

Transform versus non-transform offsets are controlled by the variation of magmatic accretion in the offset zone as well as by the offset distance

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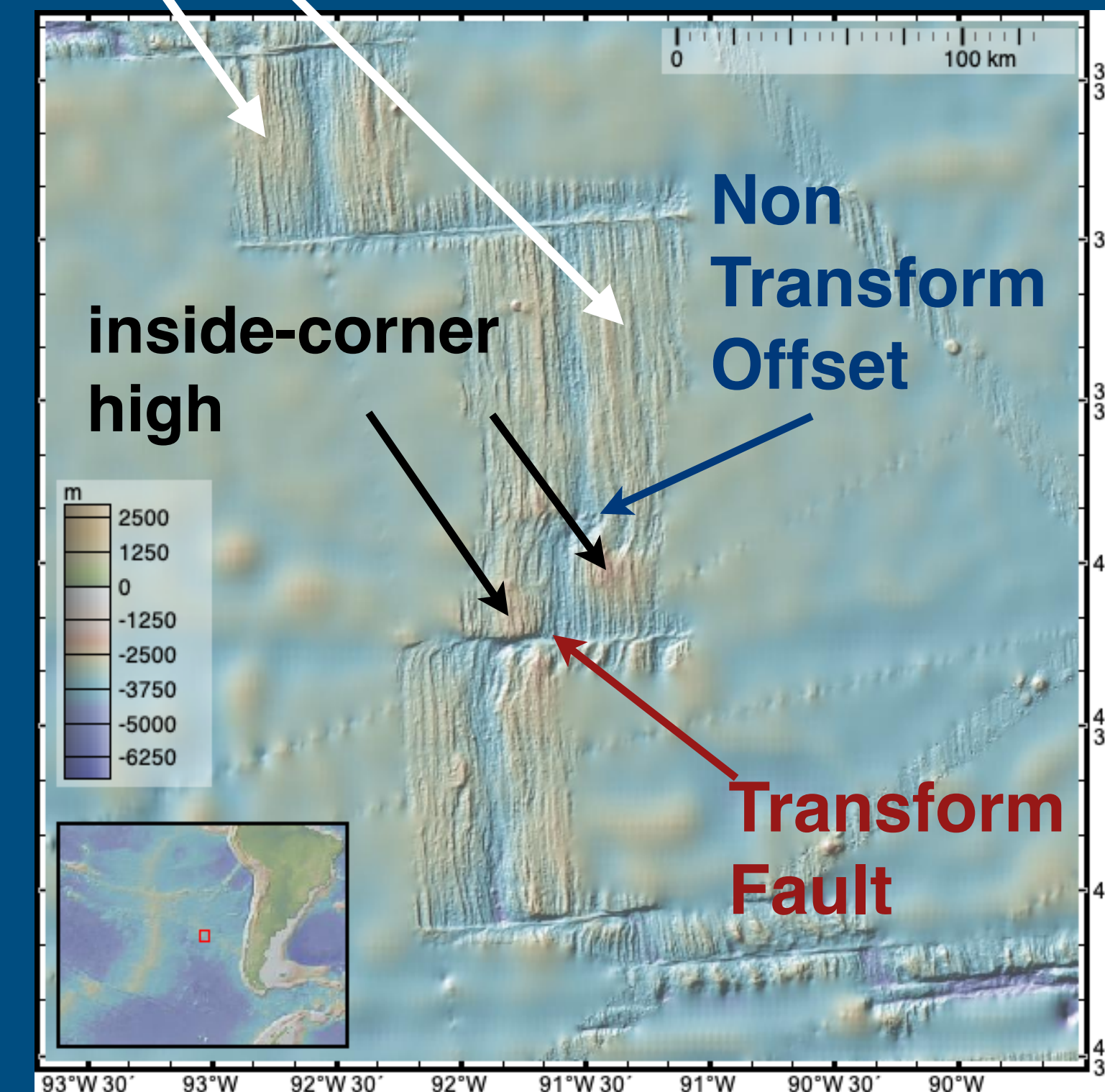
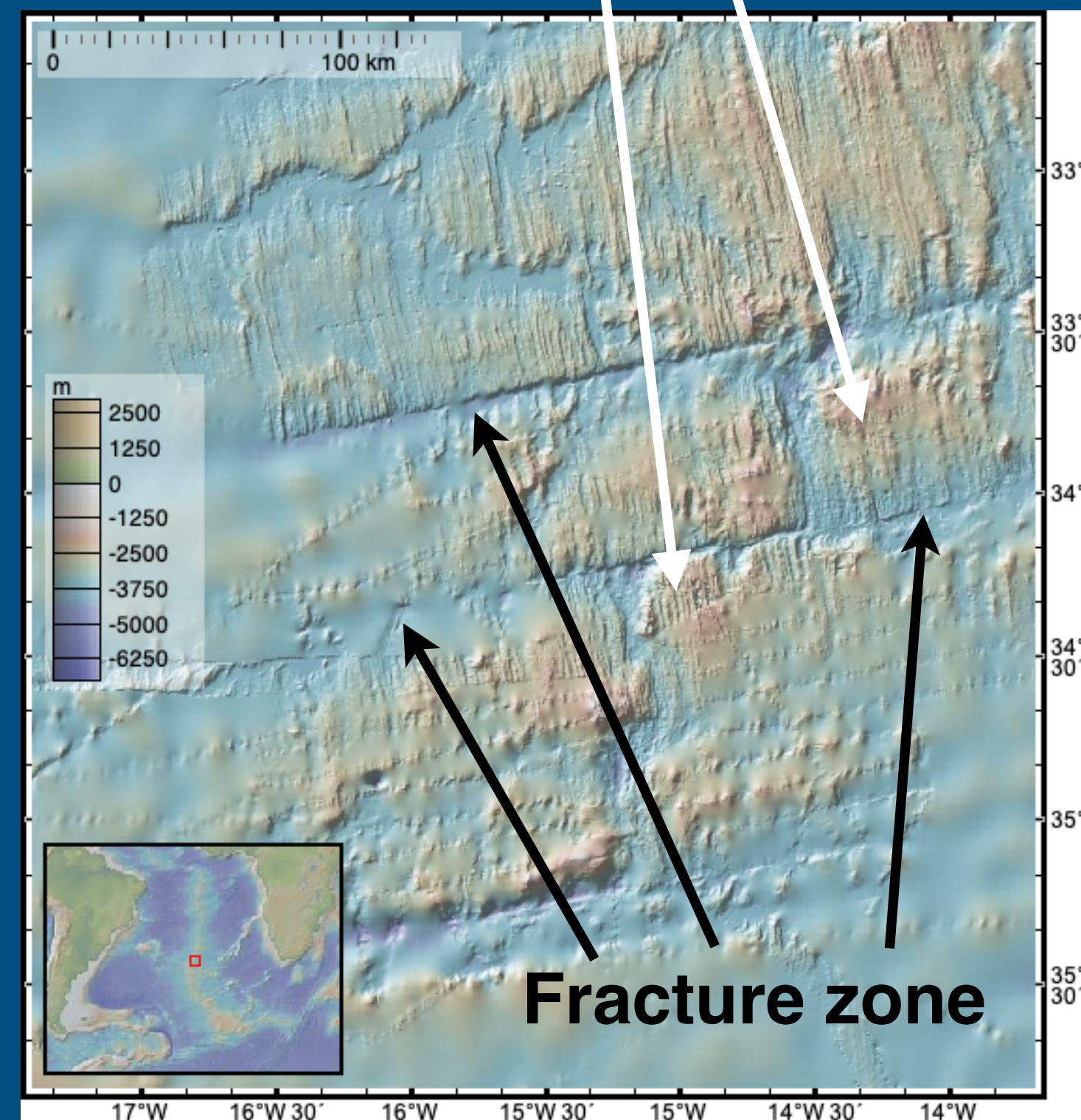
24th May, 2022

