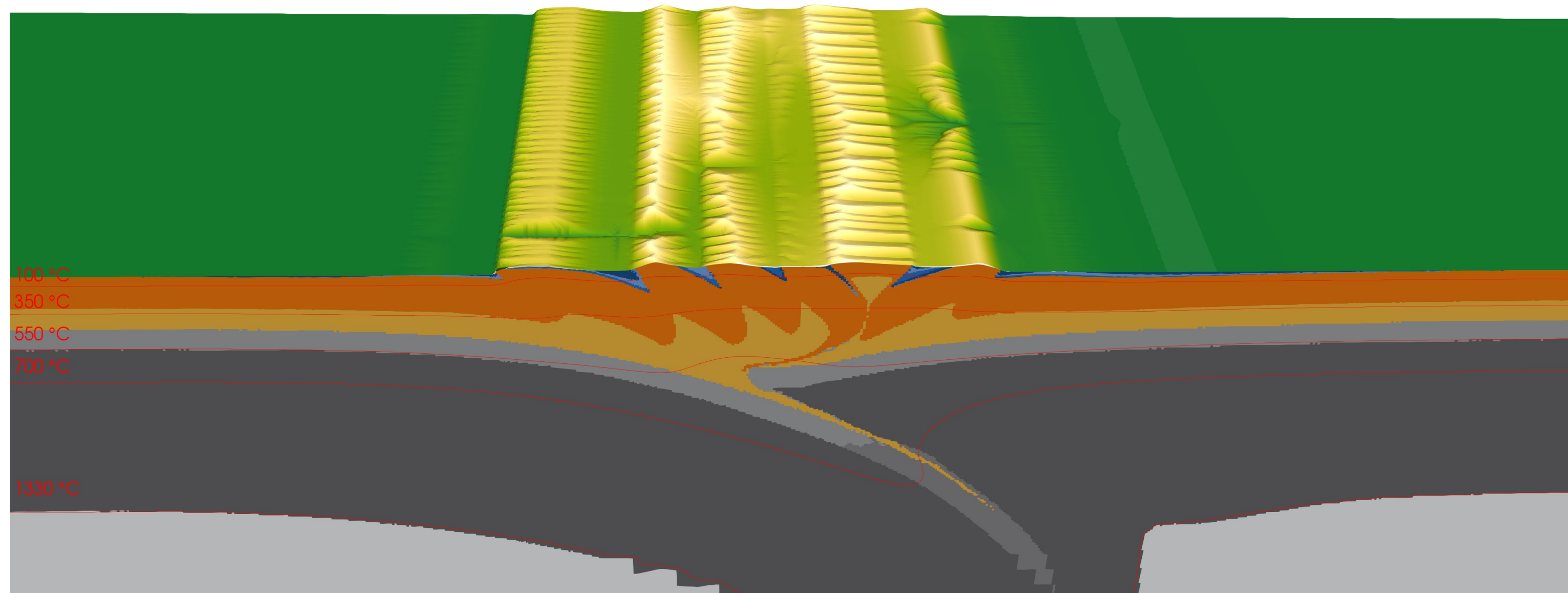


# Quantifying the growth (and decay) of topography in collisional orogens

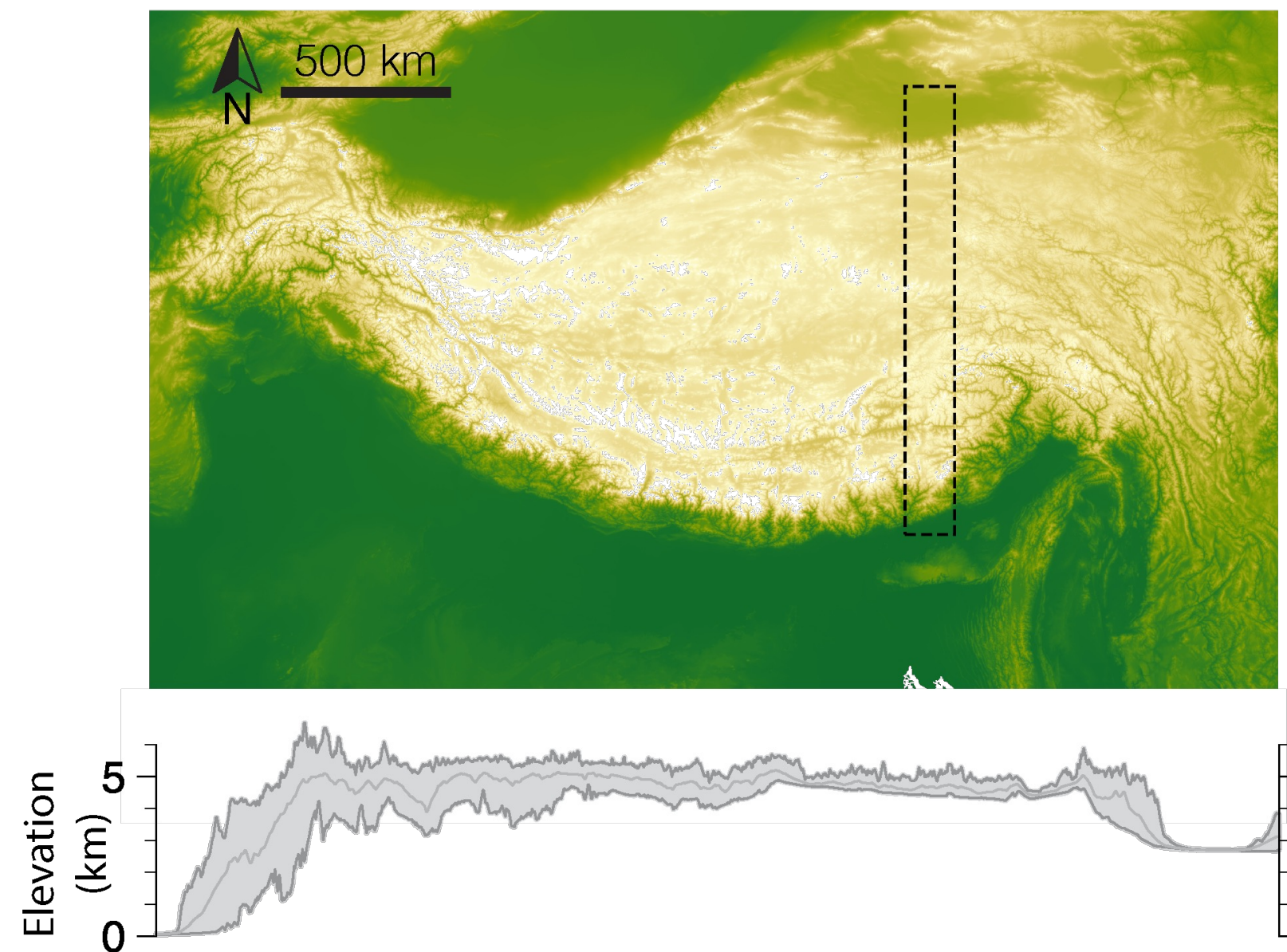
Sebastian G. Wolf, Ritske S. Huismans, Jean Braun, Xiaoping Yuan





# Tectonics or climate - what controls mountain belt height and width?

Himalaya-Tibet

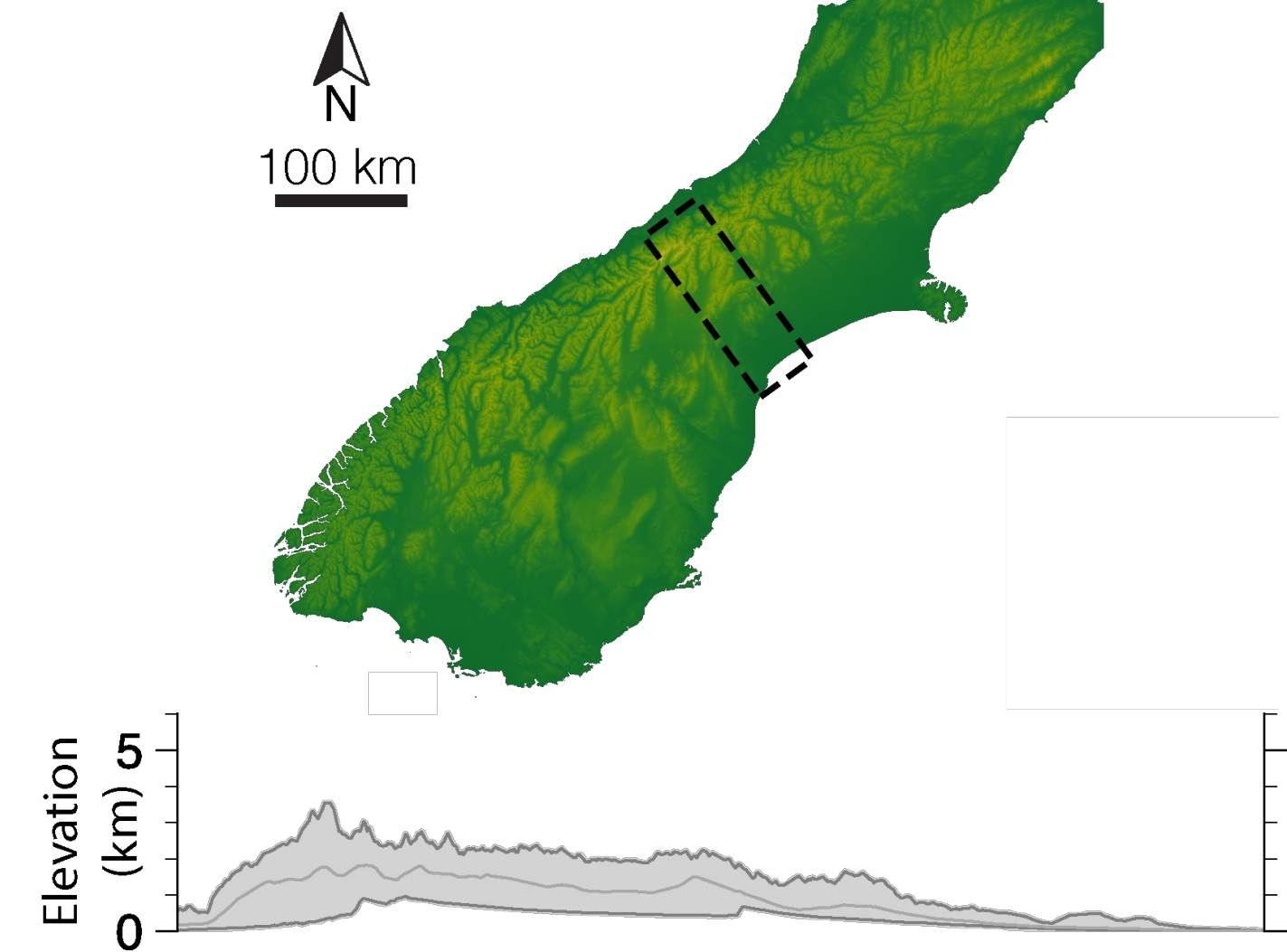


**Tectonics:**  
**Crust is too weak**

$$\text{Argand Number} = \frac{\text{Buoyancy force from crustal thickening}}{\text{Lithospheric strength}}$$

