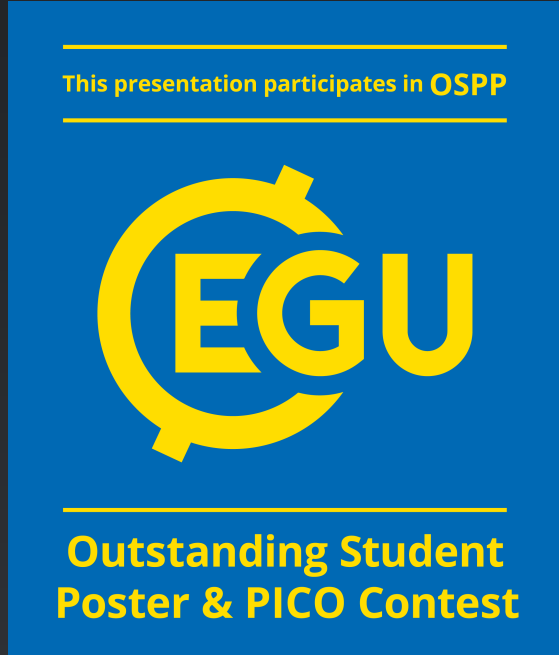
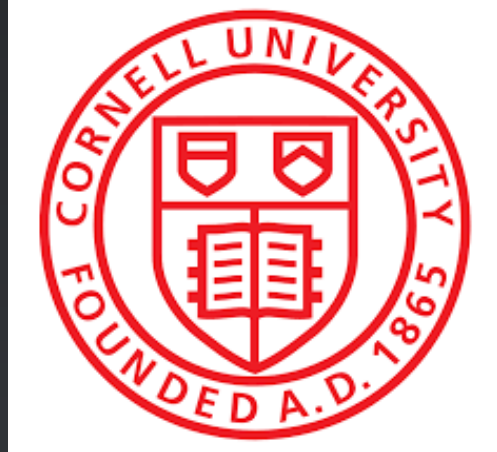




école  
normale  
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# Formation of Jupiter's envelope from supersolar gas in the protoplanetary disk

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## Abstract:

The formation mechanism of Jupiter is still uncertain. Multiple volatile accretion scenarios can reproduce its metallicity. We use a viscous disk accretion model to compute the composition of the protosolar nebula (PSN), and compare it with the recent measurement of O abundance in Jupiter's envelope by Juno. We find that the composition of the gas phase of the PSN is compatible with that of Jupiter's envelope, from which it possibly formed.

# Model overview

Our protosolar nebula (PSN) model:

- 1D  $\alpha$ -viscous disk of H<sub>2</sub>-He;
- Midplane temperature budget: viscous heating, background radiation;
- Trace species are in the form of **pebbles** ( $\mu\text{m}$  to m size) or **vapor**, radial transport due to advection and diffusion;
- With radial motion, trace specie may **sublimate** or **condense**, leading to the formation of **icelines**.