Introduction

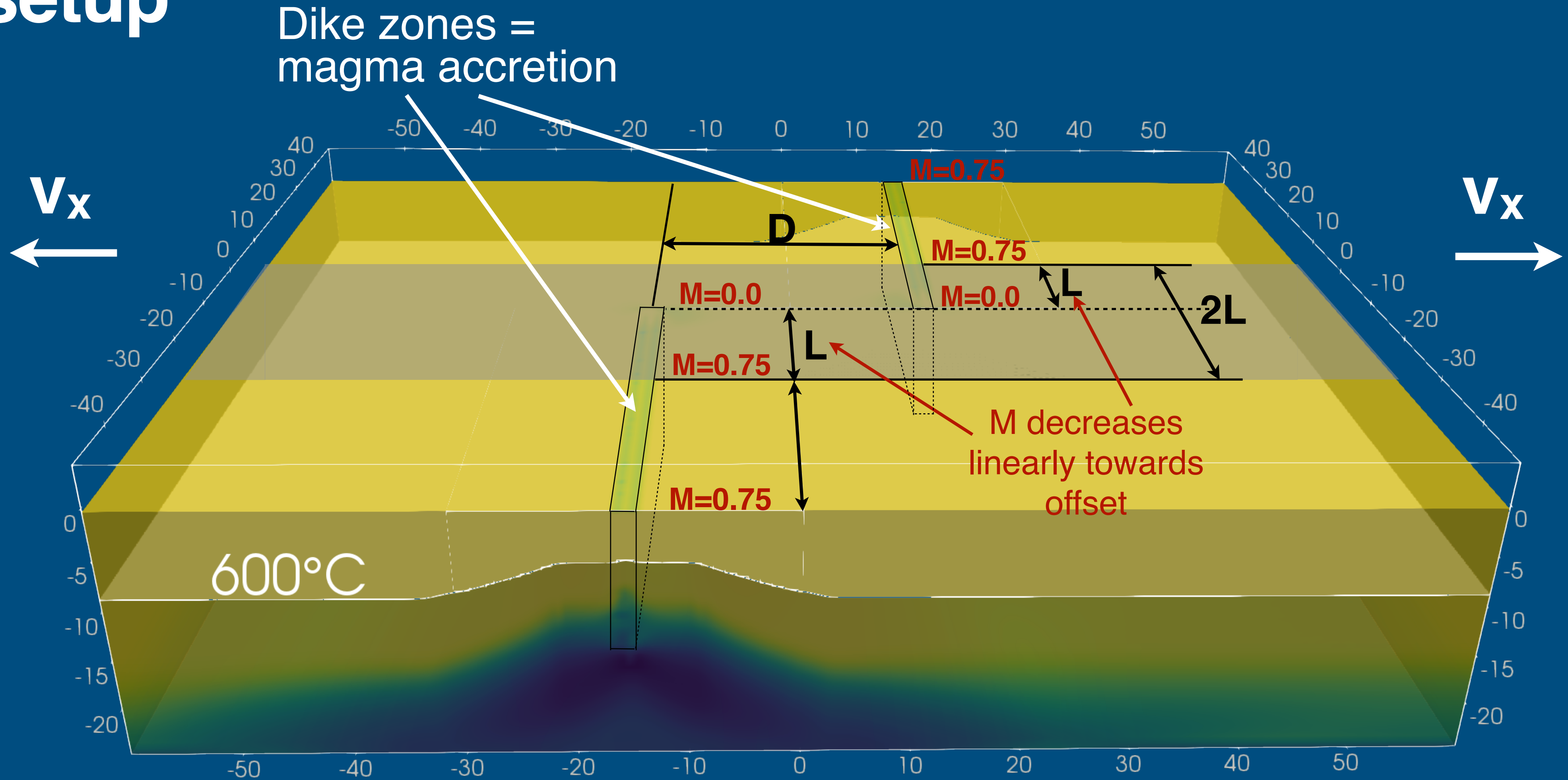
- Transform faults (TF) or non-transform offsets (NTO) separate mid-ocean ridges (MOR)
- Segment offsets can evolve from TFs to NTOs or the other way around (e.g. Hey et al., 2015)
- Fraction of spreading accommodated magmatically (M) influences fault spacing, axial topography and ridge morphology (e.g. Buck et al. 2005, Ito & Behn 2008, Howell et al. 2016)

What controls the formation of transform faults (TF) versus non-transform offsets (NTO) ?

fault scarps = abyssal hill faults



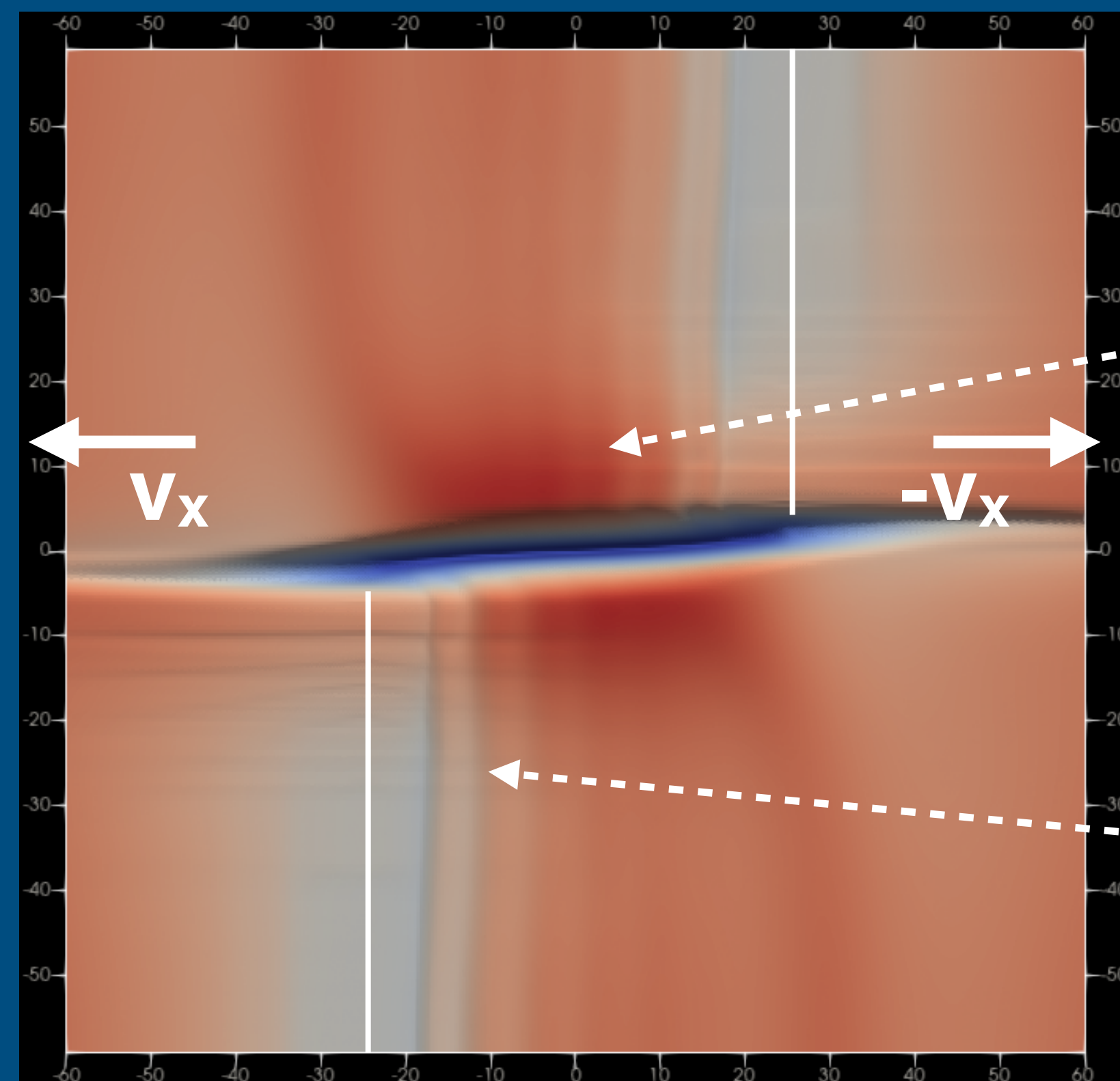
Model setup



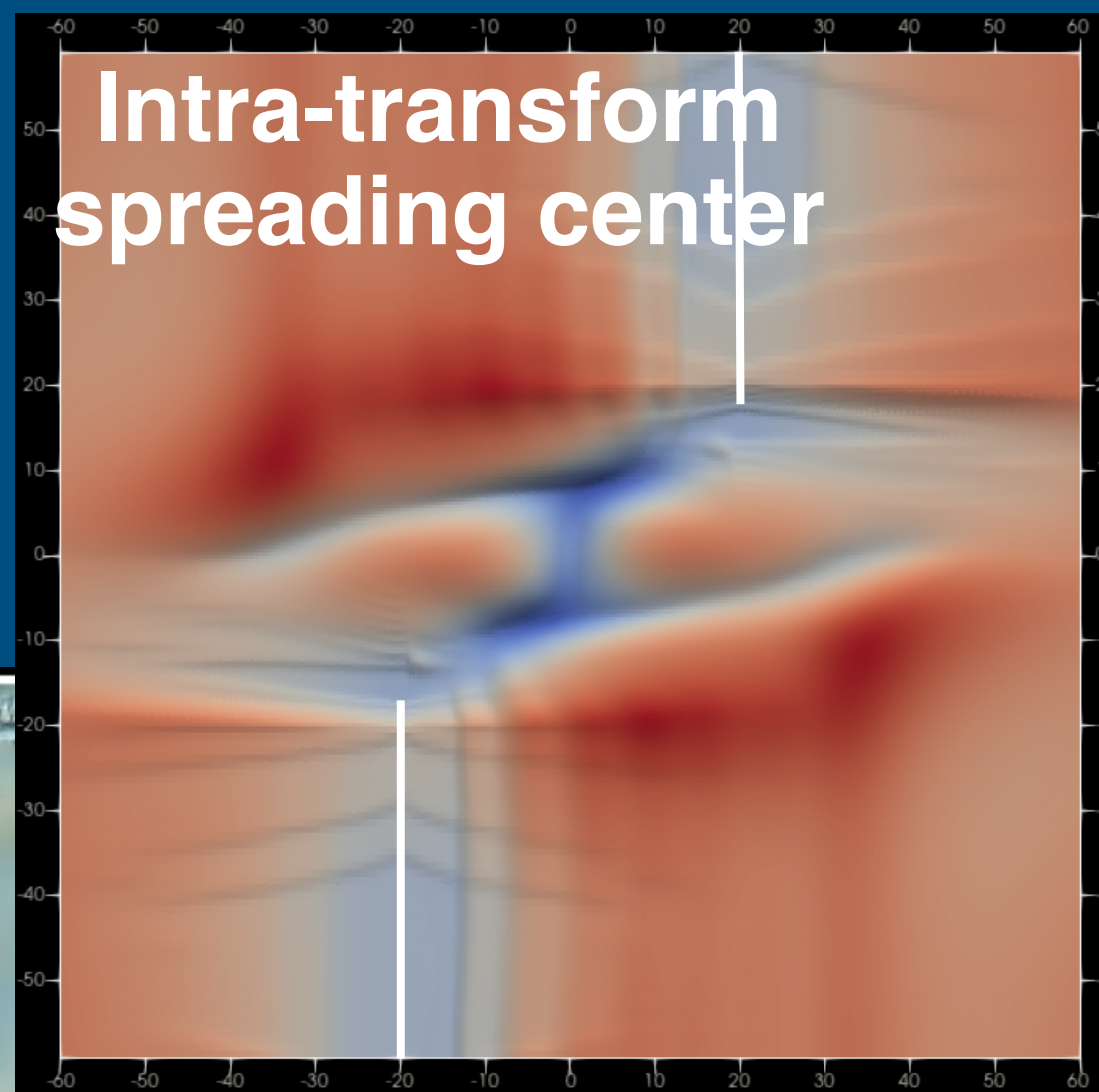
- 3D finite difference code LaMEM: solves for momentum, mass and heat conservation
- visco-elasto-plastic model allows spontaneous fault formation and evolution
- “dike zone”: imposed divergence to simulate magmatic accretion along one ridge segment

