



Effects of inclusion of adjoint sea ice rheology on estimating Arctic ocean-sea ice state

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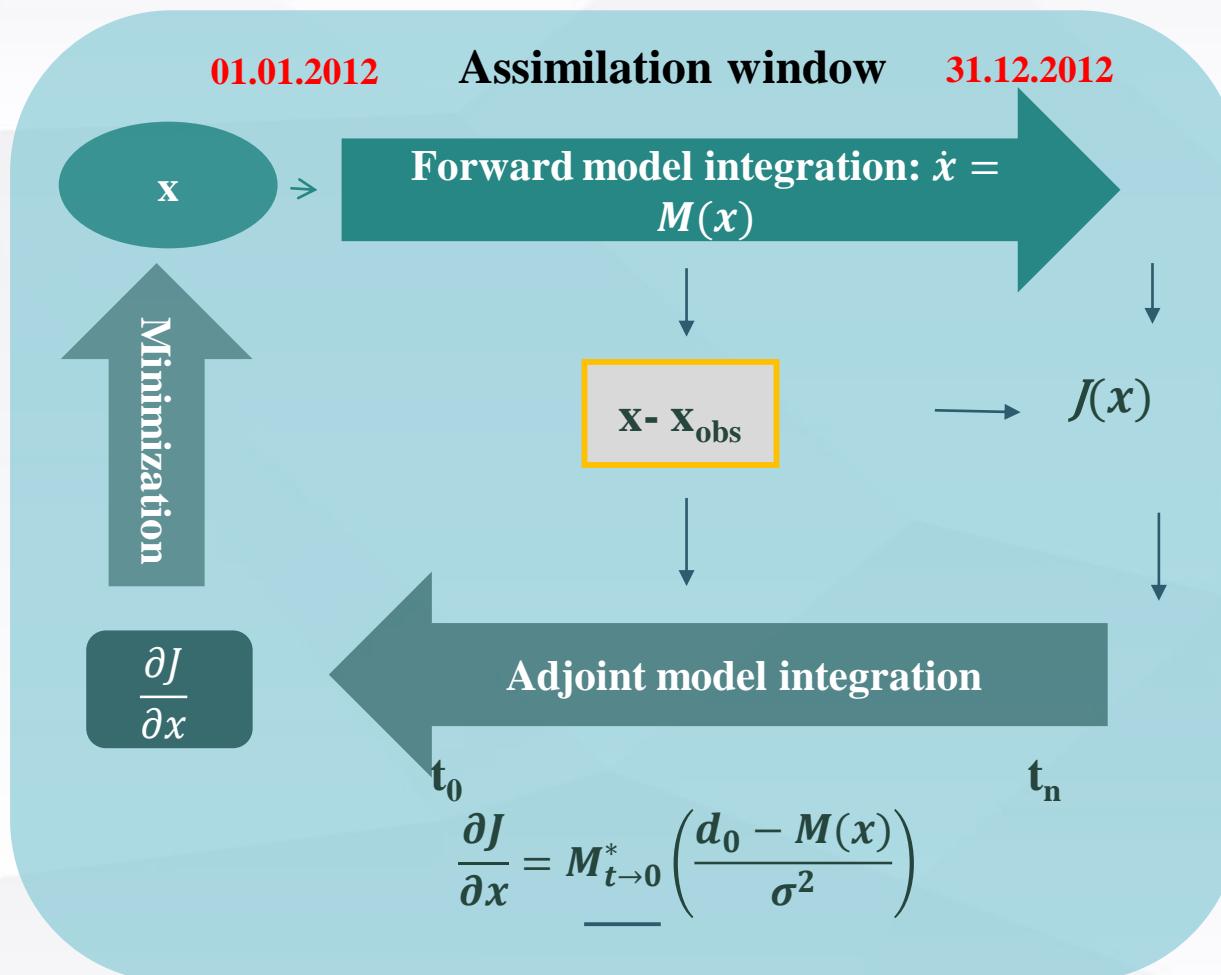
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Background and Motivation



Schematic of the adjoint method



Control variables x :

- Initial state in 2012: T/S/SIC/HEFF
- Daily atmospheric forcing:
 - 2-m air temp
 - U/V wind
 - Longwave/Shortwave
 - Specific hum





Background and Motivation

Sea-ice dynamics:

$$m \frac{d\vec{u}}{dt} = -mf\vec{k} \times \vec{u} + \tau_a + \tau_w - mg\nabla P(0) + \nabla \cdot \sigma$$

Viscous-Plastic (VP) Rheology:

Stress tensor: $\sigma_{ij} = 2\eta\varepsilon_{ij} + [\zeta - \eta]\varepsilon_{kk}\delta_{ij} - \frac{P}{2}\delta_{ij}$

Shear Viscosities: $\zeta = P/(2\Delta)$

Bulk Viscosities: $\eta = P/(2e^2\Delta)$

$$\Delta = [(\varepsilon_{11}^2 + \varepsilon_{22}^2)(1 + e^{-2}) + 4e^{-2}\varepsilon_{12}^2 + 2\varepsilon_{11}\varepsilon_{22}(1 - e^{-2})]^{\frac{1}{2}}$$

Ice pressure: $P = P^*H \times \exp(-C^* \times (1 - A))$

Ice strain rate: $\varepsilon_{ij} = \frac{1}{2}(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i})$

