

Narrow, fast, and “cold” mantle plumes on Earth explained by strain-weakening rheology in the lower mantle

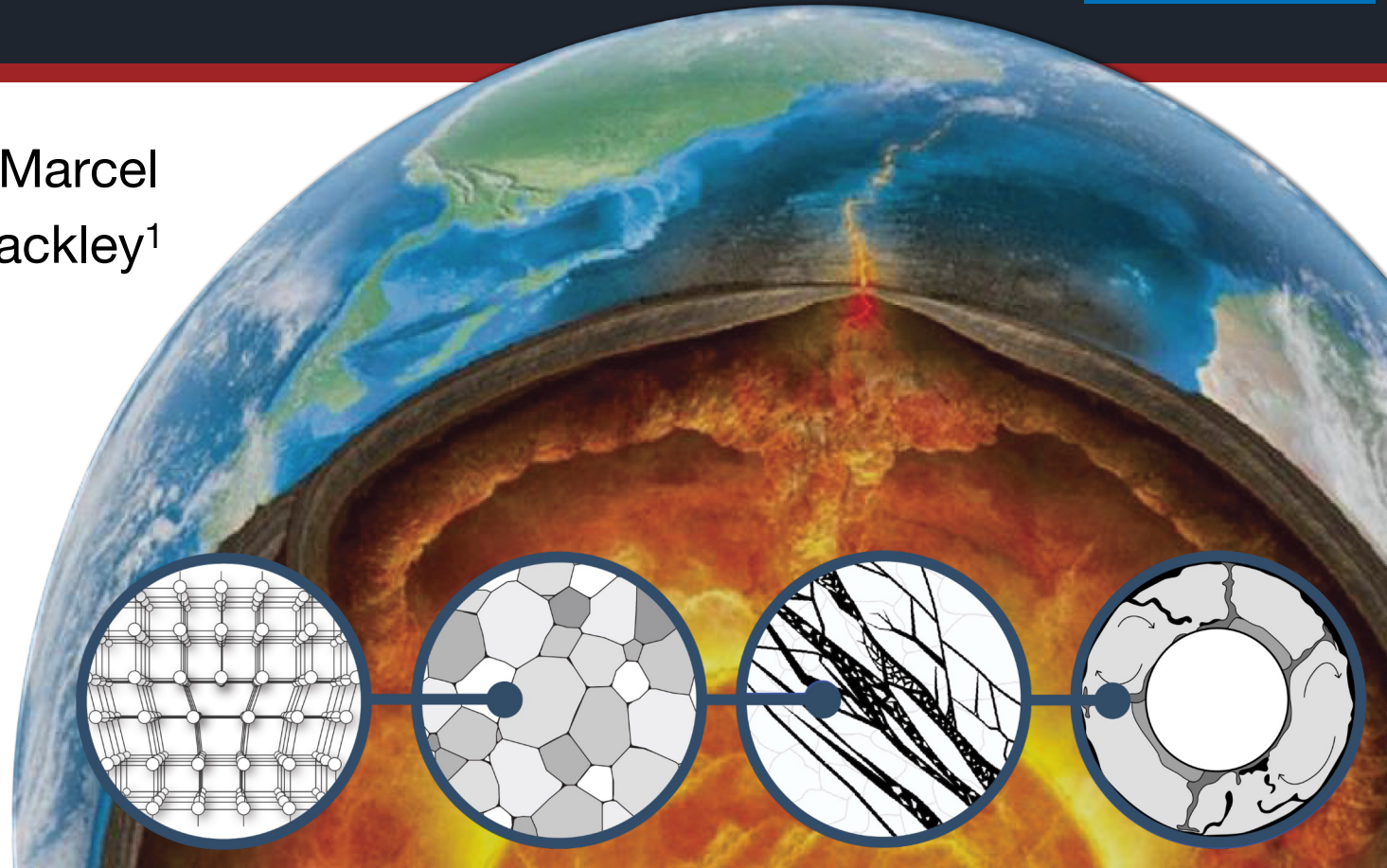


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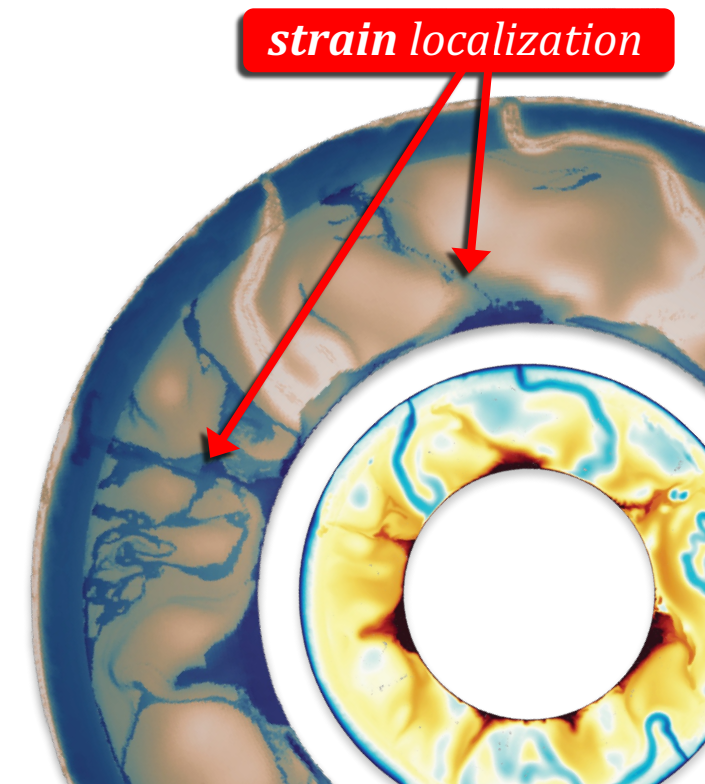
³ Dep. Of Earth Sciences, University College London, UK



Narrow, fast, and “cold” mantle plumes on Earth explained by **strain-weakening** rheology in the lower mantle

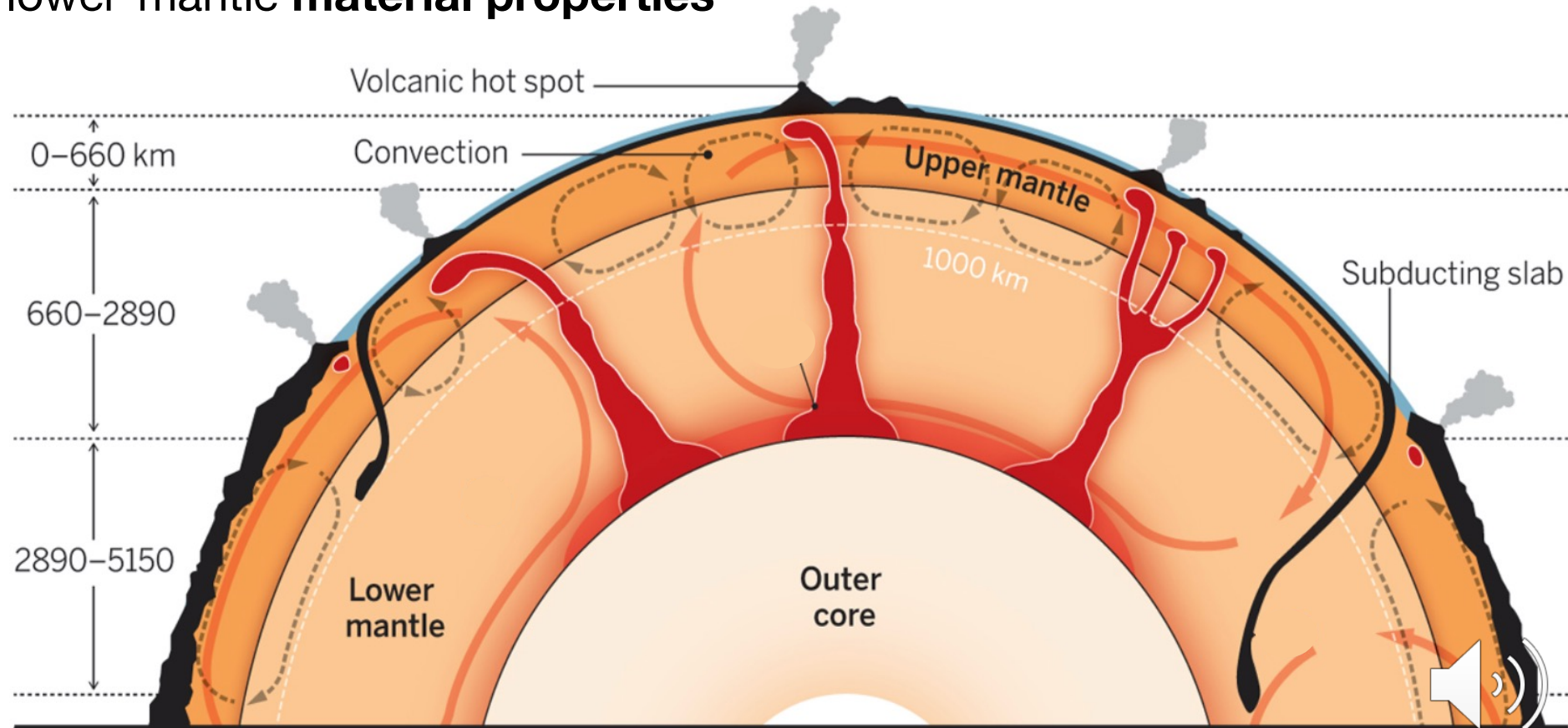
Article under review for G-cubed: <https://doi.org/10.1002/essoar.10509746.1>

- New parametrisation of lower-mantle **strain-weakening rheology** in mantle convection code
- Causes **weakening of plume conduits**, forming **narrow lubrication channels** through which hot material rapidly rises.
- Could help explain the **discrepancy** between expected and observed **thermal anomalies** of deep-seated mantle plumes on Earth?



