

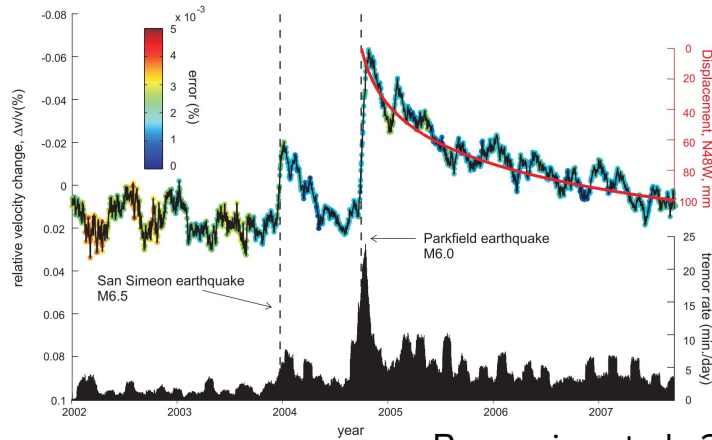
Relaxation timescales after small and large earthquakes:

Similarity and controls from seismic velocity changes estimated in Patache, Chile

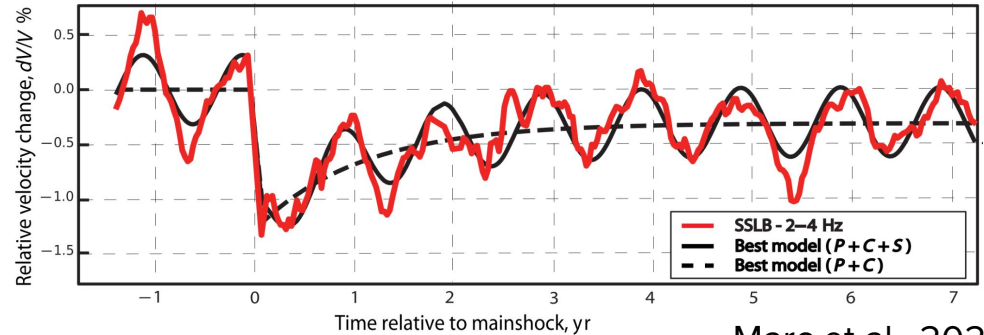
Luc Illien, *Christoph Sens-Schönfelder, Jens Turowski, Kuan-Yu Ke & Niels Hovius*

5-minute talk, EGU meeting 2022
lillien@gfz-potsdam.de

What controls the timescales of recovery/relaxation after ground shaking ?

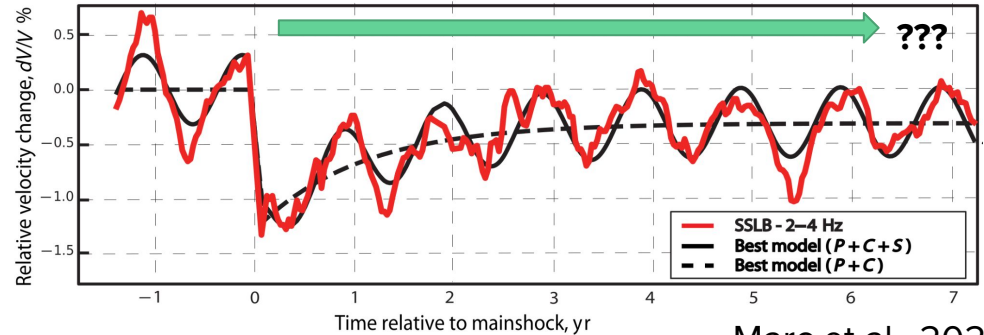
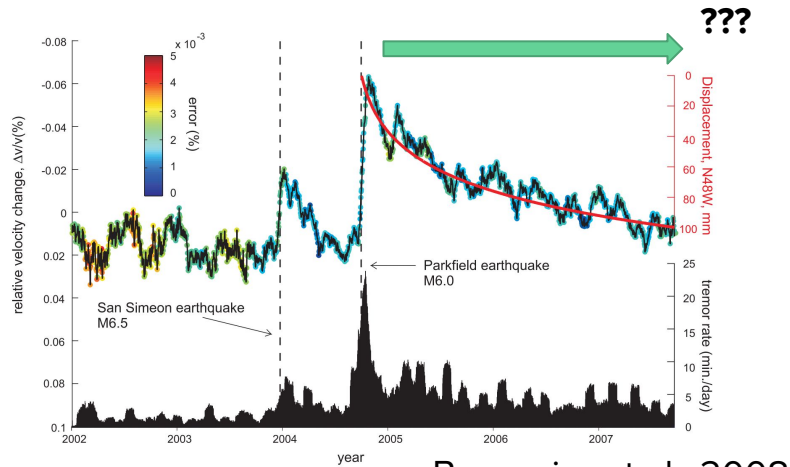


