Deciphering Neoarchean polymetamorphism and crustal melting in the northern Wyoming Province using garnet petrochronology

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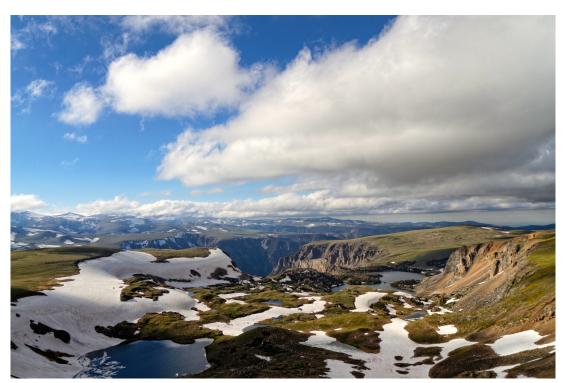
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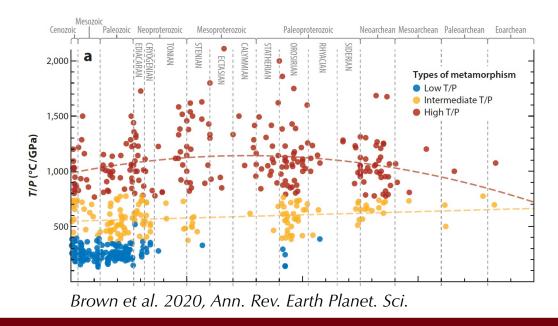
Introduction.

Constraints on depths, temperatures and rates of Archean metamorphism (esp. partial melting of crust)

Archean tectonic modes

Generation of continental crust

Craton stabilization



Archean terranes often polymetamorphic Unraveling P-T-t challenging

Complexities in both petrologic and geochronologic data



Introduction.

Garnet (and monazite) petrochronology

Thermodynamic modeling

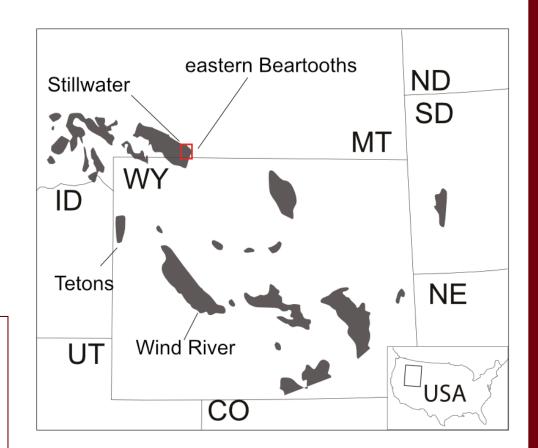
Diffusion "

1-D thermal "

Timing, timescales and mechanisms of HT metamorphism

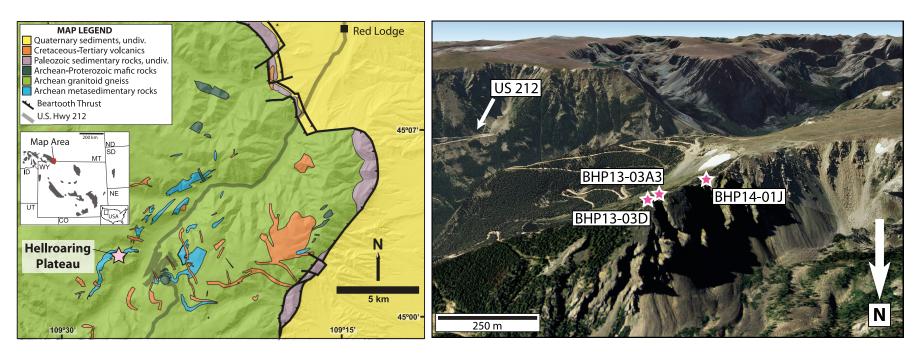
Eastern Beartooth Mountains of Montana and Wyoming, U.S.A.

*Two phases of HT metamorphism likely driven by magmatic heat advection





Eastern Beartooth Mountains.



Dominated by 2.83-2.79 Ga TTGs of the Long Lake Magmatic Complex (LLMC)

Roof pendants/xenoliths of supracrustals

Metasedimentary granulites that have been interpreted to form solely from LLMC contact heating



Petrochronology approach.