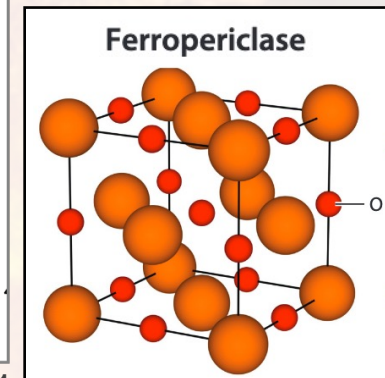
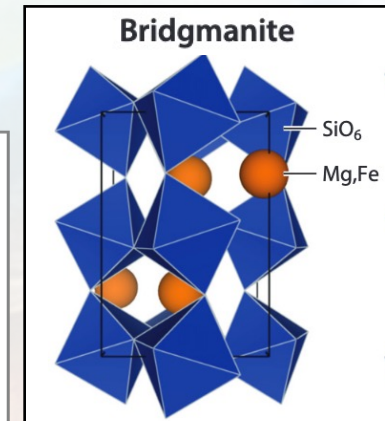
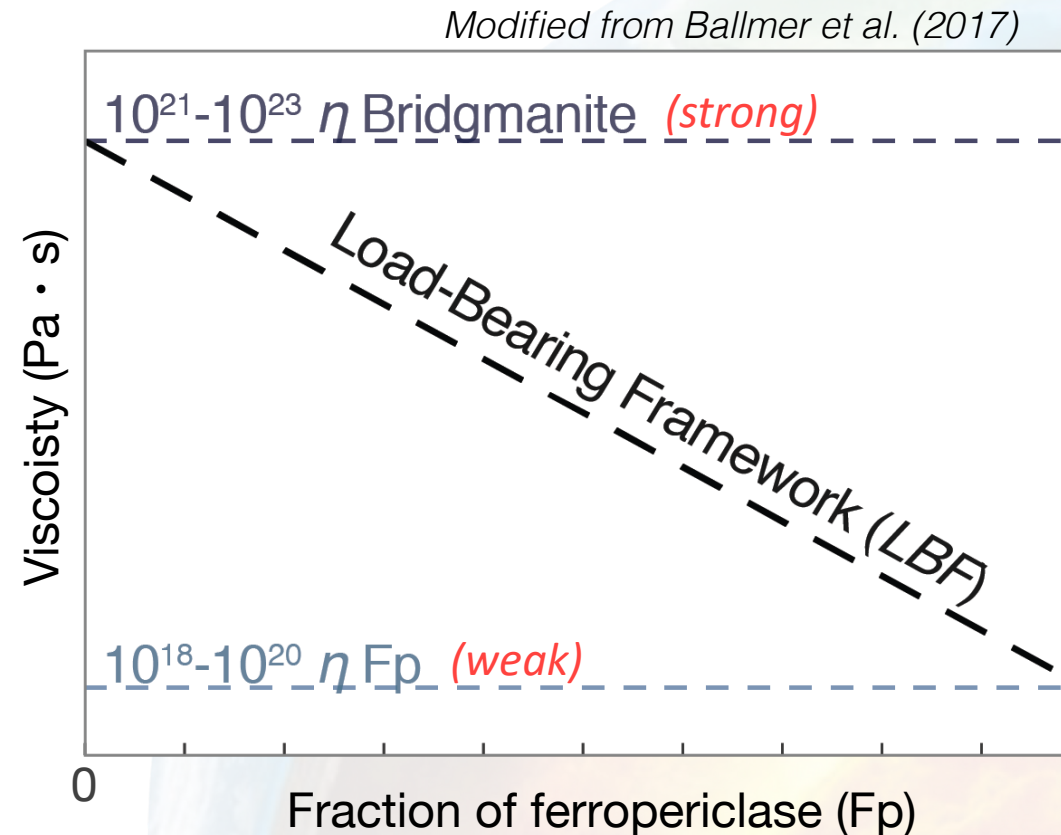
Shaping a planet by mantle flow

- Mantle: largest **geochemical reservoir** (60 vol%)
- Control of **mantle convection** and **mantle plumes**
- Dependent on lower-mantle **material properties**
(rheology)



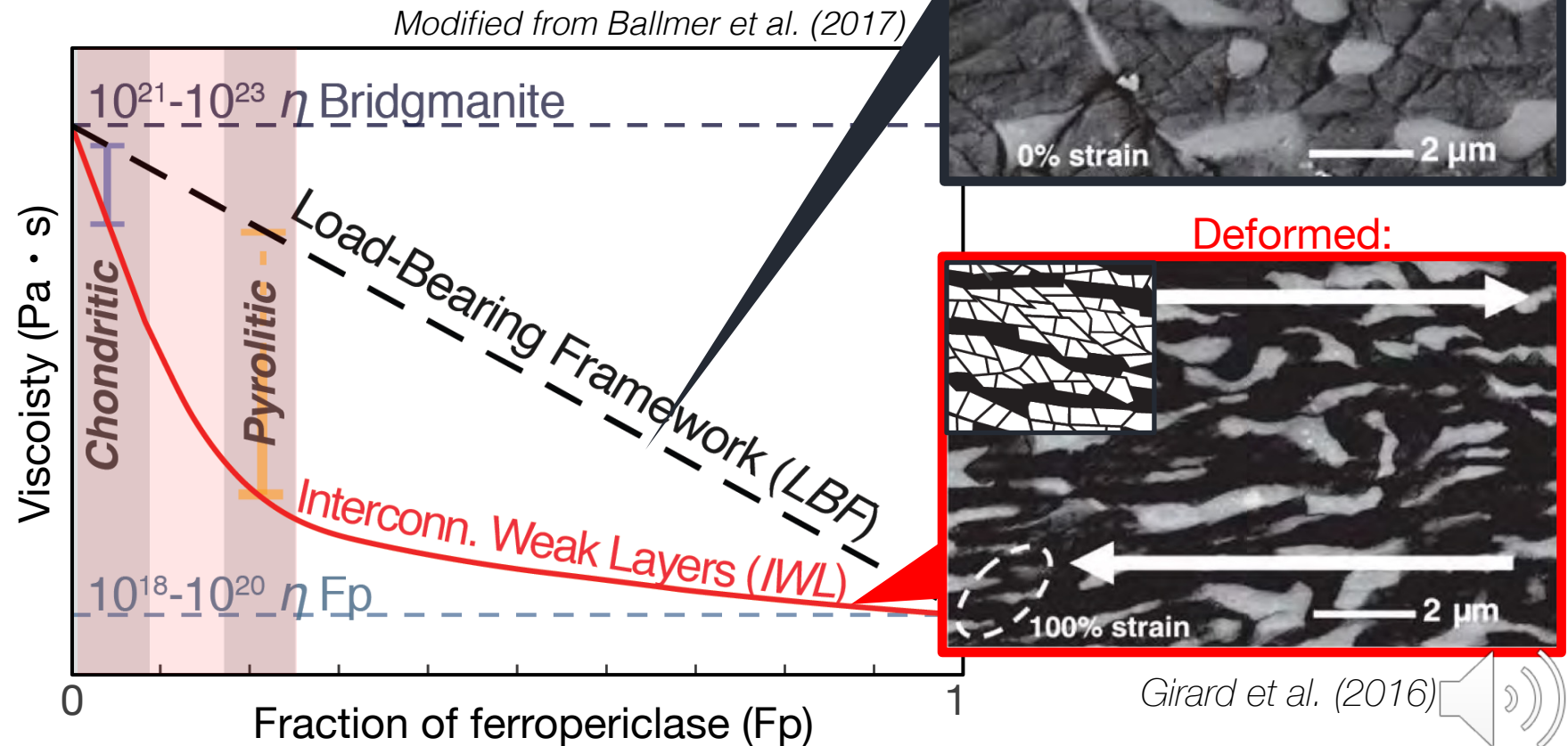
Earth's lower mantle - rheology

- >75% **Bridgmanite** $(\text{Mg,Fe})\text{SiO}_3$
- 0-25% **Ferropericlase** $(\text{Mg,Fe})\text{O}$
- few % other phases



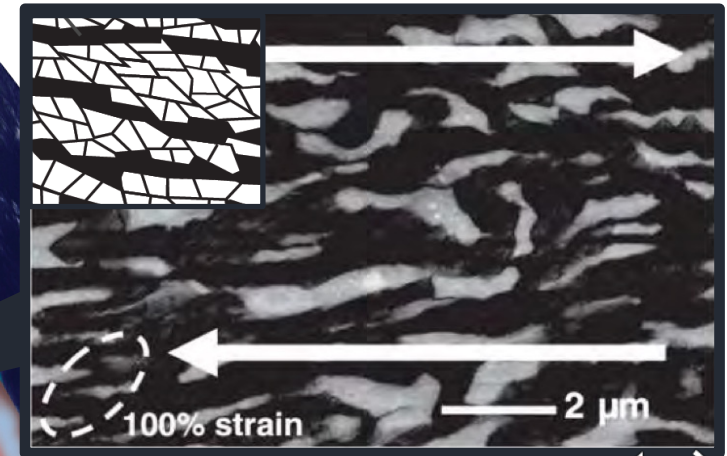
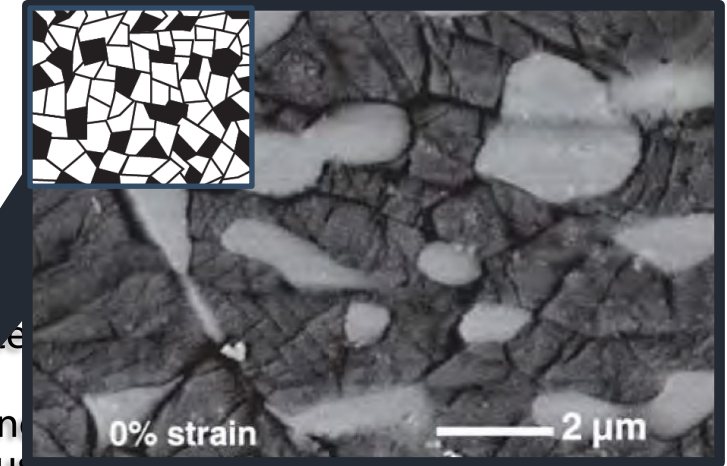
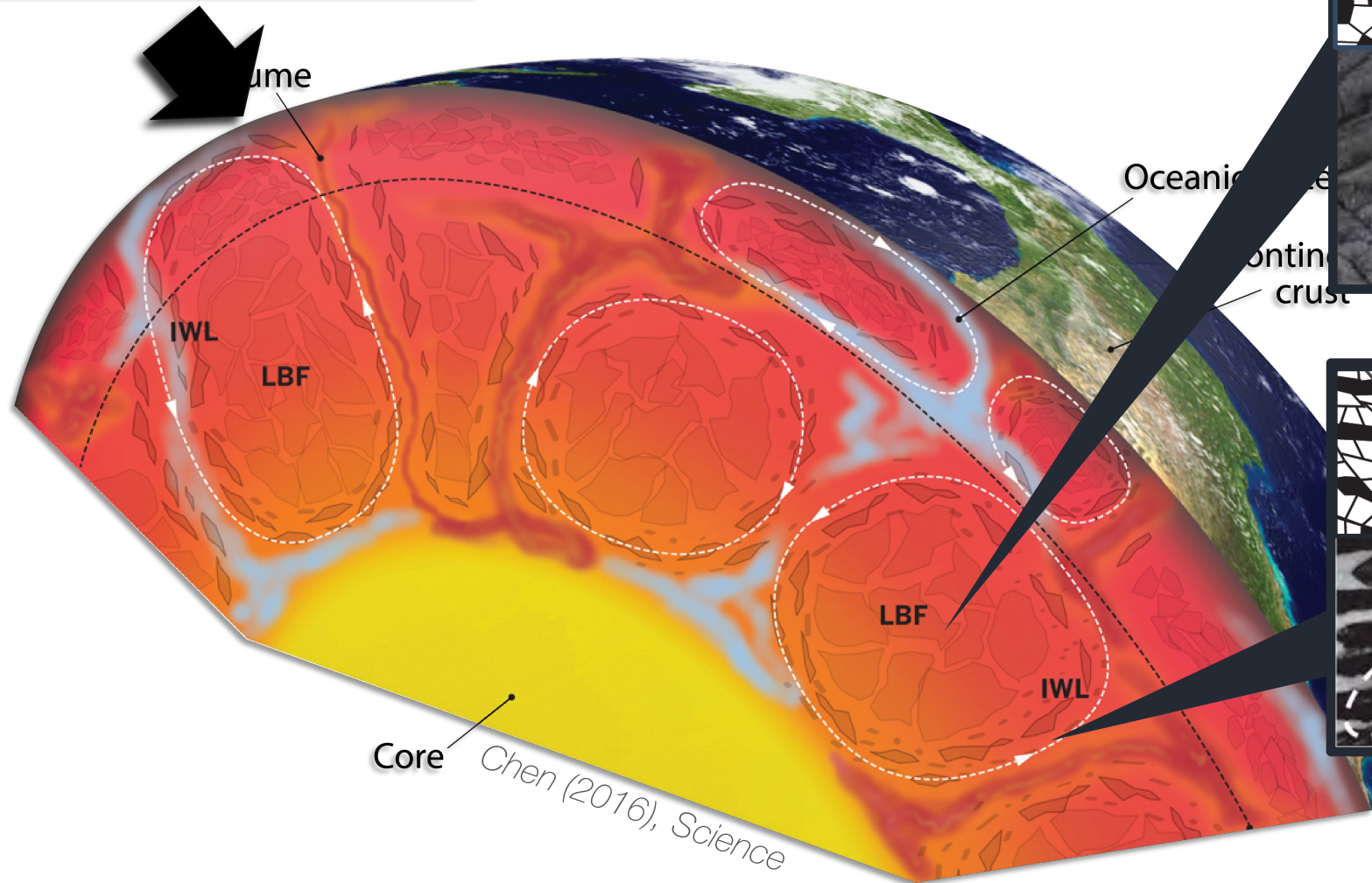
Earth's lower mantle - rheology