Brenguier et al., 2008

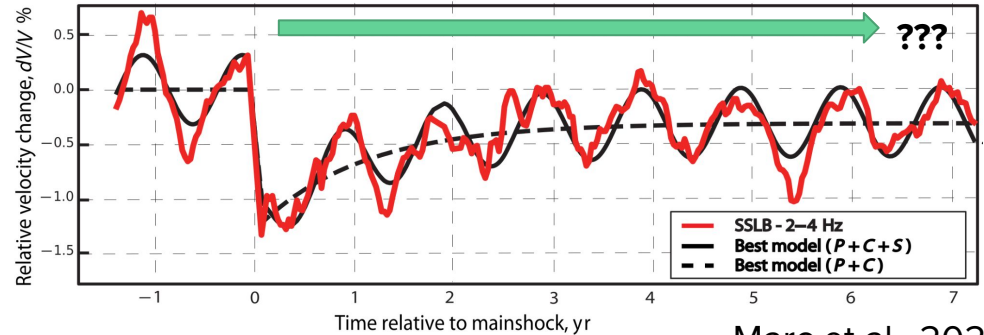
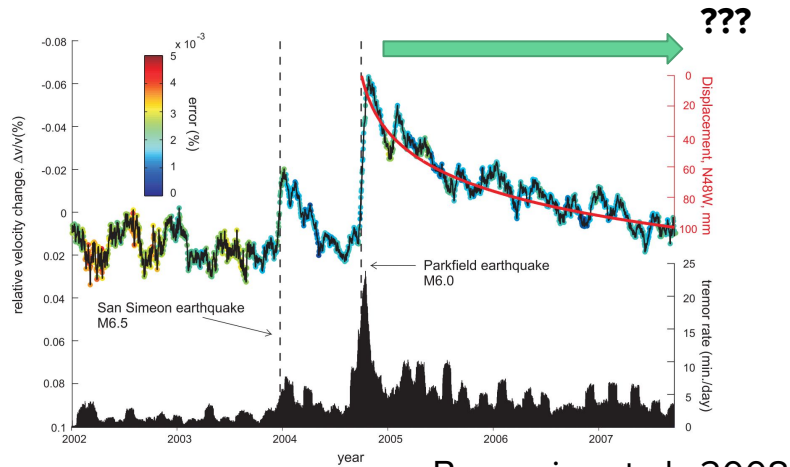


Marc et al., 2021

What controls the timescales of recovery/relaxation after ground shaking ?

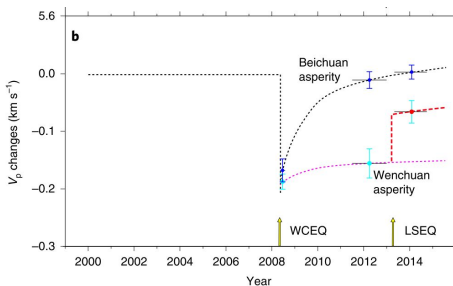


What controls the timescales of recovery/relaxation after ground shaking ?