$e=2.0$: Eccentricity of the yield curve describing the viscous-plastic rheology

$P^* = 2.5 \times 10^4$: Ice compressive strength constant

$C^* = 20.0$: Ice strength decay constant

- Sea ice adjoint sensitivity:
Heimbach et al., 2010, OM
Kauker et al., 2009, GRL
Koldunov et al., 2013, CD
Toyoda et al., 2019, MWR
- State estimation:
ECCOV2~ECCOV4
Fenty and Heimbach, 2013, JPO
Koldunov et al., 2017, TC
Lyu et al., 2021, QJRMS
Nguyen et al., 2021, JAMES

Effects of Inclusion of Adjoint Sea Ice Rheology on Backward Sensitivity Evolution Examined Using an Adjoint Ocean–Sea Ice Model

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Background and Motivation



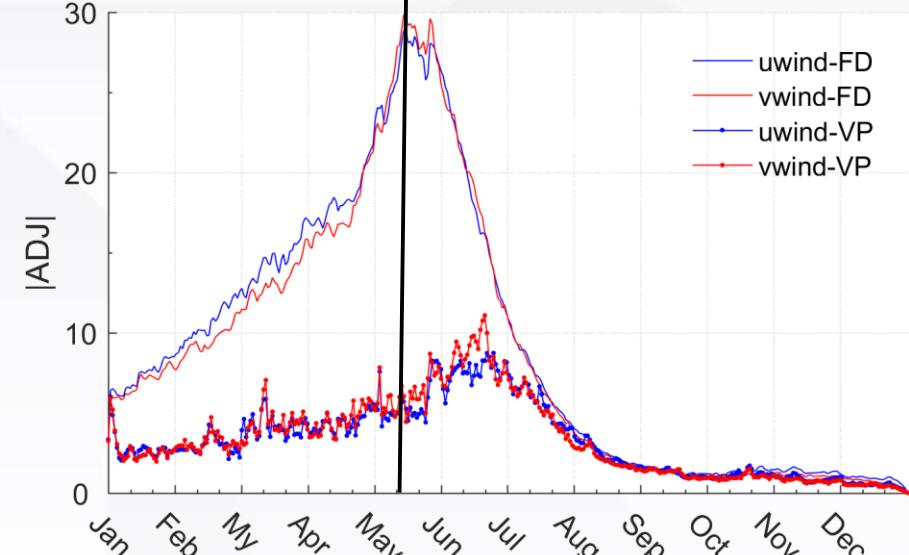
- Stabilizing the adjoint of ice Rheology :
 - Remove the adjoint sensitivity of shear/bulk viscosity (ζ/η) to strain rate (ε) (Toyoda et al., 2019, MWR)
 - Filter the adjoint of ice velocity in the adjoint of sea ice dynamics

Questions :

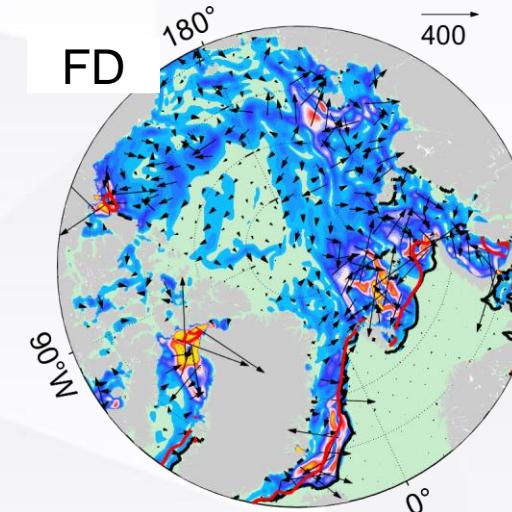
Q1: Can we further improve sea ice-ocean state estimation with adjoint of sea ice rheology?

Q2: What are the impacts on the sea ice retreat process?

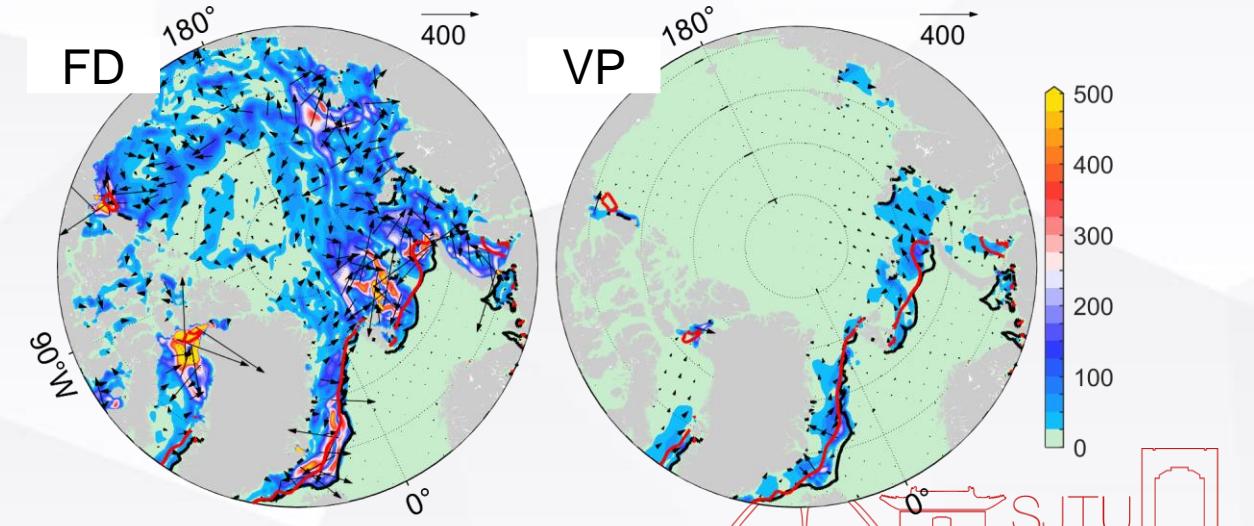
Temporal and spatial $\frac{\partial J}{\partial uwind}$ & $\frac{\partial J}{\partial vwind}$