- Lower mantle rheology is **deformation history** (**strain-**) dependent!
- **Shear localization** in the lower mantle?



Earth's lower mantle – strain-dependent rheology

$$\eta(T, P, c, \epsilon, \dots)$$

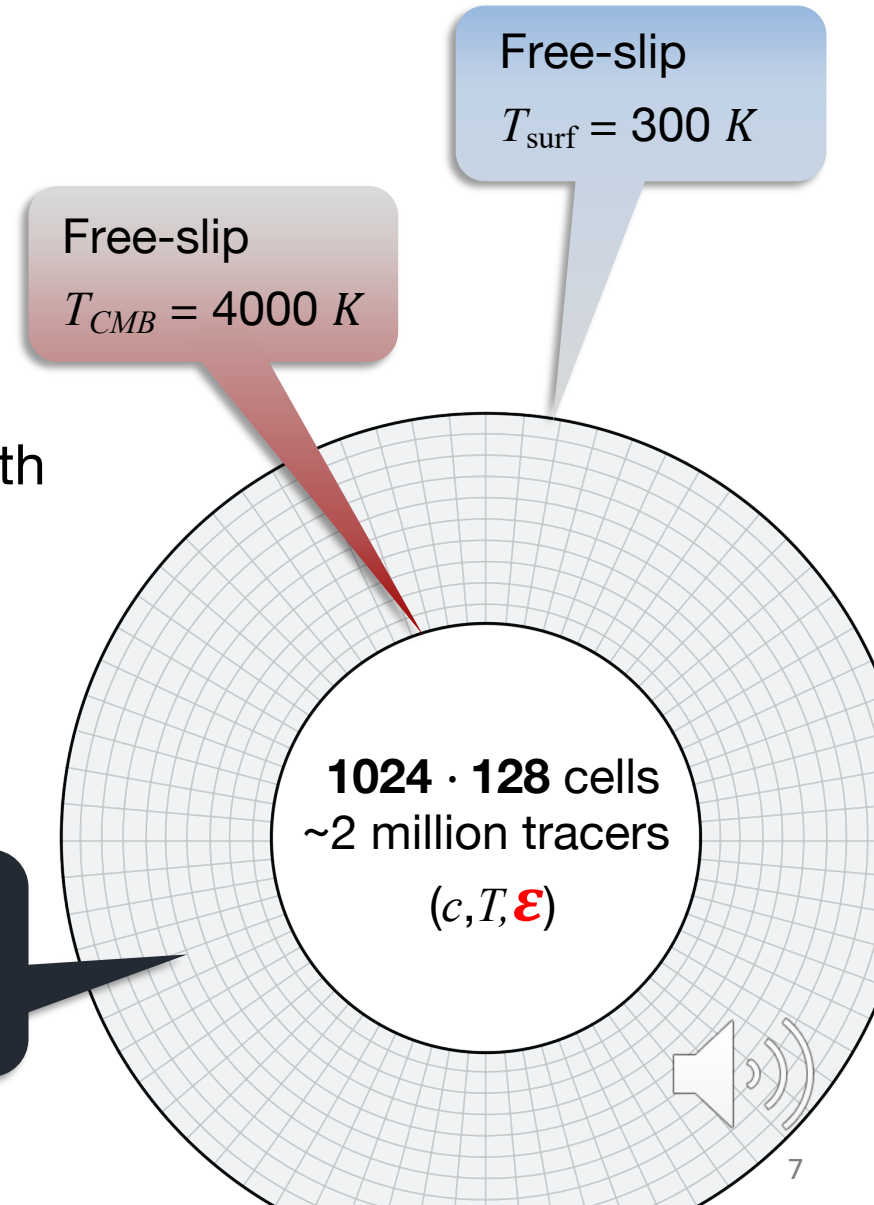


Girard et al. (2011)

Methods: geodynamic modelling (StagYY, *Hernlund and Tackley, 2008*)

- Mantle convection in **2D spherical annulus geometry**
- Pyrolitic mantle with **phase changes** and **melting**
(80% hz + 20% bs)
- **Visco-plastic rheology**: Arrhenius-type viscosity law with lithospheric yielding → '**Plate tectonic**' behavior
- Newly-implemented **strain-weakening rheology**

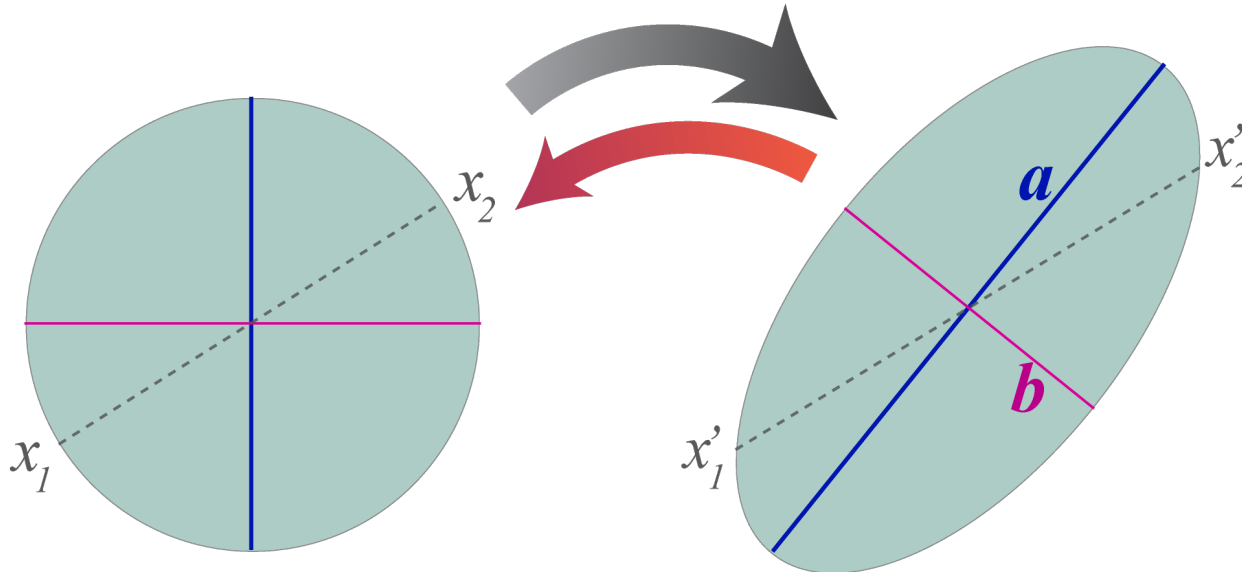
$$\eta(T, P, c, \boldsymbol{\varepsilon}) = \eta_0 \lambda_c \boldsymbol{f}_w(\boldsymbol{\varepsilon}) \exp\left(\frac{E_a + PV_a}{RT} - \frac{E_a}{RT}\right)$$



Methods: tracking strain

- The **deformation tensor** M is tracked on each tracer (**finite deformation**, McKenzie, 1987)
- $MM^T \rightarrow$ **strain ellipse**
- Strain is **reset** ($M = I$) at 660 boundary (Solomatov & Reese, 2008)