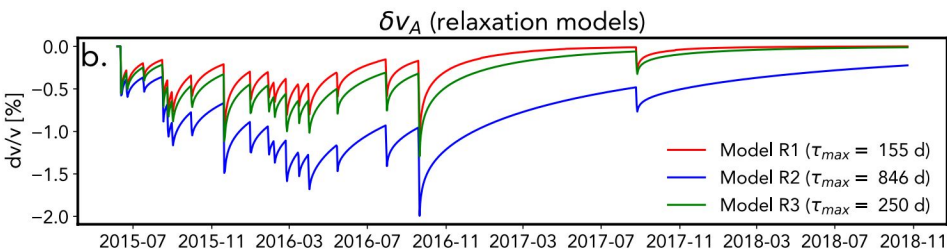


Recovery linked to other transient observations: friction/strength in fault zones, permeability disturbance, landslides rates ...

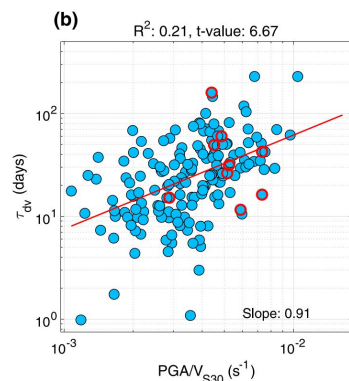
Controlled by ground
shaking intensity ?
(Viens et al, 2018)



Controlled by strain at
depth ?
(Pei et al, 2019)

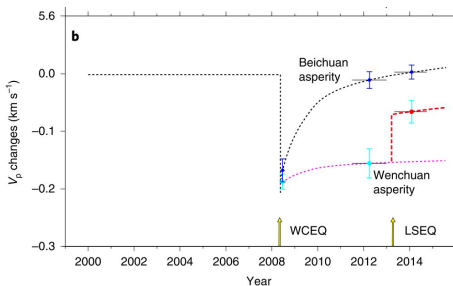


Recovery timescale independent of ground shaking ?
(Illien et al, 2022)

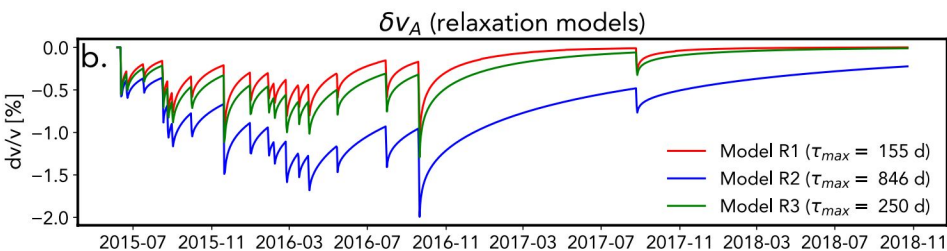
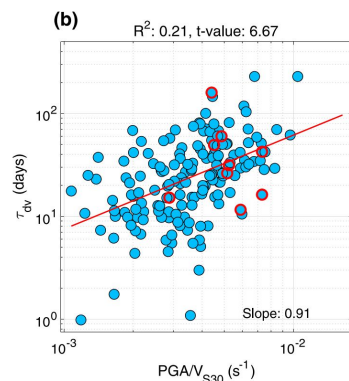


Recovery timescales are
not well constrained.

Controlled by ground
shaking intensity ?
(Viens et al, 2018)



Controlled by strain at
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Recovery timescale independent of ground shaking ?
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Recovery timescales are
not well constrained.

Need for better-resolved dv/v measurements

To cite a few...