- 1) Zoned (3 samples) and bulk (2 samples) Sm-Nd garnet geochronology*
- 2) Garnet TE zoning using LA-ICP-MS (spot analyses and mapping)

Elucidates reaction history

Informs microsampling

Validates isotope dilution work; highlights deleterious role of inclusions

3) Garnet ME diffusion modeling

Timescales at peak metamorphic conditions

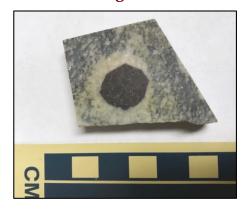
4) U-Pb monazite petrochronology**

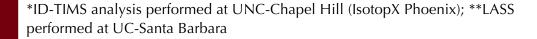




Felsic gneiss

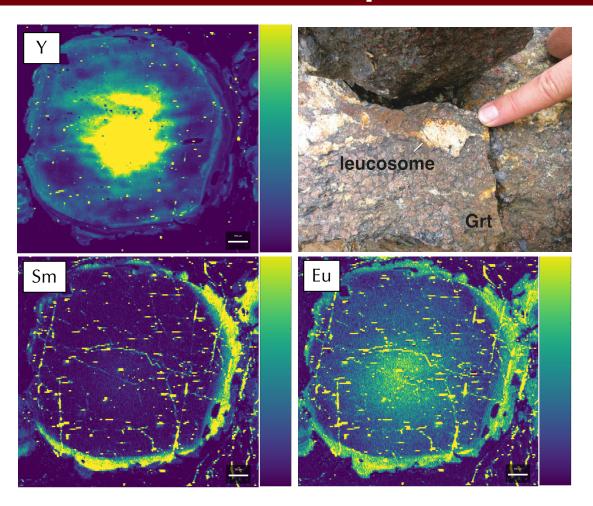
Restitic metapelites

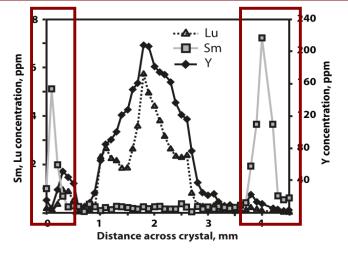






BHP14-01J – restitic metapelite.



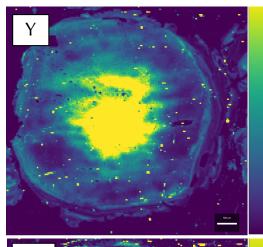


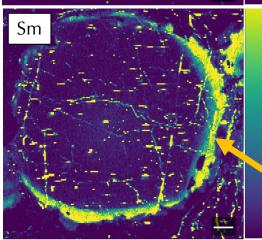
Two garnet growth stages

2nd stage shows garnet resorption and regrowth accompanied by accessory phase breakdown (?) and partial melting



BHP14-01J – restitic metapelite.





Core-WR:

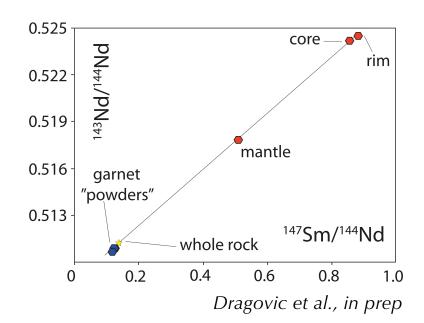
 $2761.5 \pm 5.5 Ma$

Mantle-WR:

 $2735.3 \pm 9.6 \text{ Ma}$

Rim-WR:

 $2726.1 \pm 5.6 \text{ Ma}$



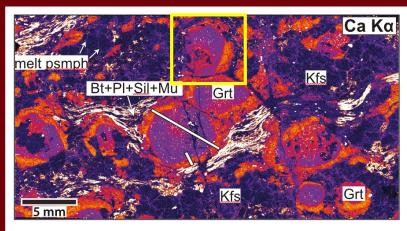
Two garnet growth stages

Gt mantle date likely mixing of two stages

Rim likely dates 2nd stage of (granulite-facies) metamorphism



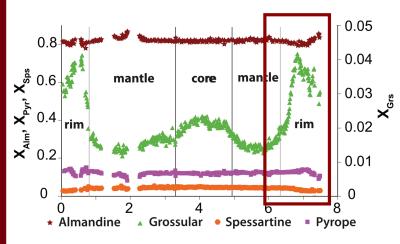
BHP13-03D – restitic metapelite.