$$dM/dt = -\mathbf{H}(M - I)$$



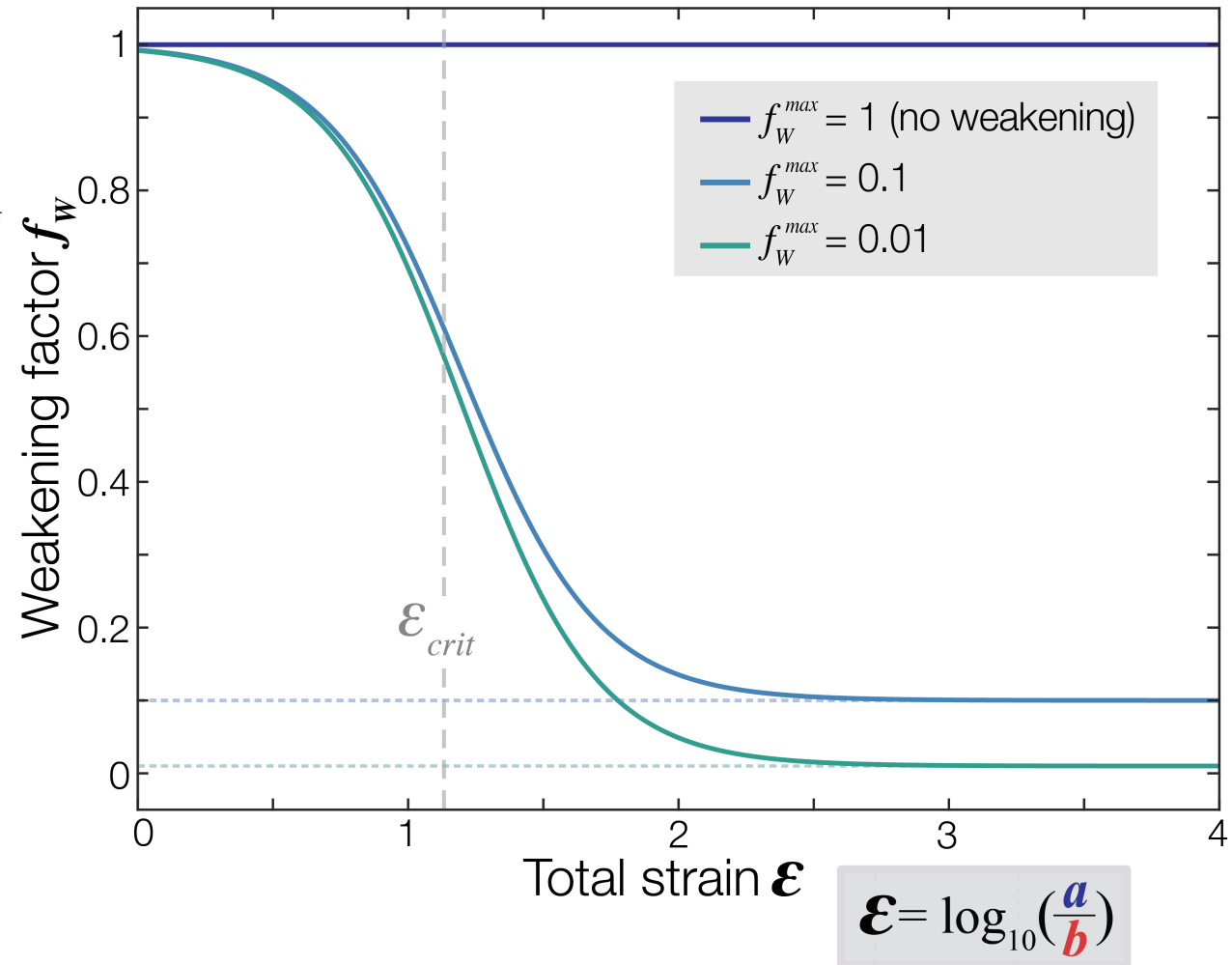
- **Strain definition:** $\mathcal{E} = \log_{10}\left(\frac{a}{b}\right)$
- **Healing term** $H(P, T)$ **relaxes** M and lowers strain (annealing, grain growth, ..)



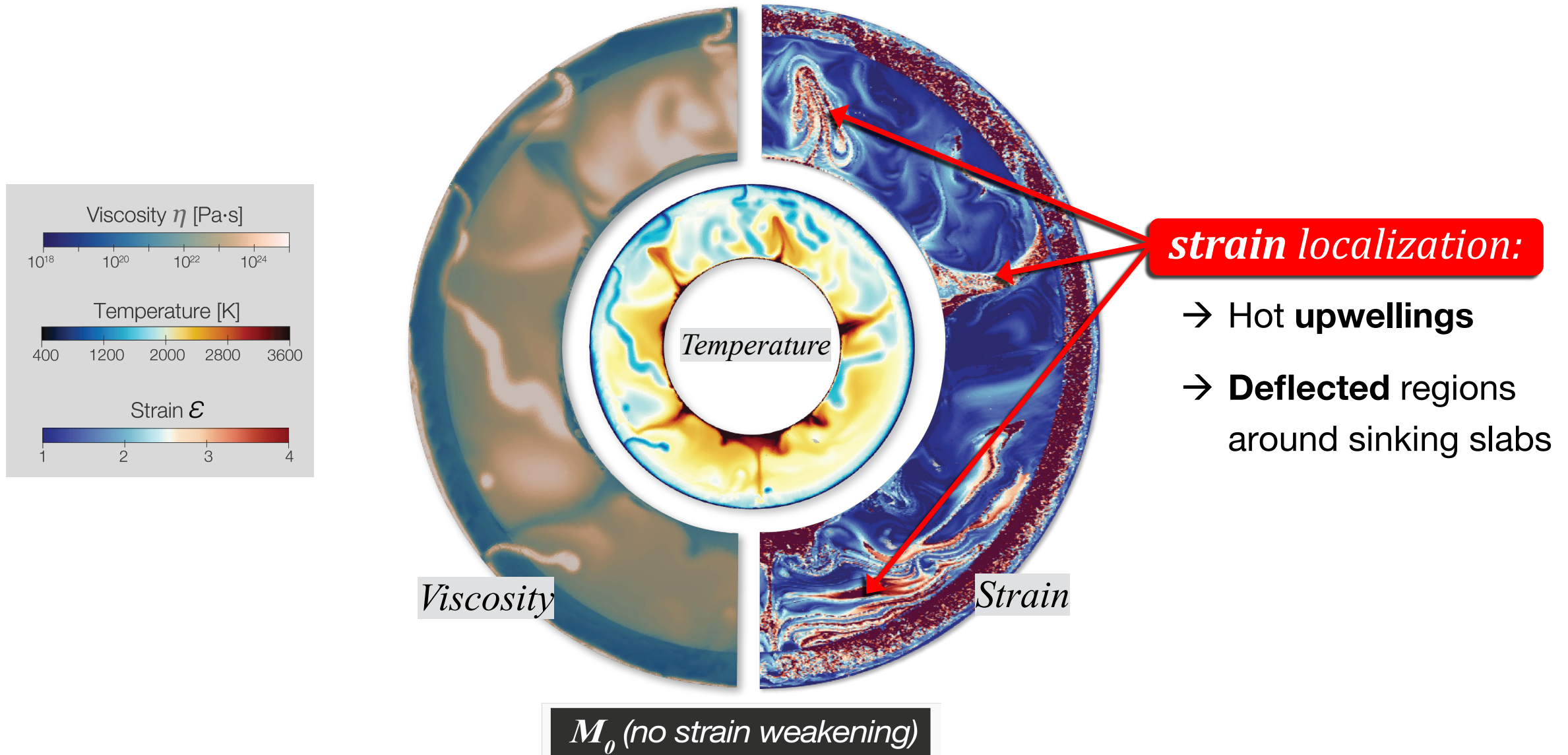
Methods: strain-dependent weakening

- Lower-mantle material may **weaken** according to:

$$\eta(T, P, c, \boldsymbol{\varepsilon}) = \eta_0 \lambda_c \boldsymbol{f}_W(\boldsymbol{\varepsilon}) \exp(\dots)$$

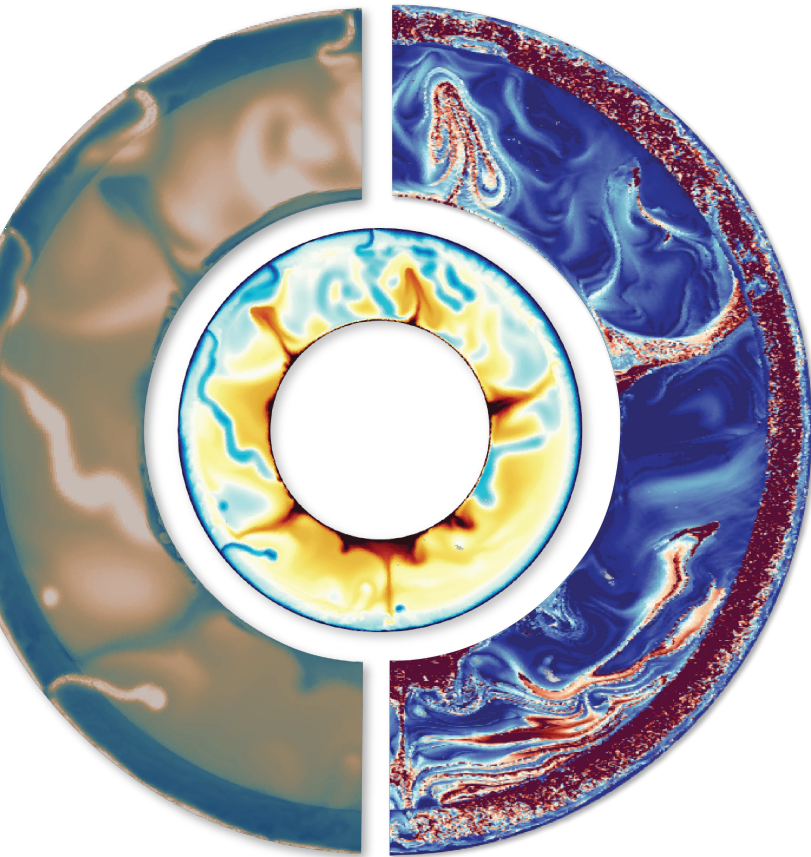


Results: reference model

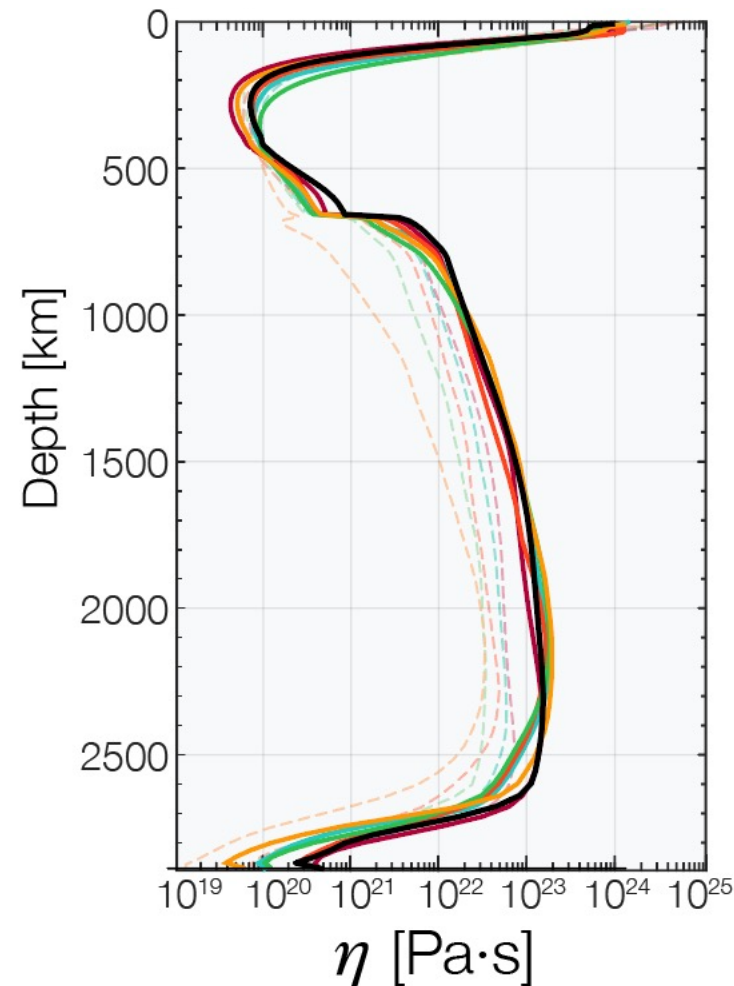


Results: SW rheology - comparison

Remove effect of **changing Rayleigh number** → Establish **first-order effect** of SW rheology