Numerical results and natural observations

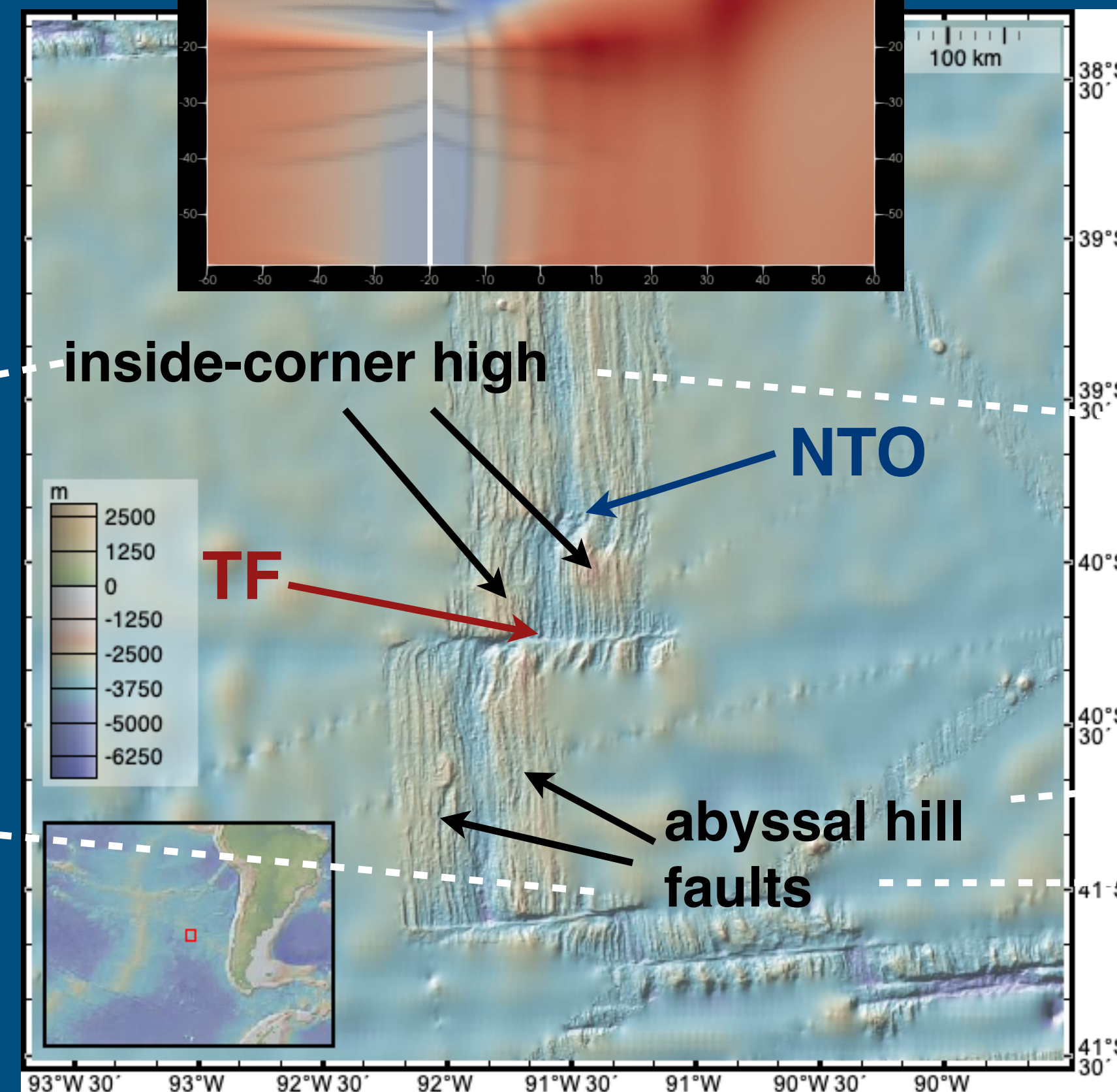
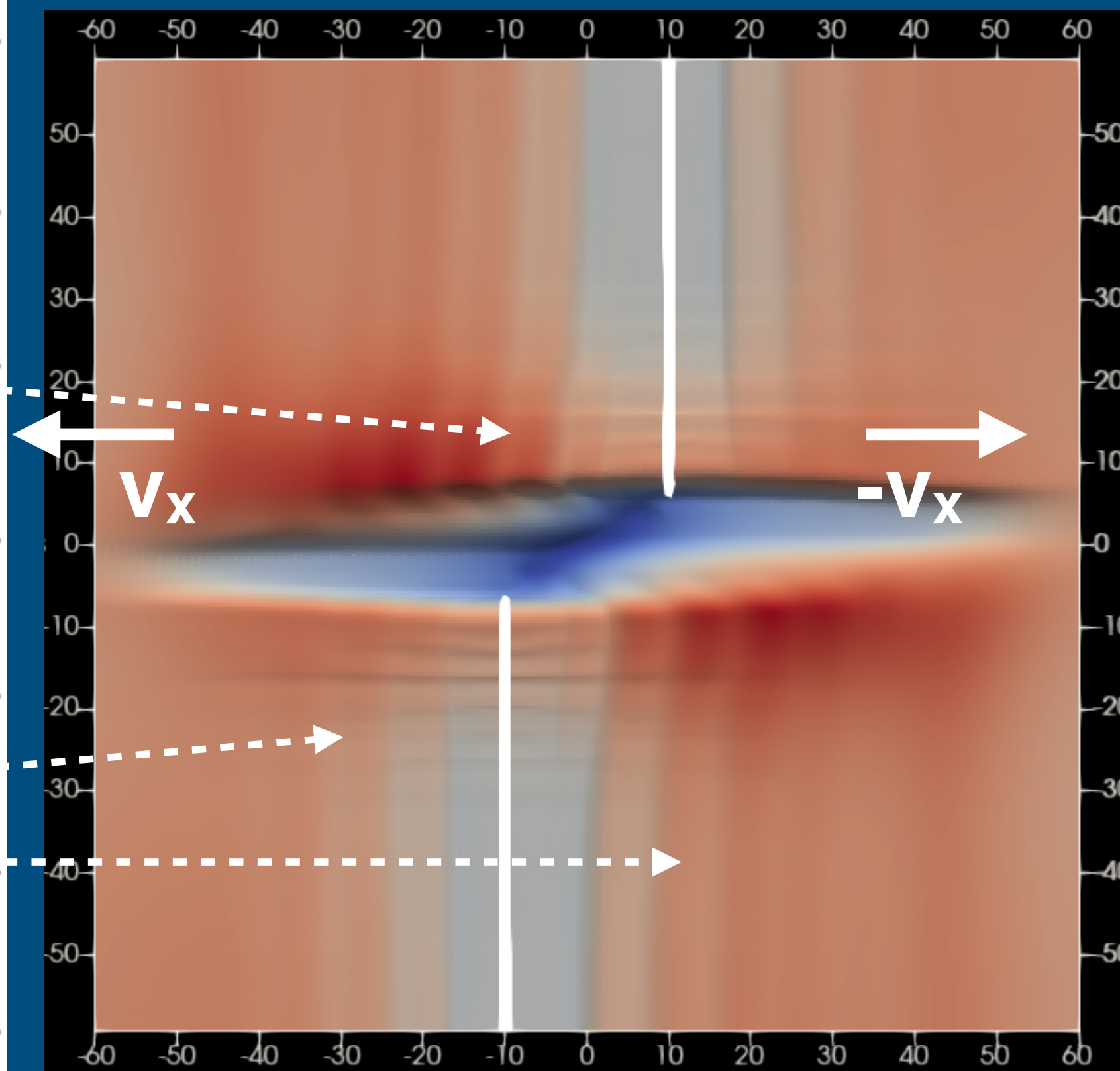
**Transform Fault
(TF)**



