

Towards a high-resolution v_s crustal velocity model for the Ivrea Geophysical Body: constraints from seismic ambient noise tomography

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Motivation

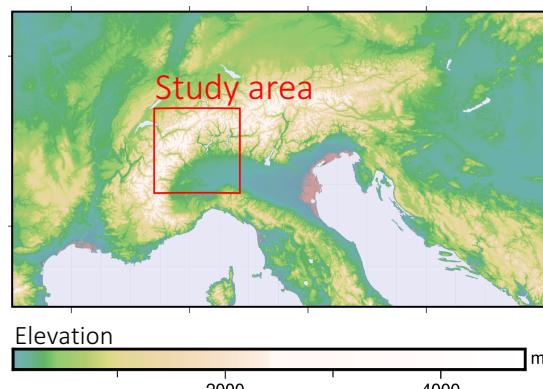
Motivation

- Heterogeneous crustal structure at the boundary between Europe and Adria, in correspondence of the Ivrea Geophysical Body (IGB);
- Target area for the scientific drilling project DIVE (www.dive2ivrea.org).



Goal

- Obtain high-resolution constraints on absolute v_s distribution in the crust, via seismic ambient noise tomography (ANT).



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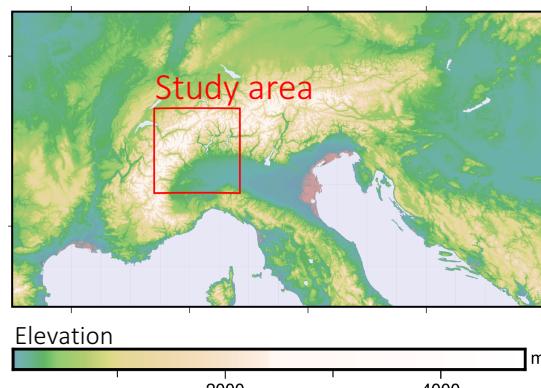
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ANT Workflow

1. Seismic ambient noise pre-processing

2. Station-pair cross-correlation (CC)

3. Surface Rayleigh wave dispersion measurement via FTAN analysis (fundamental mode / group speed)

4. 2D surface wave velocity maps at different periods (4s to 23s)

5. Inversion for local 1D v_s -depth profiles

Preliminary results

Data

▼ Seismic stations

Networks:

XK (*IA*-IvreaArray, Mobnet pool of IGCAS Prague)

Z3 (AlpArray temporary)

IV GU

CH MN YP

★ DIVE Balmuccia drilling site

PL Periadriatic Line

Adriatic side

■ Adria lower crust | Ivrea-Verbano Zone (IVZ)