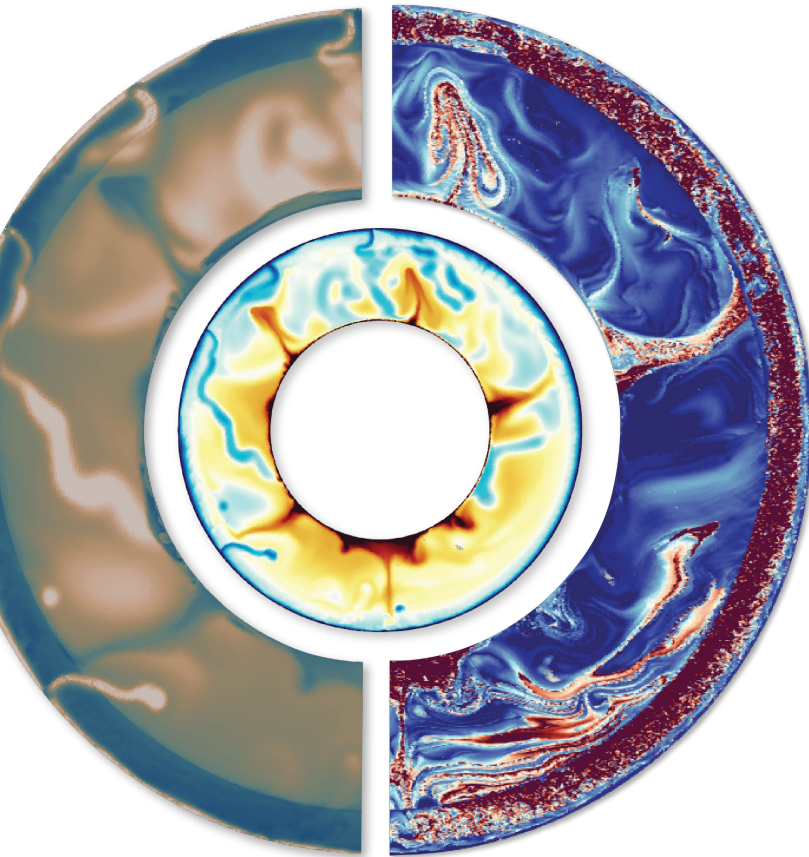
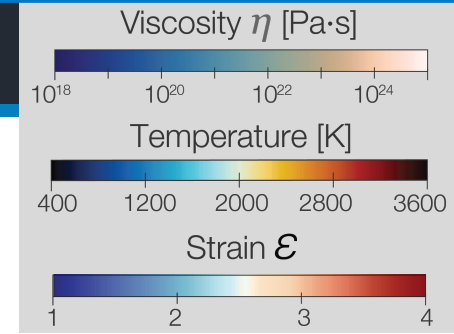
M_0 (no strain weakening)



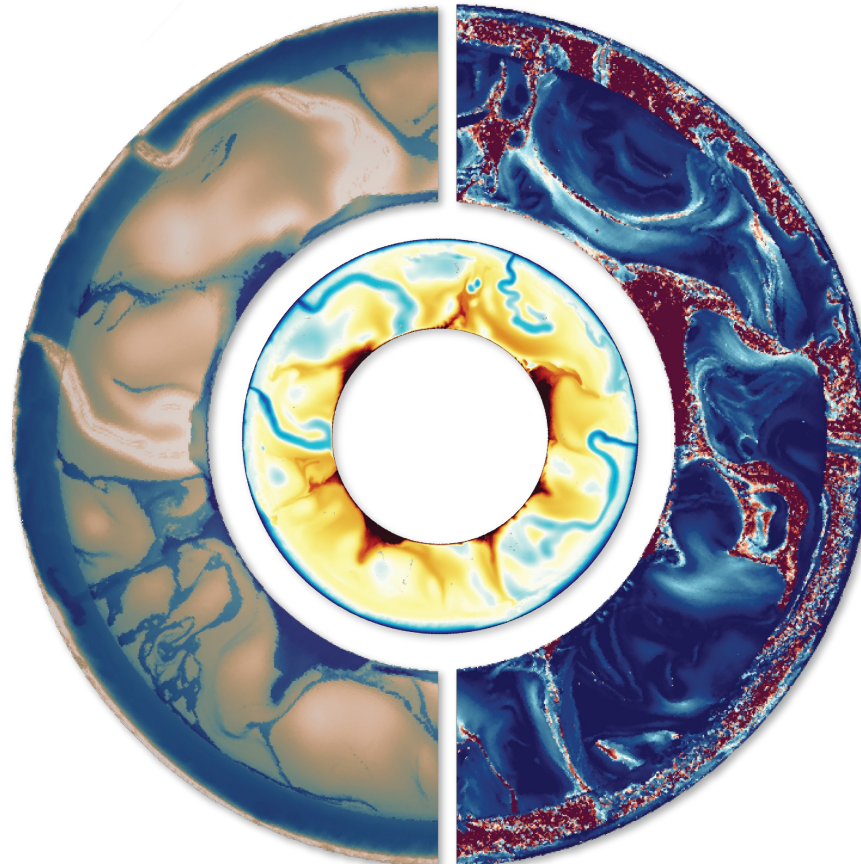
- Models with similar...
... final **viscosity profile**
... **thermal evolution**



Results: SW rheology – comparison

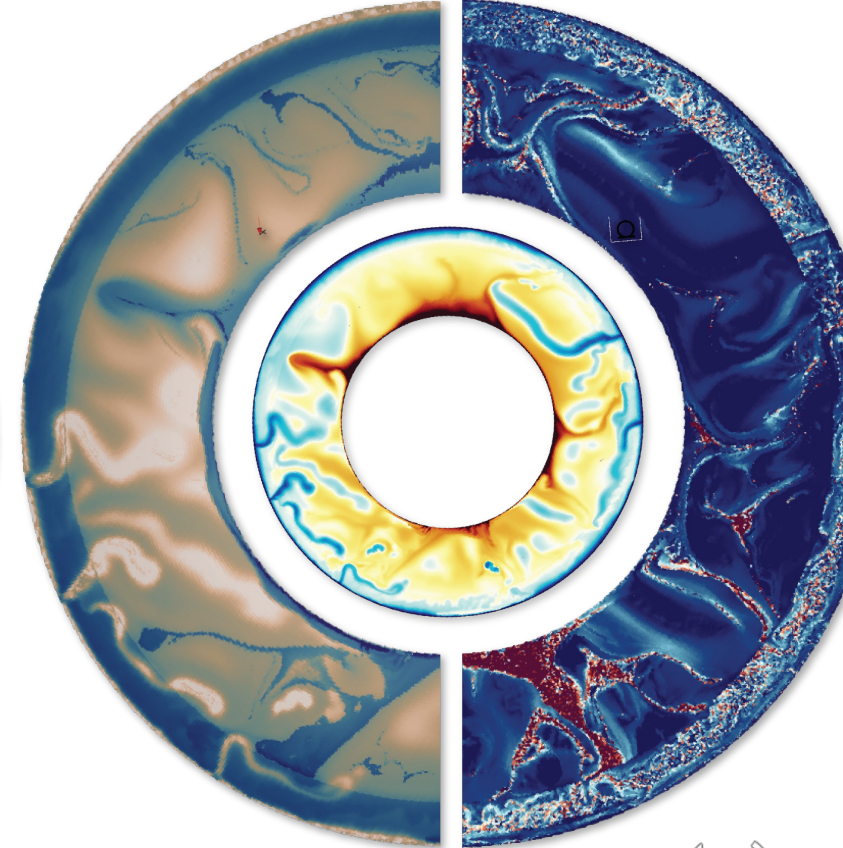


M_0 (no strain weakening)



No healing; much weakening

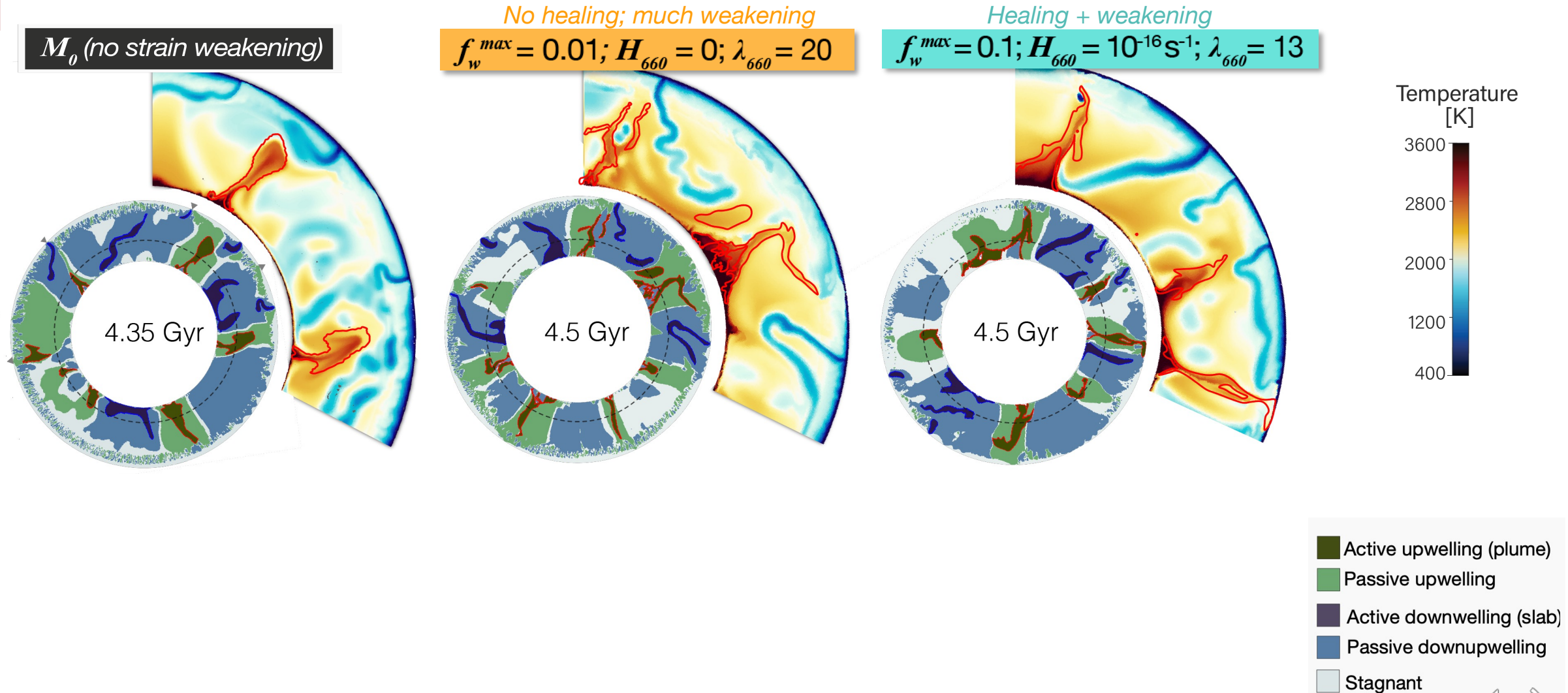
$f_w^{\max} = 0.01$; $H_{660} = 0$; $\lambda_{660} = 20$