**Intra-transform
spreading center**



**Non Transform Offset
(NTO)**



Velocity field to define regime (backup)

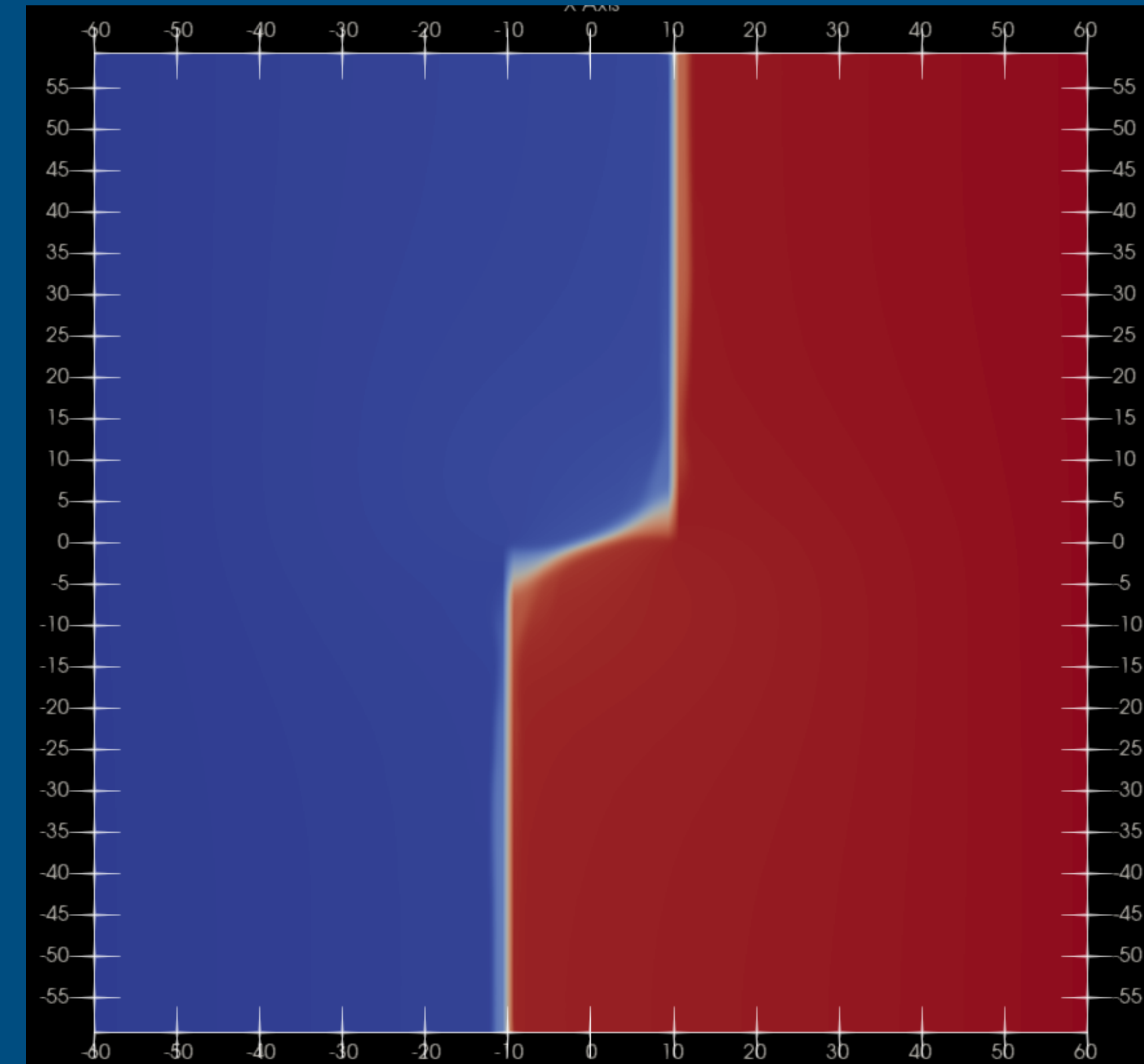
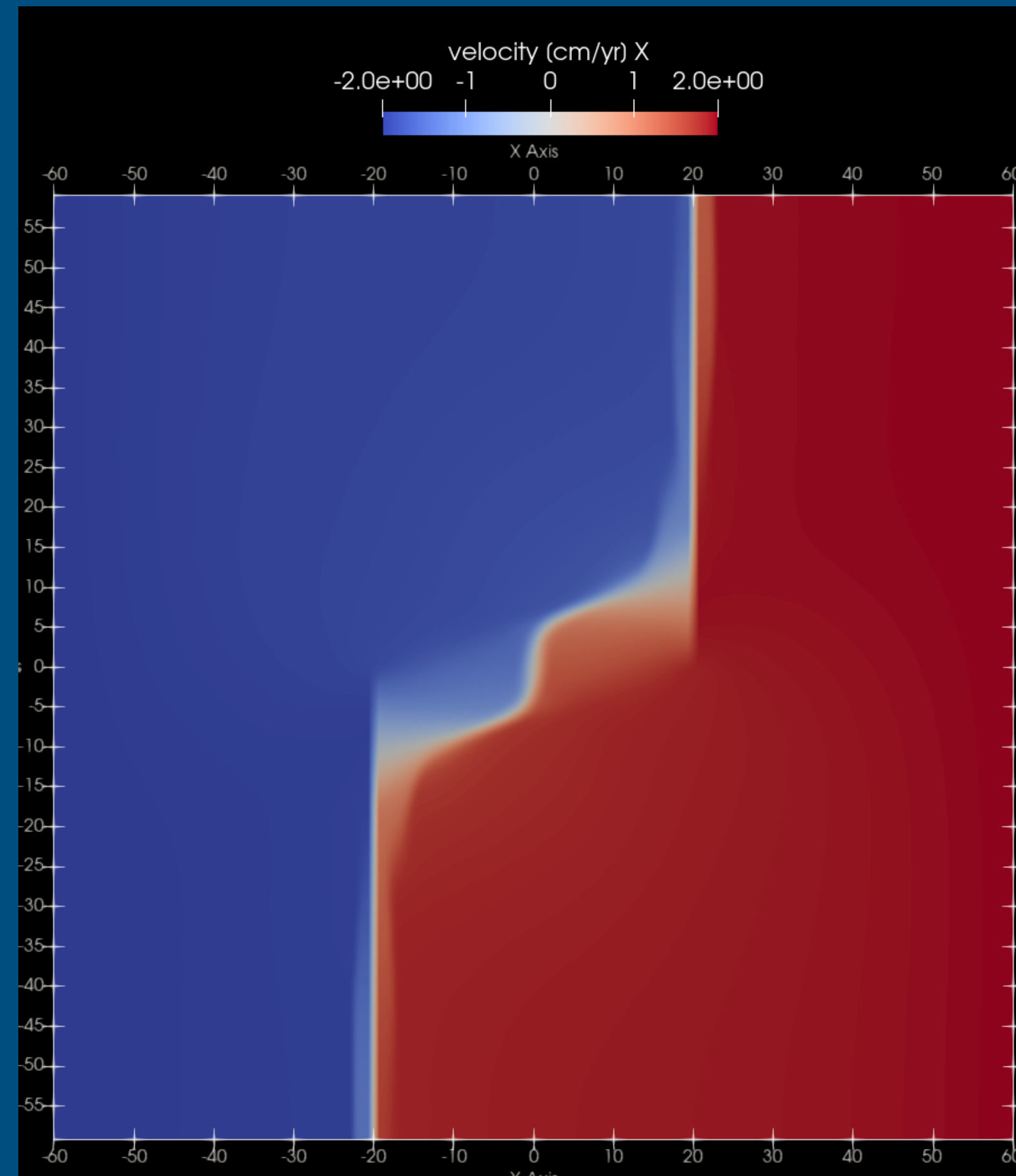
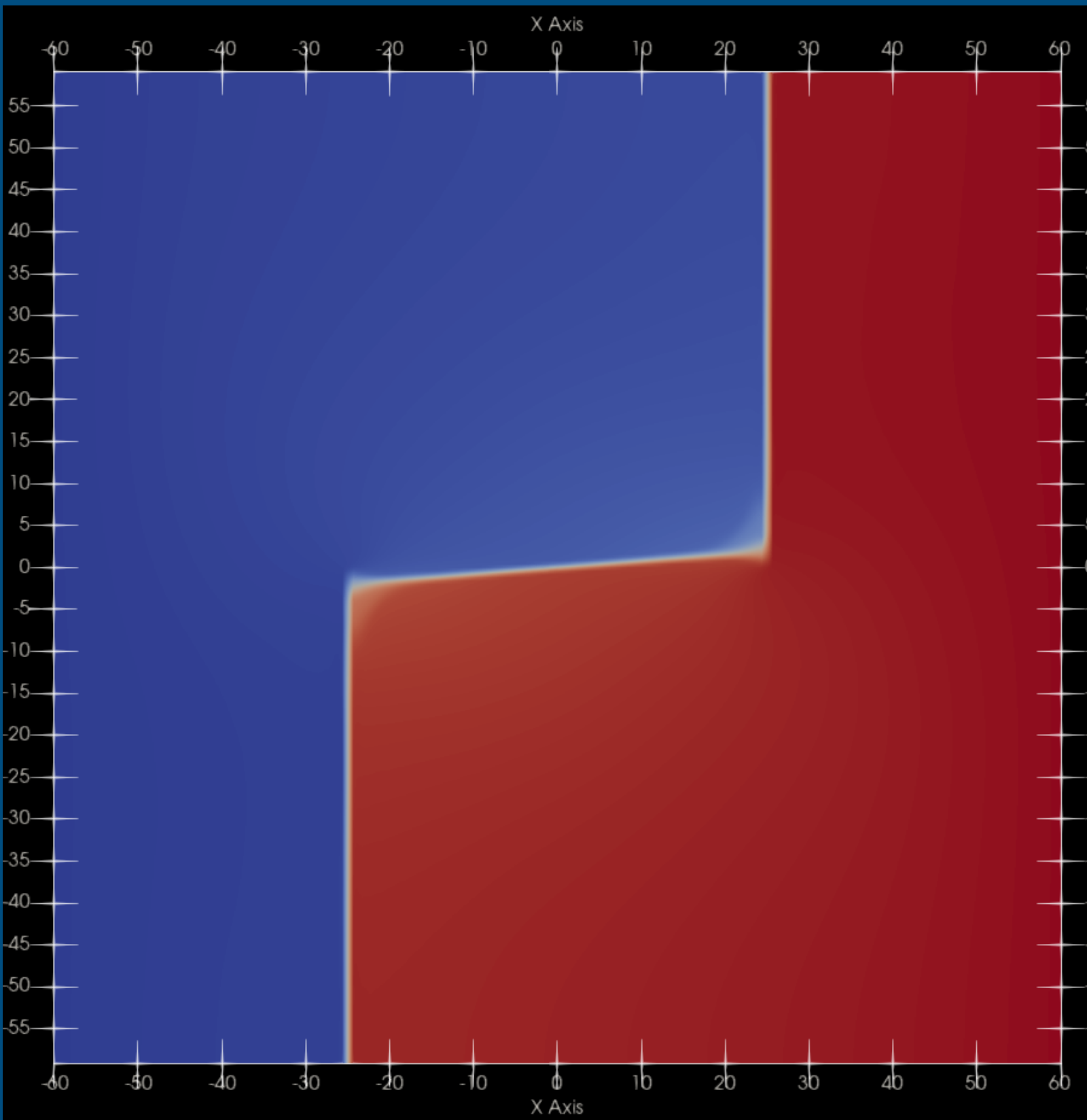
**Transform Fault
(TF)**