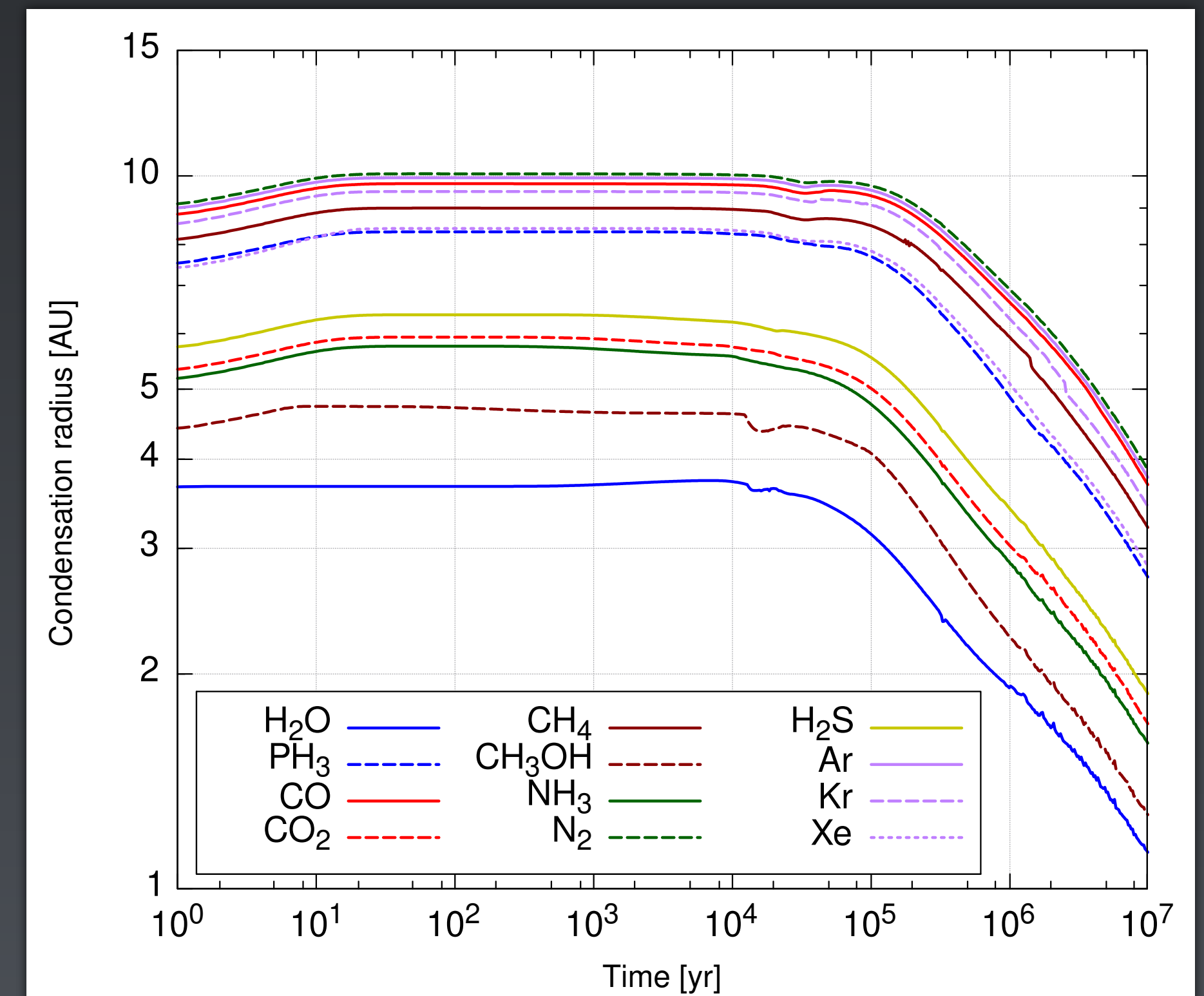


Figure 1: Position of **icelines** as a function of time during the evolution of the PSN.