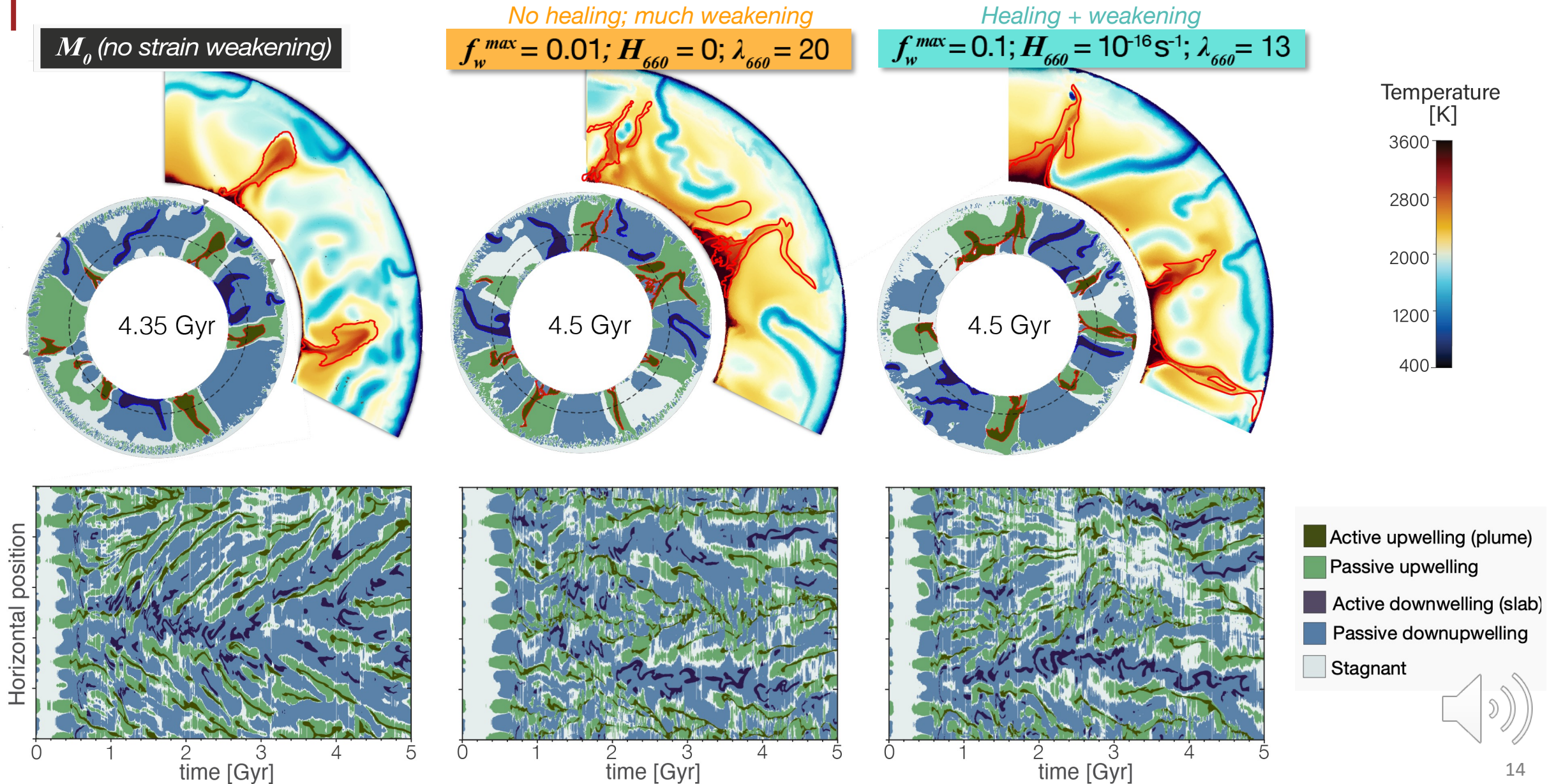
Healing + weakening

$f_w^{\max} = 0.1$; $H_{660} = 10^{-16} \text{ s}^{-1}$; $\lambda_{660} = 13$

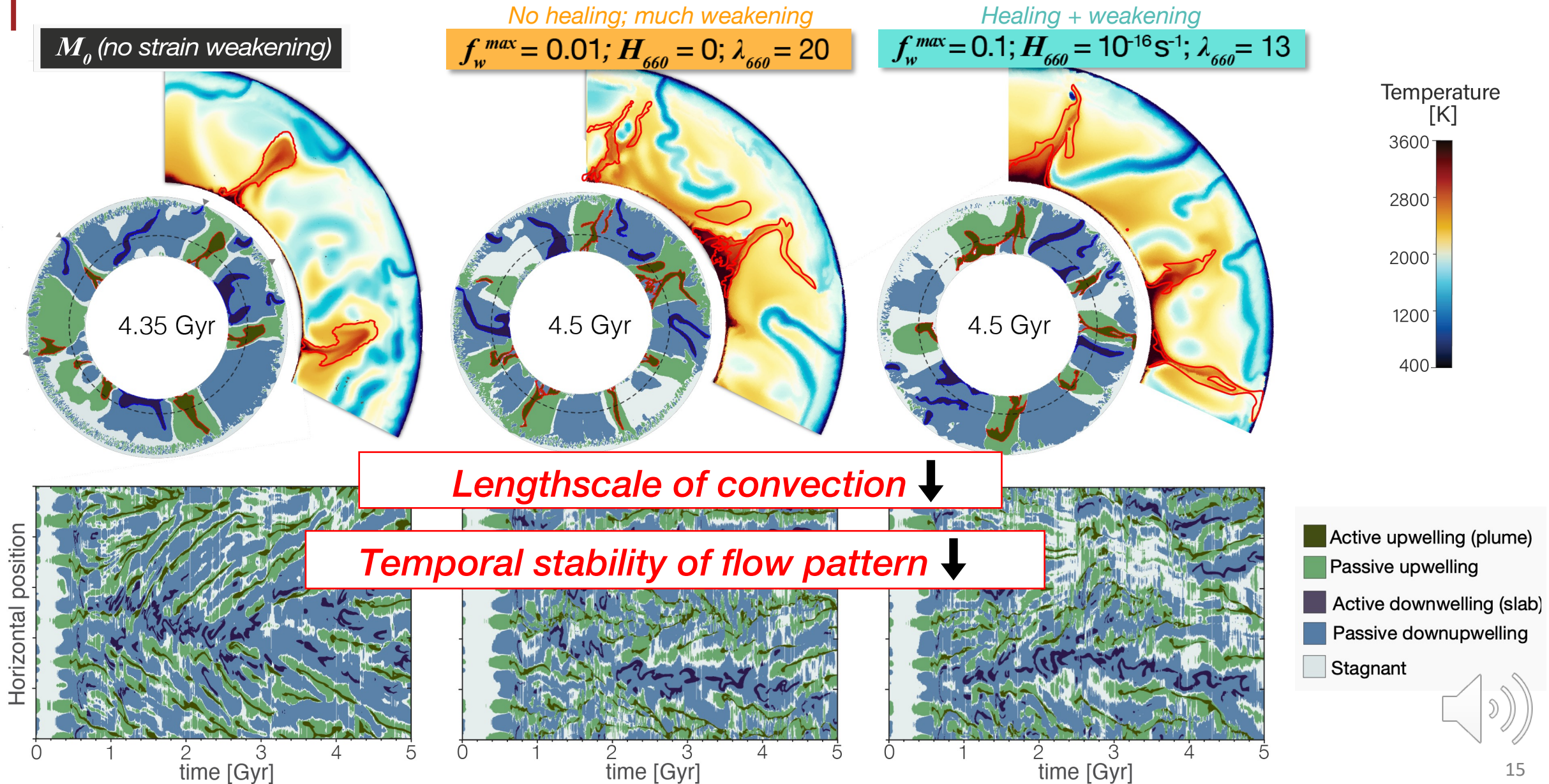
Results: SW rheology – mantle convective pattern



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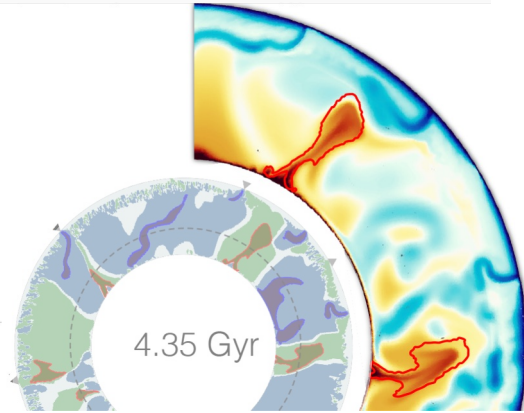


Results: SW rheology – mantle convective pattern



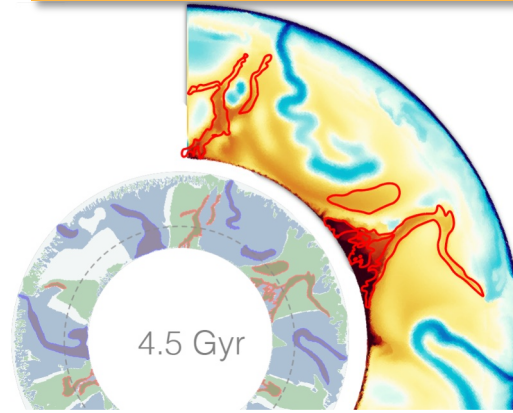
Results: SW rheology – mantle plumes

M_0 (no strain weakening)