FD



VP





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2.1

Residual Errors

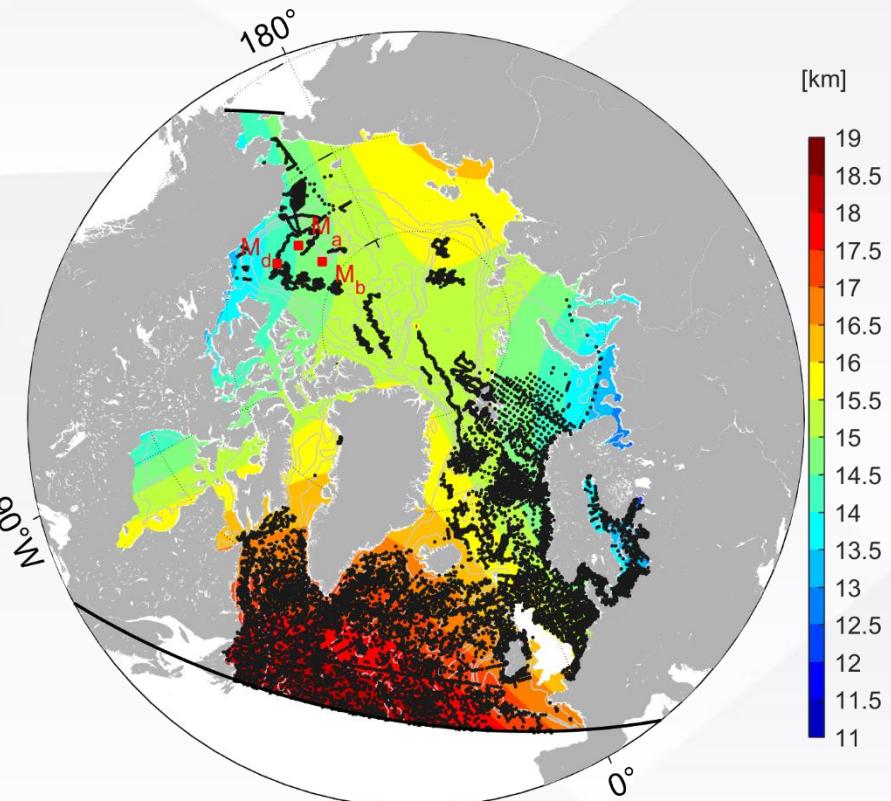
2.2

Sea Ice Retreat

3

Conclusion and Outlook

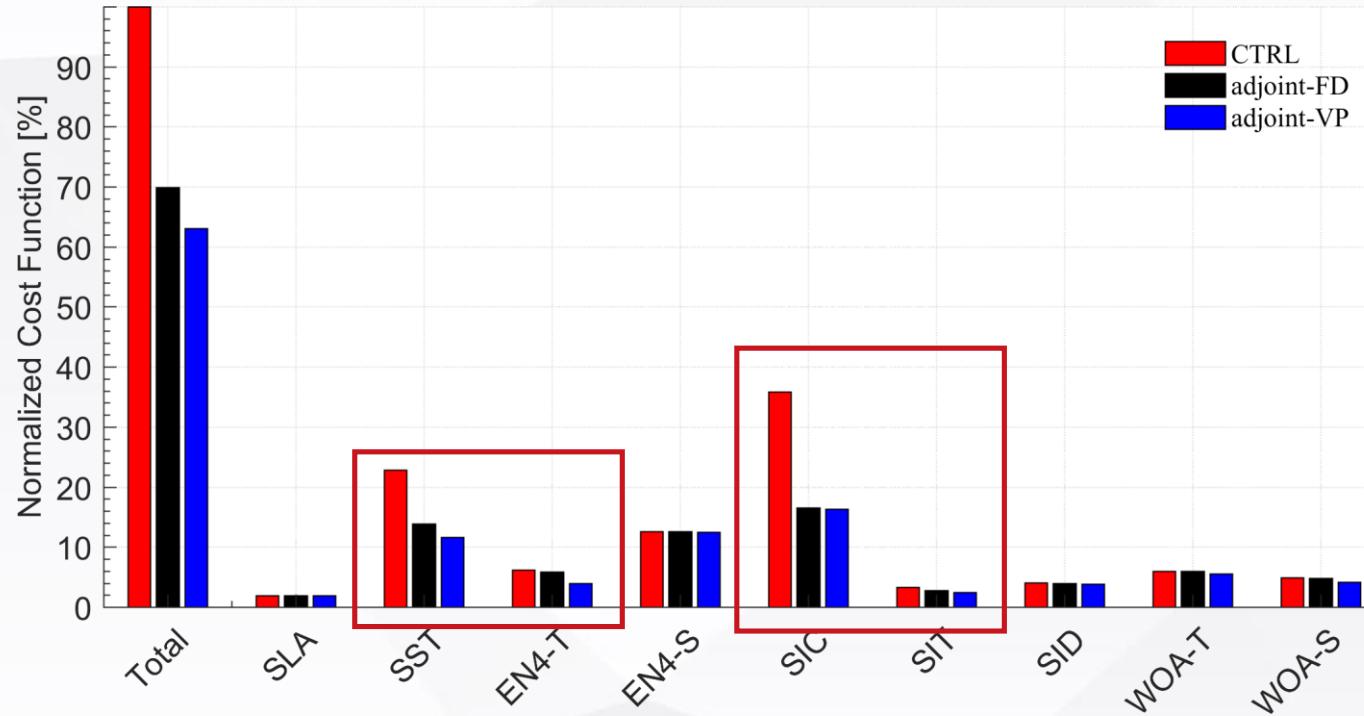
1. Experiment Setups : Models



Model domain, resolution (shading), available profiles (dots), and BGEP moorings (red rectangles)