$< 8^\circ$

**Intra Transform Spreading
Center
(ITSC)**

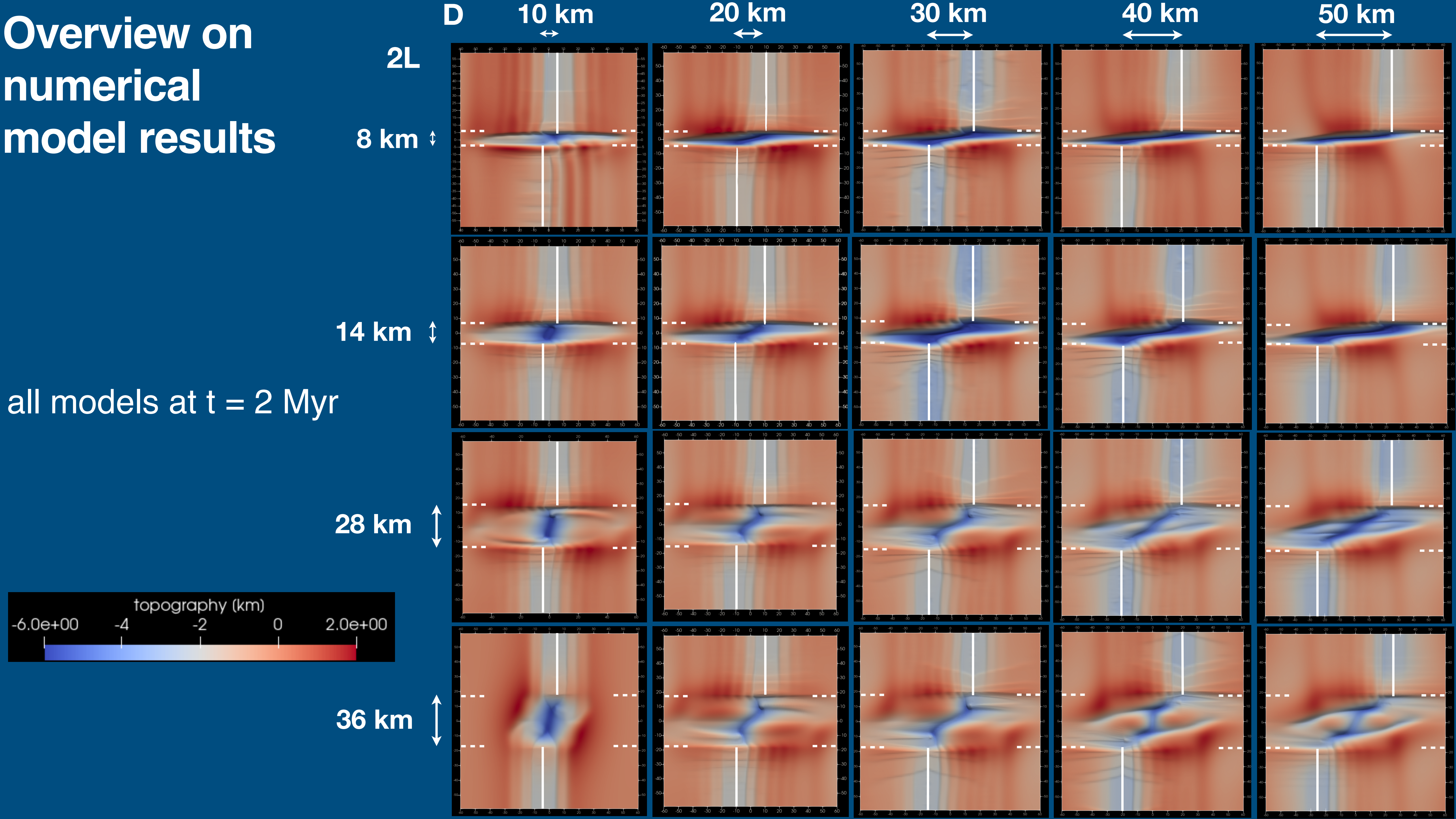
**Non Transform Offset
(NTO)**

$> 8^\circ$



Overview on numerical model results

all models at $t = 2 \text{ Myr}$

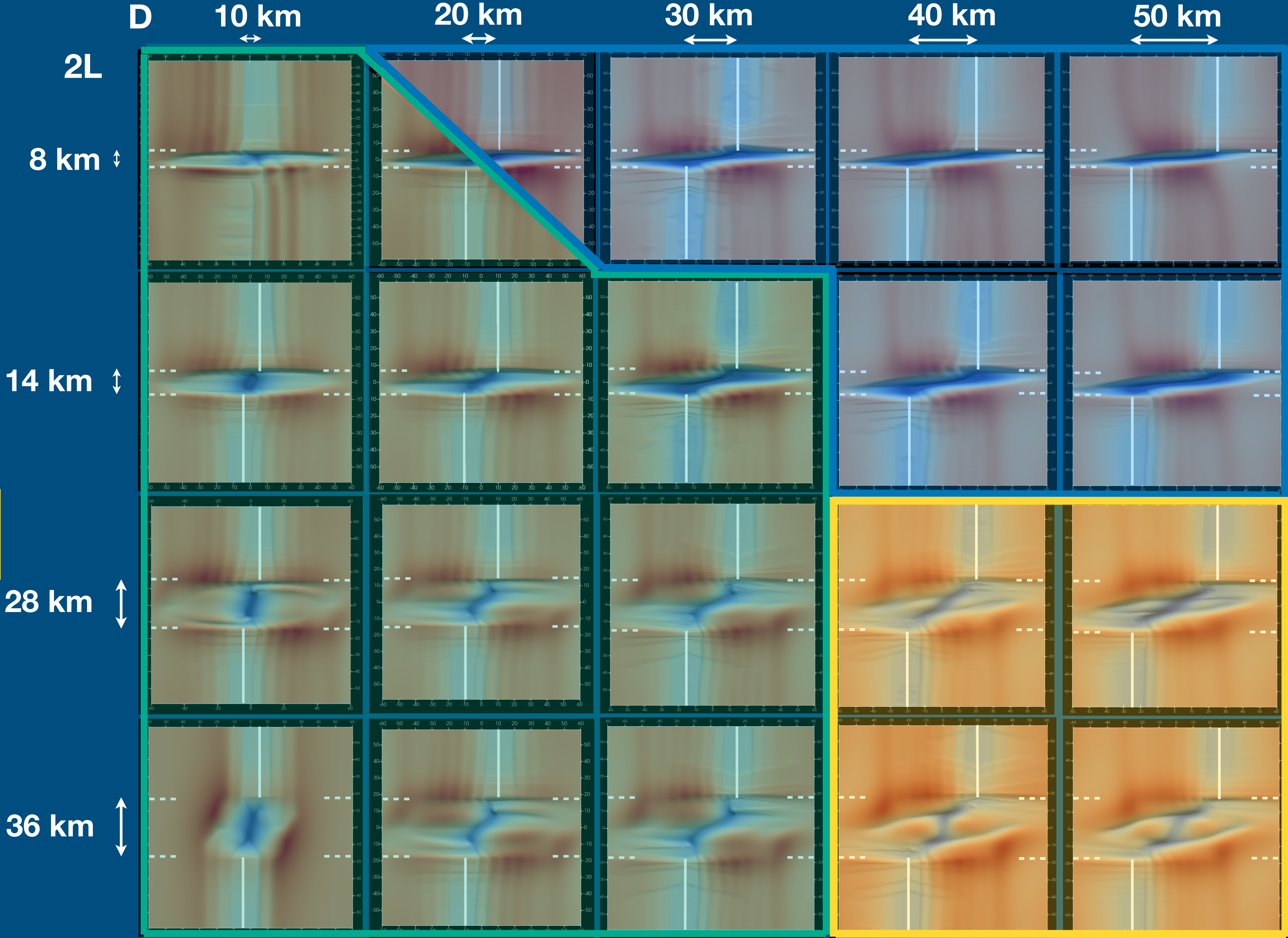


Overview on numerical model results

1) Small-medium offset, any underlap: NTO

2) medium-large offset, small underlap: TF

3) Large offset, large underlap: ITSC

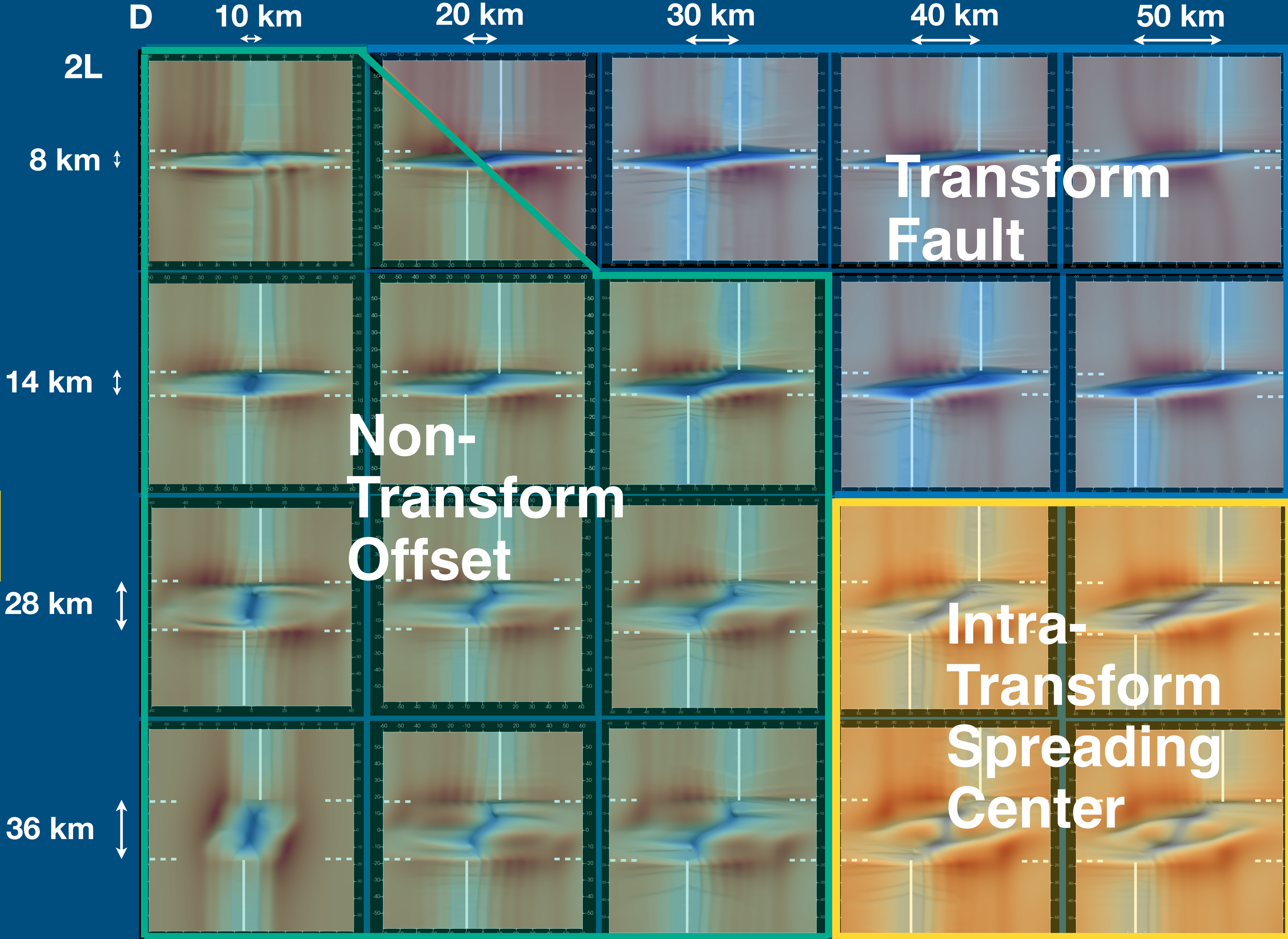