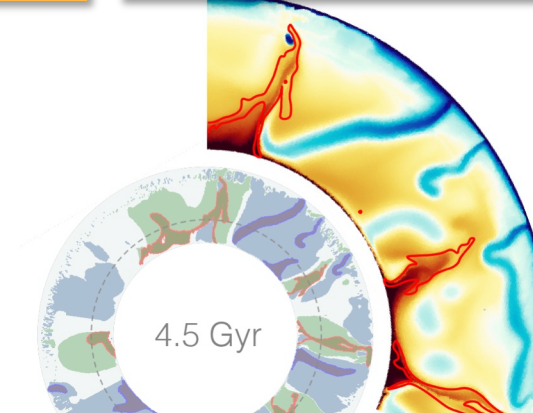
No healing; much weakening

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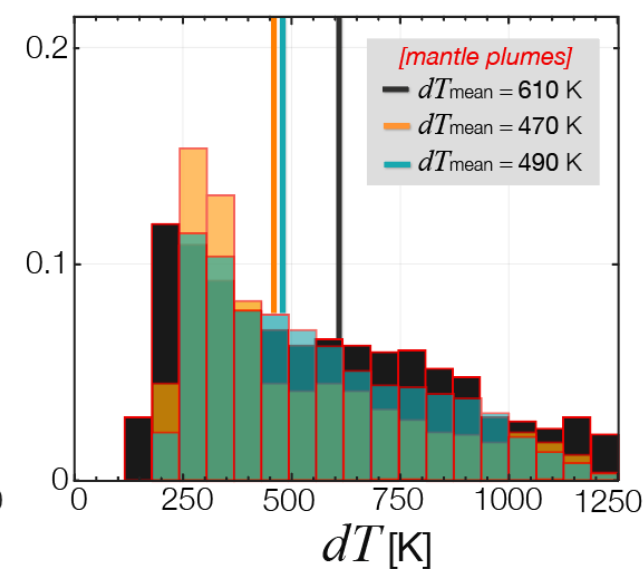
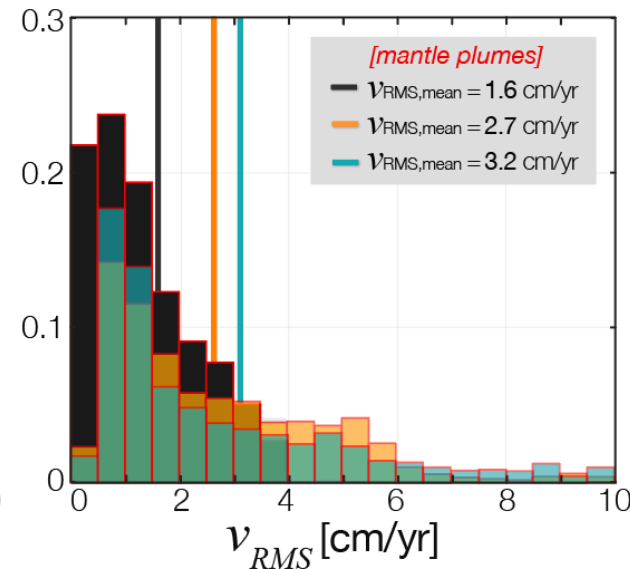
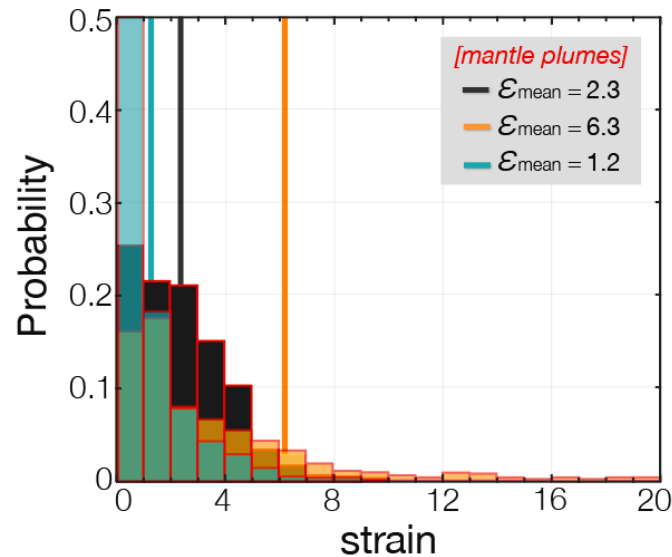
Healing + weakening

$$f_w^{\max} = 0.1; H_{660} = 10^{-16} \text{ s}^{-1}; \lambda_{660} = 13$$



Weakened plumes are:

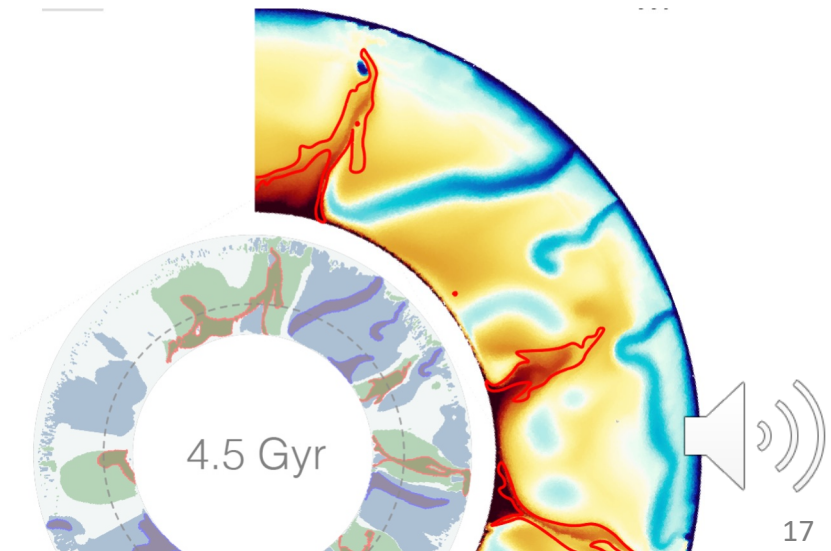
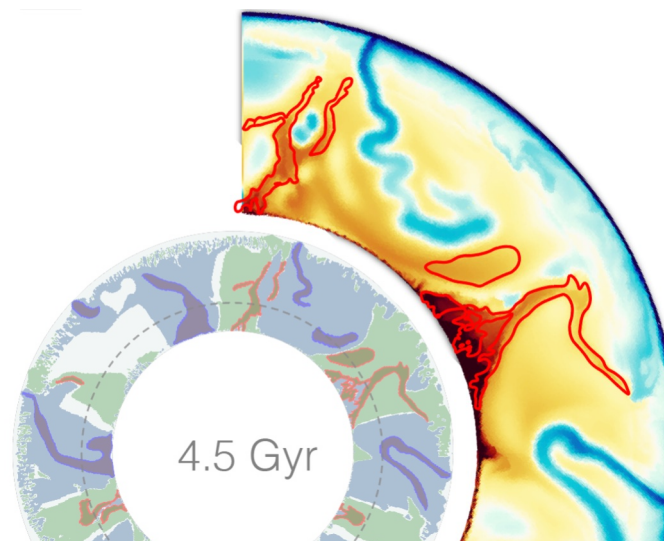
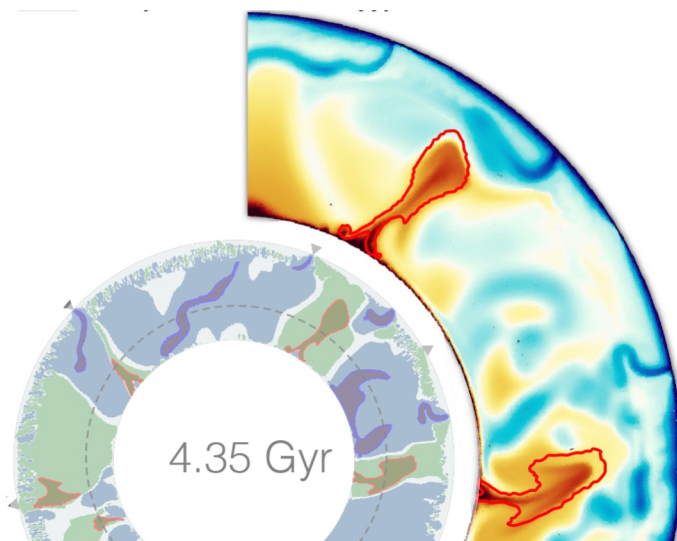
- **Thinner**
- **Faster**
- **Colder**



Results: SW rheology – summary

SW rheology in the lower mantle:

- **Reduced length-scale** of mantle convection
- Increased thermochemical **pile stability**
- Thinner, faster, and **colder mantle plumes**



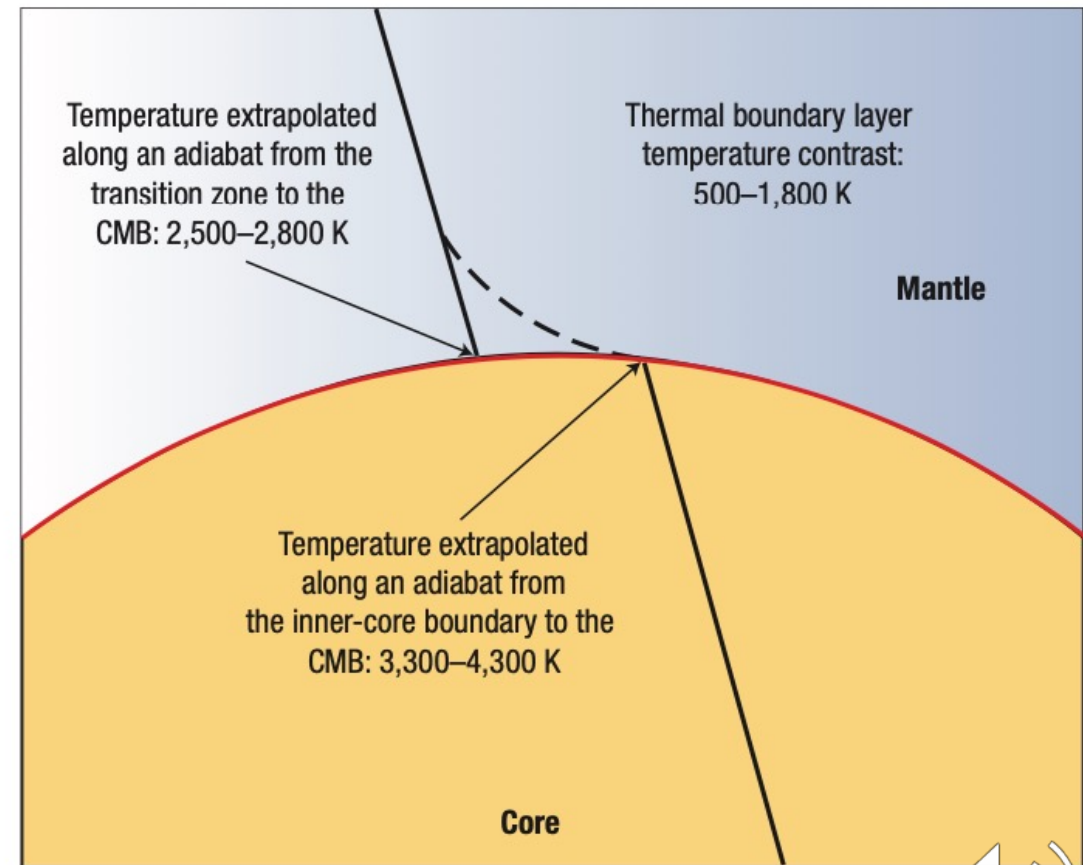
Discussion – mantle plumes

- Weakened plumes are: **faster, colder, & thinner**

- Earth observations:
 - **1000-1500 K** super-adiabatic temperature rise across CMB (Jeanloz & Morris, 1986; Lay et al., 2008)

↑↓ Mismatch!