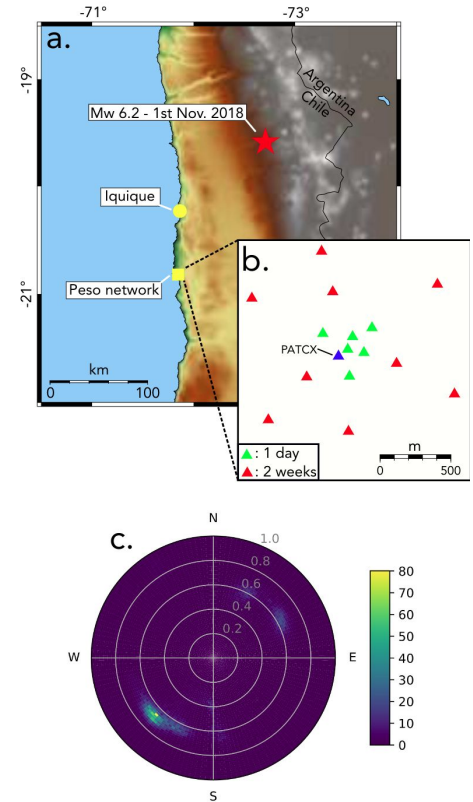
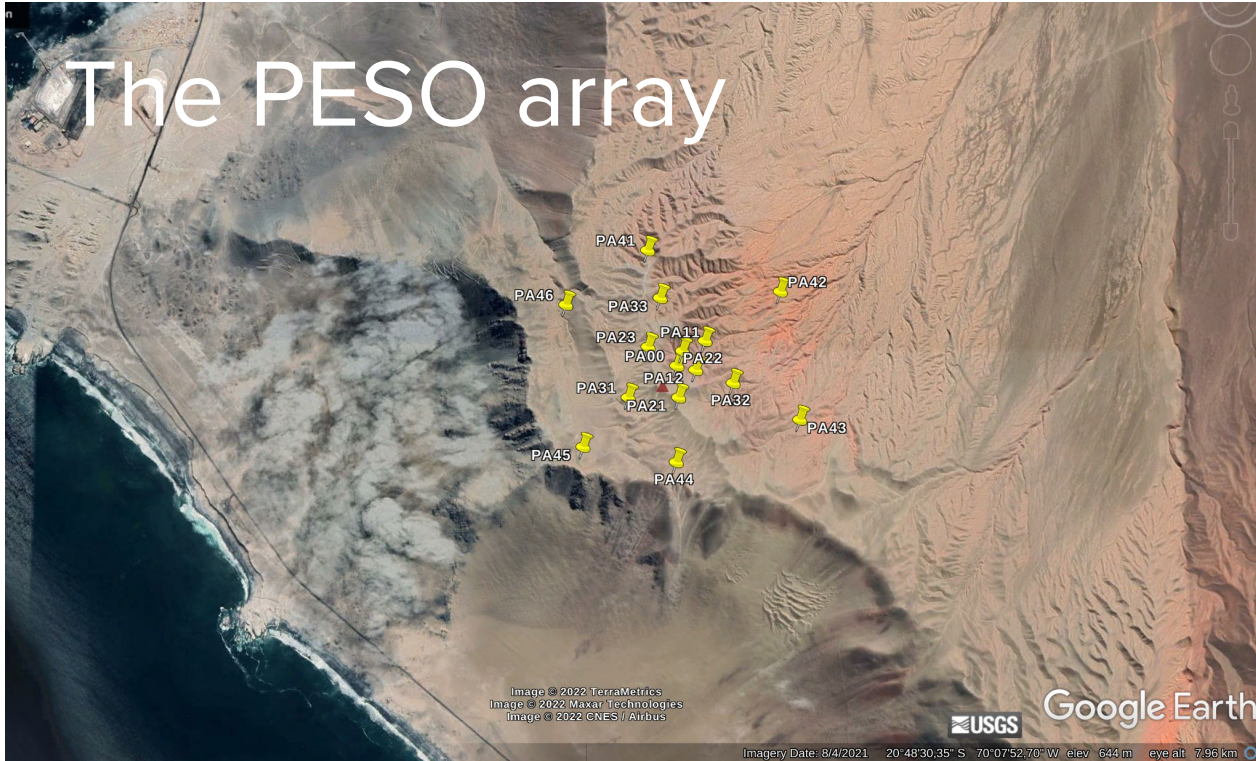
Let's use the ergodic hypothesis.

We may constrain well-resolved minute relaxations.

Principle

If the ambient noise wavefield is random, cross-correlations measurements may have the same ensemble (spatial) average than the time-average ones.

→ stack results of several stations to lower the resolution of the temporal stack and reduce the noise in dv/v measurements.