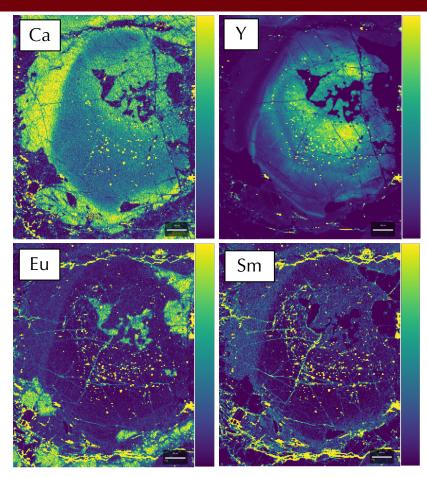
Two garnet growth stages

High Ca rim – biotite breakdown melting

Diffusion modeling of ME suggests duration at near peak T (750-800°C) of < 1 M.y.*



*from Guevara et al. 2017, JMG





BHP13-03D – restitic metapelite.



Gt1 core:

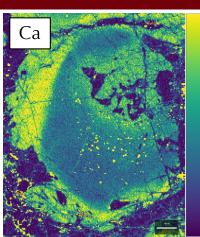
 $2759.3 \pm 6.4 \text{ Ma}$

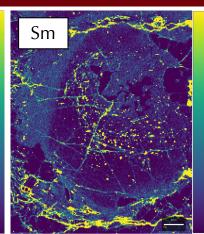
Gt1 mantle:

 $2750.1 \pm 4.3 \text{ Ma}$

Gt1 rim:

2709.2 ± 5.1 Ma





0.535 mantle1_ 43Nd/144Nd rim2 0.530 rim1 0.525 core1 0.520 whole mantle2 rock 0.515 -¹⁴⁷Sm/¹⁴⁴Nd core2 garnet "powders" 0.6 1.5 0.3 1.2 0.9

Dragovic et al., in prep

Gt2 core:

2664.4 ± 24.8 Ma

Gt2 mantle:

 $2700.0 \pm 9.2 \text{ Ma}$

Gt2 rim:

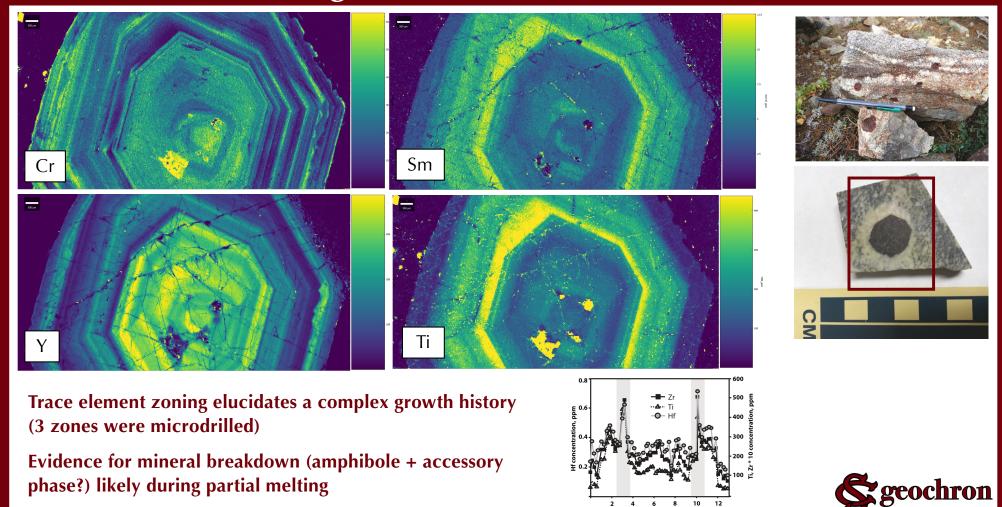
2715.1 ± 4.3 Ma

Two garnet growth stages

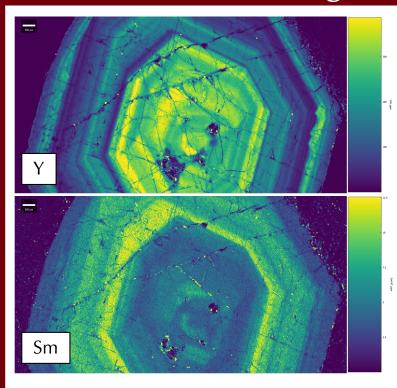
*Gt2 core and mantle – likely affected by remnant monazite inclusions – confirmed by modeling the effects of inclusion contamination



