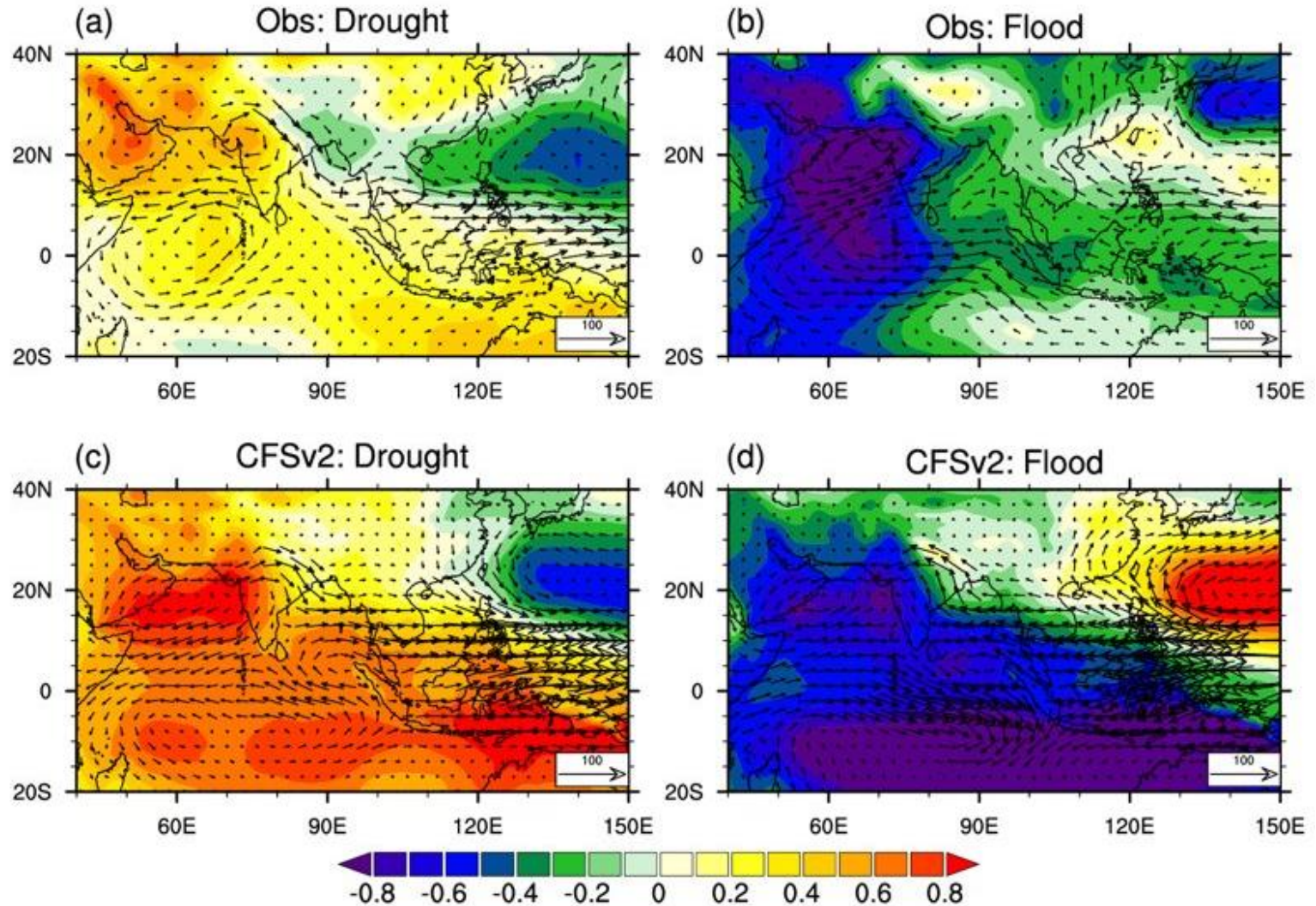
# Data & Model used

- **Datasets used:**
  - Precipitation: IMD  $1^{\circ} \times 1^{\circ}$
  - SST: ER-SST
  - Atmospheric parameters: NCEP/NCAR reanalysis product
- **Period:** 1979-2016 (June-September)
- **Model:** NCEP Climate Forecast System version 2 (CFSv2)
  - Resolution:  $0.91^{\circ} \times 91^{\circ}$  horizontal resolution & 64 vertical levels.
  - Integrated for 9 months
  - 5 Initial conditions: (00UTC of 21 April, 26 April, 1 May, 6 May and 11 May)