Overview on numerical model results

1) Small-medium offset, any underlap: NTO

2) medium-large offset, small underlap: TF

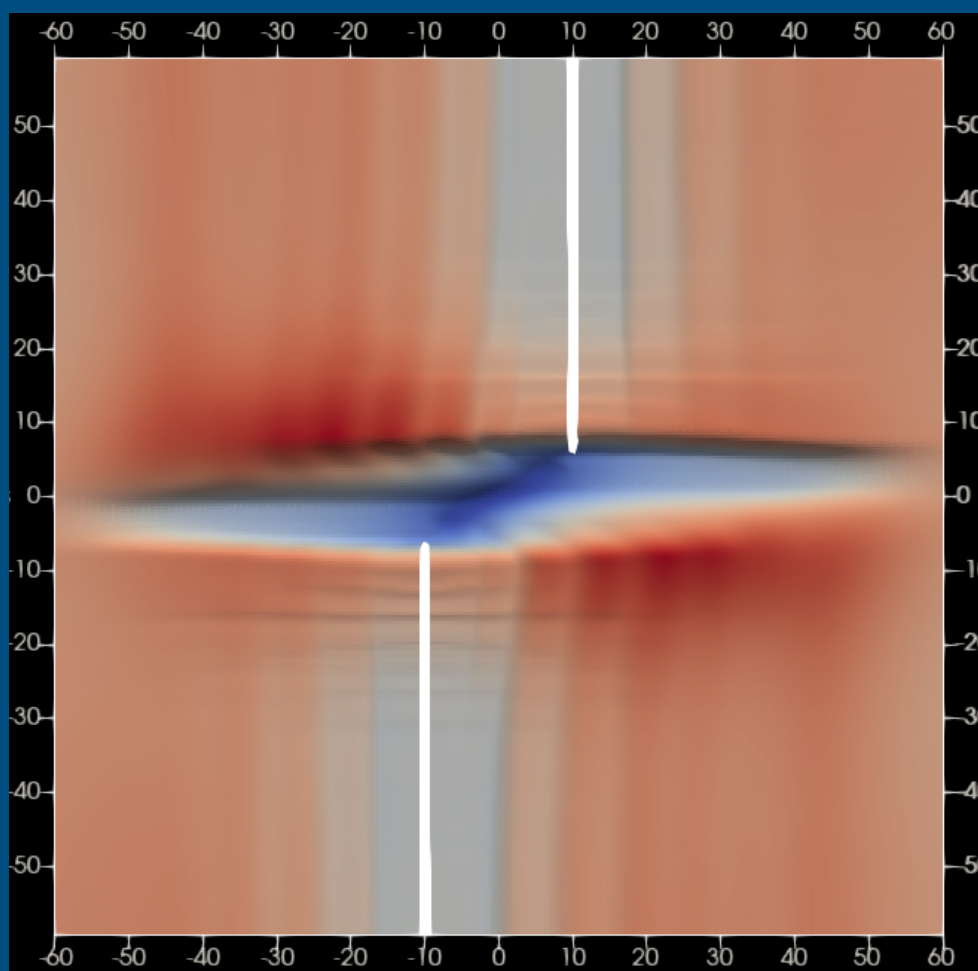
3) Large offset, large underlap: ITSC



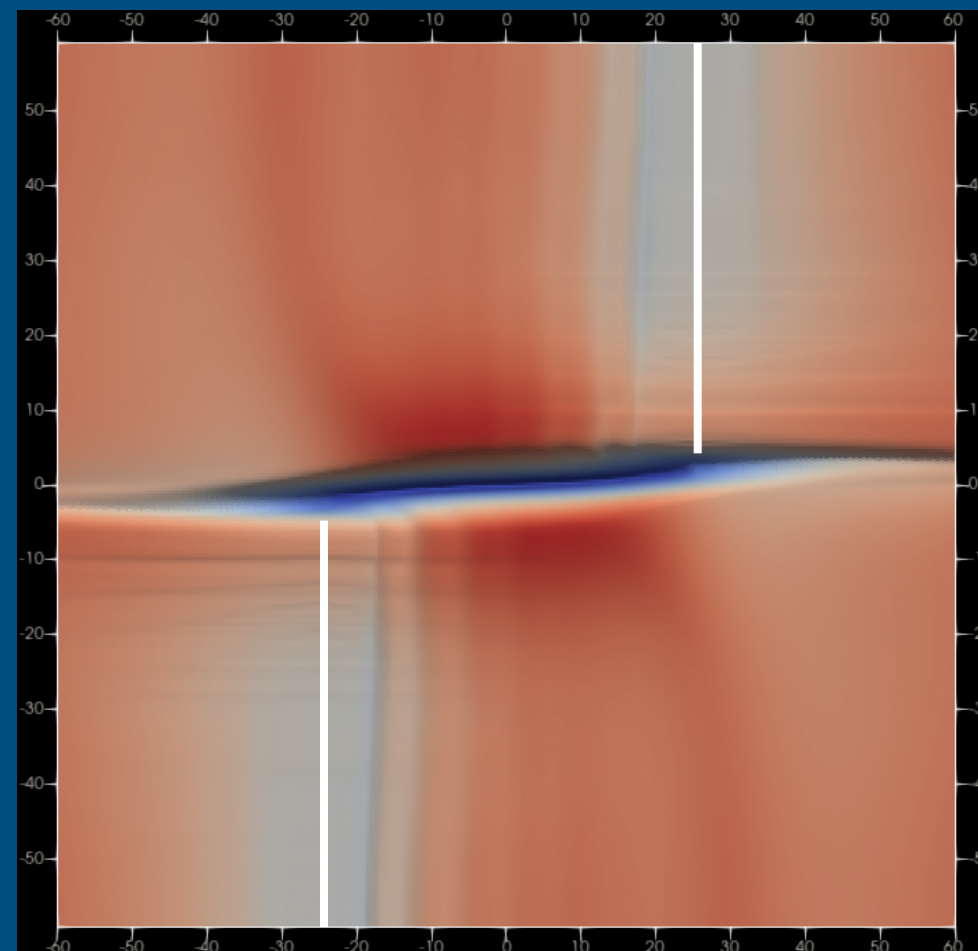
Summary Part I

- Models can reproduce a variety of offset-types between ridges
- Depending on offset D and underlap distance L , 3 different regimes can be identified:

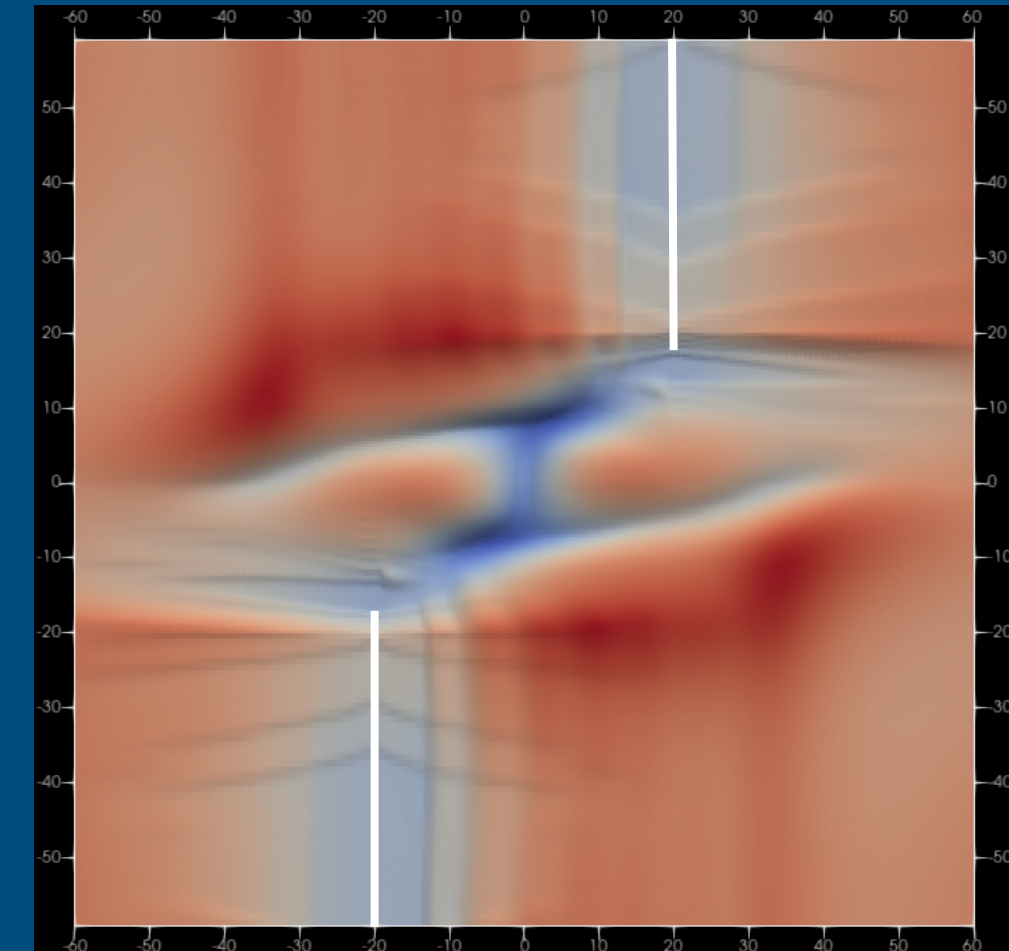
■ NTO
small-medium offset,
all distances of underlap



TF
medium-large offset
small underlap



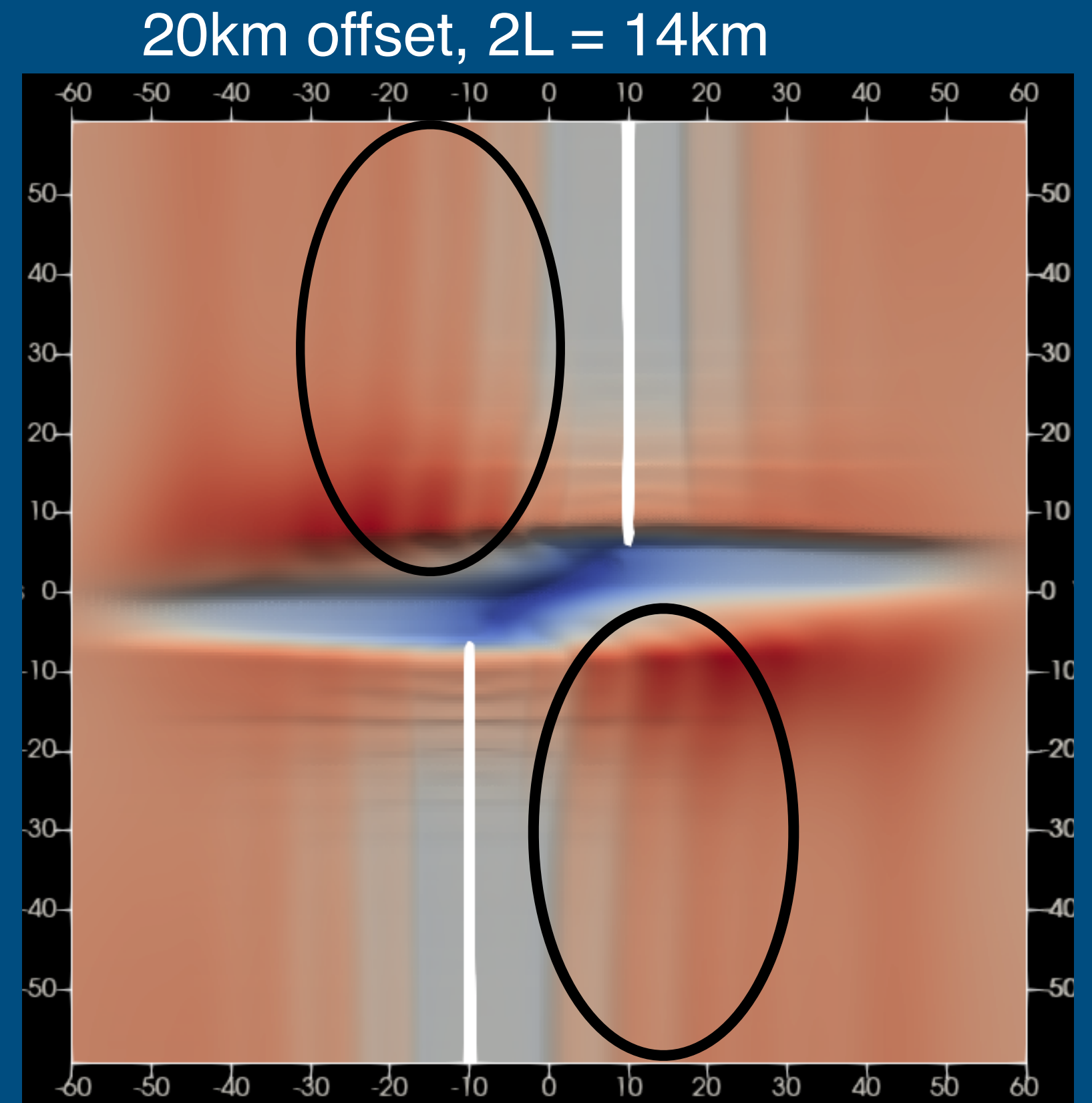
ITSC
large offset,
large underlap



Part II: Development of abyssal hills

Observation: faulting preferably on the inside where transform fault is located

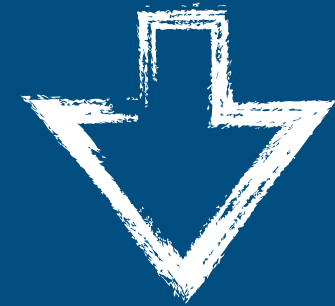
Investigate the importance of relative weakness of TF and FZ for fault development