BHP13-03A3 – felsic gneiss.



BHP13-03A3 – felsic gneiss.



Core-matrix:

 $2784.8 \pm 3.4 \text{ Ma}$

Mantle-matrix:

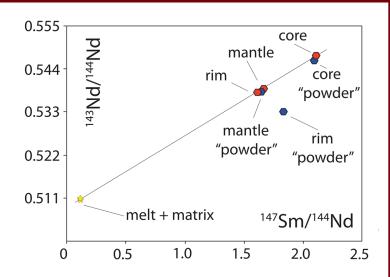
2753.6 ± 3.6 Ma*

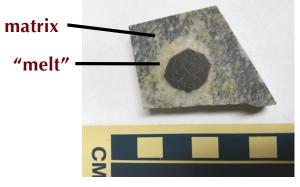
Rim-matrix:

 $2780.6 \pm 3.7 \text{ Ma}$

Rim-"melt":

 $2777.9 \pm 3.7 \text{ Ma}$





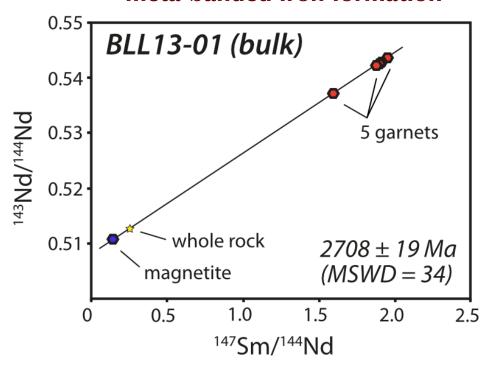


Trace element zoning elucidates a complex growth history

*Evidence for accessory phase breakdown and temporary isotopic disequilibrium (?) between garnet and surroundings during mantle growth

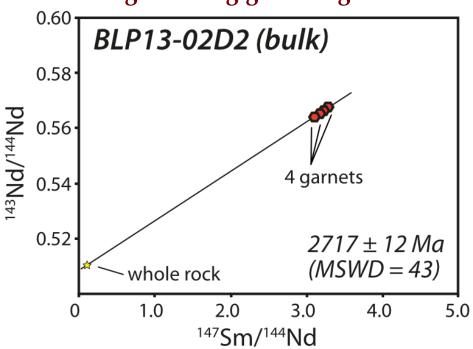
Additional bulk garnet Sm-Nd.

meta-banded iron formation



Dragovic et al., in prep

gt-bearing granitic gneiss





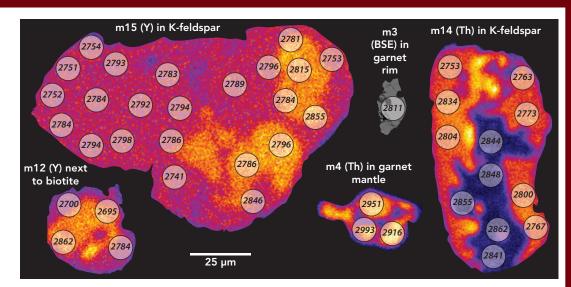
U-Pb monazite.