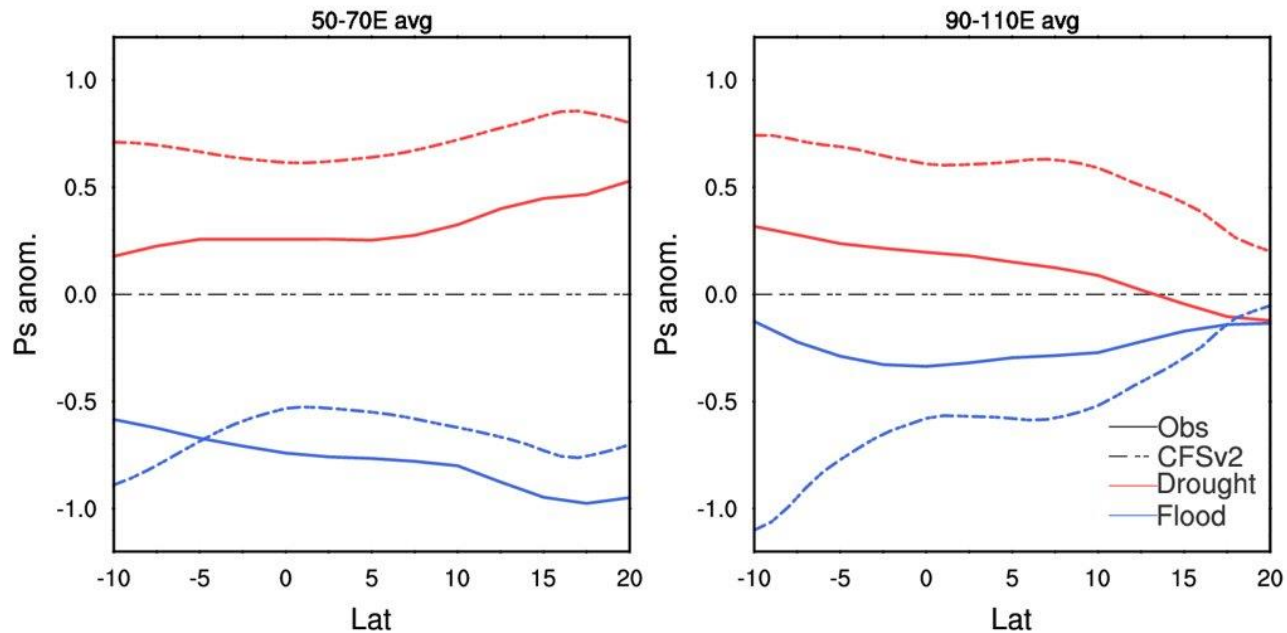


# Drought vs Flood

P<sub>s</sub> anomalies composites:



## Drought & Flood



**Arabian Sea (50-70E):**  $(dPs/dy)_{clm} < 0$

**Drought:**  $(dPs'/dy) > 0$  (weaken)

**Obs:** Weakening of Ps gradient

**CFSv2:** Ps remain unchanged

**Flood:**  $(dPs'/dy) < 0$  (strengthen)

**Obs:** Strengthening of Ps gradient

**CFSv2:** Ps remain unchanged

**Bay of Bengal (90-110E):**  $(dPs/dy)_{clm} < 0$

**Drought:**  $(dPs'/dy) < 0$  (Strength)

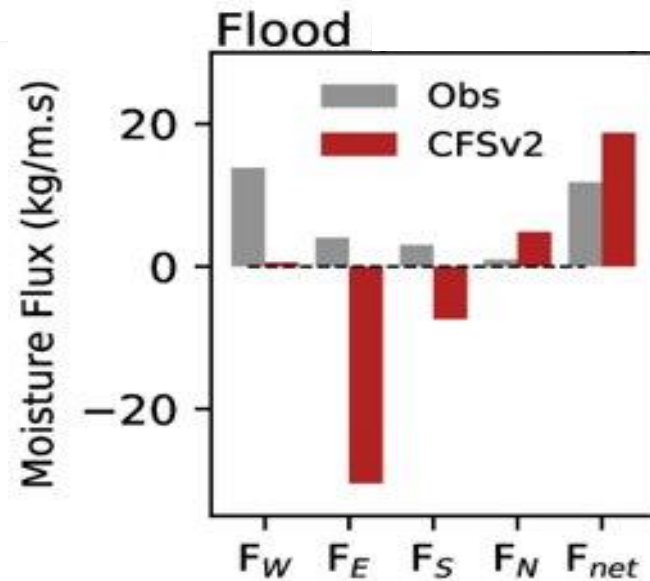
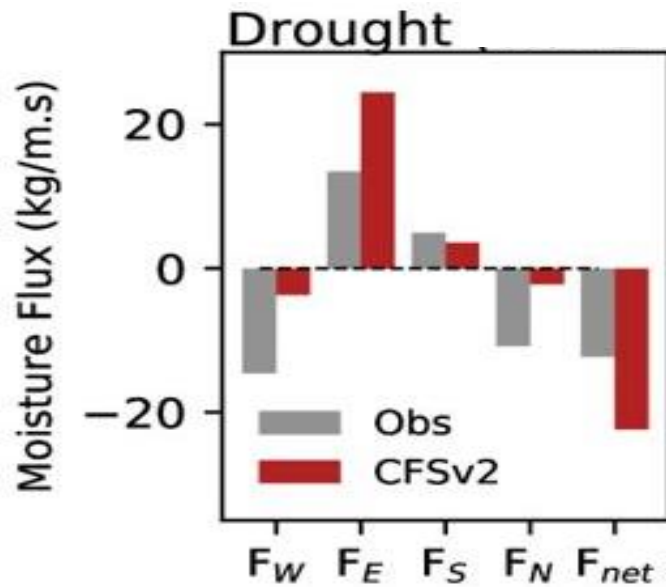
**Obs:** Strengthening of Ps gradient

**CFSv2:** Strengthening of Ps gradient

**Flood:**  $(dPs'/dy) > 0$  (Weaken)

**Obs:** Weakening of Ps gradient

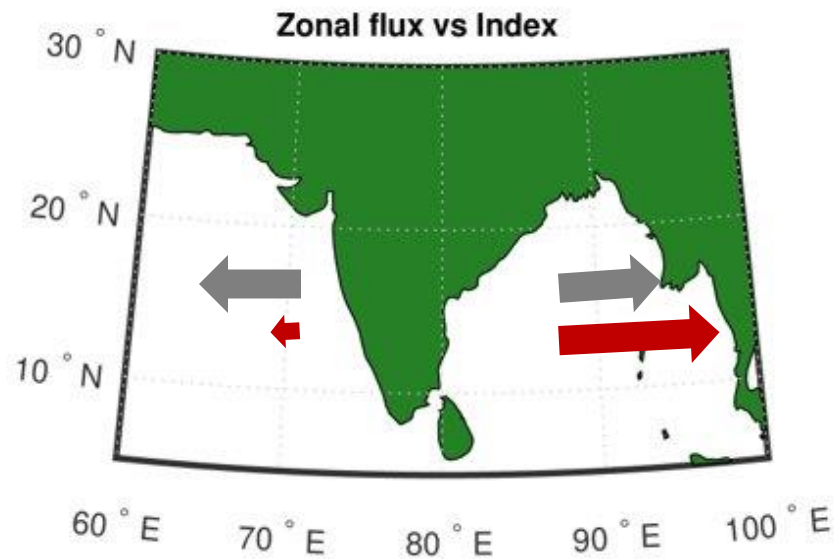
**CFSv2:** Strong weakening of Ps gradient



The corresponding change in moisture flux and its components and precipitation during drought and flood years over the Indian region in observation and model.