England & McKenzie, 1982

Southern Alps  
of New Zealand



**Geomorphology:**  
**Erosion is too fast**

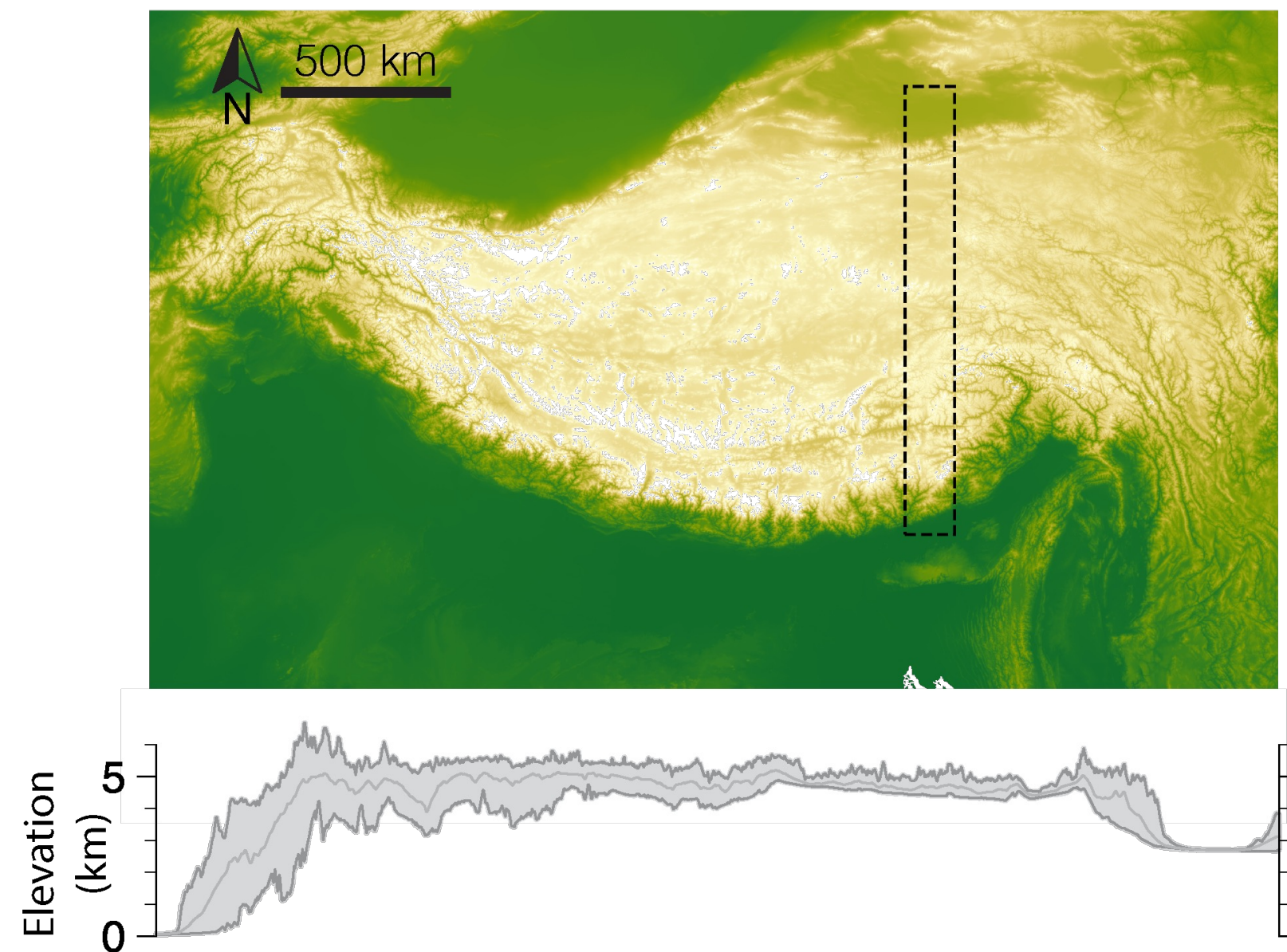
$$\text{Uplift/erosion Number} = \frac{\text{Time scale to reach max height}}{\text{Time scale to erode orogen}}$$

Whipple & Tucker, 1999; Willett 1999



# Tectonics or climate - what controls mountain belt height and width?

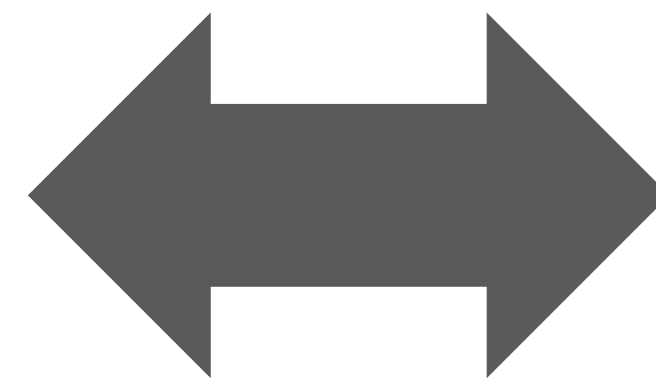
Himalaya-Tibet



**Tectonics:**  
Crust is too weak

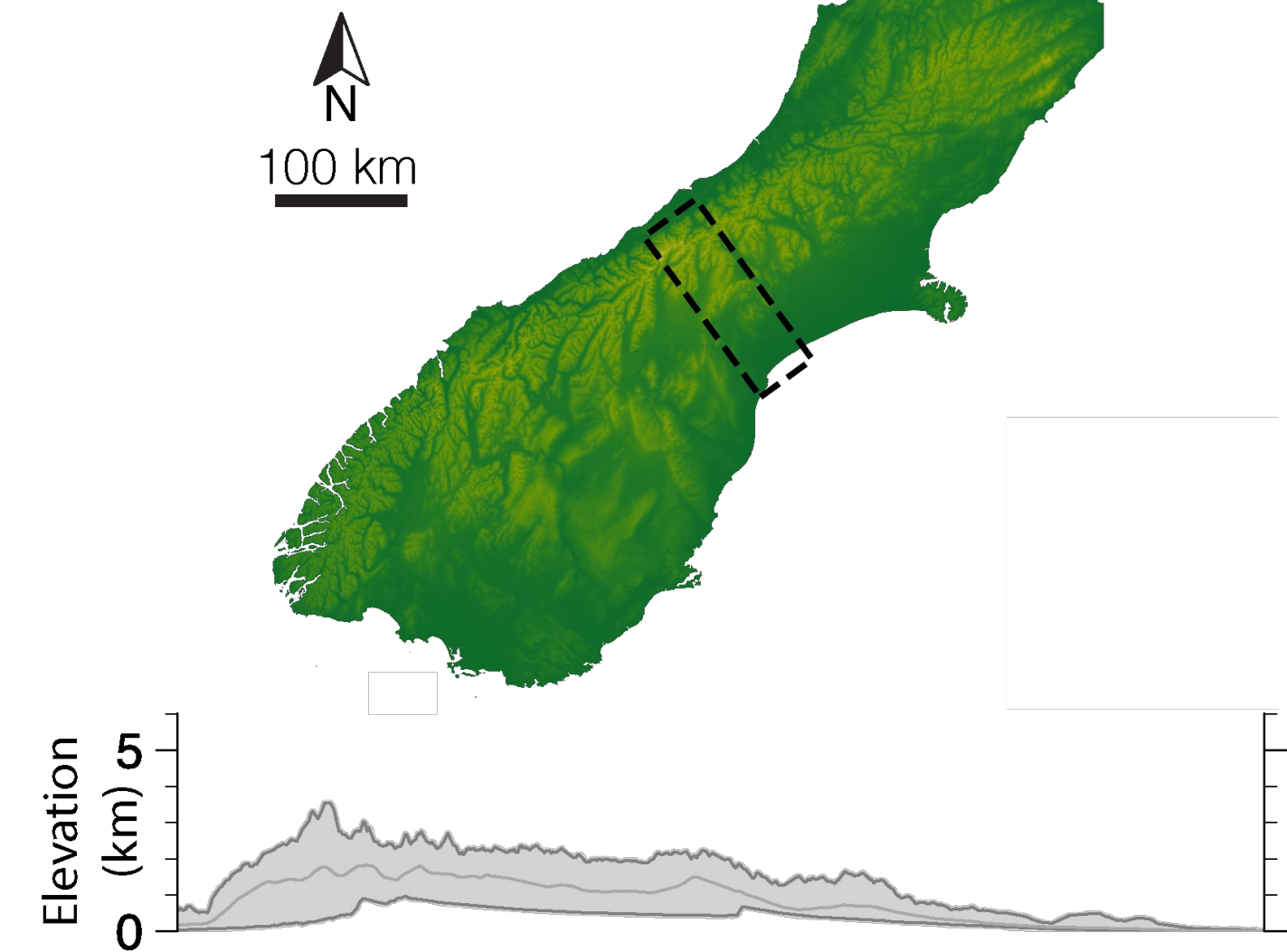
$$\text{Argand Number} = \frac{\text{Buoyancy force from crustal thickening}}{\text{Lithospheric strength}}$$

England & McKenzie, 1982



**LINK**

Southern Alps  
of New Zealand

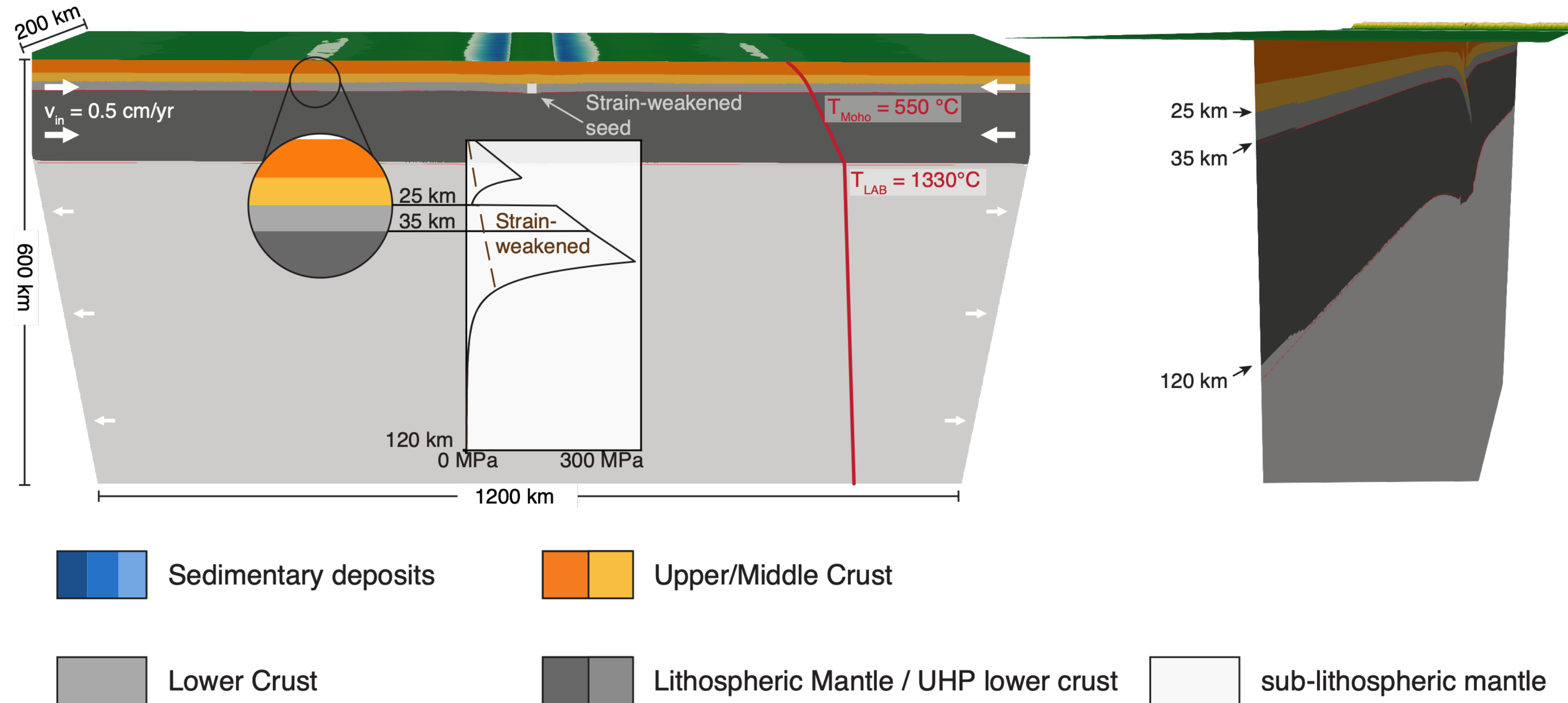


**Geomorphology:**  
Erosion is too fast

$$\text{Uplift/erosion Number} = \frac{\text{Time scale to reach max height}}{\text{Time scale to erode orogen}}$$