- Based on **ECCO framework**
- A Pan-Arctic Ocean-Sea ice model configuration:
 - ✓ MITgcm+dynamic-thermodynamic sea ice model (Hilber et al ,1979)
 - ✓ Resolution: ~16 km + 50 levs
 - ✓ The adjoint model: generated by TAF (Gering et al., 1998)
- **Assimilated measurements:**
 1. Ice concentration: AMSRE, AMSE2, SSMI;
 2. Ice thickness : SMOS/CRYsat
 3. Along-track SLA, MDT
 4. Profiles from EN4, UDASH, moorings from BGEP, Bering Strait
 5. SST; WOA18
- **Assimilation experiments:**
 - ✓ Control run: **CTRL**
 - ✓ Adjoint of free drift model: **adjoint-FD**
 - ✓ Adjoint of VP rheology : **adjoint_VP**

2.1 Results: cost function reduction



Adjoint_VP 31 iterations

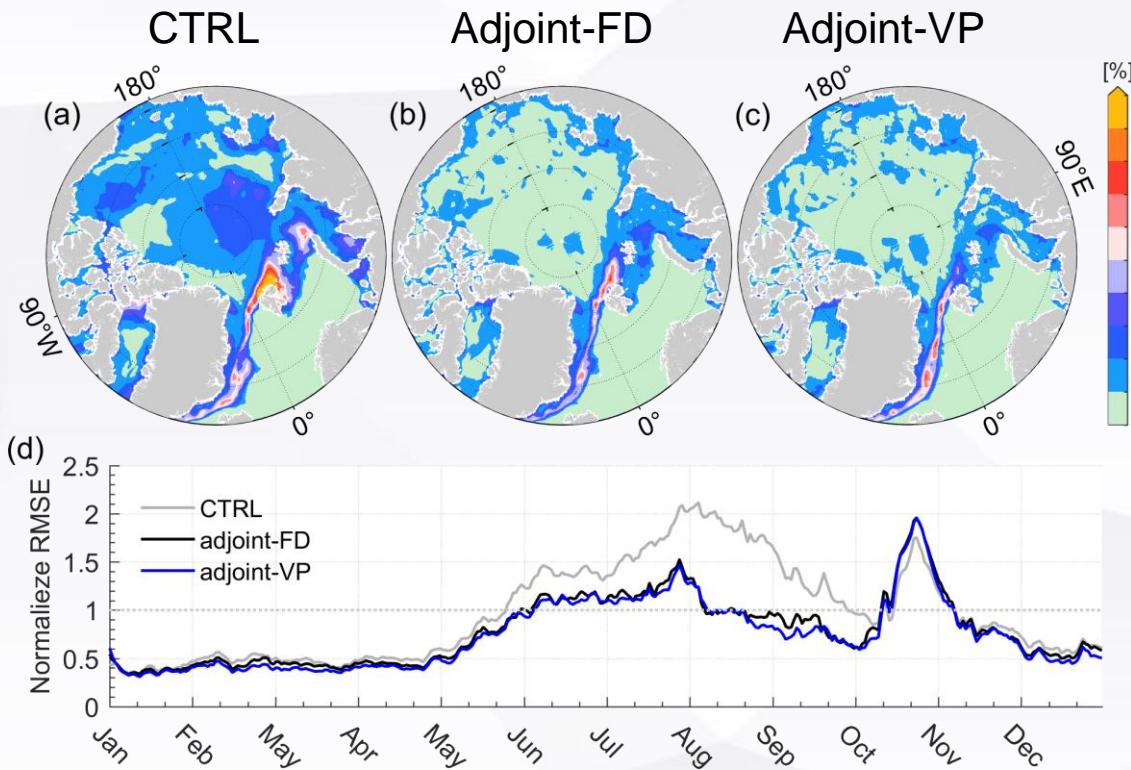
Adjoint_FD 12 iterations

- more iterations can be performed, resulting in more cost reduction (7-8%).