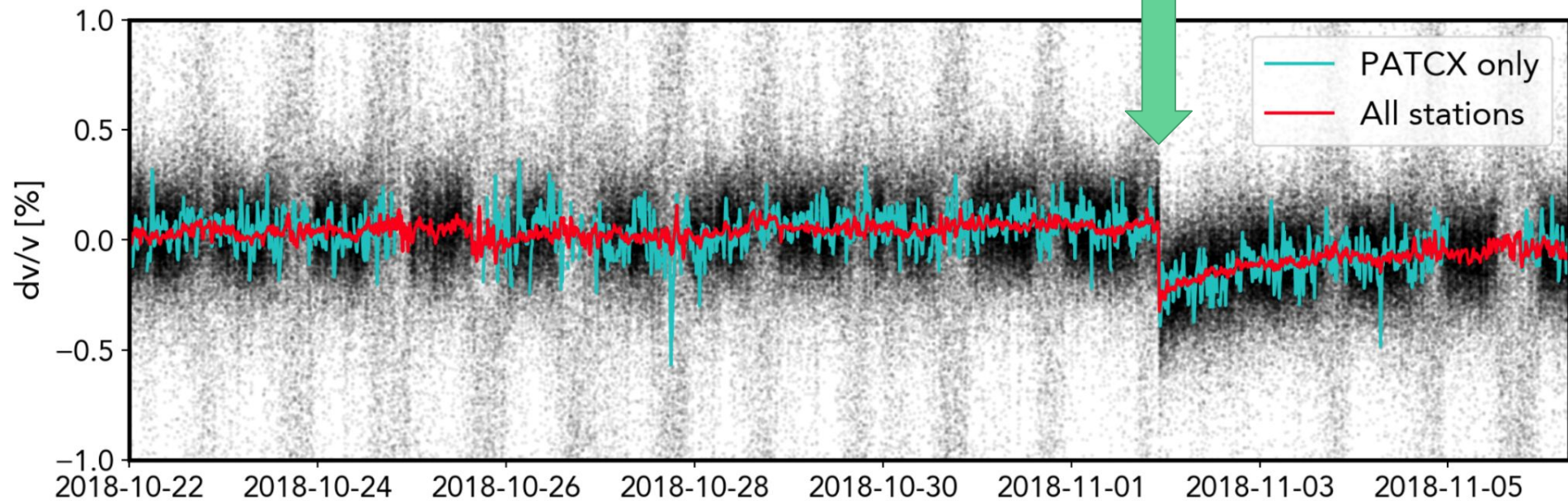
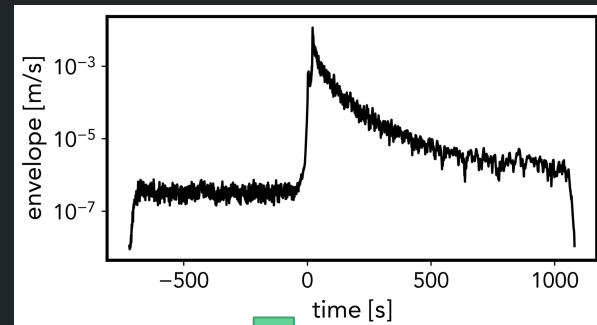
- 10 3-components BB stations set-up for 15 days
 - Noise coming from the shore at 3-6 Hz
 - Recorded one seismic event

Seismic processing

From seismic traces to relative velocity changes (dv/v)