Whipple & Tucker, 1999; Willett 1999

# Thermo-mechanical-landscape-evolution model

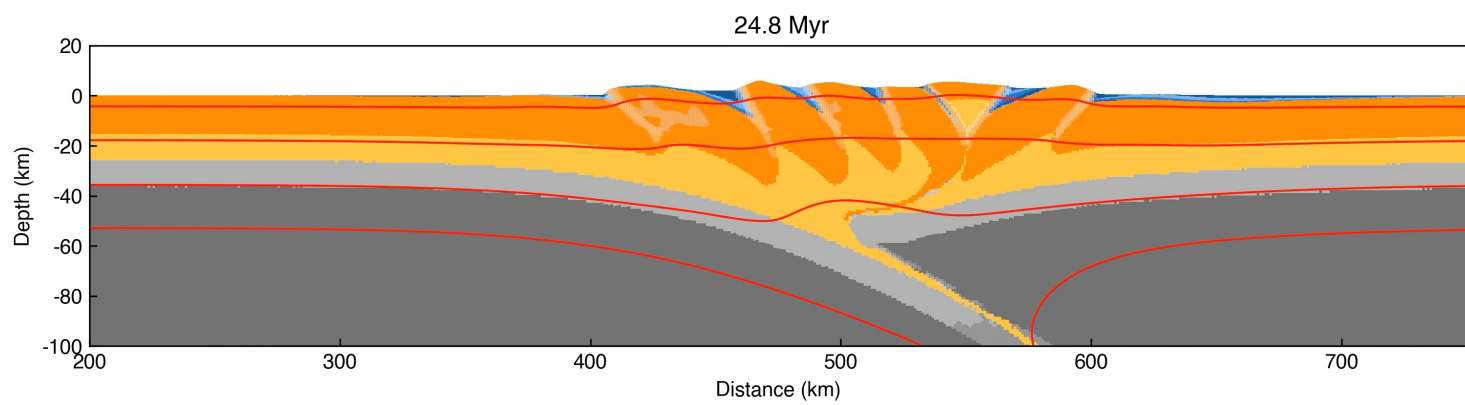
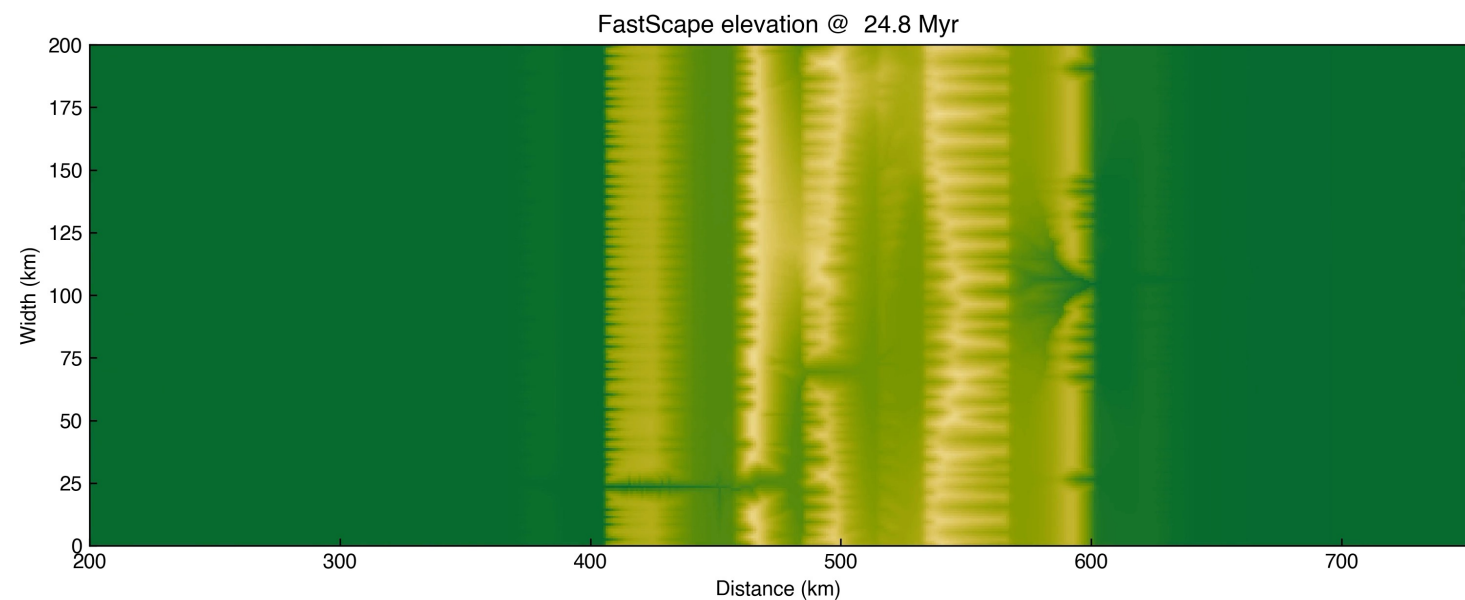
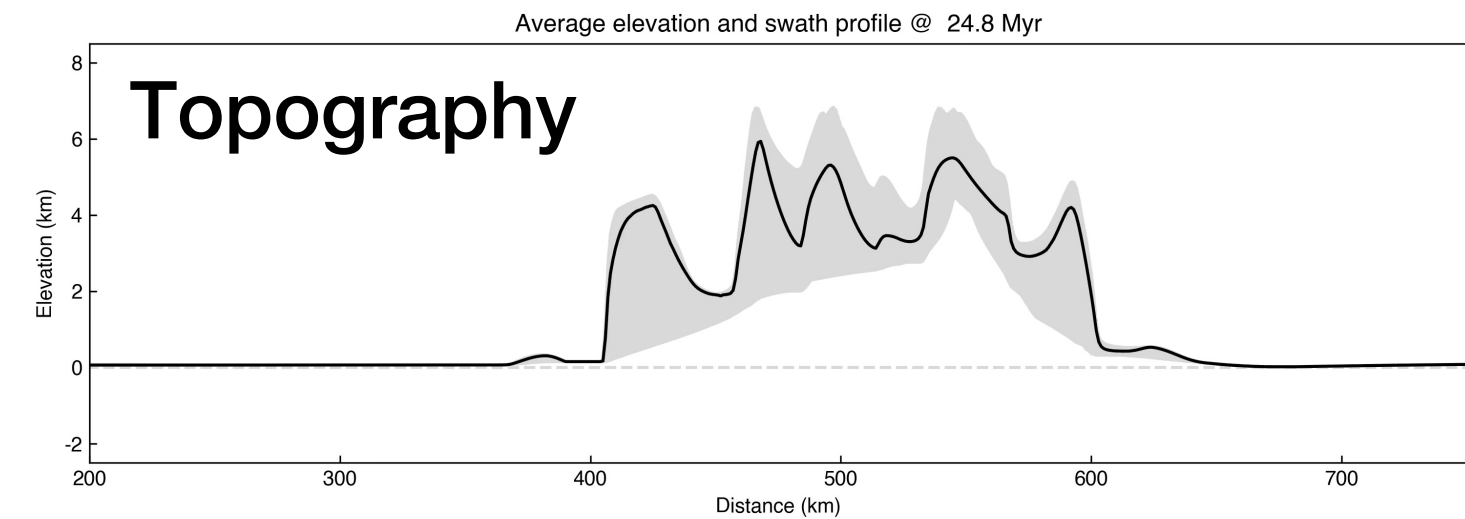
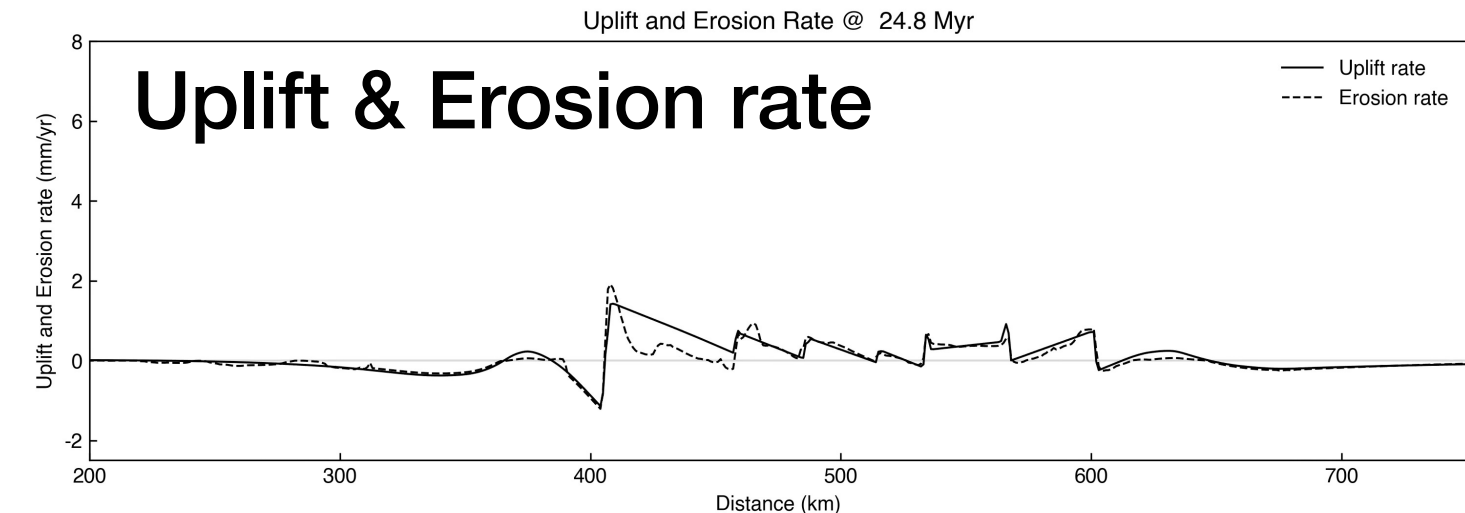


Physics based forward models:

- 2D thermo-mechanically coupled tectonic model (FANTOM, Thieulot 2011, Wolf et al. 2021)
- 2D surface process model (FastScape, Braun and Willett 2013, Yuan et al. 2019)
- Both models are tightly coupled in T-coupling manner
- Simple setup with crustal thickening and one-sided subduction of lithospheric mantle