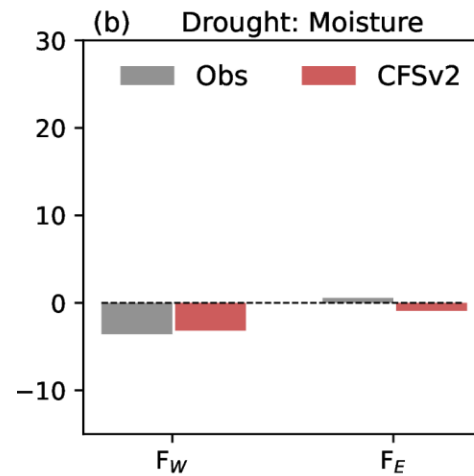
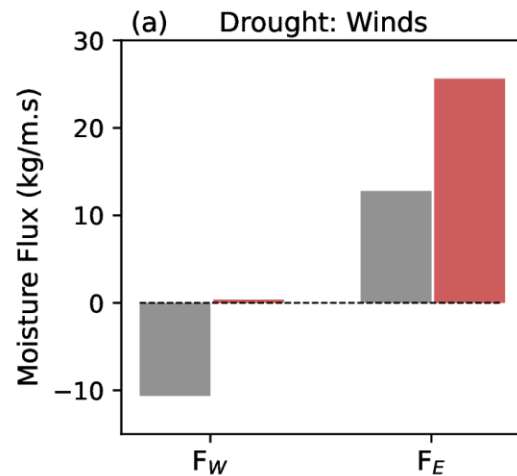
Drought:  
Zonal divergence

Flood:  
Zonal convergence



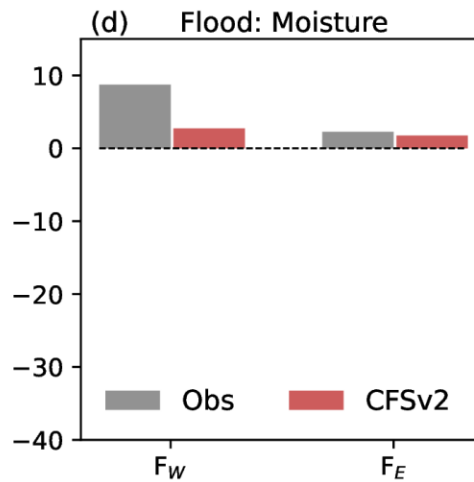
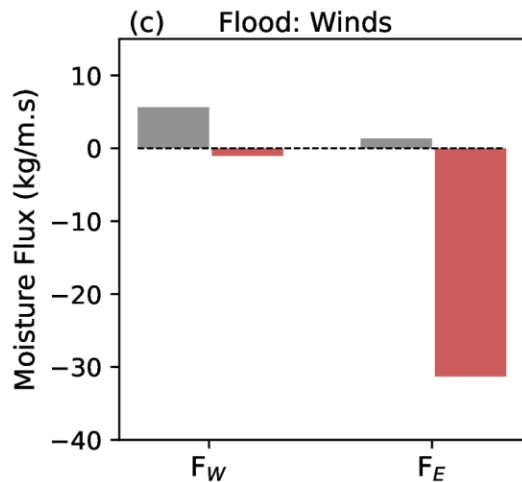
$$\vec{F}' = \langle \vec{V} q \rangle' = \underbrace{\langle \vec{V}' \bar{q} \rangle}_{\text{Wind component}} + \underbrace{\langle \vec{V} q' \rangle}_{\text{Moisture component}} + \underbrace{\langle \vec{V}' q' \rangle}_{\text{Eddy flux component}}$$