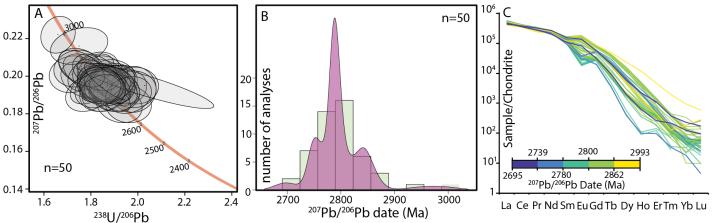
Dominant population @ ~2780 Ma

TE shows growth in the presence of garnet

Minor population @ ~2700 Ma

Relative enrichment in HREE suggests growth during gt breakdown

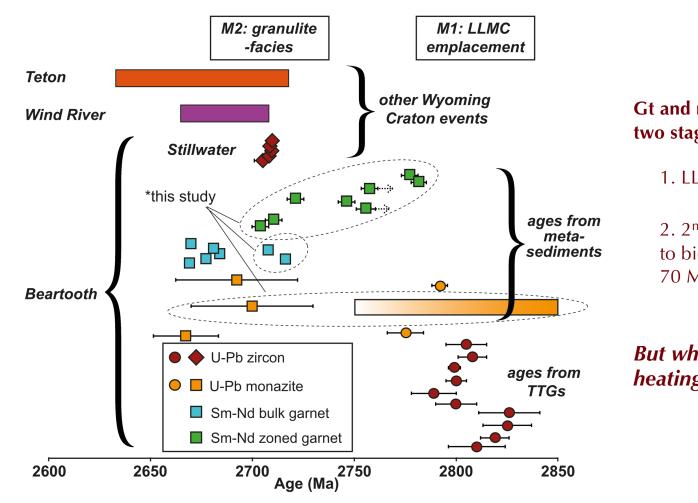




Dragovic et al., in prep



Age summary.



Gt and mnz show correspondence to two stages of HT metamorphism

- 1. LLMC emplacement
- 2. 2nd metamorphic, related to biotite breakdown melting, 70 Myrs later....

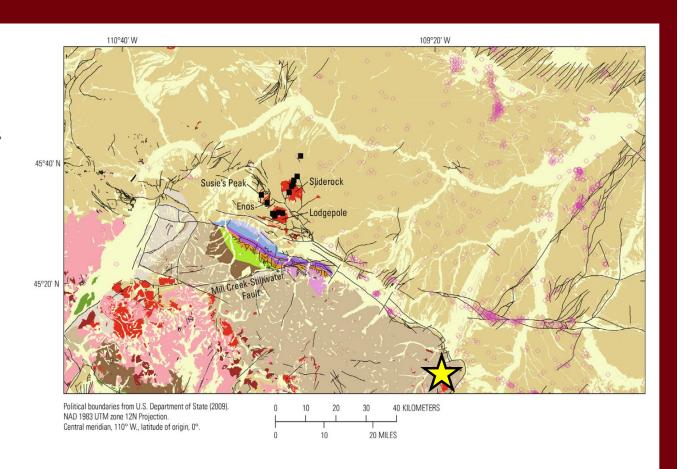
But what could have driven heating of the crust at this time?



Stillwater Complex?

Recent geophysical modelling by Finn et al. (2020) shows that subsurface extent of Stillwater Complex is ~10 times greater than is exposed, dipping ~25-30° to the north.

If Stillwater extended south with the same geometry, it would have been ~15 km above the current exposure of the eastern Beartooth granulites.





1-D thermal modeling.

Does a model for heat advection from the Stillwater recreate P-T-t estimates?

Stillwater thickness estimates range from 7-12 km

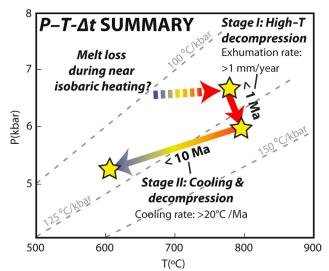
Model: 10 km

Depth of the base of Stillwater, based on thermobarometry of aureole (Thomson, 2008) – **10 km**

Stillwater T – 1250°C

Radiogenic heat production:

 $Q_{rad} = 2 \mu W/m^3$



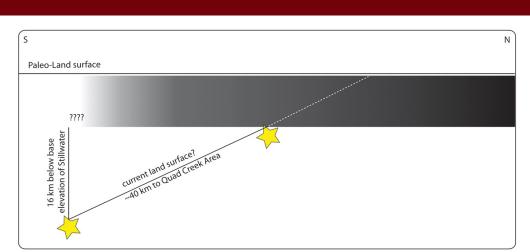
Granulite-facies *P-T* path

(thermodynamic modeling, mineral chemistry and petrographic observations)