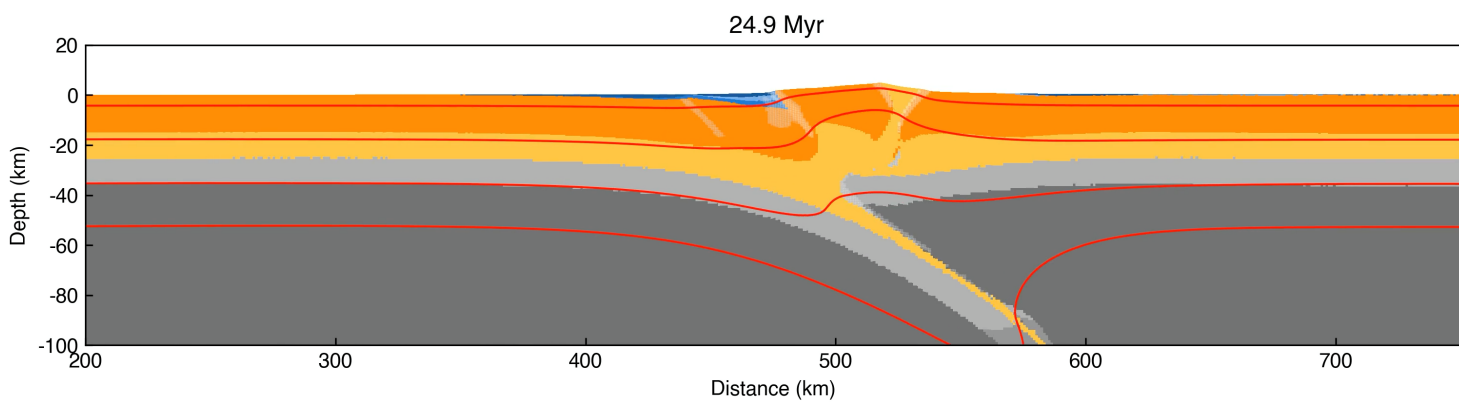
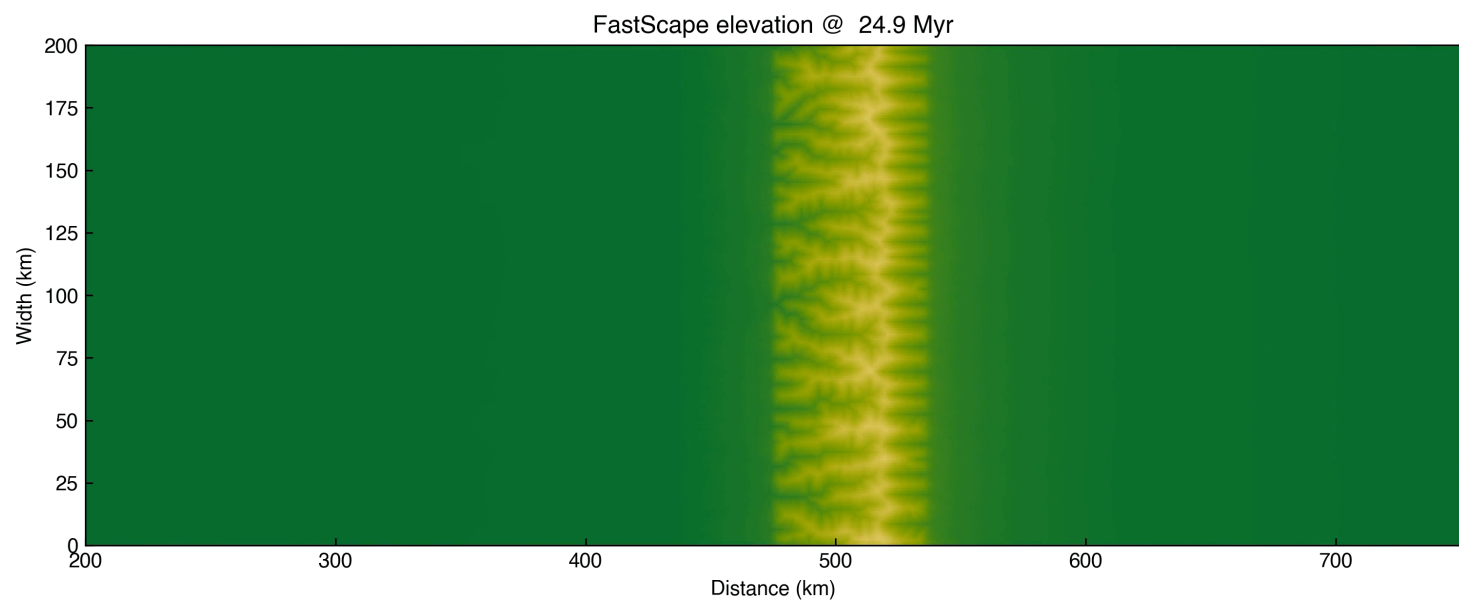
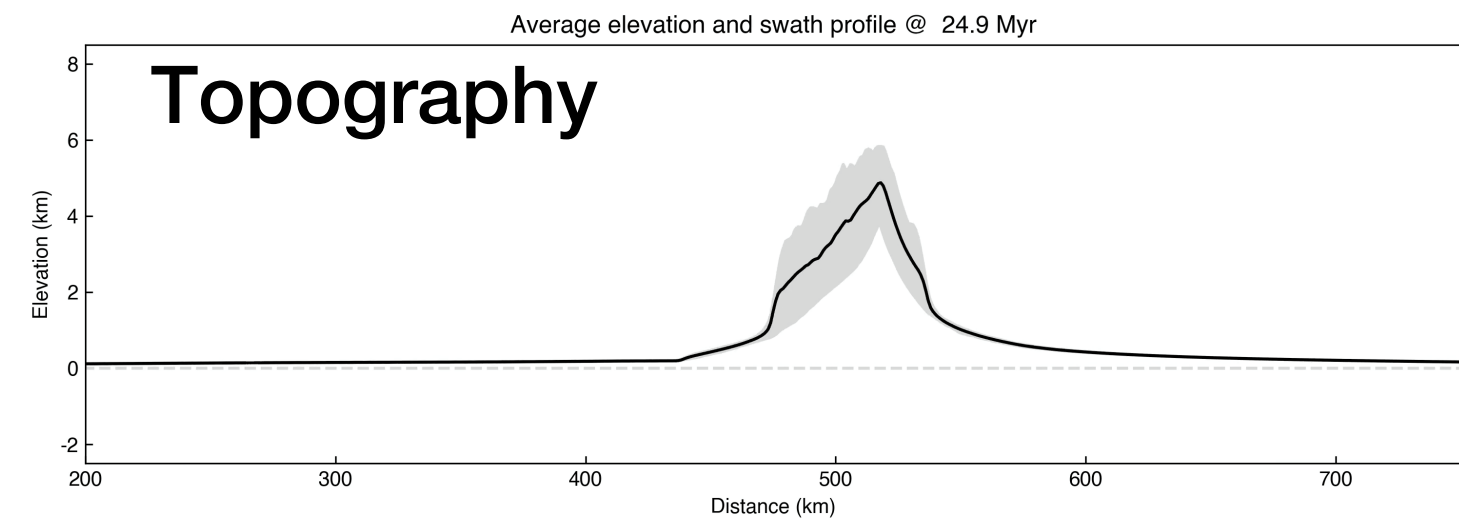
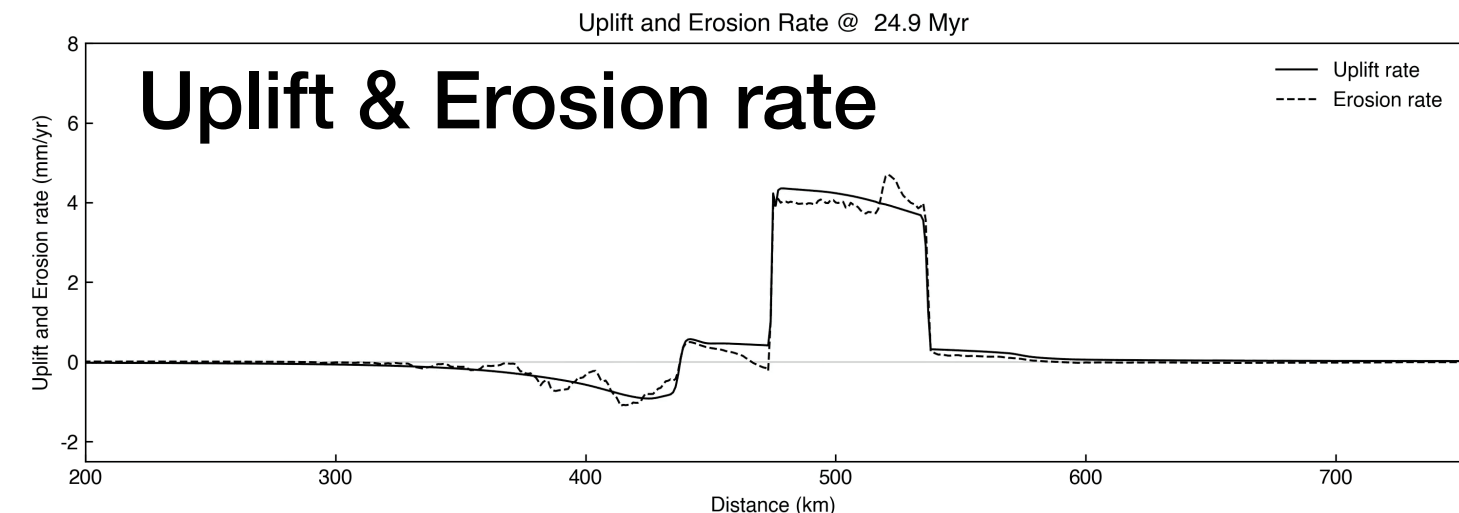