Wind component    Moisture component    Eddy flux component



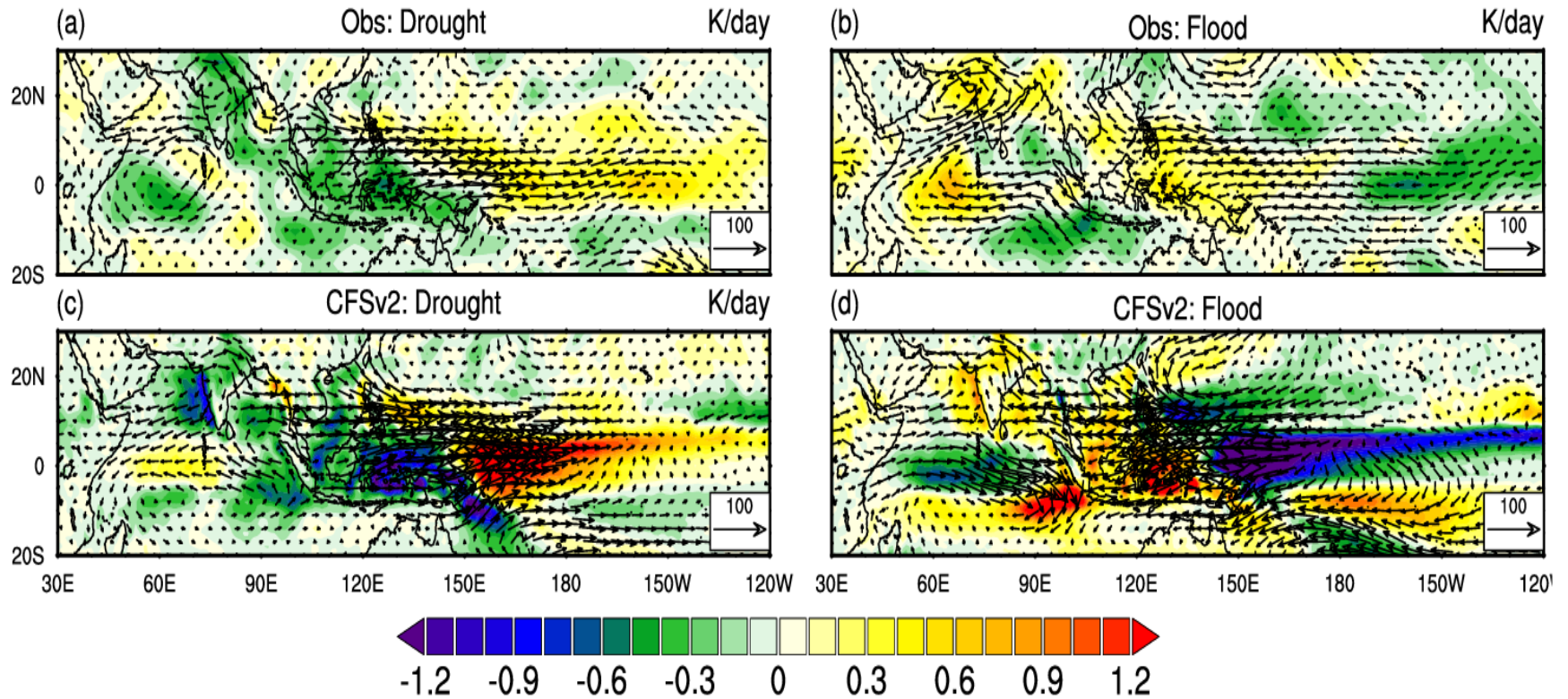
**Obs:** Winds and moisture components of  $F_W$  both play an important role

**CFSv2:** Winds component over the Bay of Bengal dominates.





## Q1: Extreme



Composite change in the anomalous heating rate during drought and flood for observation and model.

# Conclusions & Significance

- Outgoing eastward moisture flux over the Bay of Bengal is the only important factor determining ISMR extremes in the CFSv2, in contrast to observation where both outgoing and incoming eastward fluxes play a major role.
- Characterized the nature of moisture fluxes during extreme events.

## **Significance:**

- Open new avenue for the improvement of general circulation model deployed for seasonal prediction.
- Teleconnection patterns in changing climate scenarios.

$$\text{VMS} = \text{MSE}_{\{500-100 \text{ hPa}\}} - \text{MSE}_{\{\text{Psfc}-500 \text{ hPa}\}}$$

$$\text{VMS} \propto 1/\text{Prec}$$