# Enrichment profiles

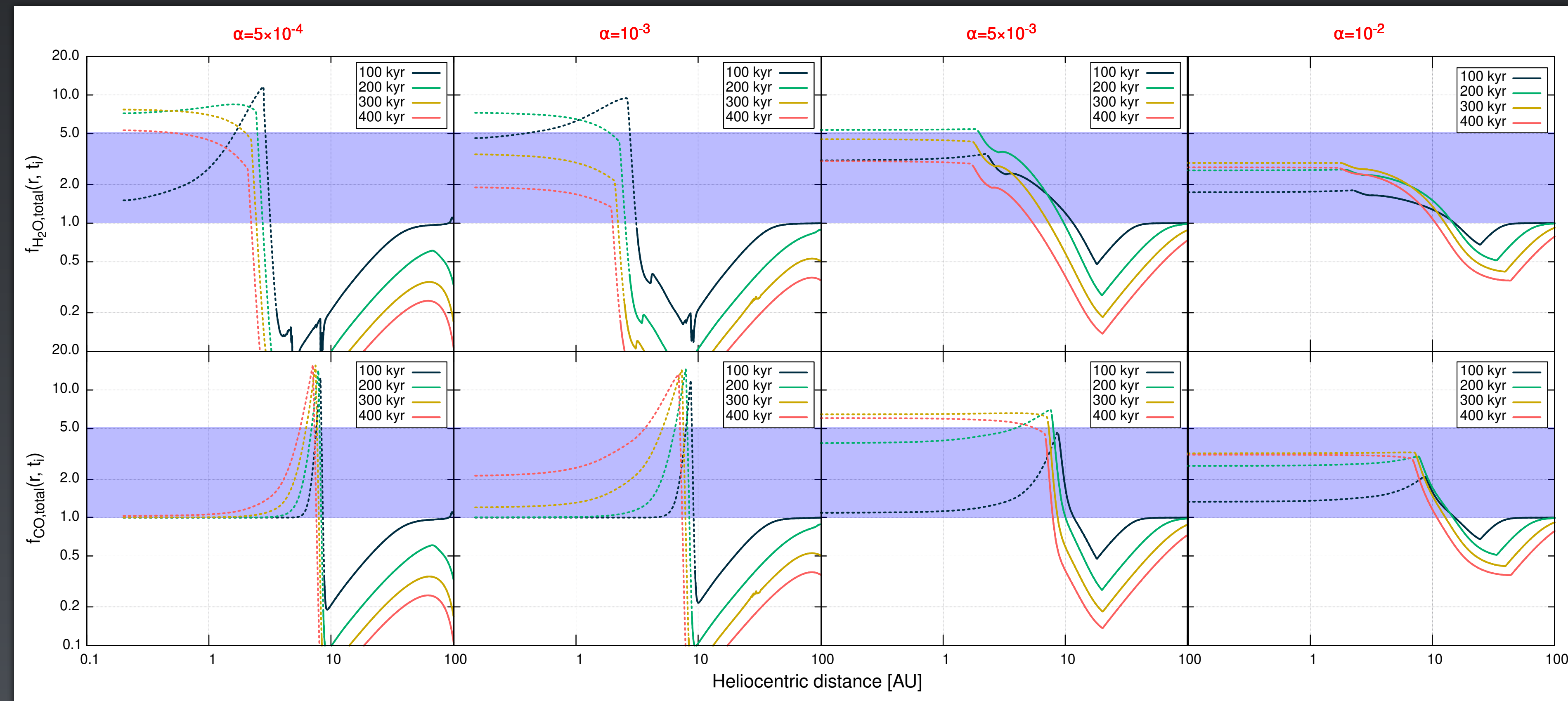


Figure 2: **H<sub>2</sub>O** (top) and **CO** (bottom) abundance normalized to protosolar.

The **viscosity parameter  $\alpha$**  is indicated above.

**Solid or dashed** lines correspond to **solids or vapors** (resp.).

Blue rectangles correspond to 1- $\sigma$  abundance of O in Jupiter's envelope (Li et al. 2020).

Pebble inward drift: depletion of the disk in solids, efficient delivery to icelines.  
Vapor diffusion: supersolar abundances, uniform enrichments.