# Type 1: Low surface process efficiency



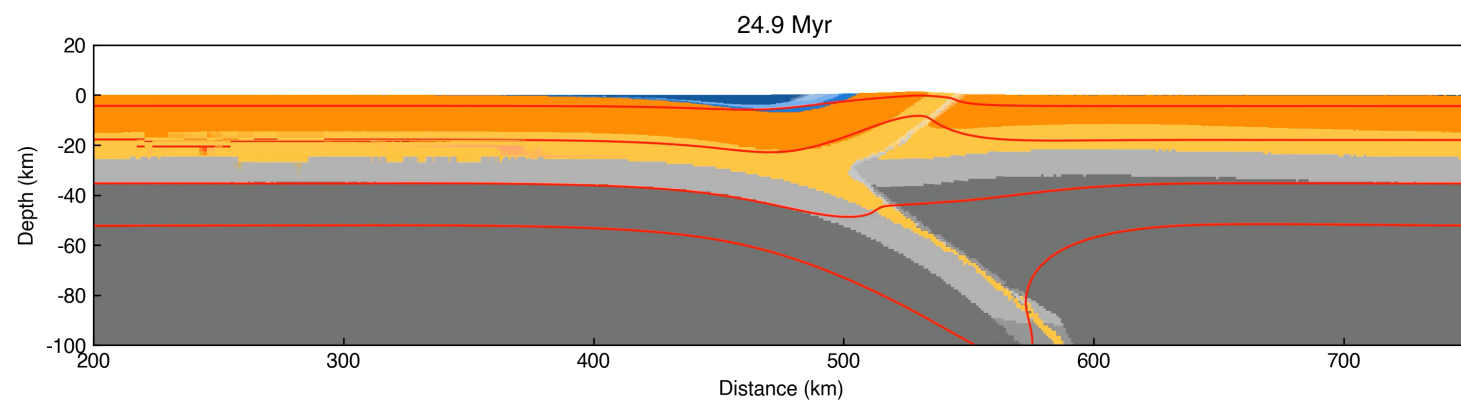
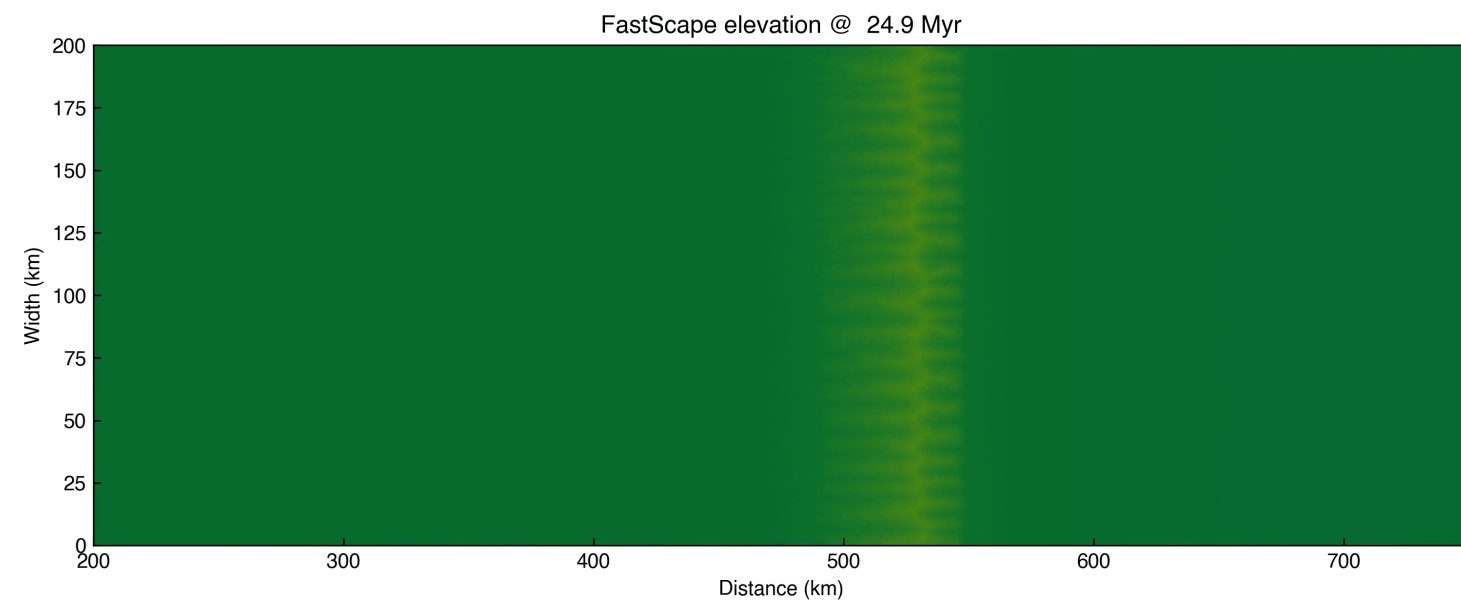
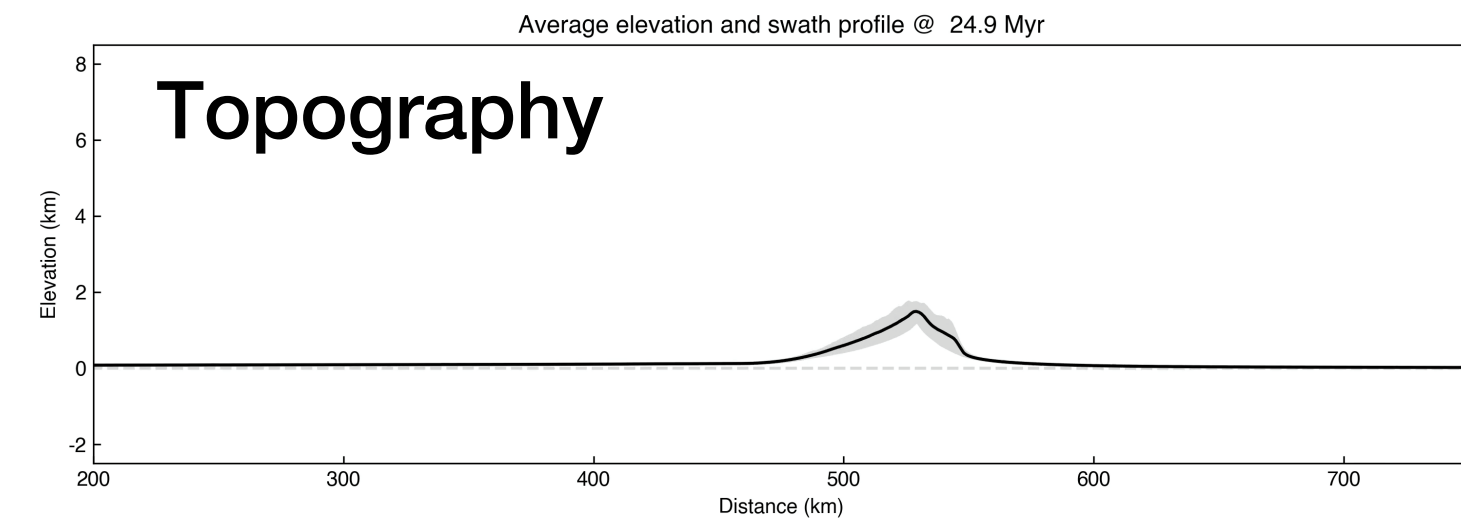
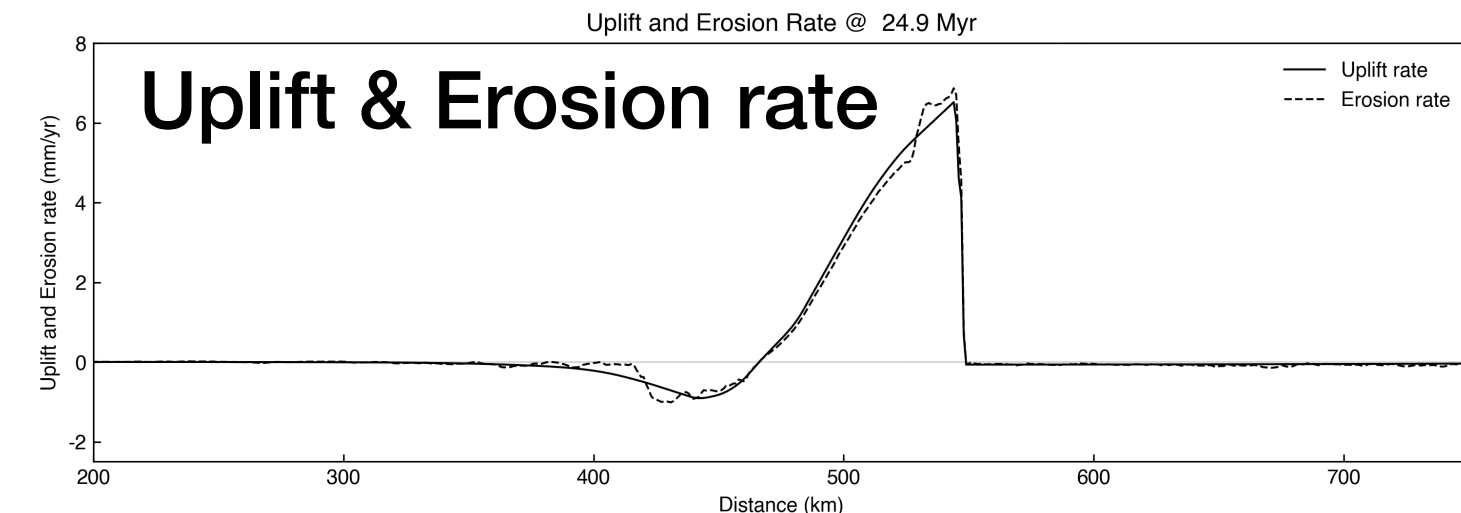
- Growth in height followed by constant widening.

# Type 2: High surface process efficiency



- Growth in height, then width, followed by steady state with bivergent thrusting.

# Type 3: Very High efficiency



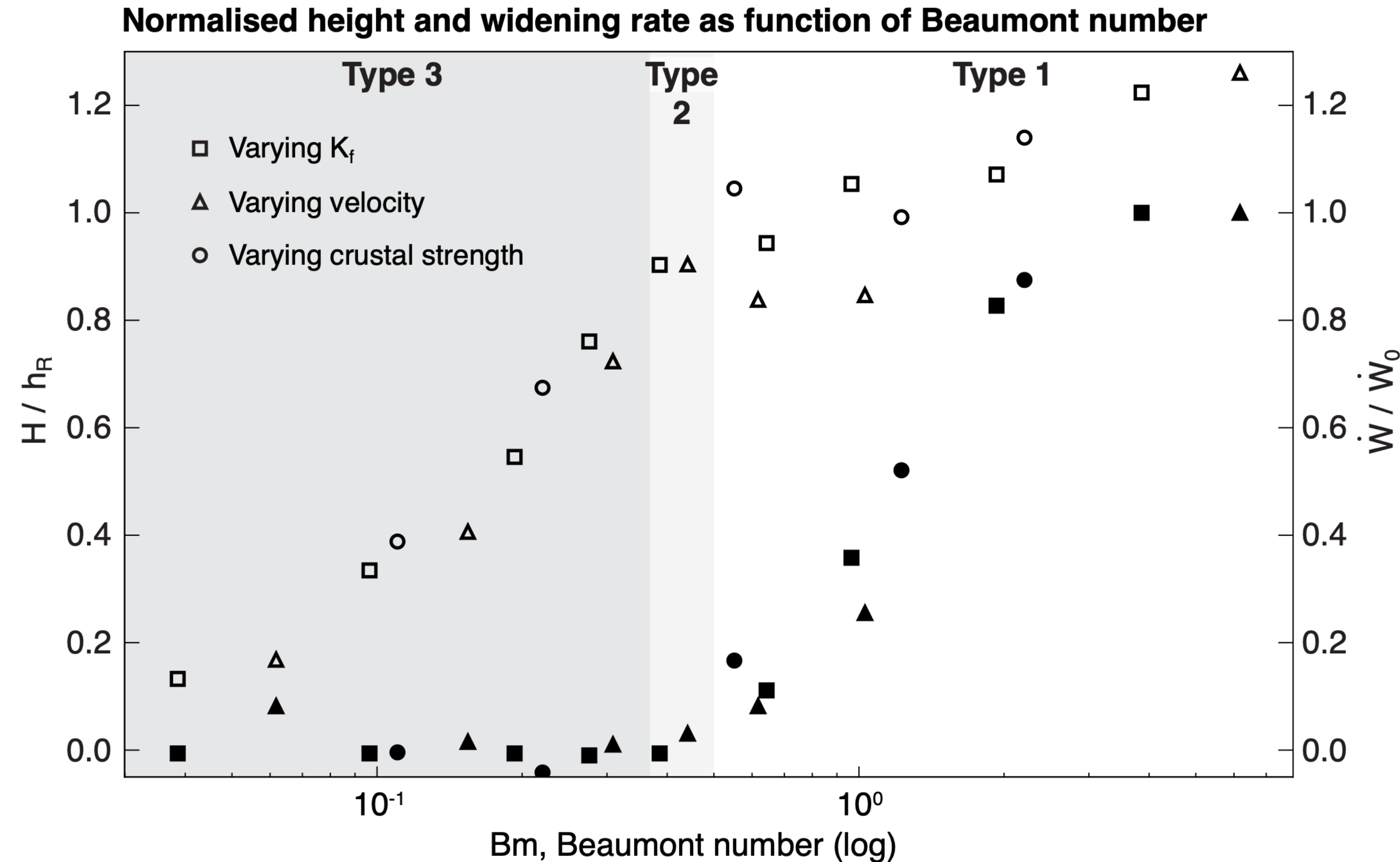
- Growth in height limited by erosion; thrusting on single retro-shear zone.