■ Adria upper crust

■ Adria Mesozoic

European margin

■ Helvetic Nappes

■ European basement

■ European upper crust

■ High-pressure Subpenninic Units

Other units

■ Valaisan oceanic domain

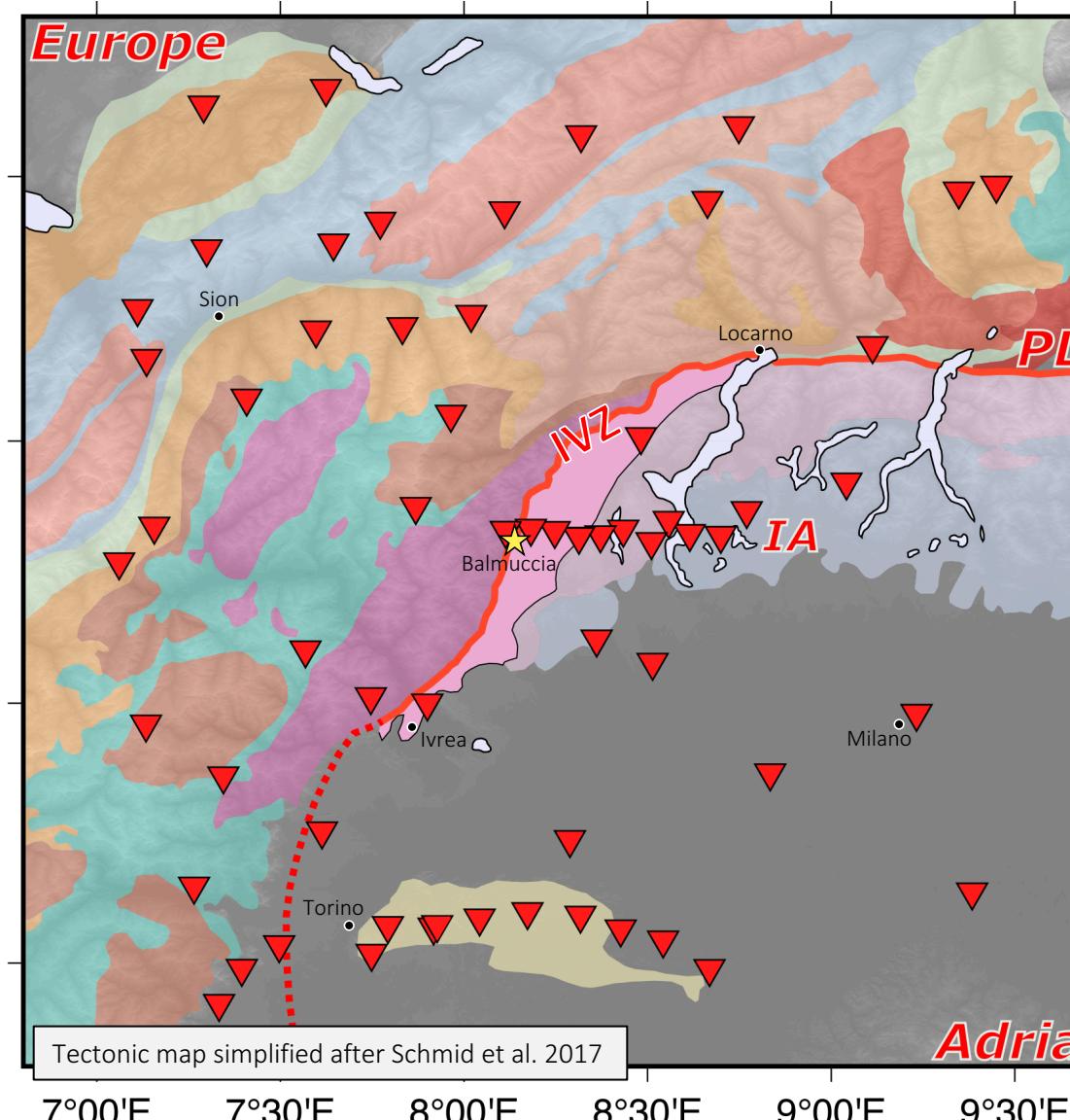
■ Brianconnais micro-continent

■ Sesia, Dent Blanche

■ Piedmont-Liguria Ocean

■ Cenozoic plutons

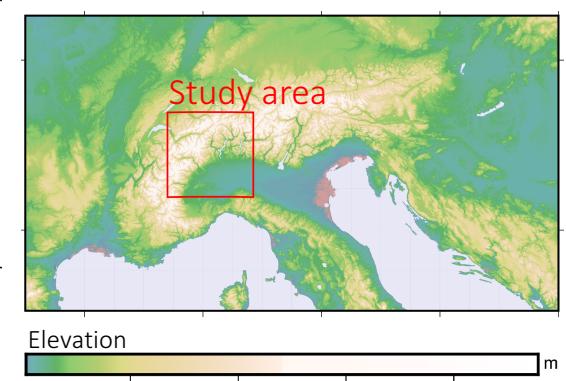
■ Cenozoic



Up to 9/12 months of broadband summer data (June to September) across 2+ years of recording

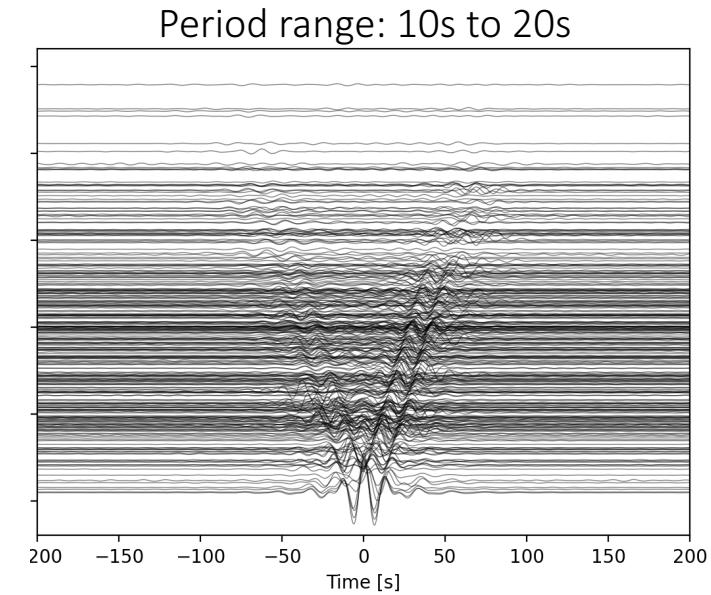
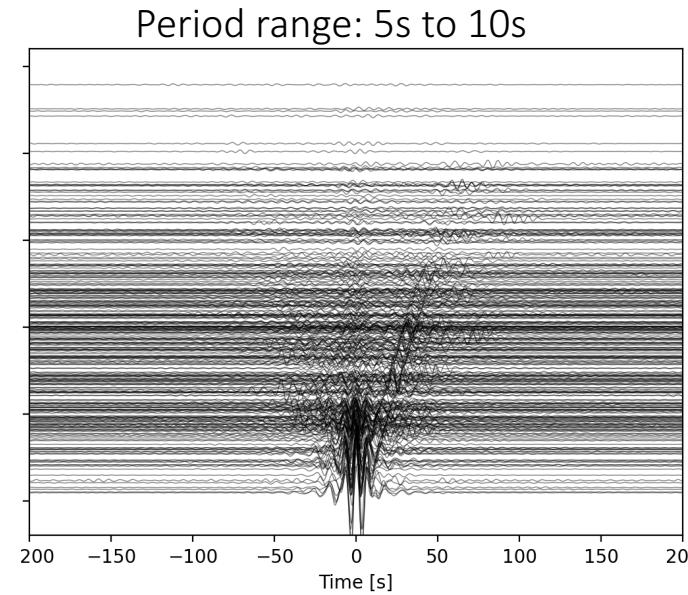
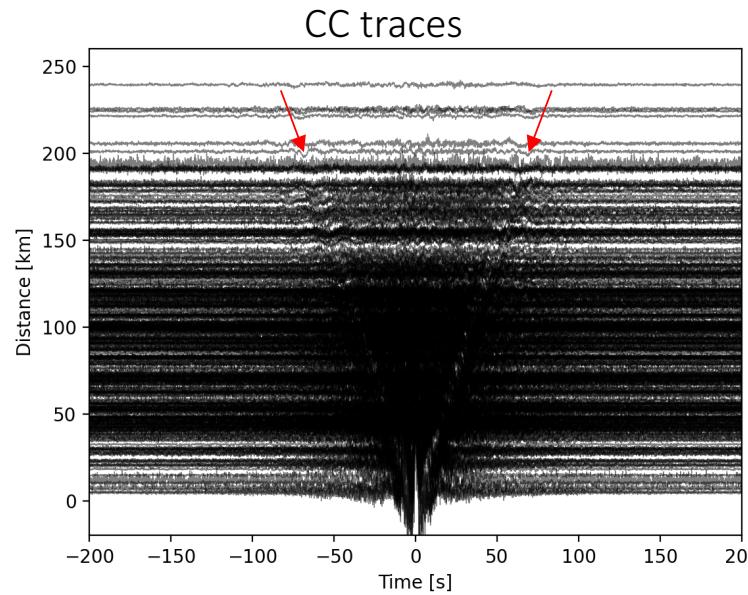
Homogeneous pre-processing:

- Pre-filtering
- Instrument response removal
- Amplitude normalization
- Spectral whitening



Seismic ambient noise cross-correlation (CC)

Subset of the data (ca. 300 out of 1400 traces)



CC traces:

- 10Hz sampling rate
- Max lag: 480s
- Length of the correlated segments: 3600s
- Overlap: 0.5

MSNoise package (Lecocq et al. 2014)

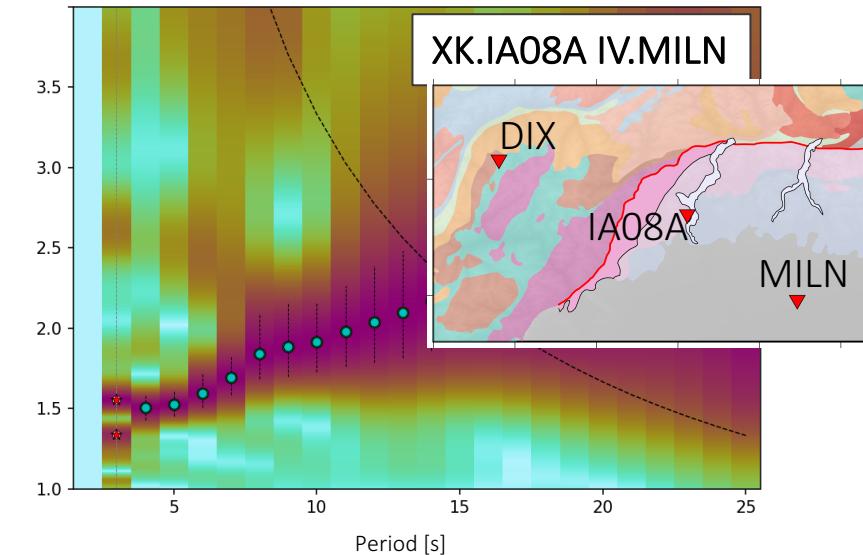
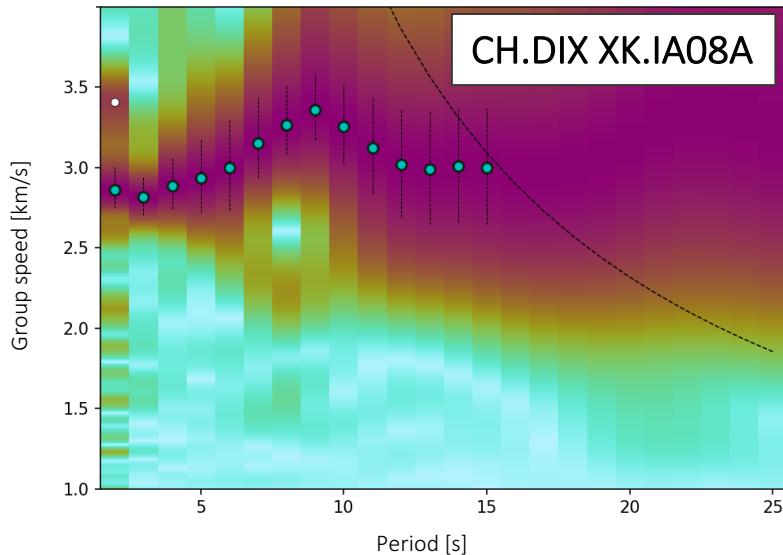
CC traces → Input for dispersion measurement of surface Rayleigh waves (group speed / fundamental mode)

Dispersion measurement via frequency-time analysis (FTAN)

Dispersion measurement via Frequency-Time analysis, for each station-pair (FTAN, e.g. Levshin and Ritzwoller 2001):

1. Green's Function envelope computation
2. Narrow band-pass filtering
3. Rayleigh group velocity dispersion measurement (fundamental mode)