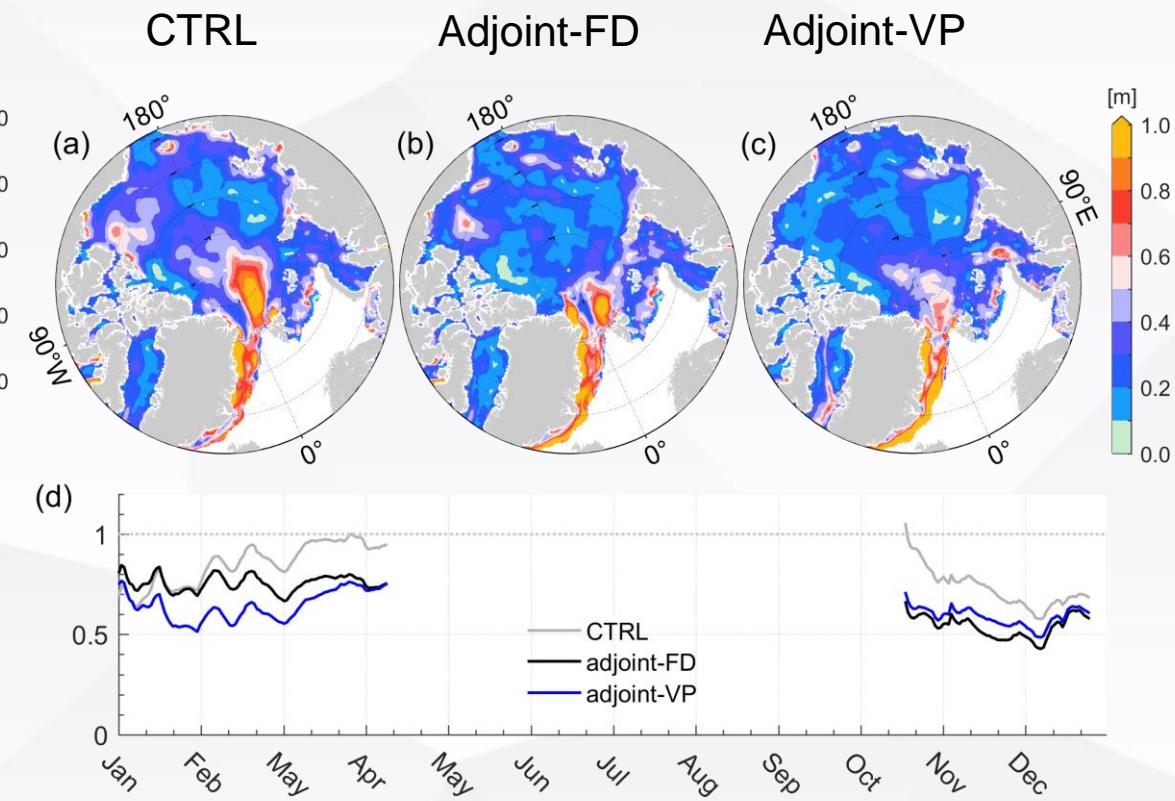
2.1 Results: RMSEs of SIC & SIT



Root Mean Square Errors of SIC



Root Mean Square Errors of SIT



RMSEs of
Ctrl
Adjoint-FD
Adjoint-LSR

SIC (%) and SIT (m)
6.22 0.37
4.30 0.34
4.28 0.34

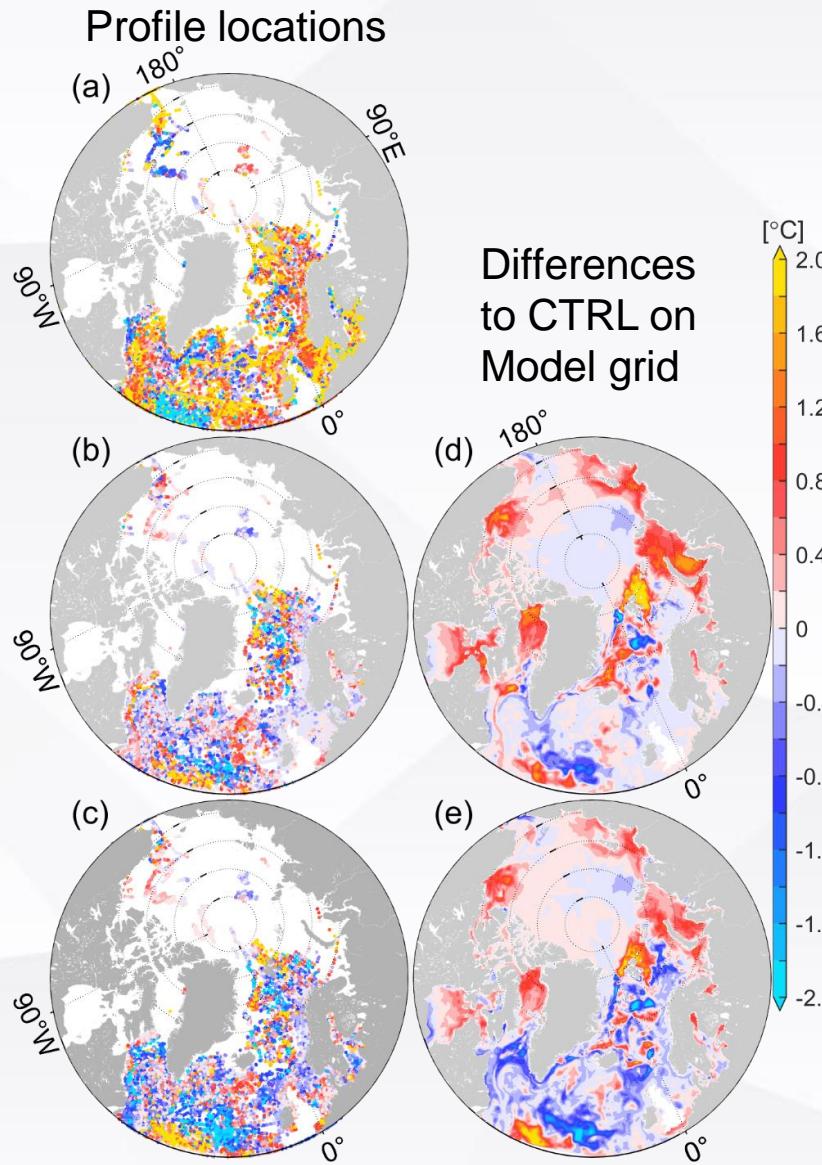
2.1 Results: Water Temperature over top 50m