We define a new non-dimensional number, the Beaumont number that relates tectonics and surface processes.

$$Bm = \frac{Ar}{N_e} \propto \frac{v_c}{F_{int} K_f}$$

$$Ar = \frac{\rho_c \rho' g h_{dec}^2}{F_{int}}$$

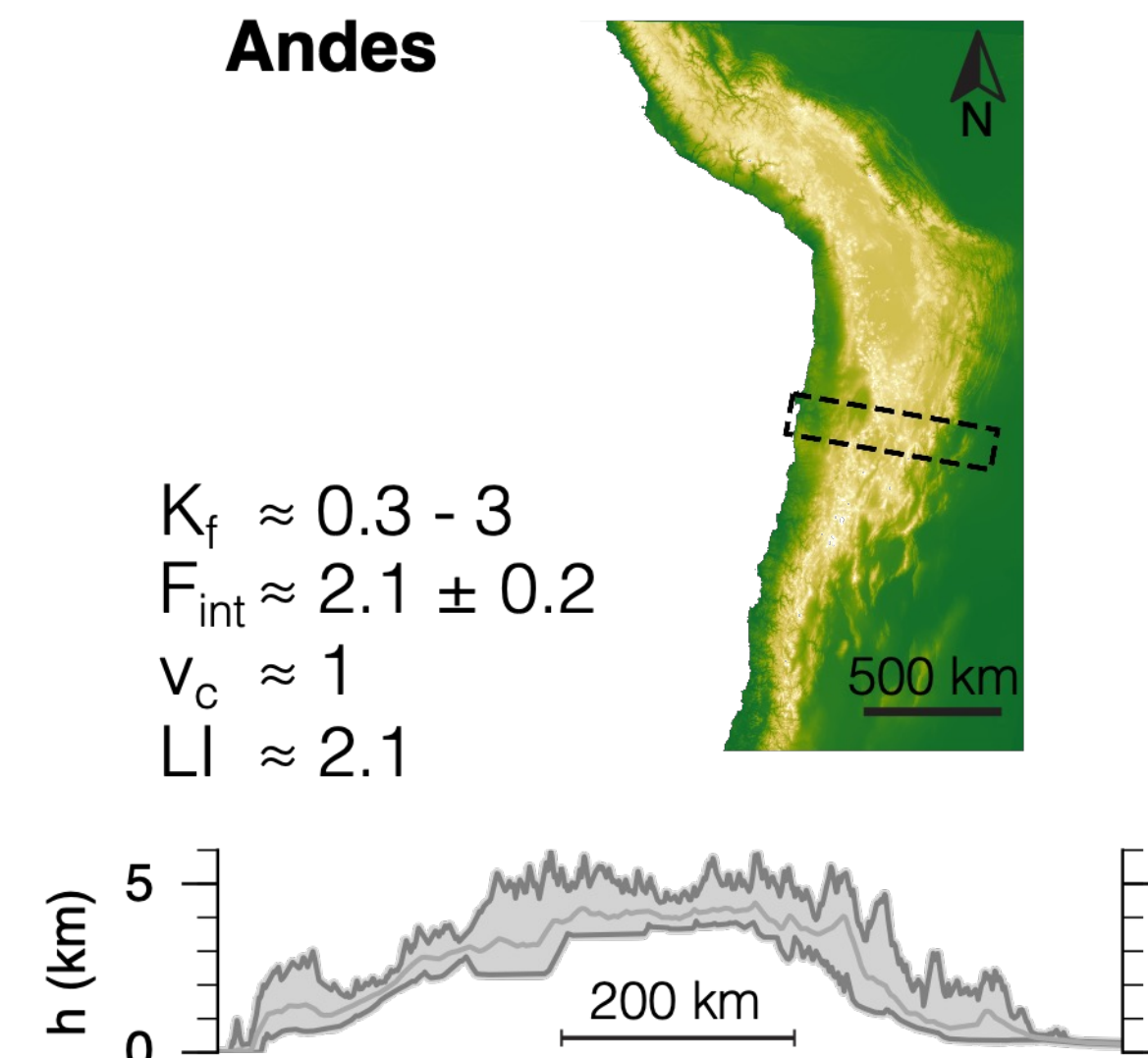
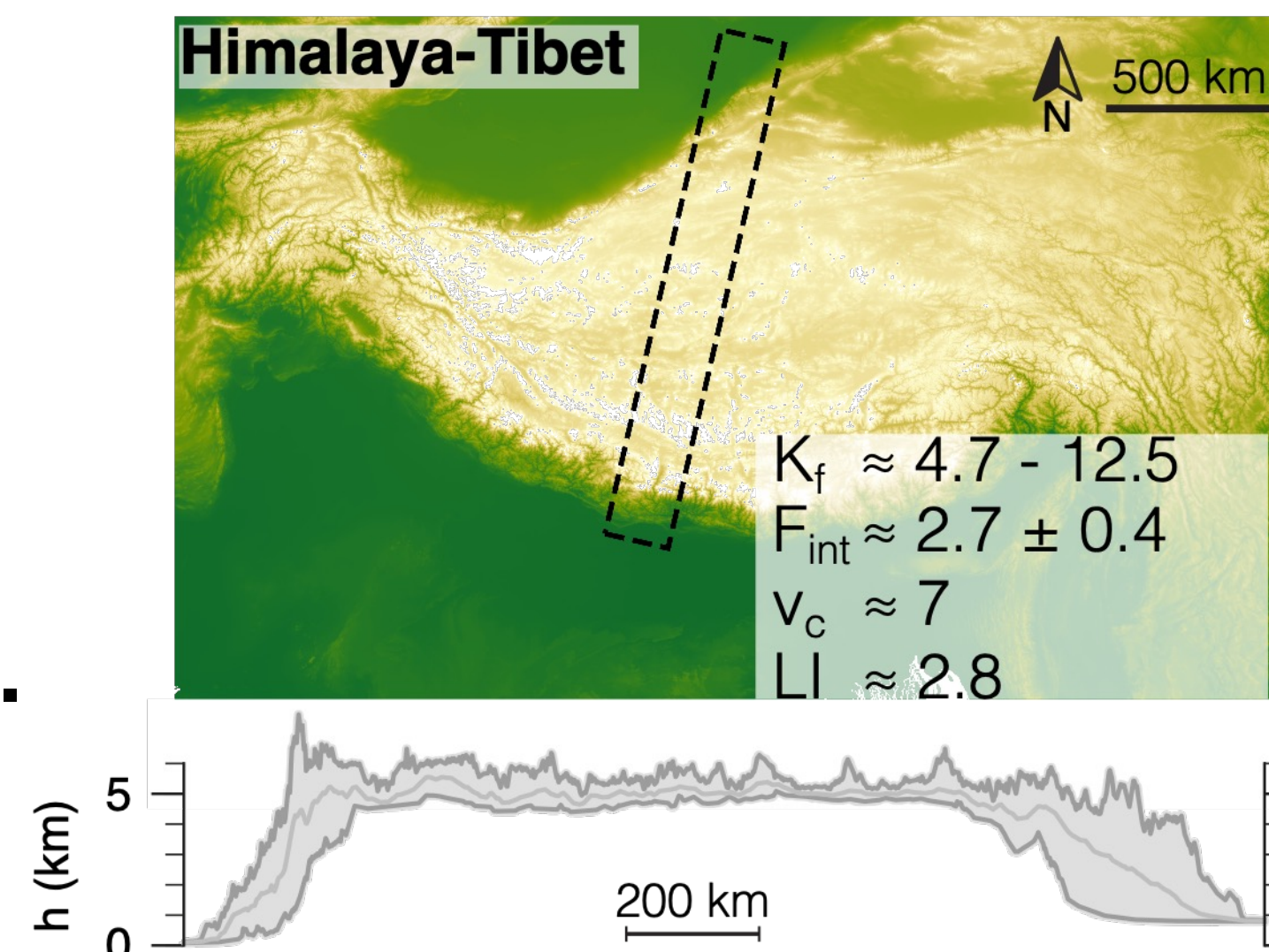
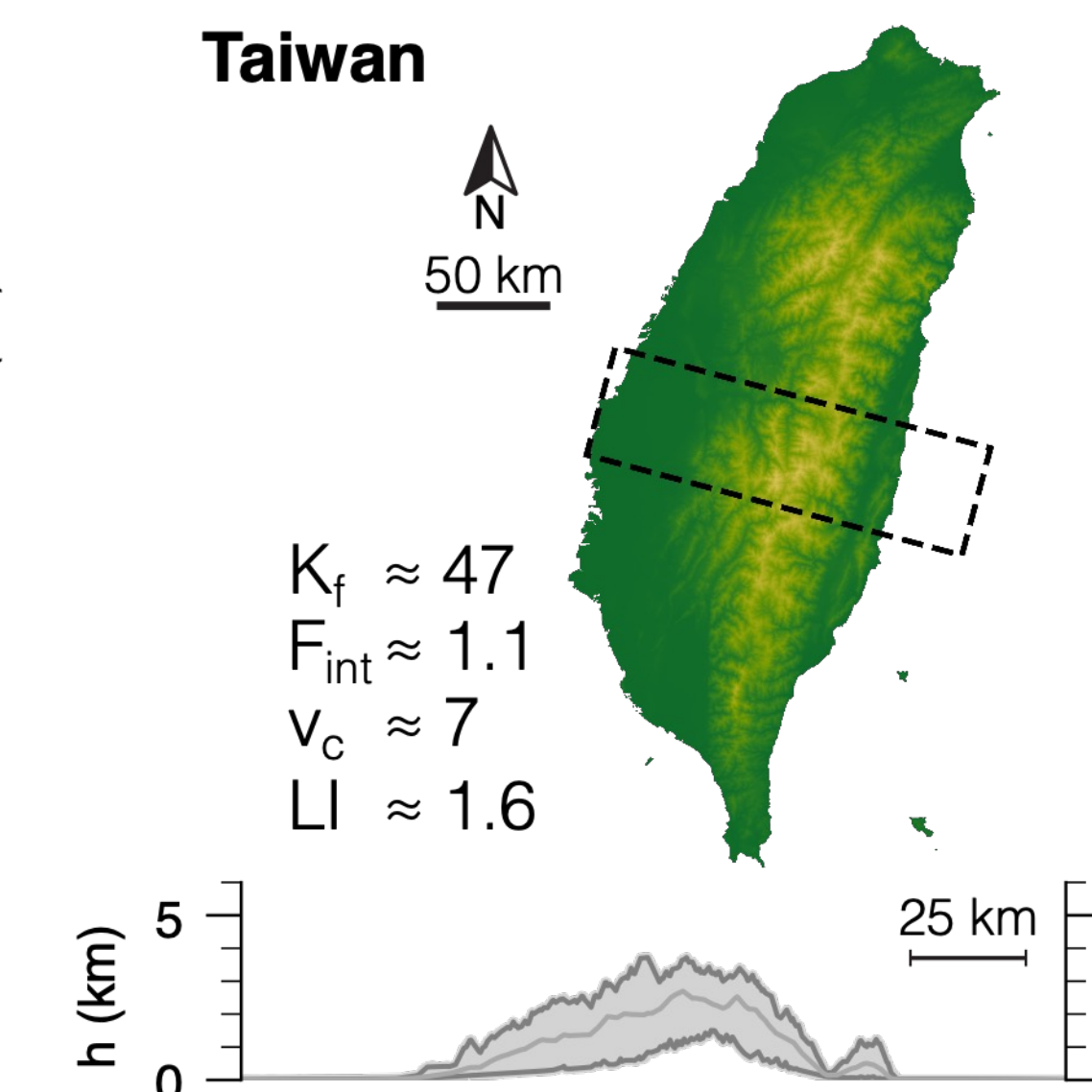
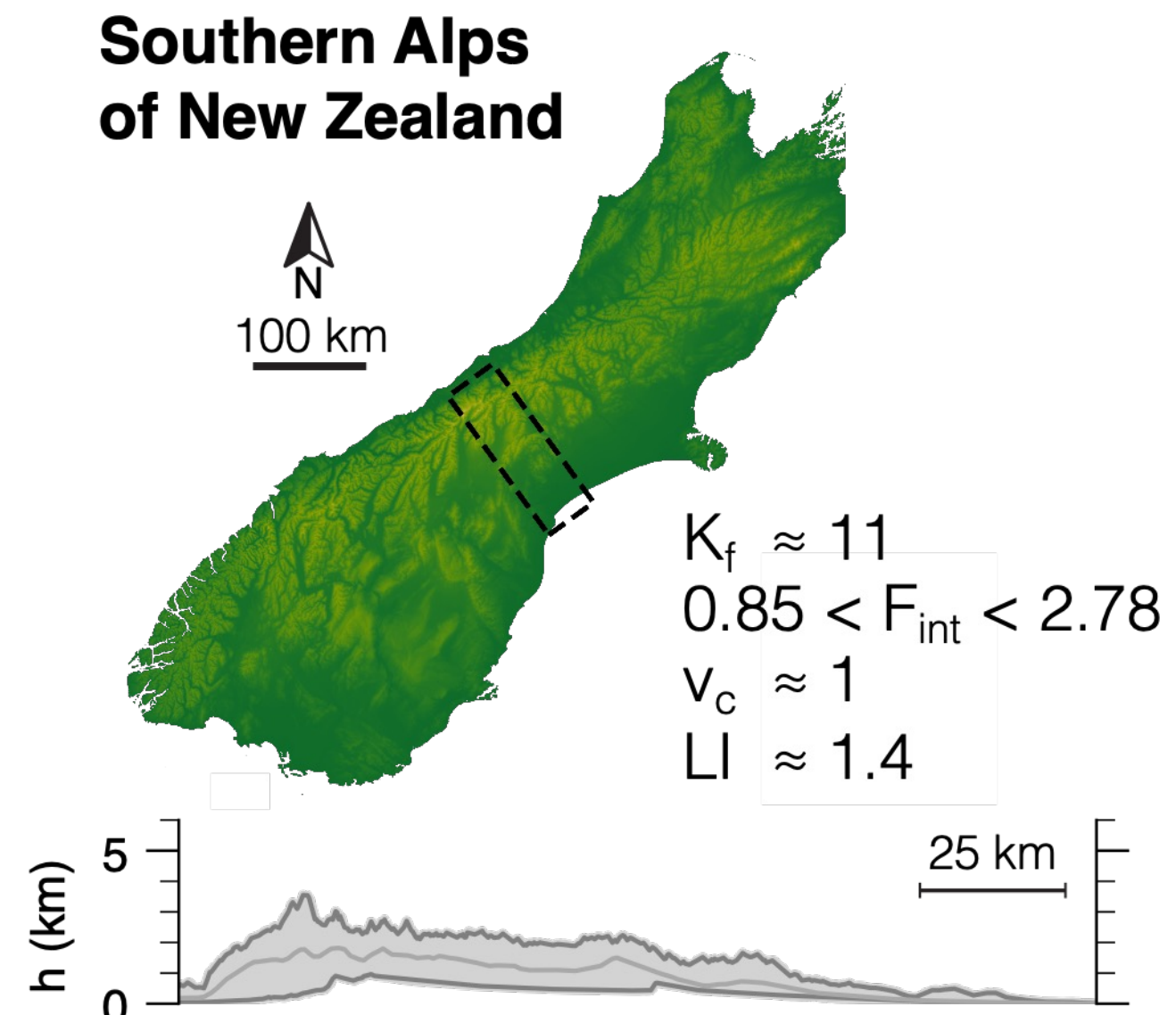
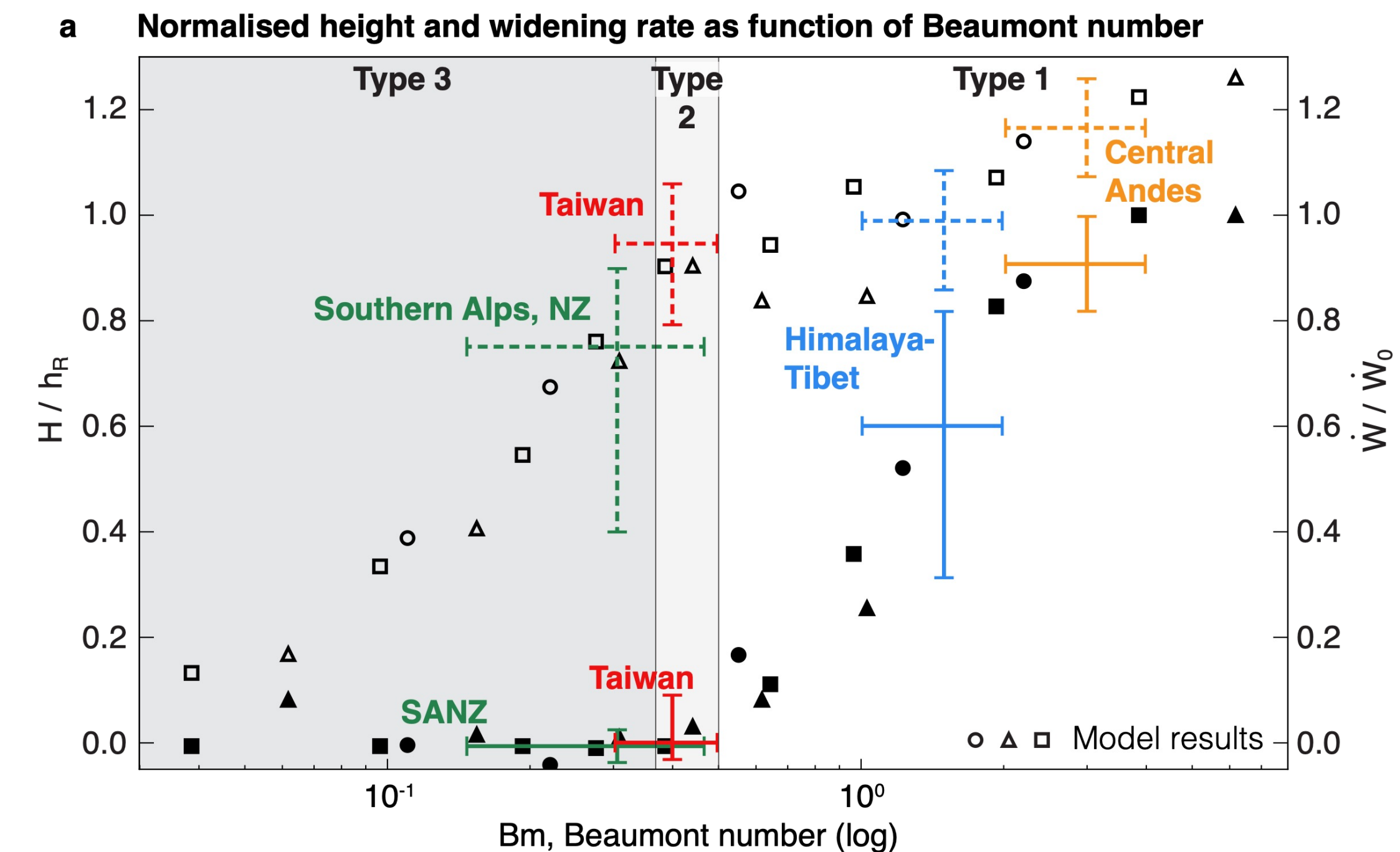
$$N_E = \frac{K_f L^{-0.2} H}{U}$$