Period range 2s to 24s

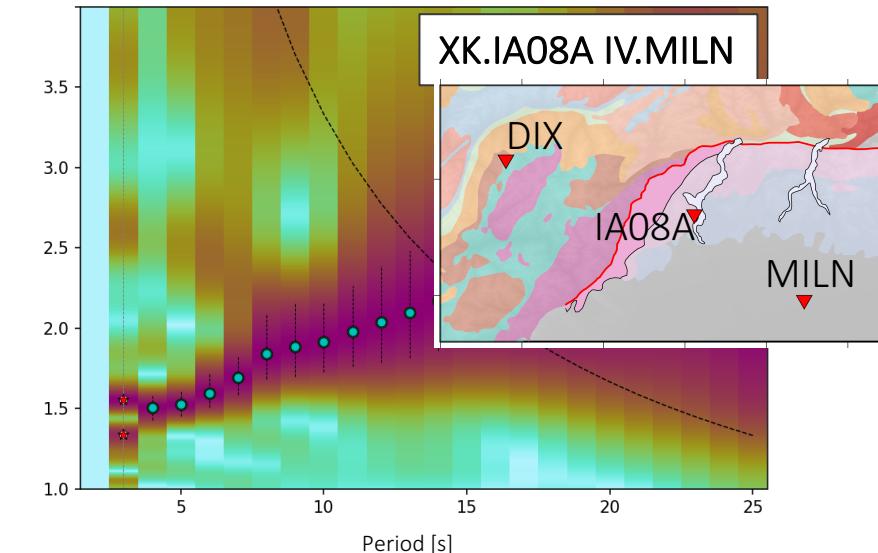
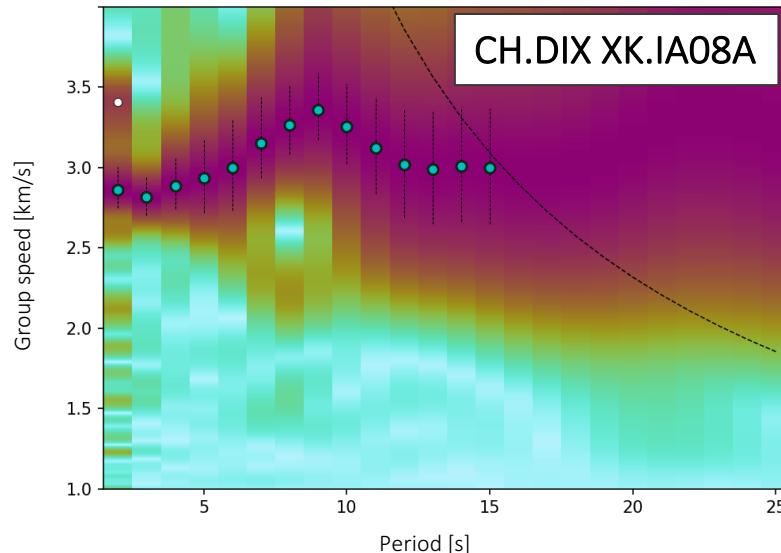


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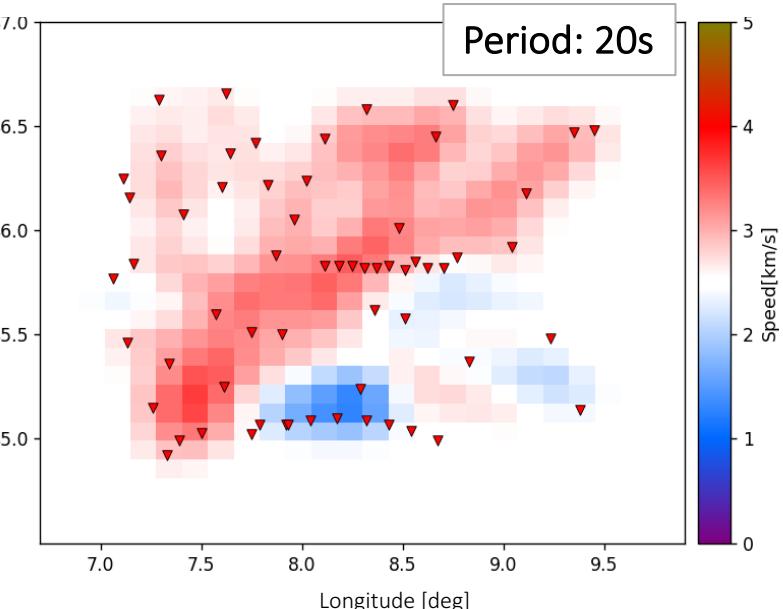
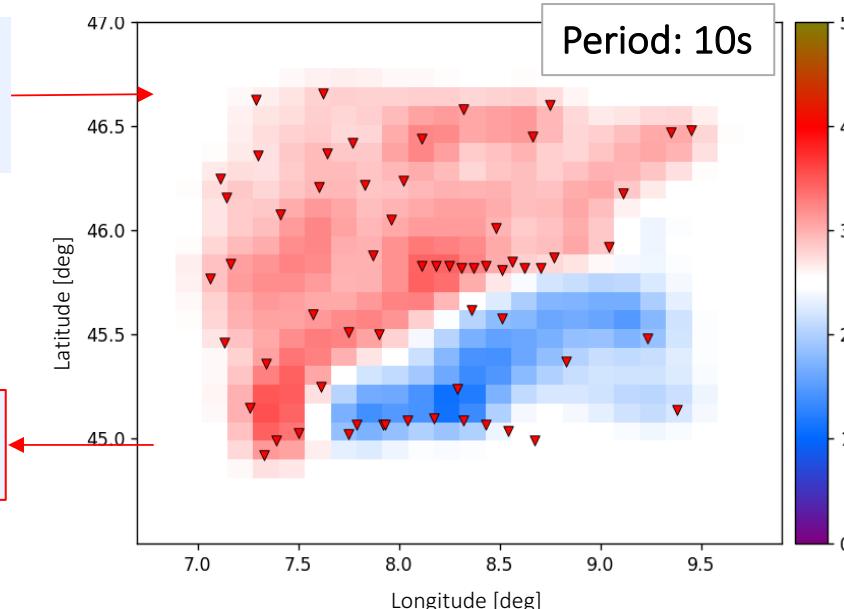
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2D surface group velocity maps via fast-marching surface tomography (FMST, Rawlinson, 2005)