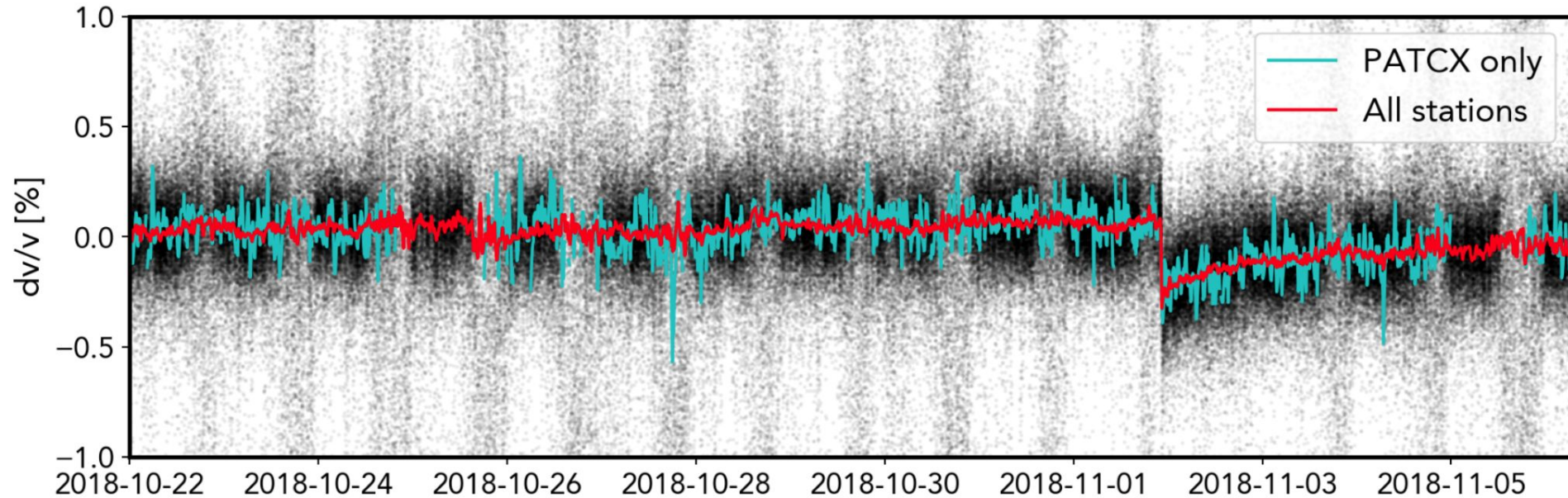
- Pre-processing: muting of segments that do not plot on a theoretical random amplitudes distribution
 - Compute all CC, SC, AC (465 combinations)
 - Estimate dv/v with the stretching technique
 - Scheme to jointly average the stretching estimates
-