# Elemental abundances

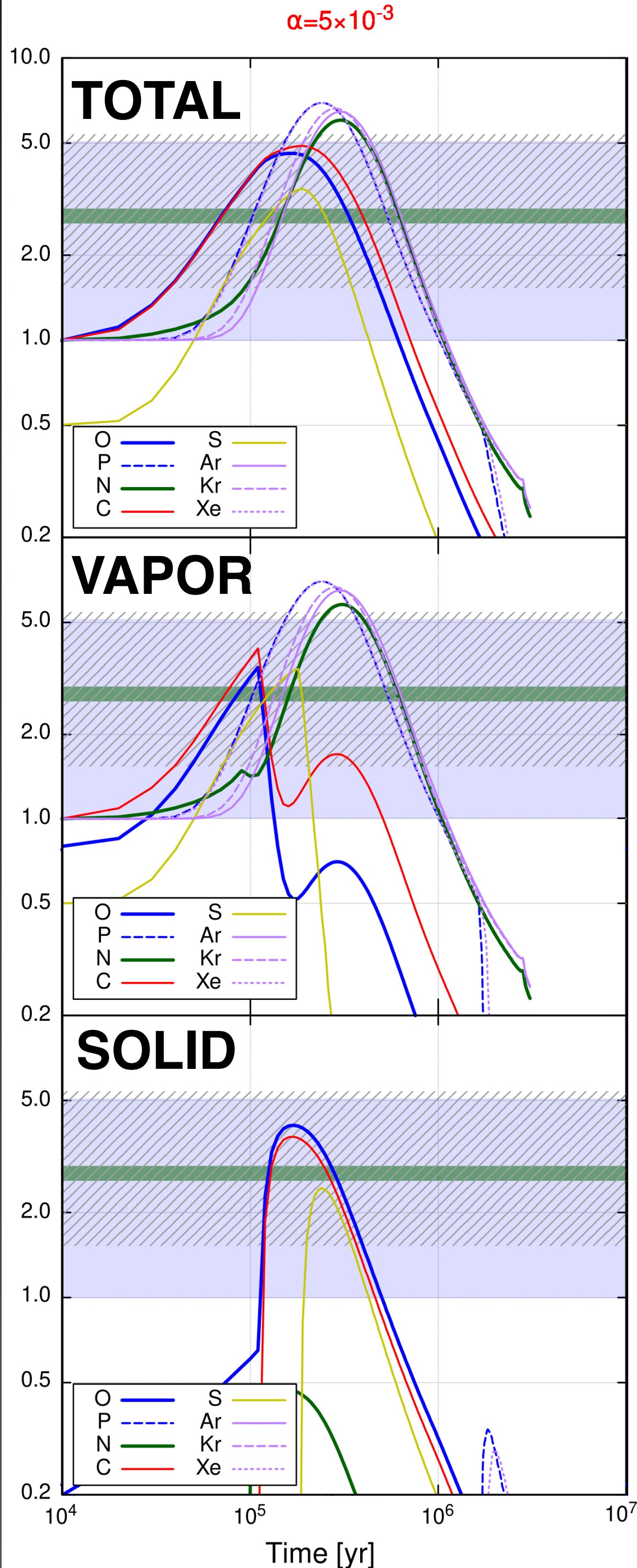


Figure 3: PSN to protosolar elemental abundances at 4 AU.  
Top to bottom : total, vapor, and solid phases.  
Blue, green, shaded regions: O, N, and volatile abundances in Jupiter.

- Elemental abundances at 4 AU: between icelines of H<sub>2</sub>O and CO<sub>2</sub>
- CO<sub>2</sub>/CO = 3 as in 67P/C-G
- CO<sub>2</sub> moves to 4 AU at 100 kyr
- **PSN = Jupiter** at 100–300 kyr, from vapors or vapors+solids, depending on  $\alpha$

First stage: fast delivery of solids to icelines, increase of the enrichment in vapor phase.  
Second stage: depletion of the PSN in trace species due to accretion.