- Bm uniquely characterizes mountain belt type and relative importance of tectonics and surface processes
- Bm allows computing crustal strength ( $F_{int}$ ) and fluvial erodibility ( $K_f$ )



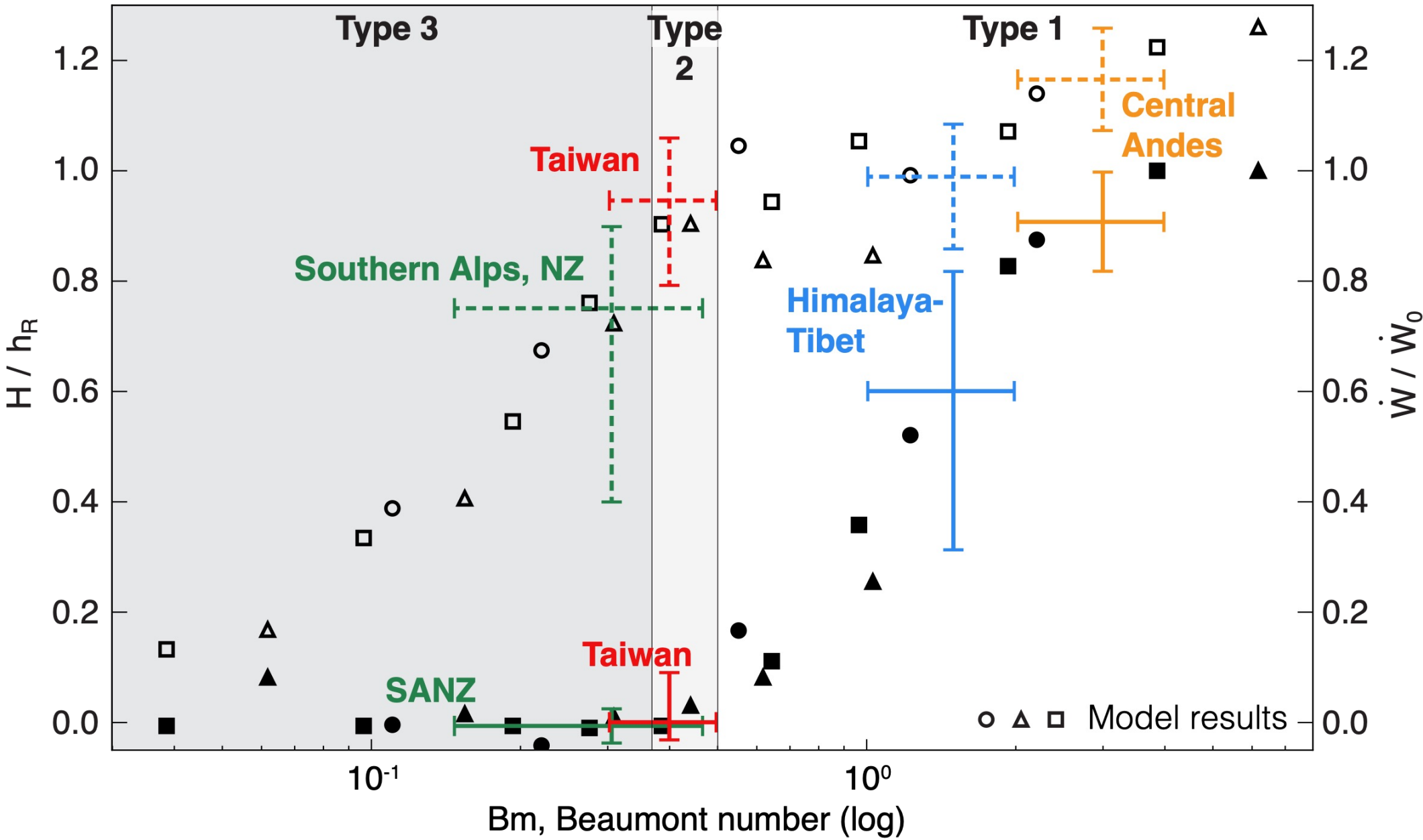
# The Beaumont number of models and orogens



- Type 1: Himalaya-Tibet and Central Andes
- Type 2: Taiwan
- Type 3: Southern Alps of New Zealand.
- Crustal strength varies only by a factor 2-3.
- Convergence rate and fluvial erodibility primarily determine orogen type.