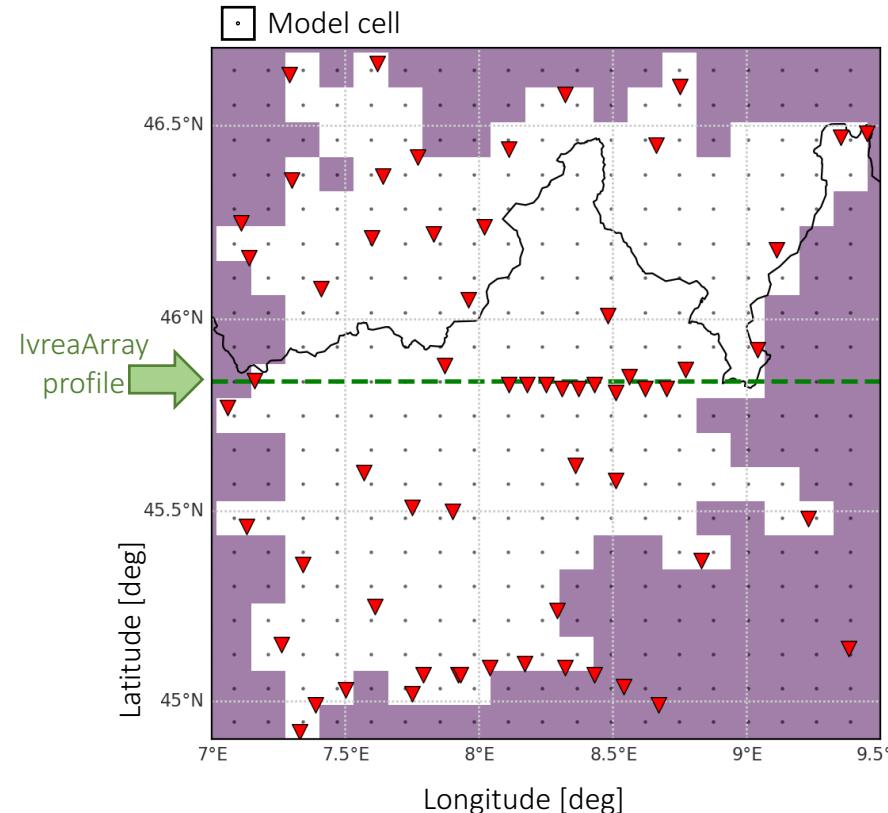
2D maps (one for each period) -> input for local 1D v_s -depth inversion



Preliminary results

Neighborhood Algorithm (NA) inversion
for local 1D v_s -depth profiles.
Dinver package (Wathelet et al. 2008)

- 234 cells ($\sim 10\text{km} \times 10\text{km}$)
- 6 horizontal layers over a half-space
- Each layer has a 5 sub-layers gradient
- 50500 explored models for each cell