Seismic velocity changes



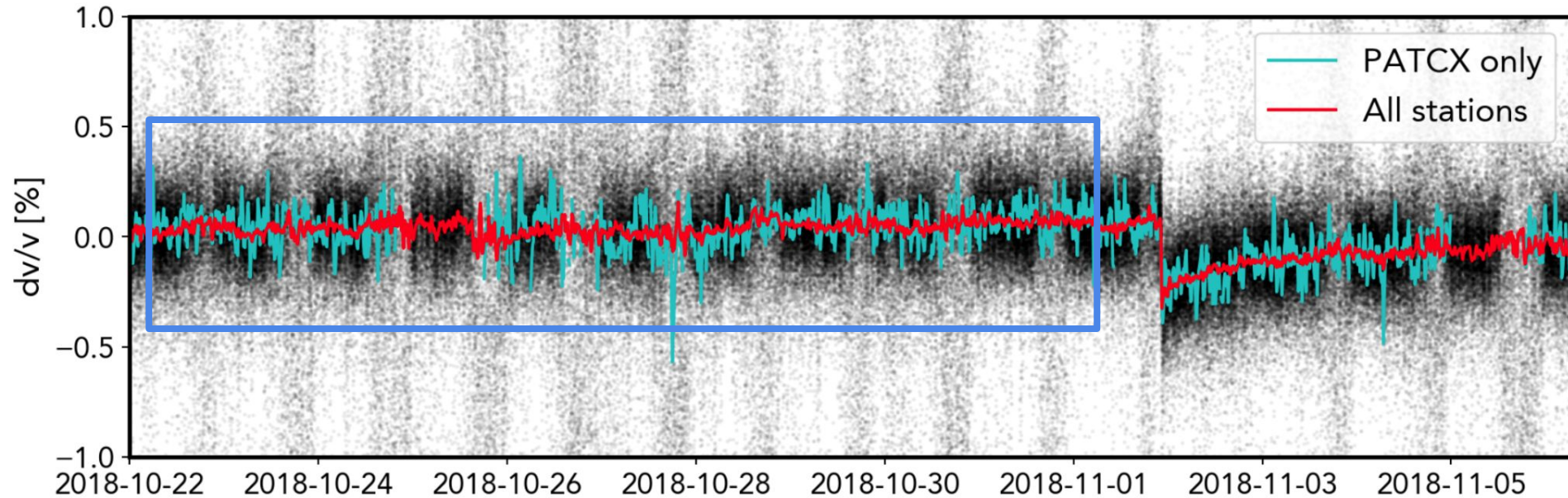
Seismic velocity changes

- well resolved 0.4 % drop (3-6 Hz)
- recover to 50% of the initial value in 2 days
- 10-minutes resolution



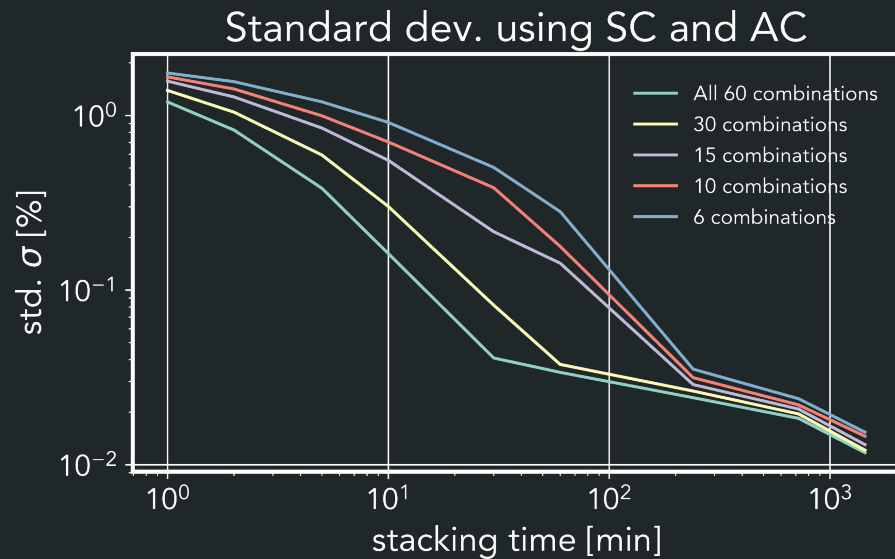
Seismic velocity changes

- What can we hope to achieve using more stations ?
- Let's use the standard deviation of the week before the earthquake



Using more seismic stations ?

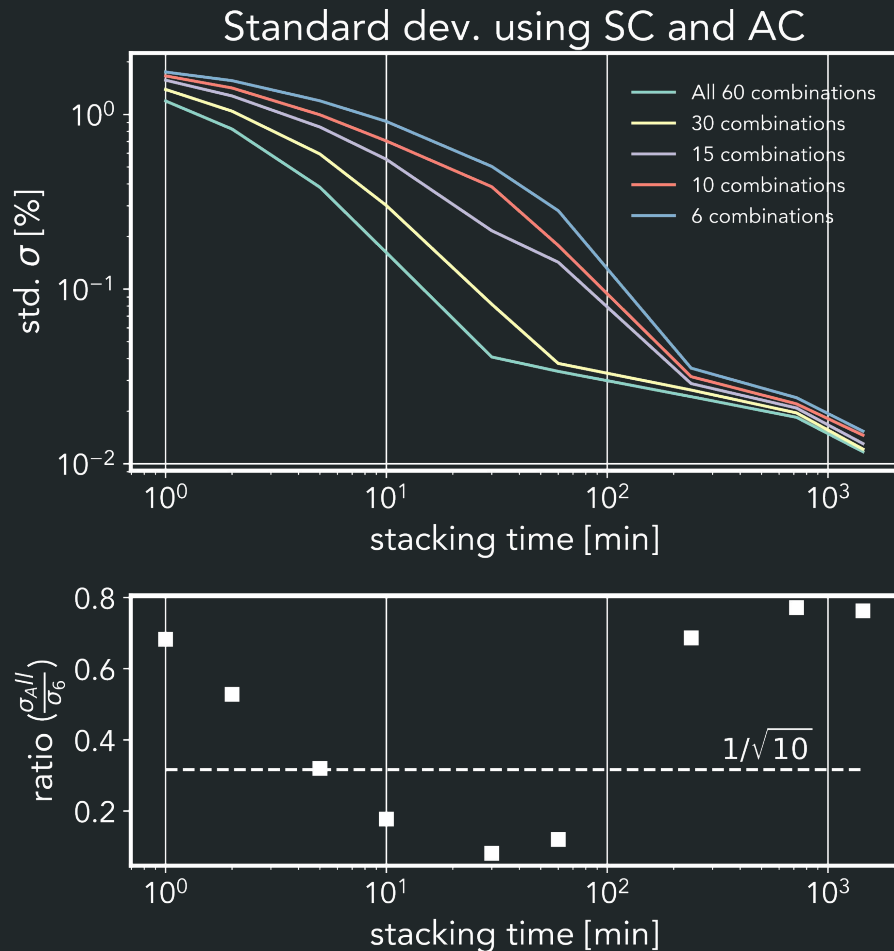
Random shuffle of the 60 single
station combinations (AC + SC)