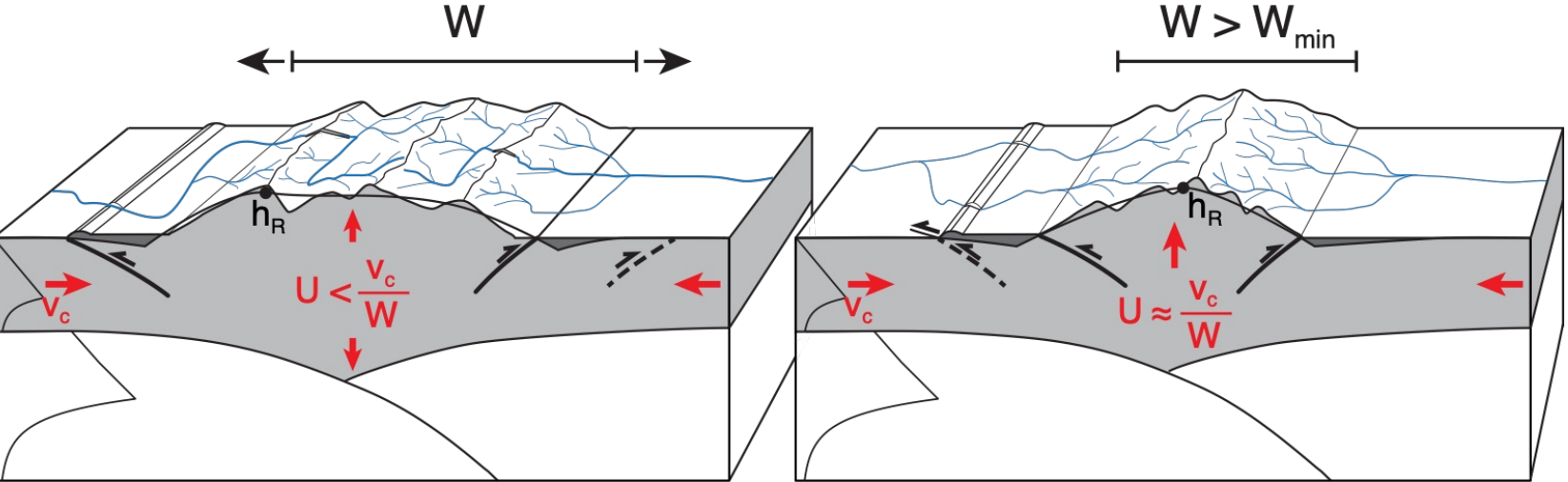


# Conclusions

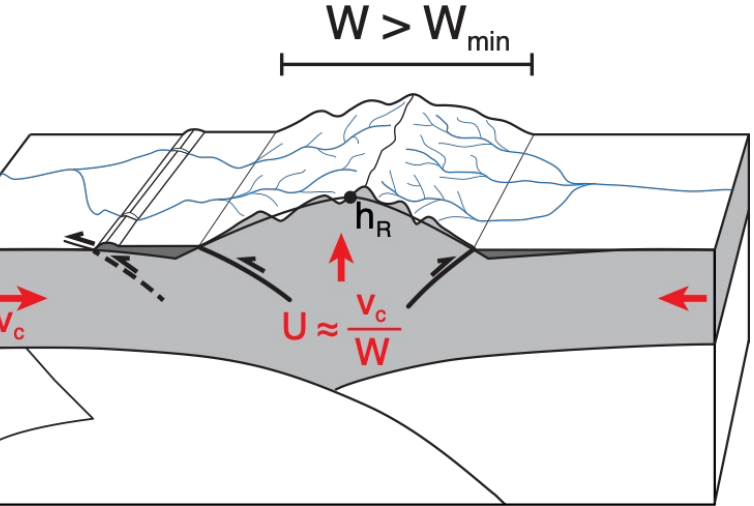


- 1) We investigate the interplay between **surface processes and tectonics** during collisional **mountain building**.
- 2) We define a **new non-dimensional number**, the **Beaumont number,  $B_m$** , that **quantifies** the interaction between **surface processes and tectonics**.
- 3)  $B_m$  allows estimation of **crustal strength** and **fluvial erodibility**.
- 4) Based on  $B_m$  and model results, we define **three types** of mountain belts.
- 5) We **classify** 4 different **orogens**, find their controlling factors, and their crustal strength and fluvial erodibility

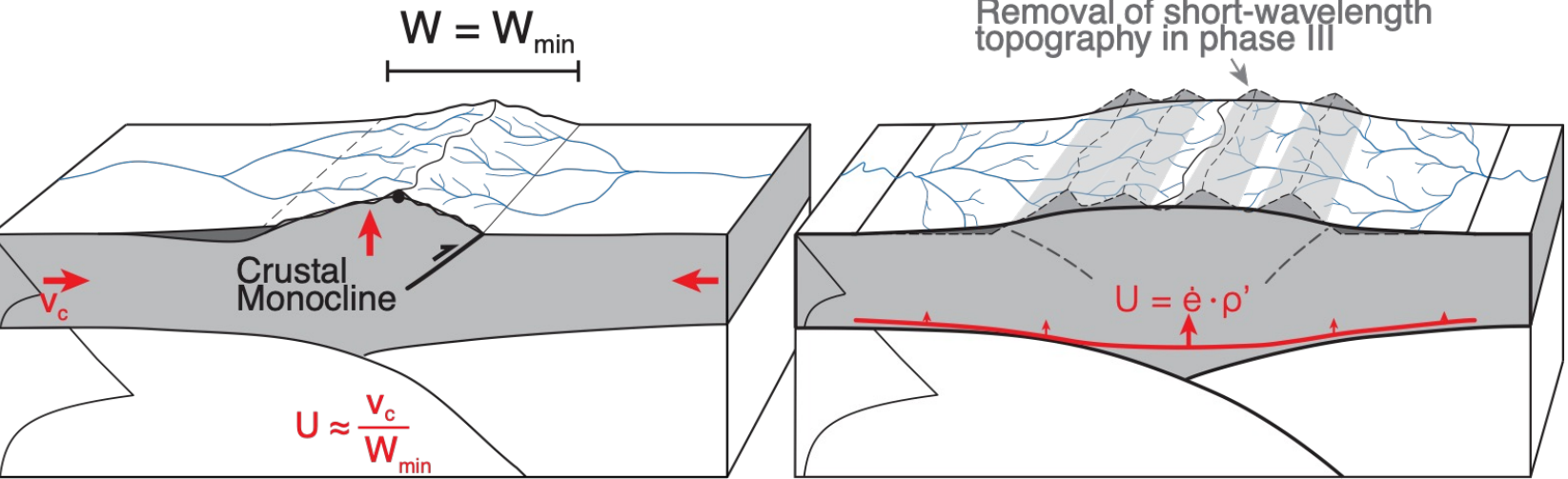
**b Type 1: non steady state, strength limited,  $B_m > 0.5$**



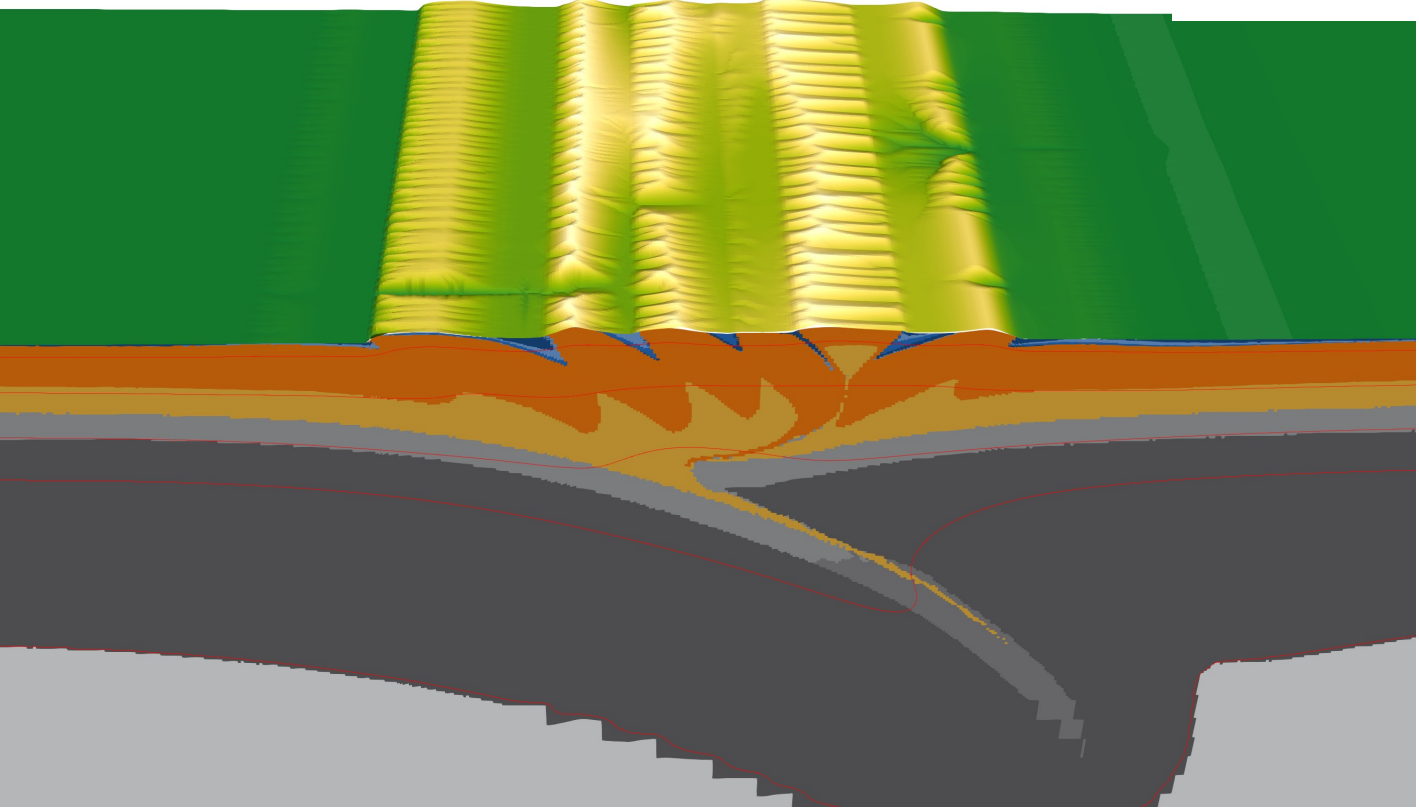
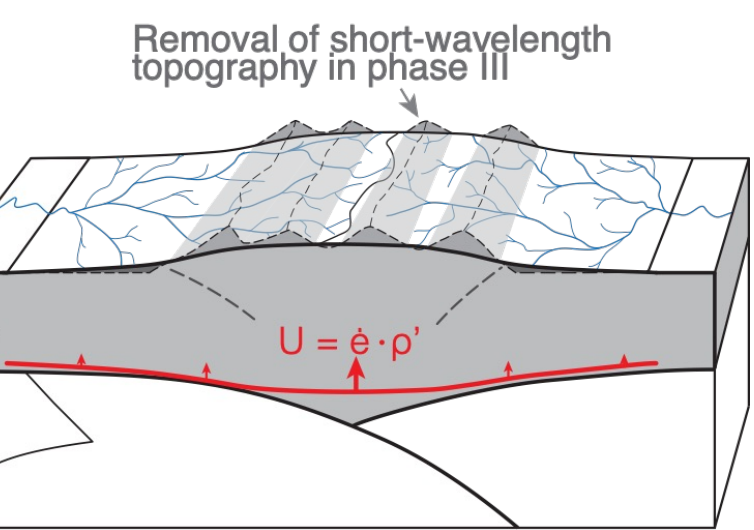
**c Type 2: flux steady state, strength limited,  $B_m \approx 0.4-0.5$**



**d Type 3: flux steady state, erosion limited,  $B_m < 0.4$**



**e Decay**



**nature**

Wednesday, 1<sup>st</sup> of June