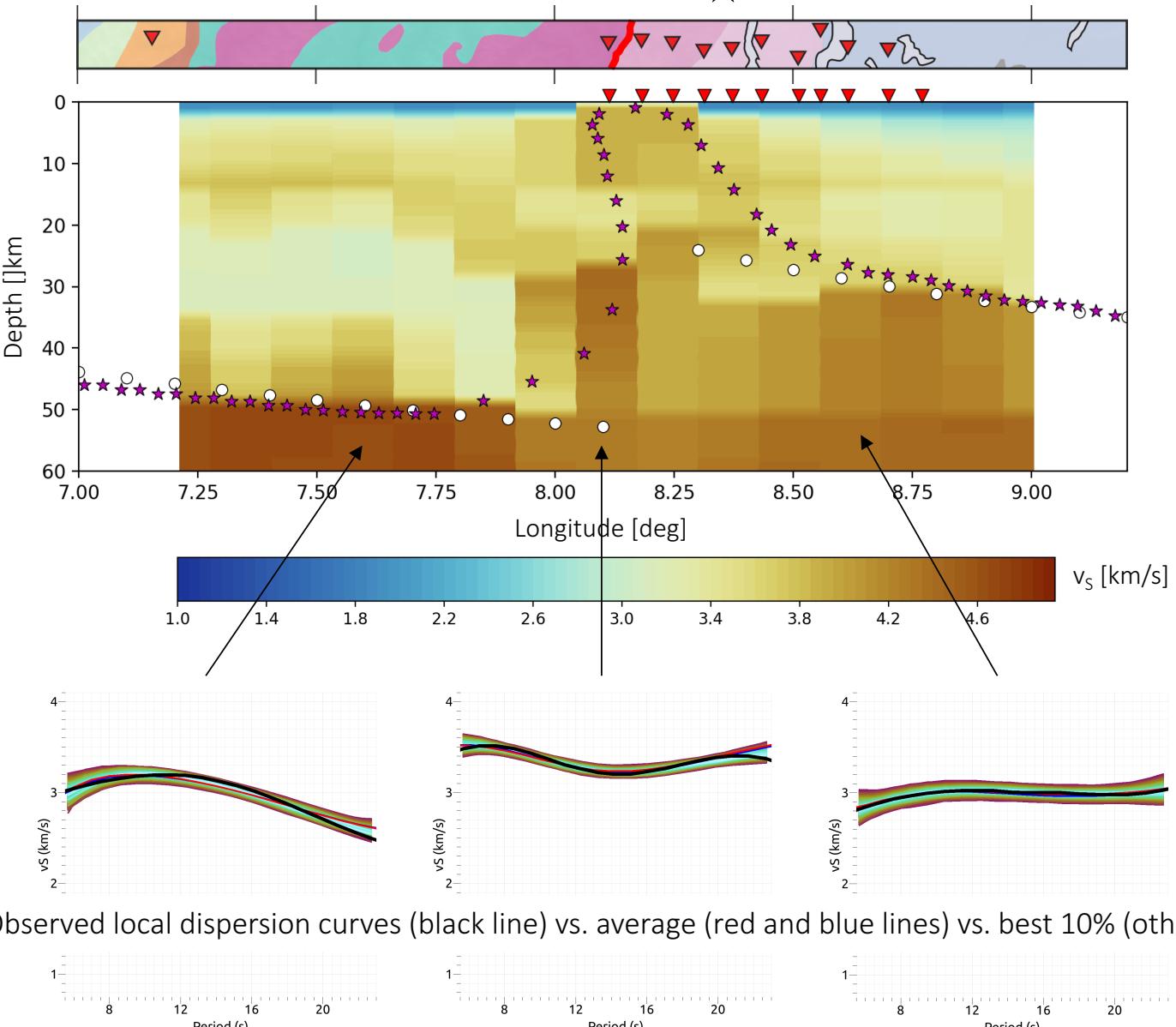
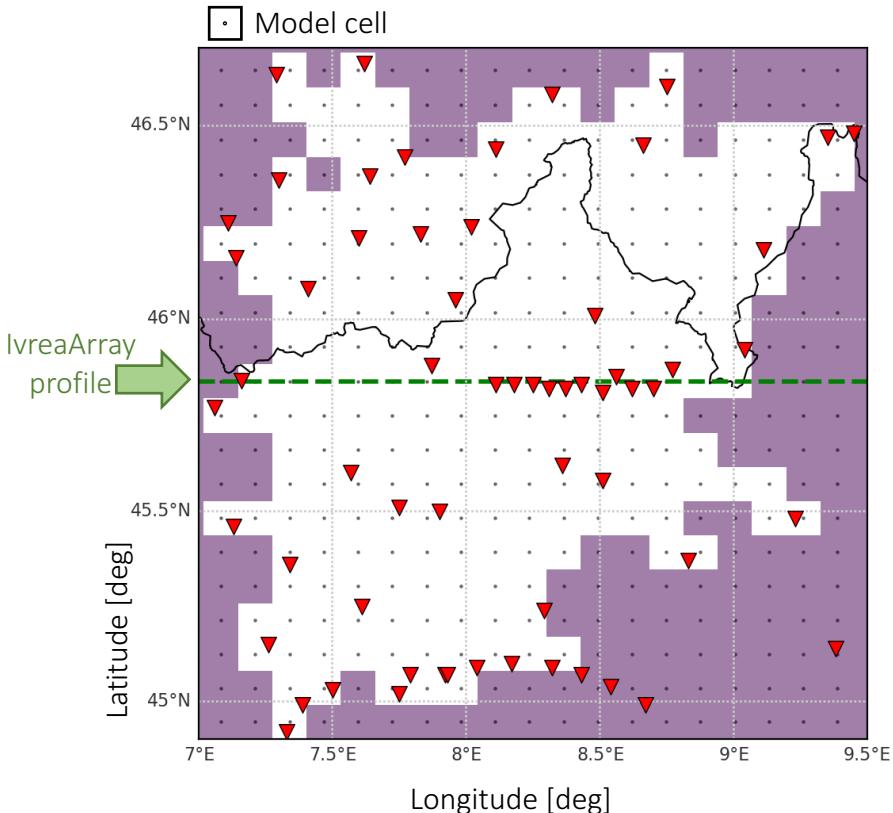


