

# Energy flux quantification in the oceanic internal wavefield

**Giovanni Dematteis<sup>(1)</sup>, Kurt Polzin<sup>(2)</sup>, and Yuri Lvov<sup>(1)</sup>**

<sup>(1)</sup>Rensselaer Polytechnic Institute, Dep. of Mathematical Sciences – Troy, NY

<sup>(2)</sup>Woods Hole Oceanographic Institution – Woods Hole, MA

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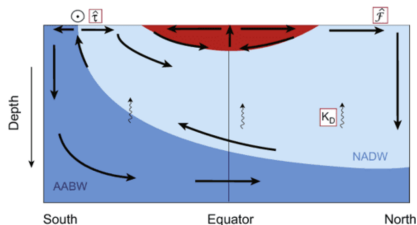
Session on *Internal Gravity Waves*

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# Oceanic circulation, mixing and internal waves

- **Dense water** sinking to the abyss at high latitudes **must eventually rise**
- Mechanism: **Turbulent diapycnal mixing** of “dense bottom water”/“light water above”
- Modelled as **diffusivity**  $K_D$  in advection-diffusion closure (*Munk '66*)
- One of the main drivers of the **Meridional Overturning Circulation (MOC)**



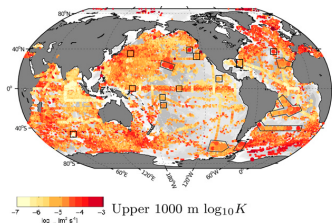
from *Sévellec, Fedorov, Deep Sea Res. '11*  
see also *Whalen et al., Nat. Rev. '20*

**Finescale Parameterization:**  $\mathcal{P} = \mathcal{P}_0 \frac{f}{f_0} \frac{N^2}{N_0^2} \frac{E^2}{E_0^2}$

**Turbulent production**  $\mathcal{P}$  has 2 components:

- **mixing:**  $K_D \simeq \frac{R_f}{\rho_0 N^2} \mathcal{P}$
- **Dissipation:**  $\epsilon = (1 - R_f) \mathcal{P}$

and is **quantified heuristically** in terms of parameters of **internal wave spectrum!**



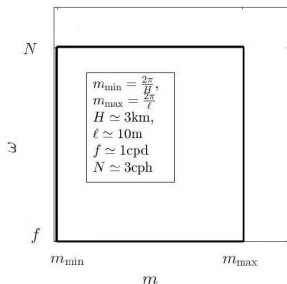
*Waterhouse et al., JPO '14: finescale estimates of global patterns of mixing*

# Internal Wave Turbulence

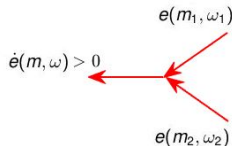
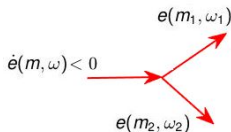
- Boussinesq approximation for ocean with **stable stratification**
- Assume isopycnal coordinates, zero total potential vorticity, hydrostatic balance, spatial homogeneity, horizontal isotropy.
- Assume “Random phase” statistics: energy spectrum  $e(m, \omega)$  evolves according to the **Wave Kinetic Equation** (Lvov & Tabak, *Phys. D* '04, Olbers '74)

$$\dot{e}(m, \omega) = \mathcal{I}(e(m, \omega)).$$

- $\mathcal{I}(e(m, \omega))$ : **Collision Integral** quadratic in  $e(m, \omega)$



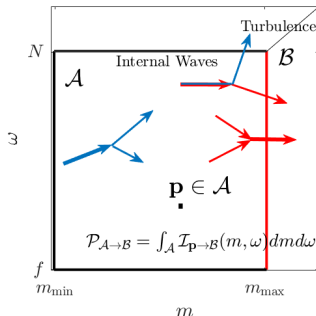
The Collision Integral quantifies the **inter-scale energy exchanges via 3-wave resonances**



# Calculation of inter-scale energy fluxes

Method to compute **energy flux between “control volumes” in Fourier space** from Collision Integral, *Dematteis and Lvov, JFM '21*