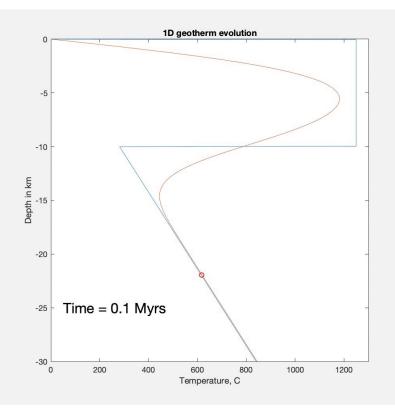
Timescales at peak *T* and during the initial stage of cooling (diffusion modeling)

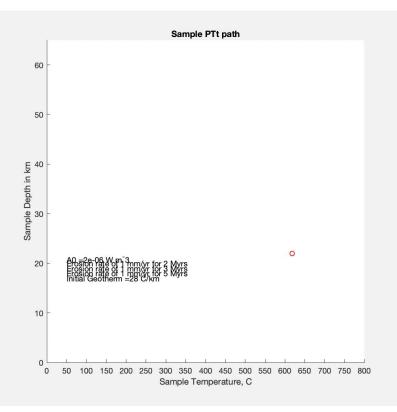
Sgeochron

*based from Guevara et al. 2017, JMG



28°C/km; 1 mm/yr exhumation.

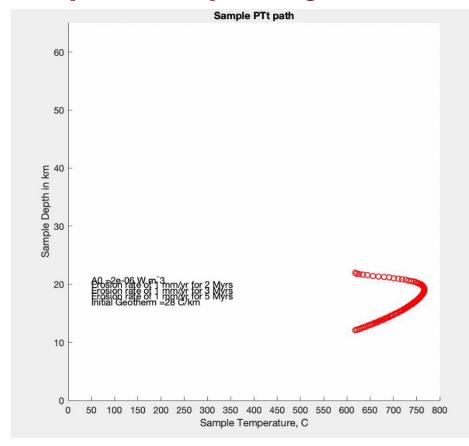


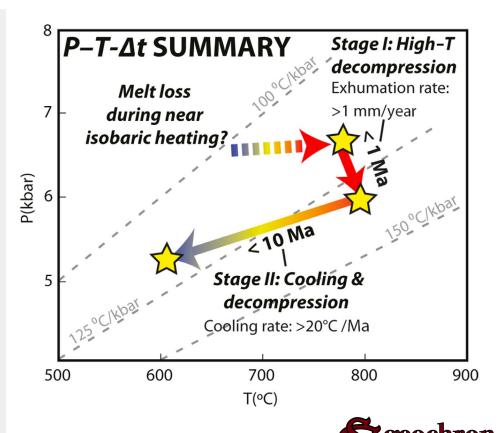




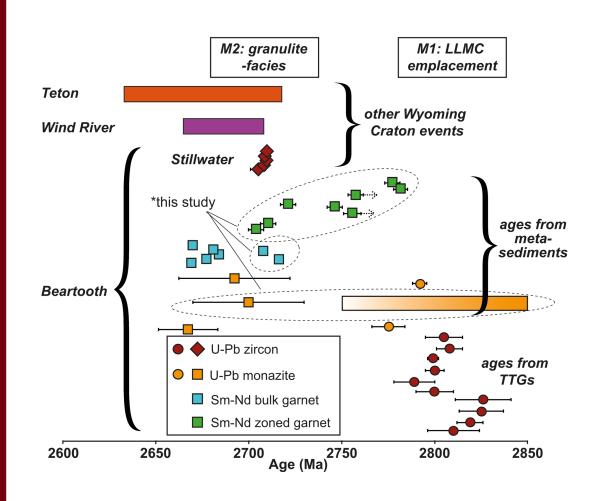
28°C/km; 1 mm/yr exhumation.

Comparison to petrologic constraints





Conclusions.



Two stages of metamorphism, though cryptic or absent in some lithologies

The extent and significance of this second event for craton assembly and stabilization thus remain unclear, despite its synchronicity with metamorphism and magmatism in other portions of the Wyoming Province.

High-temperature metamorphic terranes still remain our clearest window into the tectonic mechanisms operating in the Archean.



Thank you!

Questions?



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