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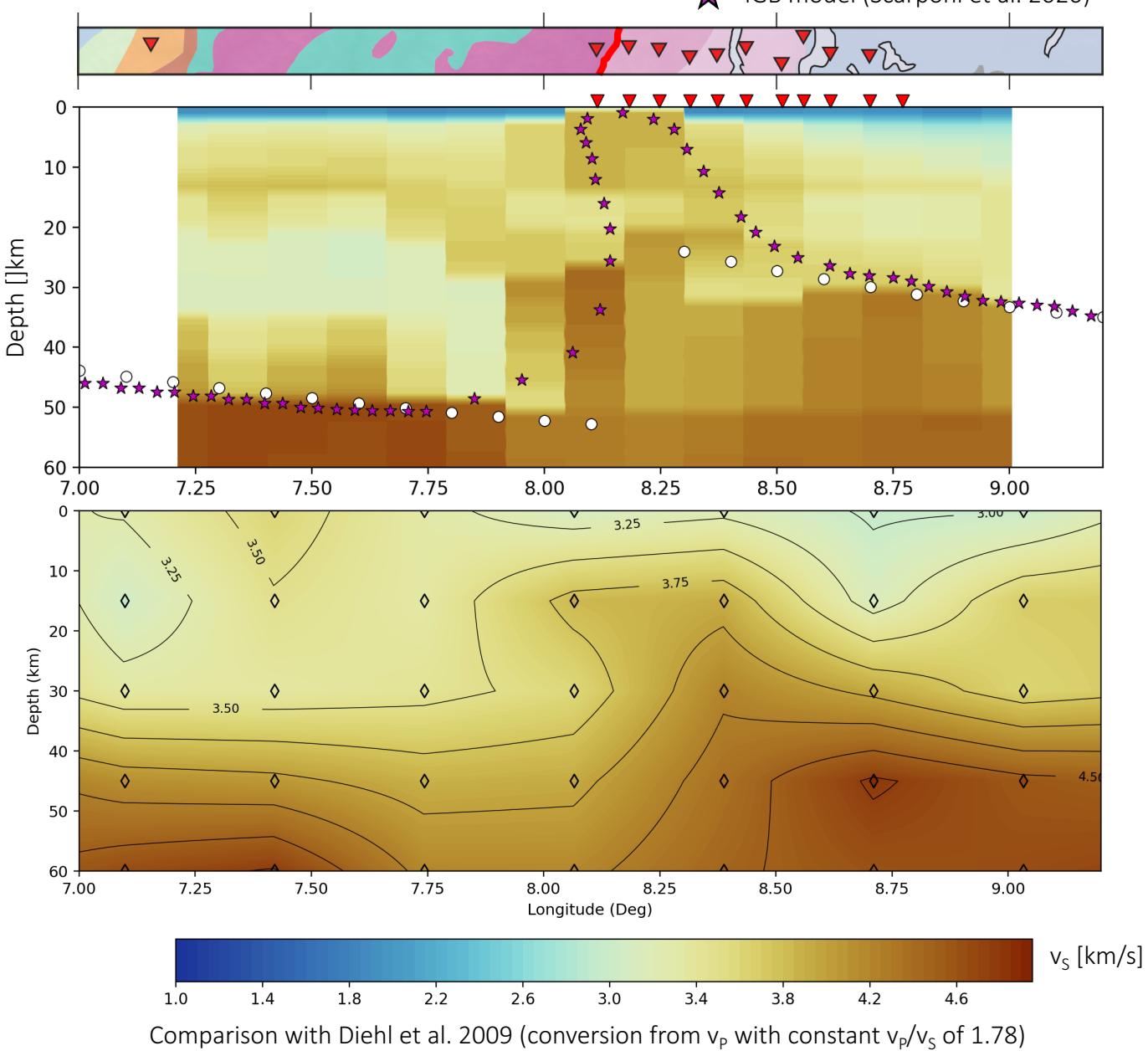
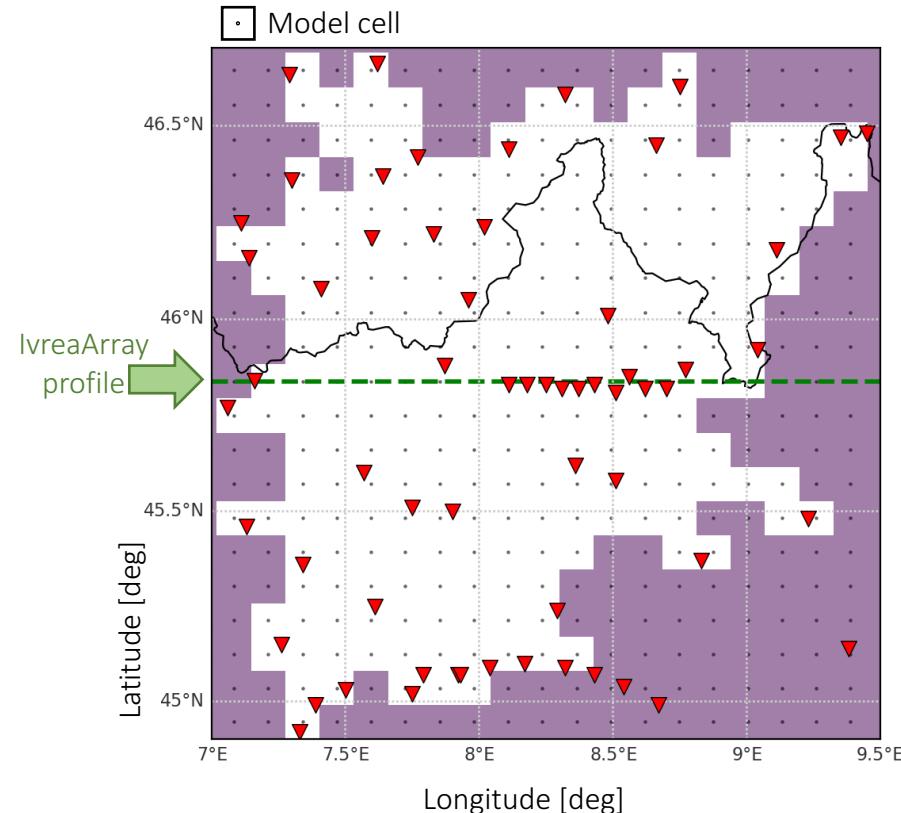