- $\mathcal{I}_{\mathbf{p}}$ : energy budget of mode  $\mathbf{p}$  in relation to all other modes
- $\mathcal{I}_{\mathbf{p} \rightarrow \mathcal{B}}$ : energy budget of mode  $\mathbf{p}$  such that the outcome of the interaction  $\mathbf{p}$  is in region  $\mathcal{B}$
- Consider **control volume** ( $\mathcal{A}$ ) containing internal-wave band and integrate  $\omega_{\mathbf{p}} \mathcal{I}_{\mathbf{p} \rightarrow \mathcal{B}}$  over  $\mathbf{p} \in \mathcal{A}$



Method paper to appear: *Dematteis and Lvov '22*

Wave turbulence analogy to Kraichnan's '59 DIA calculation of turbulent fluxes using control volumes in Fourier space

# Comparison with finescale parameterization

$$\mathcal{P}_{\text{out,h}} = \frac{\Gamma C_h}{1-\nu} \left[ 1 - \left( \frac{\ell}{2b} \right)^{1-\nu} \right] f^{1+\nu} N E^2,$$

$$\mathcal{P}_{\text{out,v}} = \frac{\Gamma C_v}{\nu} \left[ 1 - \left( \frac{f}{N} \right)^\nu \right] f N^{1+\nu} E^2,$$

$$\Gamma = \frac{1}{\pi^3} \left( \frac{2\ell}{b} \right)^\nu \frac{N_0^{1-\nu} b^3}{c^{1-\nu} \ell}, \quad \nu = 2a - 7 = 0.38, \quad \ell = 10 \text{ m},$$

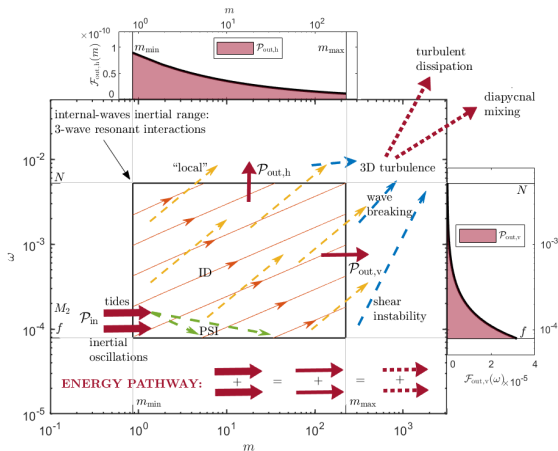
$$\mathcal{P}_{\text{out,h}} \simeq -3.8 \times 10^{-9} \text{ W kg}^{-1},$$

$$\mathcal{P}_{\text{out,v}} \simeq -5.2 \times 10^{-9} \text{ W kg}^{-1}.$$

This amounts to a total dissipated power

$$\mathcal{P}_{\text{out}} = \mathcal{P}_h + \mathcal{P}_v \simeq -9.0 \times 10^{-9} \text{ W kg}^{-1}.$$

$$\mathcal{P}_{\text{finescale}} \simeq 5.9 \times 10^{-9} \text{ W/kg}.$$



For GM scaling ( $\nu = 1$ ),

$$\mathcal{P}_{\text{out,v}} \propto f N^2 E^2$$

**= scaling of finescale parameterization formula**

We **calculated** the constant  $\mathcal{P}_0$  from theory, i.e. the **"Kolmogorov constant"** of the problem: within a factor 1.5 from empirical value!

*Dematteis, Polzin, Lvov, JPO (2022)*

# Conclusions

- Open question on **how to calculate the energy dissipated by internal waves**, with important implications for climate.
- **Finescale parameterization** (synthesis of observation, dimensional analysis and intuition, from 3 decades ago: *Gregg '89, Henyey '91, Polzin et al. '95*) is the **state-of-the-art understanding** of dissipation and mixing:

$$\mathcal{P} = \mathcal{P}_0 f N^2 E^2$$

with  $\mathcal{P}_0$  **empirically determined**.

- We have **explained how to obtain the formula from the first principles**, with no adjustable parameters.
- **Computation** of  $\mathcal{P}_0$  (analogue of “Kolmogorov constant” of turbulence) **in agreement with empirical value** up to factor 1.5.
- Future plan is to extend the result to any internal wave spectrum and to test the results directly against oceanic field data and high-resolution numerical models.

Thank you for your attention!

**The END**

