

Stochastic Modelling and Location Uncertainty (LU) formalism: hydrostatic Boussinesq equations and their implementation in NEMO

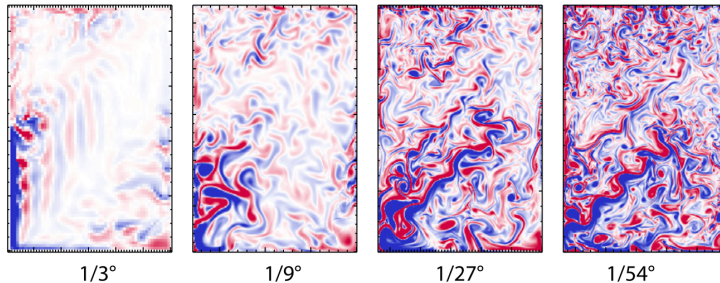
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The accuracy of a simulation is intrinsically dependent on the chosen resolution¹.



Fixed δ representing the resolution, there will always be a distance between the real solution \mathbf{u} and its approximation \mathbf{u}_δ .

This distance exists also in the physics that is representable by the simulation, due to non-linear interactions between scales that propagate information to scales smaller than the computed resolution.

Can we improve the solution at a coarse resolution \mathbf{u}_Δ by modelling what is left behind from a finer information \mathbf{u}_δ ?

¹Image taken from Levy 2012, representing vorticity for an idealised jet stream configuration.

Location Uncertainty [1] decomposition principle:

$$dX_t^i(\mathbf{x}_0) = \underbrace{u_i(\mathbf{X}_t, t) dt}_{\text{Resolved}} + \underbrace{\eta_t^i(\mathbf{X}_t)}_{\text{Unresolved}},$$

where the resolved component is dependent on the resolution chosen, the unresolved component accounts for turbulent effects, truncation and approximation effects, uncertainty in the forcing and initial conditions;

This introduces the **stochastic material derivative**:

$$\mathbb{D}_t \theta = d_t \theta + \underbrace{[\mathbf{u}^* dt + \boldsymbol{\eta}] \cdot \nabla \theta}_{\text{Noise advection}} - \underbrace{\frac{1}{2} \nabla \cdot (\mathbf{a} \nabla \theta) dt}_{\substack{\text{Diffusion by Unresolved} \\ \text{scales}}}.$$

where

$$\mathbf{u}^* = \mathbf{u} - \mathbf{u}_s,$$

with \mathbf{u}_s accounting for heterogeneities.

LU Boussinesq

Horizontal momentum:

$$\mathbb{D}_t \mathbf{v} + f \mathbf{e}_3 \times \left(\mathbf{v} dt + \frac{1}{2} \boldsymbol{\eta}_t^H \right) = -\nabla_H (\pi dt + dp_t^\sigma)$$

Vertical momentum:

$$\mathbb{D}_t w = -\partial_z (\pi dt + dp_t^\sigma) + b dt$$

Tracer (Temperature and salinity):

$$\mathbb{D}_t \tau = \kappa_\tau \Delta \tau dt,$$

Incompressibility:

$$\nabla \cdot [\mathbf{u} - \mathbf{u}_s] = 0, \quad \nabla \cdot \boldsymbol{\eta} = 0,$$

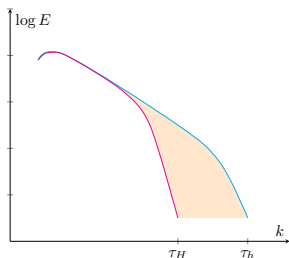
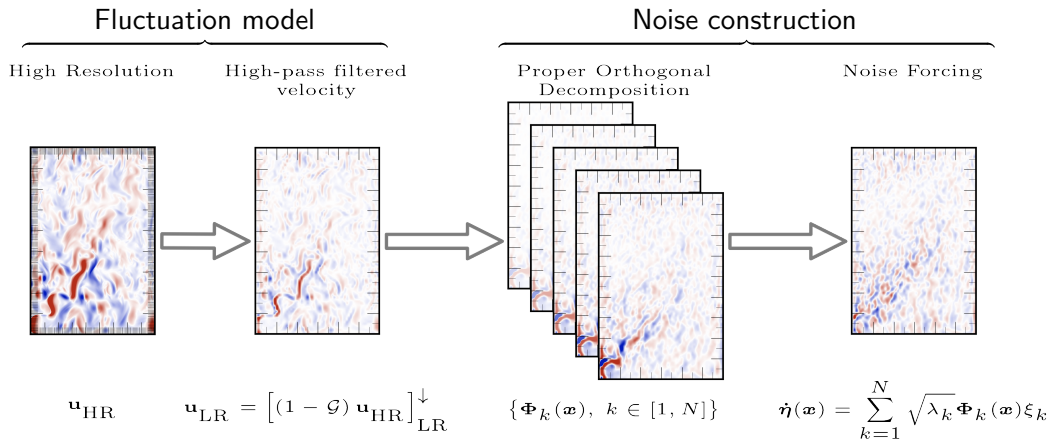
Equation of state:

$$b = b(T, S, z).$$

with

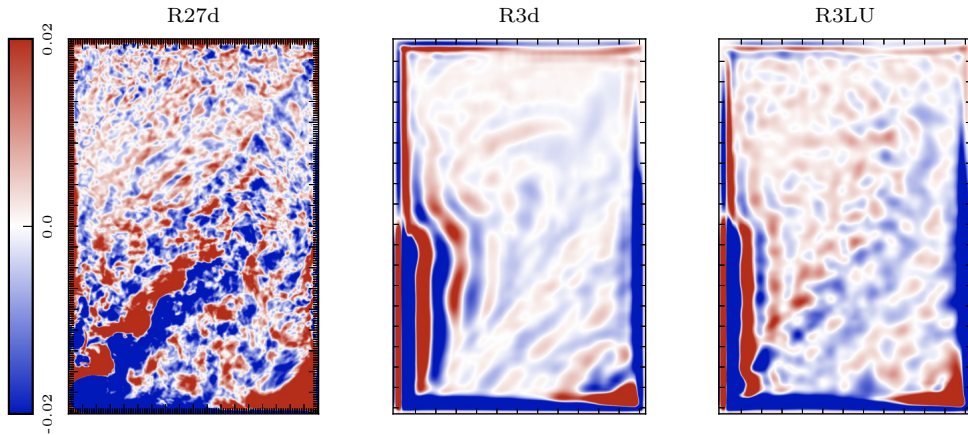
$$\pi = p' - \frac{\nu}{3} \nabla \cdot \mathbf{u}_s$$





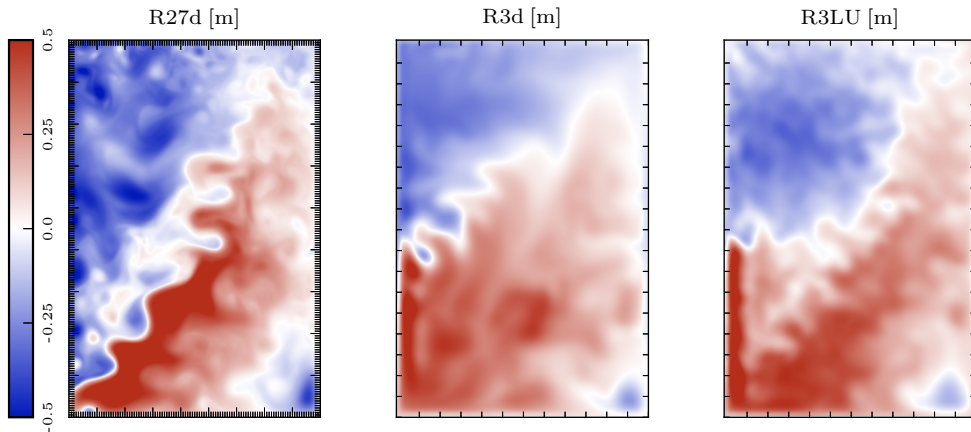
The **fluctuation model** should carry the information that cannot be represented with a coarse simulation, i.e. the fraction of energy represented with the shaded area. The noise should be constructed to carry this partial information rather than the whole velocity field information

10-years averaged relative vorticity $\zeta = (\partial_x v - \partial_y u) / f$ at the surface layer of the model.



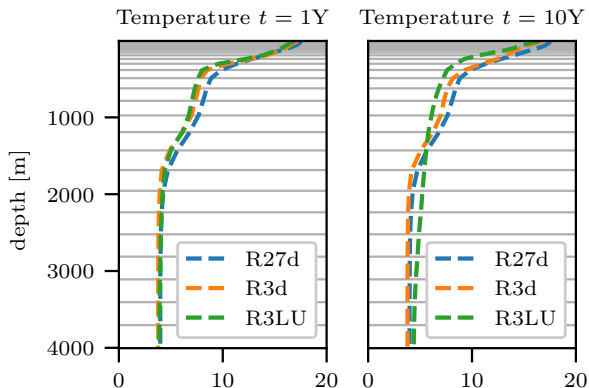
Improvements: Better resolution of Vorticity.

5-days averaged sea surface height.



Improvements: Better resolution of Sea Surface Height.

Drawbacks: Increase in vertical diffusion and transport in tracers, leading to fields far from the reference R27 simulation.



Possible motivations to investigate:




1. No feedback from u^t velocity;
2. Two-dimensional noise not adequate;

Future works:

1. different types of noise (3D POD, DMD, Local decompositions...);
2. hydrostatic stochastic pressure model;
3. barotropic-baroclinic splitting for the noise;
4. realistic configurations.

Thank you for your attention.

References

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-  Tucciarone, F. L., Mémin, E., Li, L.: Primitive equations under location uncertainty: analytical description and model development. STUOD proceedings.
-  Lévy, M., Resplandy, L., Klein, P., Capet, X., Iovino, D., Ethé, C.: Grid degradation of submesoscale resolving ocean models: Benefits for offline passive tracer transport. Ocean Modelling, 1–9, (2012).