Using more seismic stations ?

Random shuffle of the 60 single
station combinations (AC + SC)

*If the measurements at the 10 stations can be
considered independent with same variance, we
may expect a reduction of the deviation \propto
 $\text{sqrt}(10)$*



Use more stations.

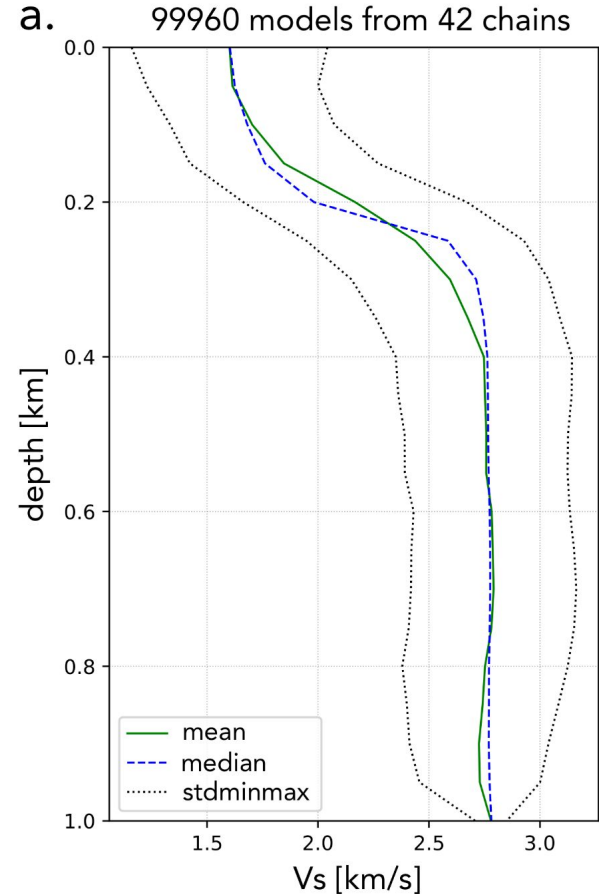


Bonus

Inversion of a local 1D shear-velocity profile

- Performed with the focal spot imaging technique and a MCMC transdimensional Bayesian inversion scheme
- Suggests that relaxation occurred in a hard-rock site

Contact me if you want to know more about this part



Conclusions

- Averaging station measurements is worth it in a specific stacking-time duration window (although with some deviation of the ergodic hypothesis)
- We observe a 0.4 % drop after a PGV of 1 cm/s with half of the drop recovered in 2 days
- Our observation of relaxation happened in a hard-rock site

More in 'Resolving minute relaxation-induced seismic velocity changes: The more station, the merrier ?' Illien et al., in prep

Ongoing work

We will compare our results with a longer-term time-series to derive an effective field-based law (or attempt to...)

This talk