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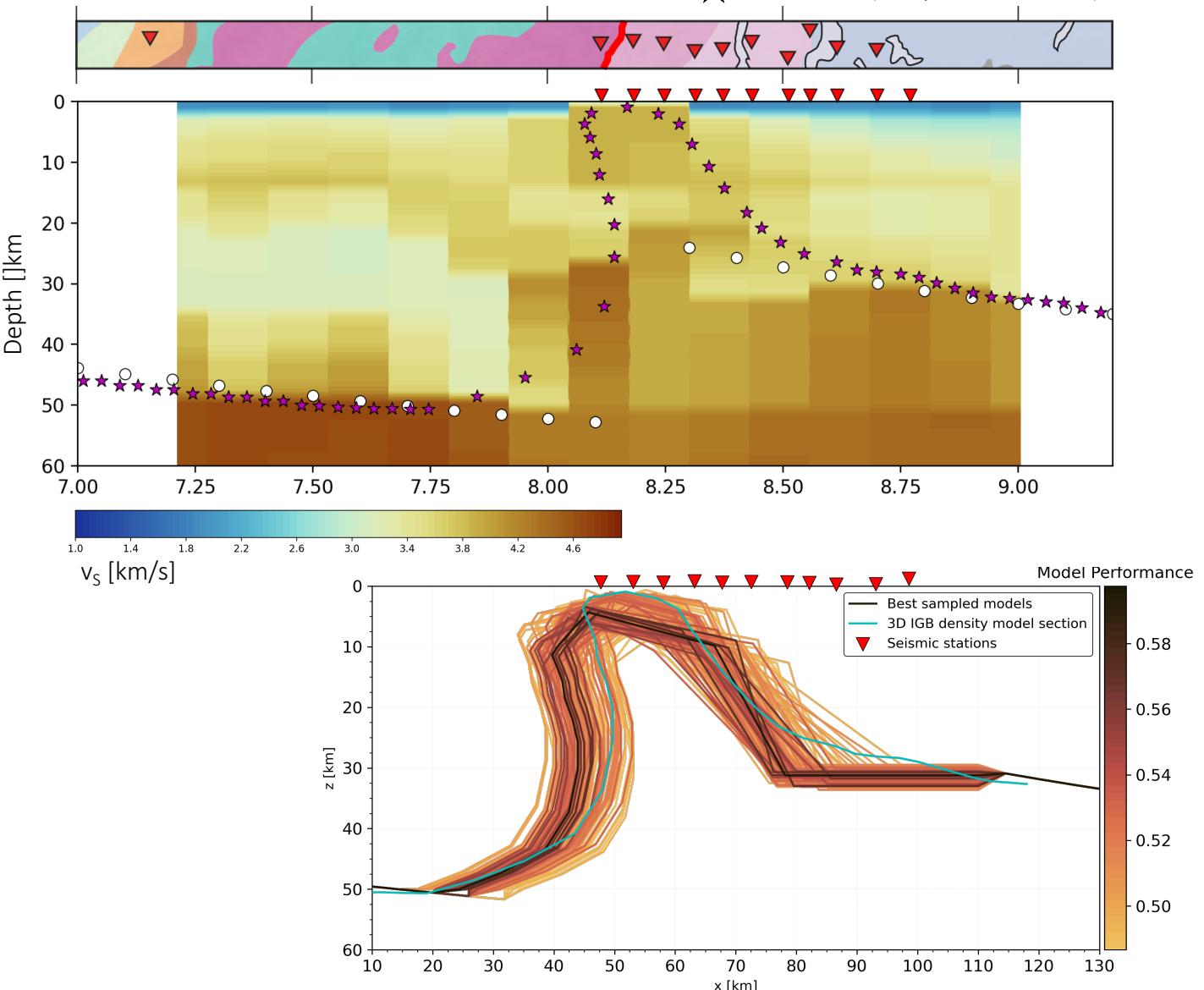
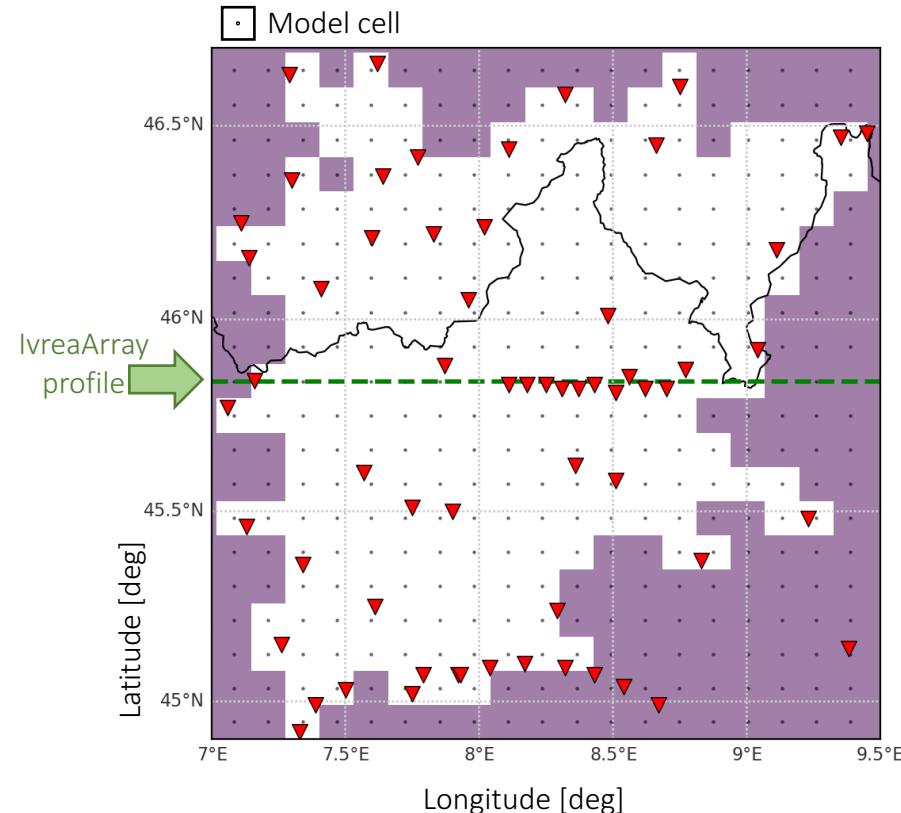


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Comparison with Scarpa et al. 2021 (Receiver functions + gravity joint inversion)

Current status

Preliminary conclusions

- ANT images the IGB at very shallow depths
- Agreement with previous gravity and joint RF + gravity modelling (Scarponi et al. 2020, 2021)
- Agreement with LET (Diehl et al. 2009), with IGB detected at even shallower depths
- Moho depth coherent in the Europe and Adria plates with respect to LET

Future steps

- Evaluate a priori constraints to improve parameterization for v_s
- Consider including longer periods (i.e. higher sensitivity in the lower crust)

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