CTRL- Obs
RMSEs 1.25 °C

FD-CTRL
RMSEs: 1.18 °C

LSR-CTRL
RMSEs: 0.98 °C

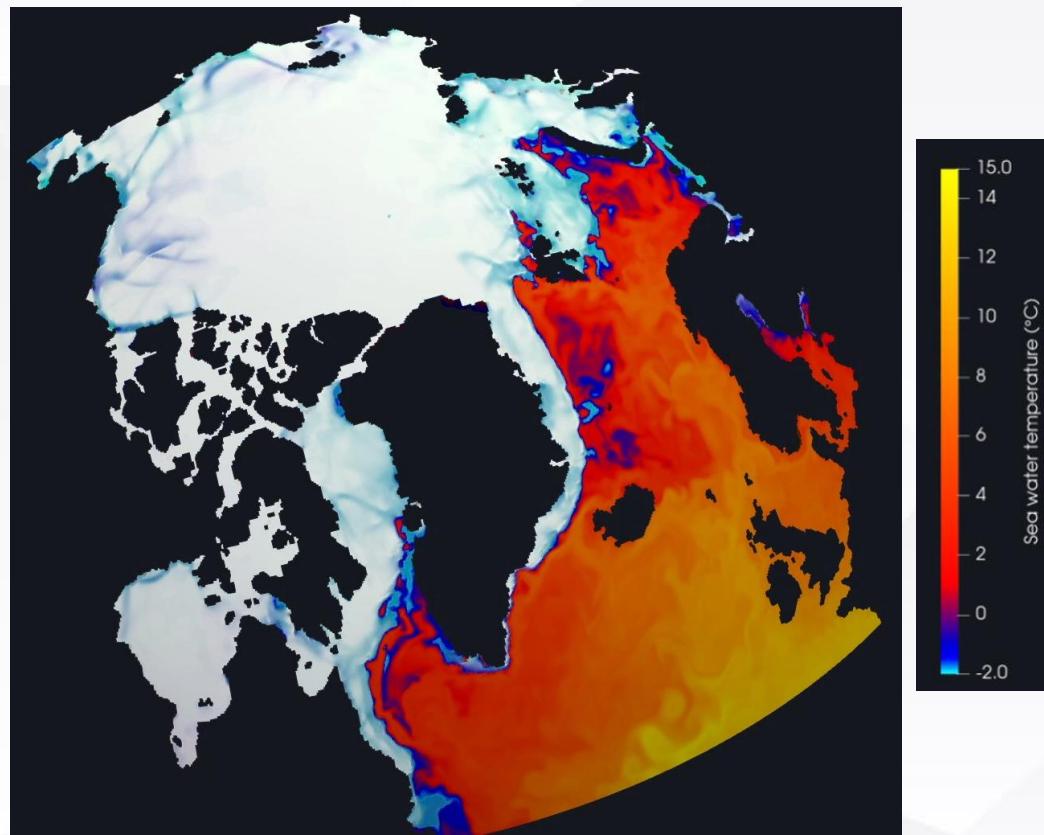


- Adjoint-VP results smaller temperature errors
- Ice-covered region: they are similar, due to the very-well representation of SIC changes.
- Ice-free regions, north Atlantic Ocean, LSR results in further error reduction, due to the more iterations.

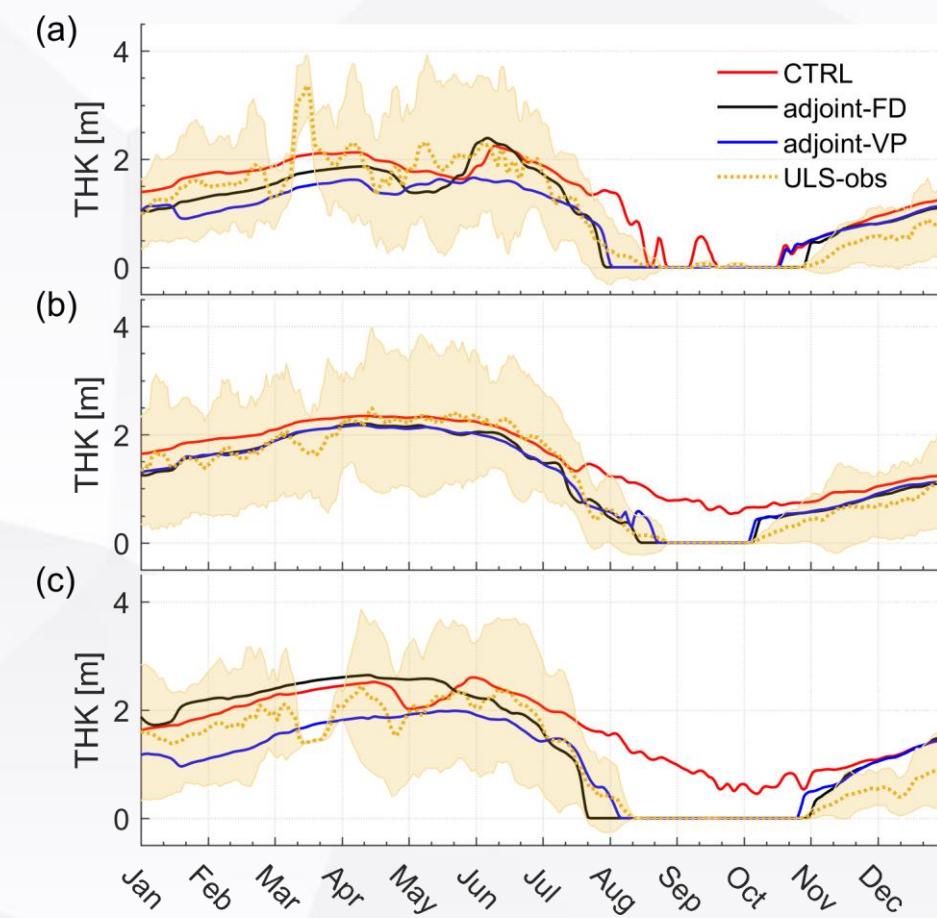
2.2 Results: Spatial-temporal sea ice variability