# Elemental abundances

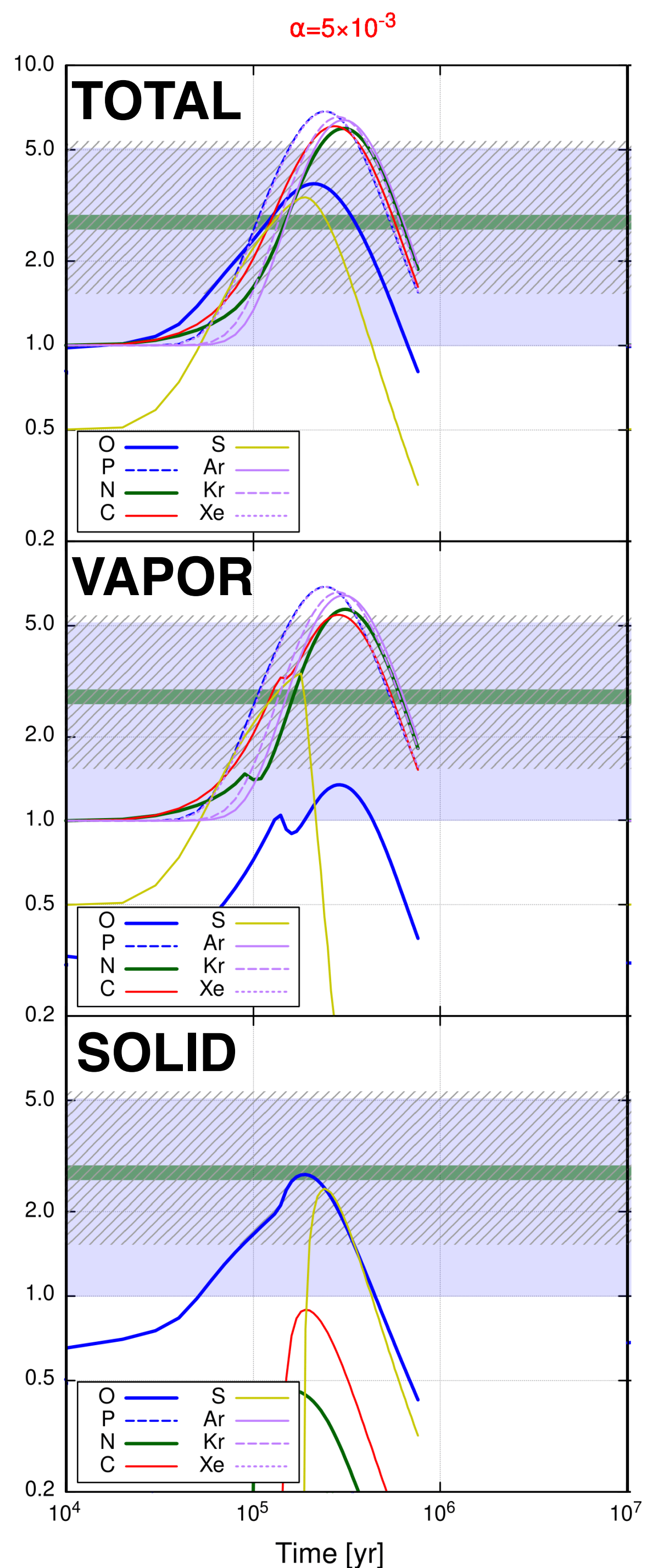


Figure 3: PSN to protosolar elemental abundances at 4 AU.

Top to bottom : total, vapor, and solid phases.

Blue, green, shaded regions: O, N, and volatile abundances in Jupiter.

- Elemental abundances at 4 AU: between icelines of H<sub>2</sub>O and CO<sub>2</sub>
- CO<sub>2</sub>/CO = 0
- **PSN = Jupiter** possible at ~1 Myr with **subsolar O**.



# Formation mechanism

If O in Jupiter is **supersolar**

1- $\sigma$  Juno measurement ([Li et al. 2020](#)).

Possible formation scenario:

- **Gravitational instability**
- 3.5 — 5.5 AU
- 100 — 300 kyr
- Accretion of gas enriched in O

If O in Jupiter is **subsolar**

Galileo measurement ([Atreya et al. 2003](#)),  
or 2- $\sigma$  Juno measurement ([Li et al. 2020](#)).

Possible formation scenario:

- **Core or pebble accretion**
- 3.5 — 10 AU
- ~1 Myr
- Accretion of gas depleted in O

# Conclusions

- Volatile content accreted from the vapor phase during formation
- Origin of volatiles constrain the formation location and timescale
- Possibly reconsider:  $\text{CO}_2/\text{CO}$  ratio, dust dynamics
- Similar results for Saturn, but not ice giants

Detailed results and discussion: Aguichine A., Mousis O., Lunine I. J. 2022, PSJ.

Amorphous ice and clathrates: next talk by Schneeberger A.