- Deep-seated plume T excess of only **100-400 K** (e.g., Albers & Christensen, 1996)

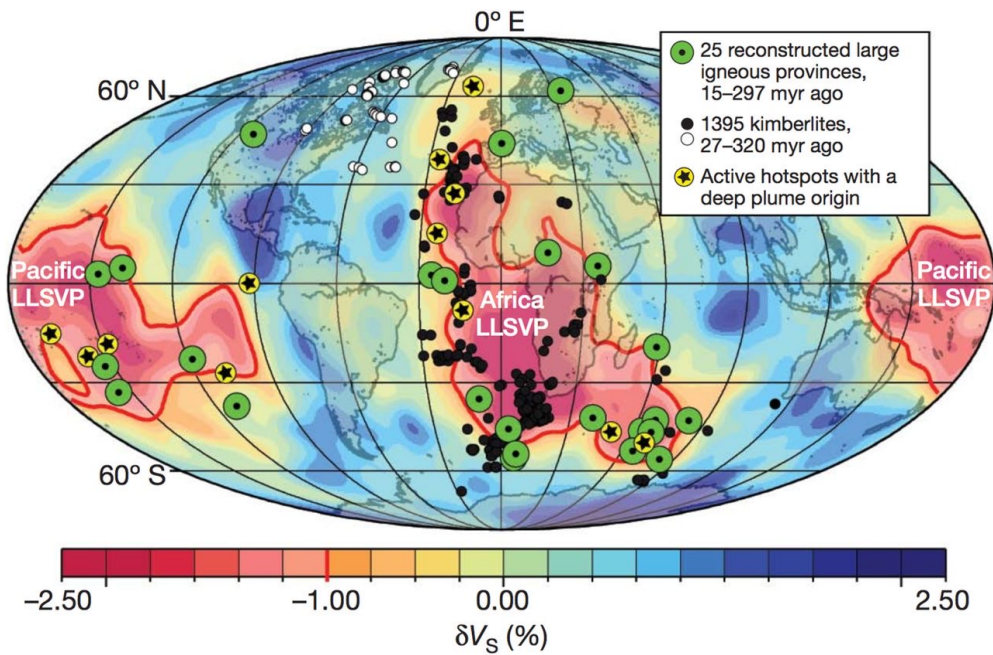


Lay et al., (2008)

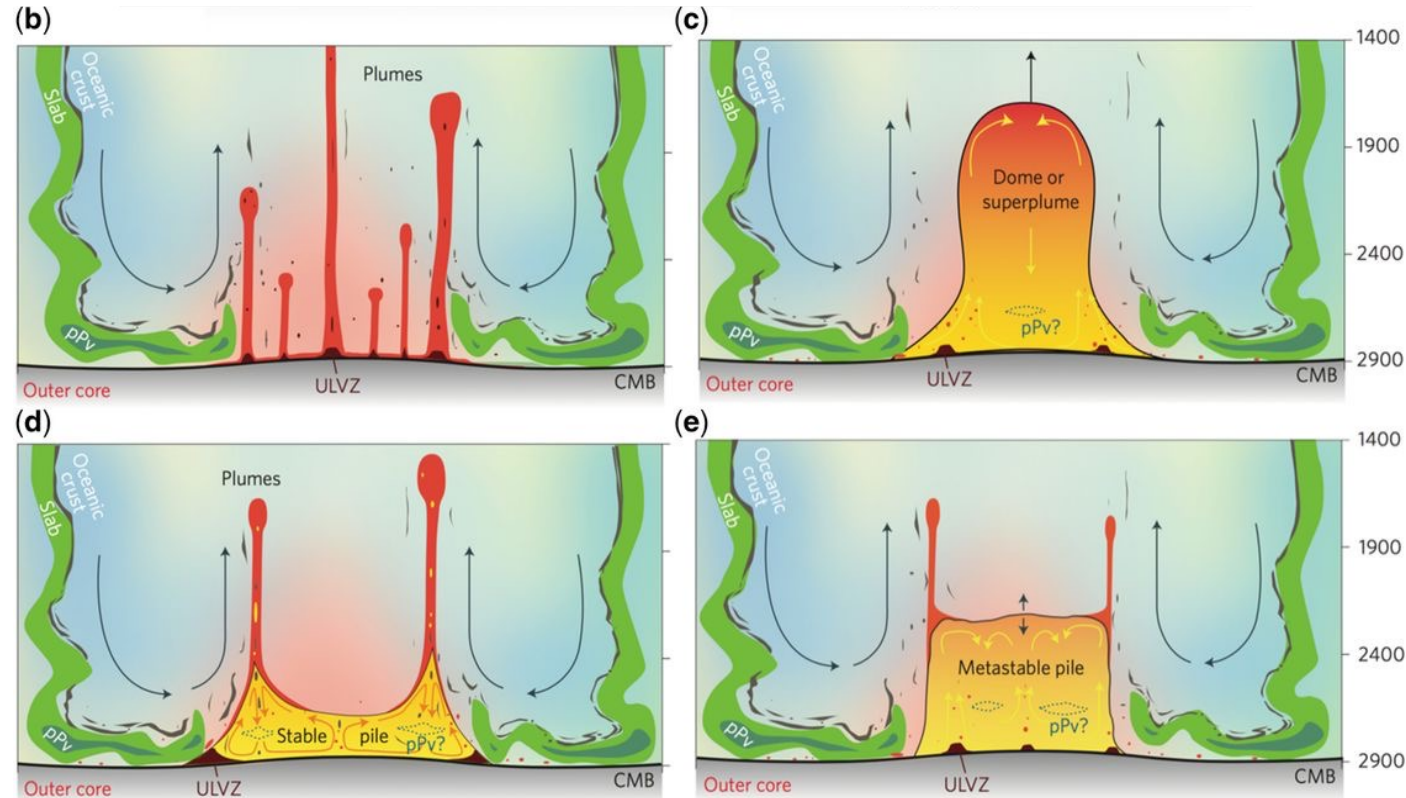


Discussion – mantle plumes

- Weakened plumes are: **faster, colder, & thinner**



Torsvik et al. (2010)



Garnero et al. (2016)



Discussion – mantle plumes

- Weakened plumes are: **faster, colder, & thinner**

LETTER

doi:10.1038/nature14876

Broad plumes rooted at the base of the Earth's mantle beneath major hotspots

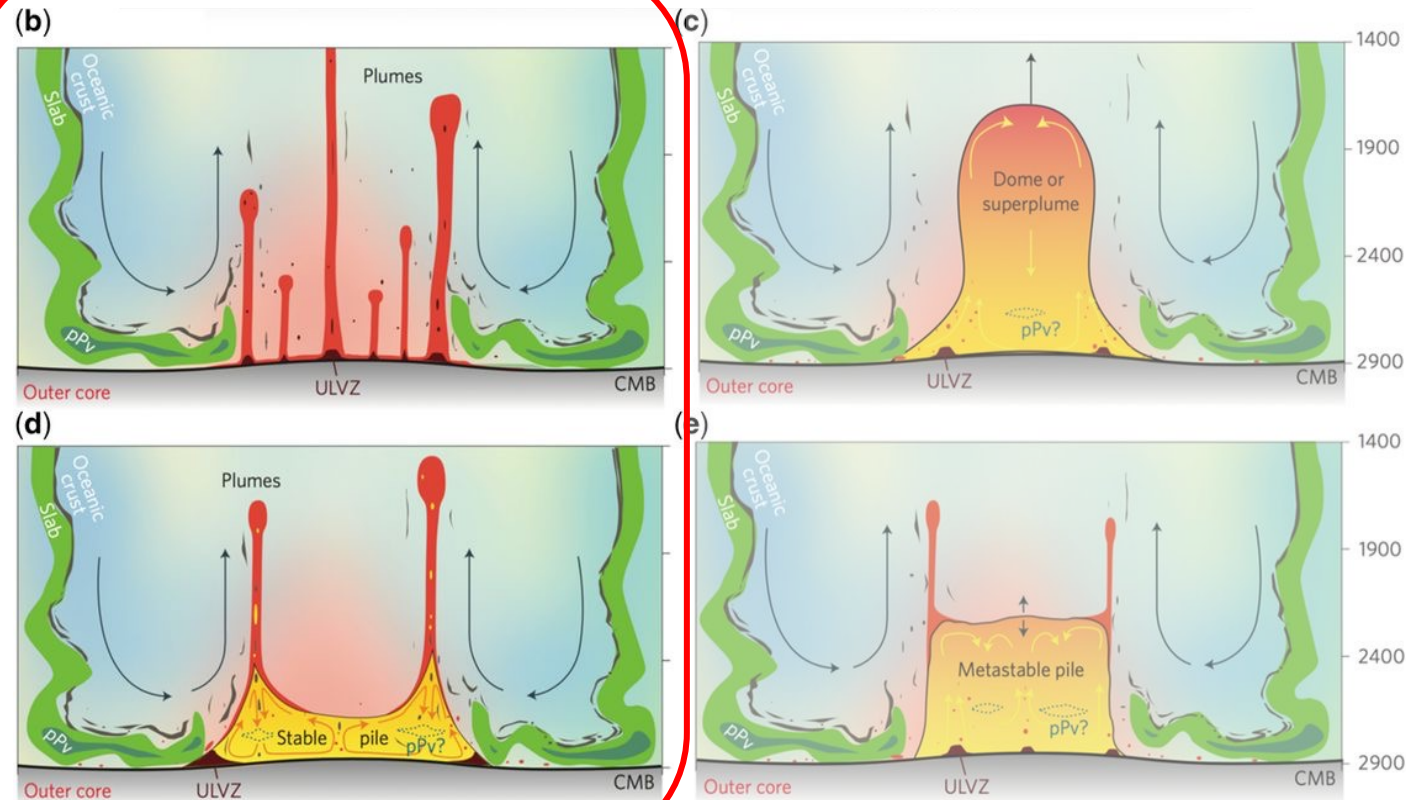
Scott W. French^{1†} & Barbara Romanowicz^{1,2,3}

Tectonics

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Deflating the LLSVPs: Bundles of Mantle Thermochemical Plumes Rather Than Thick Stagnant “Piles”

Anne Davaille¹ and Barbara Romanowicz^{2,3,4}



Garnero et al. (2016)



Key points

<https://doi.org/10.1002/essoar.10509746.1>

- A new parametrisation of **strain-dependent rheology** for lower mantle materials is implemented in *StagYY*
- Such rheology particularly causes **weakening of plume conduits**, forming **narrow lubrication channels** through which hot material easily rises.
- Could help explain the **discrepancy** between expected and observed **thermal anomalies** of deep-seated mantle plumes on Earth?