Sea ice concentration & thickness



Up-looking-Sonar SIT in M_a , M_b , M_d from Beaufort Gyre Exploration Project

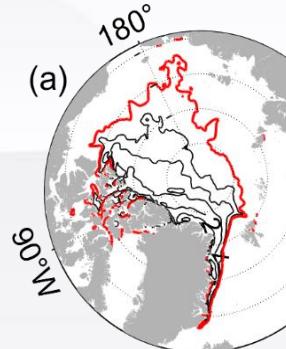


2.2 Results: Spatial pattern of sea ice retreat

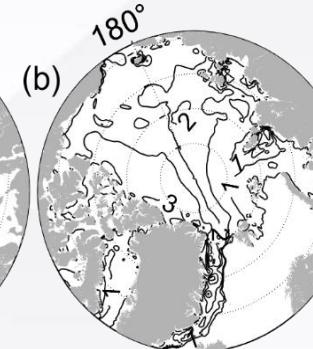


Budget Terms

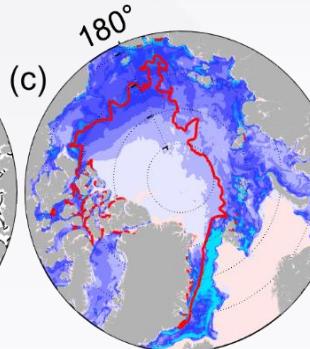
$$H_{Sep} =$$



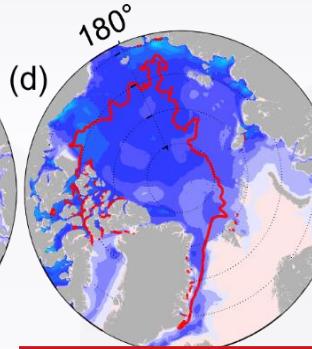
$$H_{Apr}$$



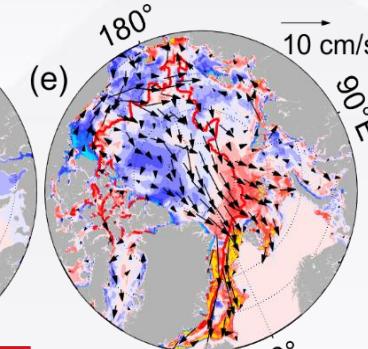
$$+F_{oi}$$



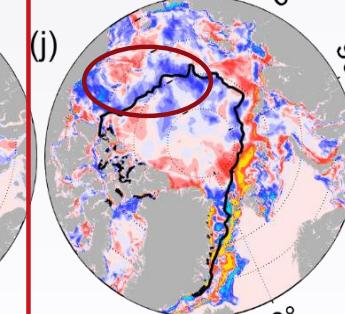
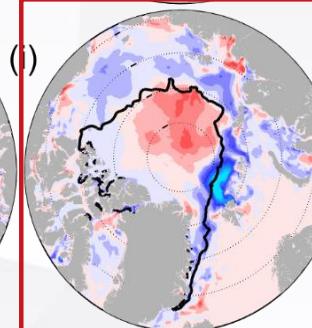
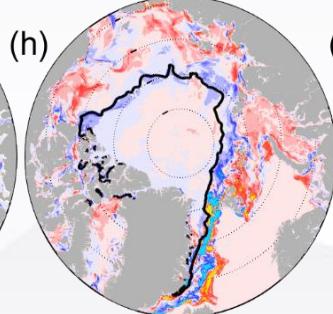
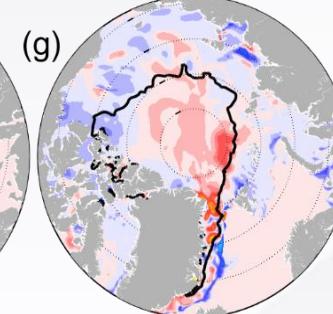
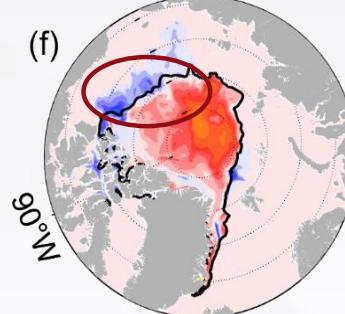
$$+F_{ai}$$



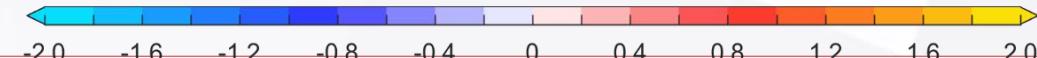
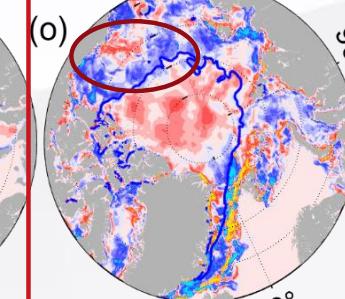
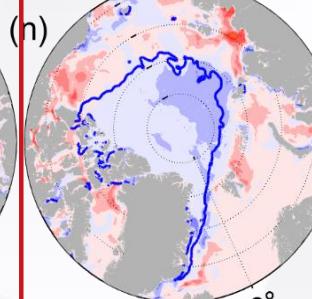
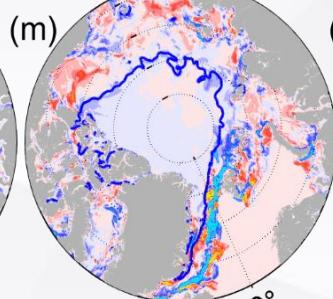
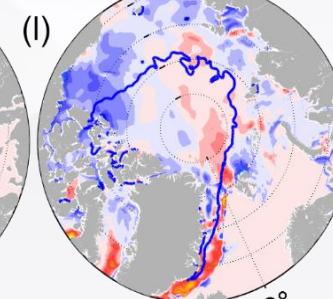
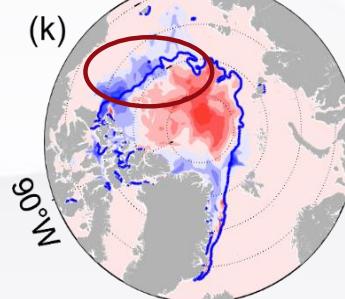
$$-\nabla \cdot (\vec{u} h)$$



FD-CTRL



VP-CTRL

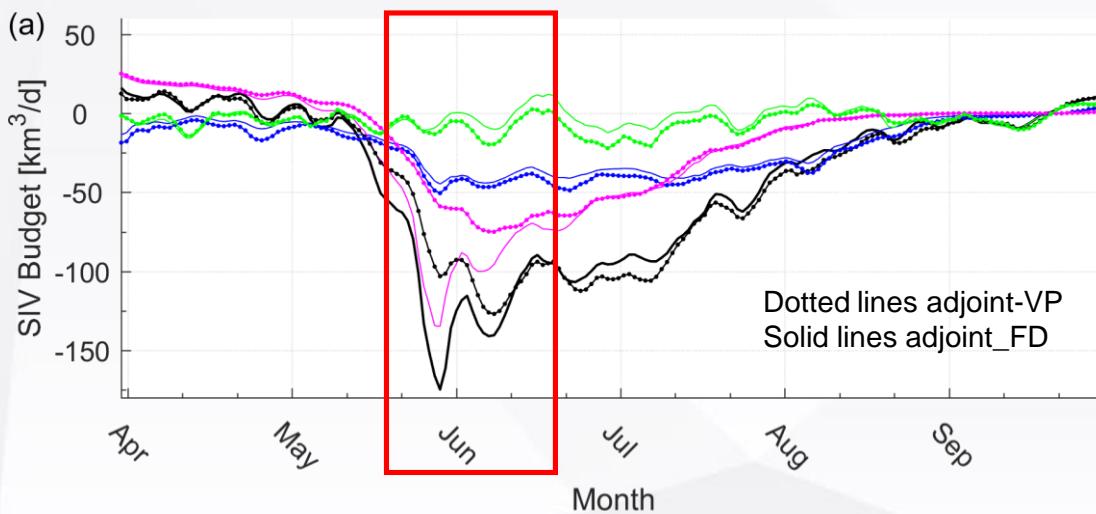


- Ice advective flux shapes September effective thickness pattern
- Differences of central Arctic Ocean ice thickness is related to F_{ai} and F_{adv}

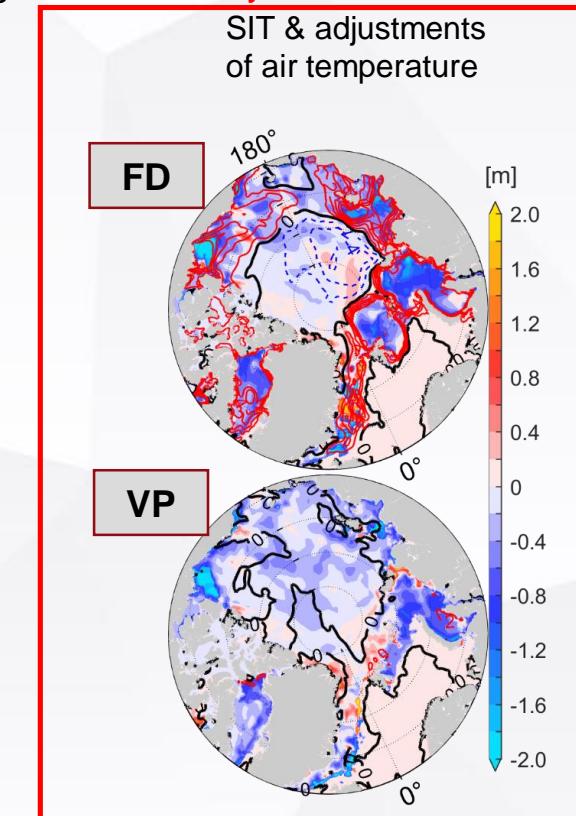
2.2 Results: sea ice retreat processes



$$\frac{\partial SIV}{\partial t} = -\nabla \cdot (\vec{u}h) + F_{oi} + F_{ai} + F_{res}$$



Changes from 20.May-15.June



- adjoint-FD melts more ice over the marginal seas through strong temperature adjustments (>10 °C)
- Adjoint-VP leads to temperature adjustments <2 °C, is more reasonable.



3. Conclusions and Outlook

We improved the adjoint of sea ice rheology (VP), tested its impacts on ocean-sea ice state estimation.

- Including adjoint of ice rheology allows more iterations to be performed, and further reduces the total cost (7-8%) .
- Analyzing the sea ice retreat process, we note that is more reasonable.

Enables estimating parameters, e.g.,
 $e=2.0$: Eccentricity of the yield curve describing the viscous-plastic rheology
 $P^* = 2.5 \times 10^4$: Ice compressive strength constant
 $C^*=20.0$: Ice strength decay constant

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