Part II: Development of abyssal hills

Observation: faulting preferably on the inside where transform fault is located

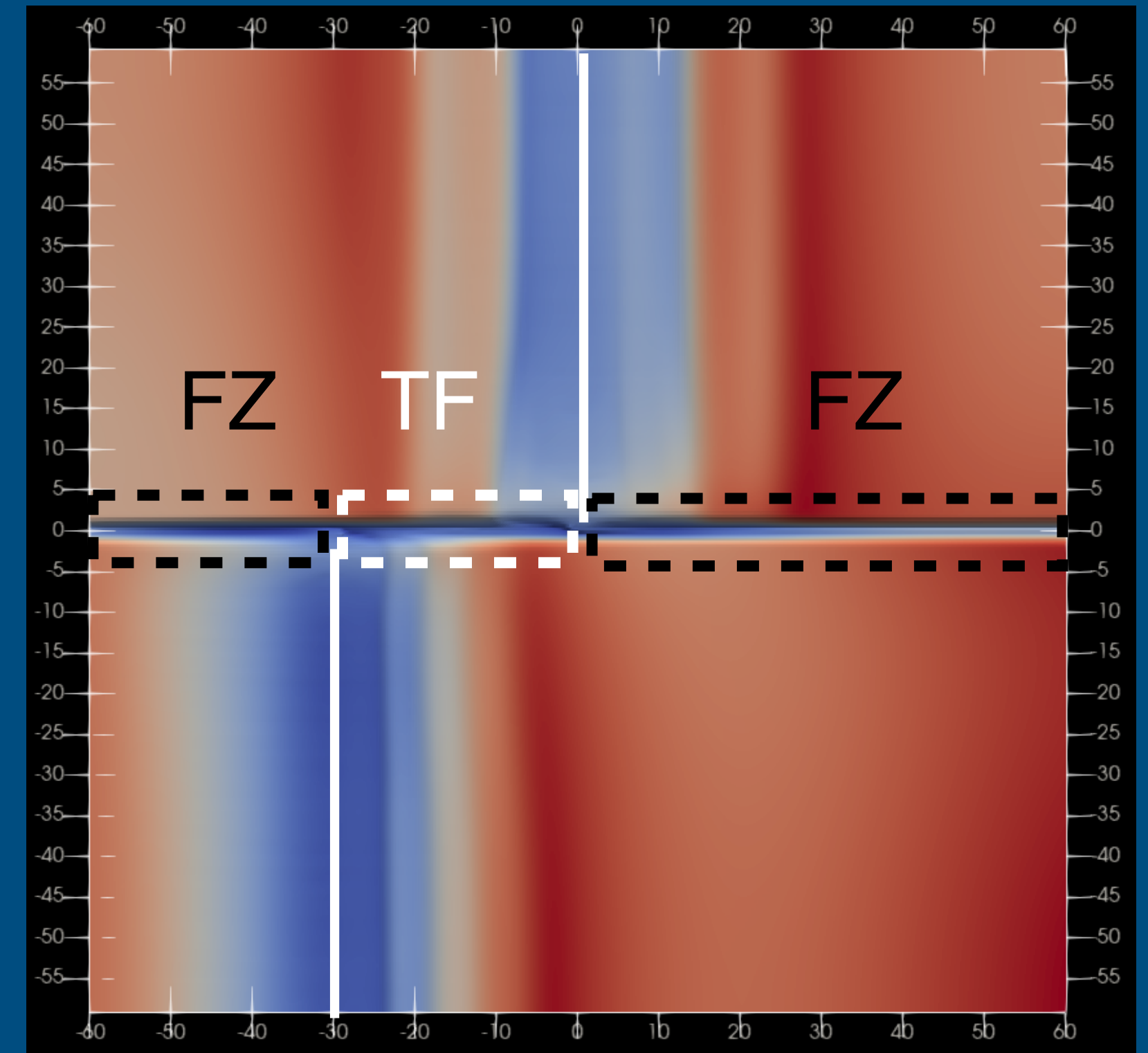
Investigate the importance of relative weakness of TF and FZ for fault development



Isolate behavior of faulting: constant $M=0.75$

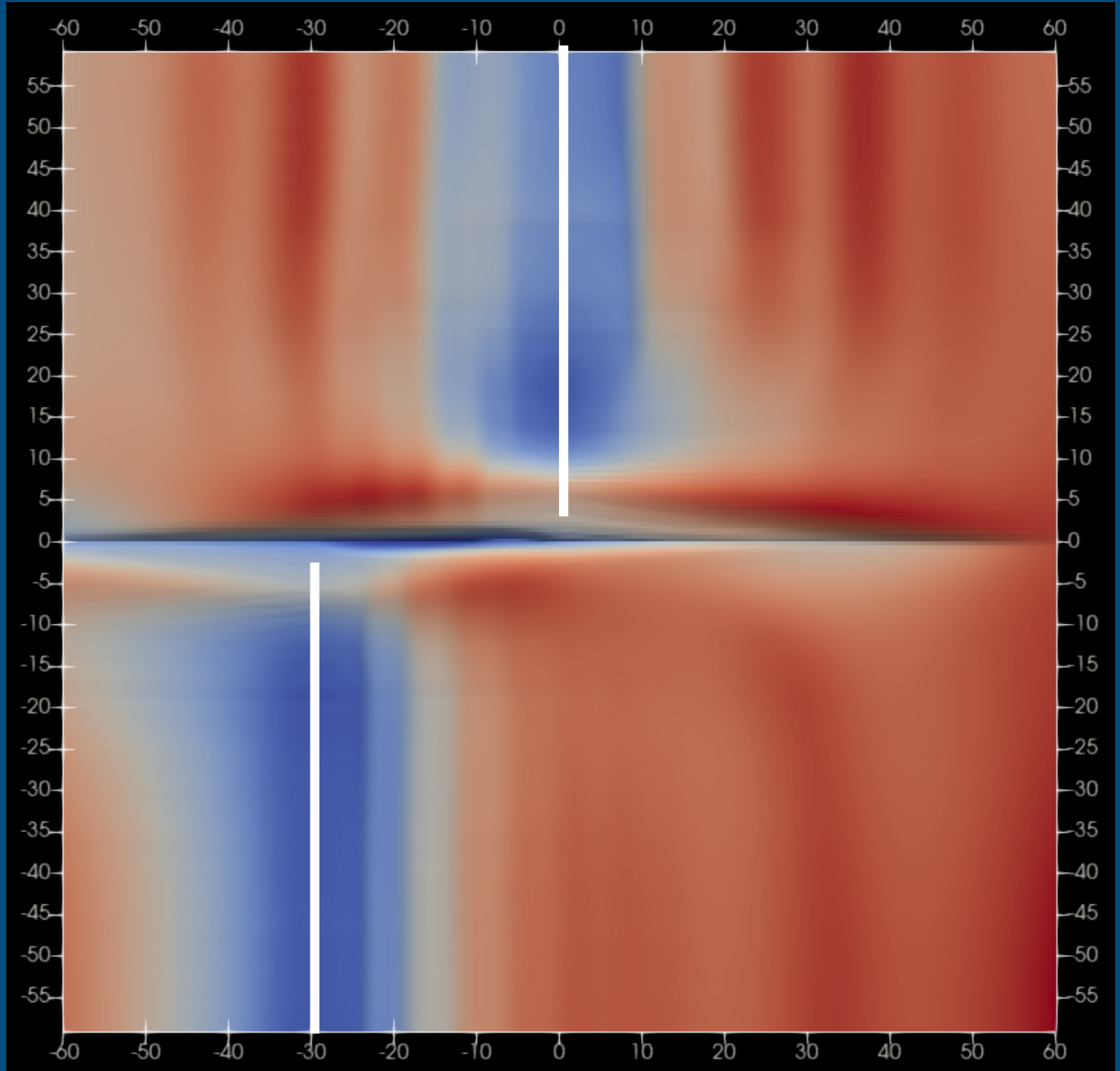
Impose lower initial friction angle and low cohesion in transform fault (TF) and/or fracture zone (FZ)

30km offset, $2L = 0\text{km}$



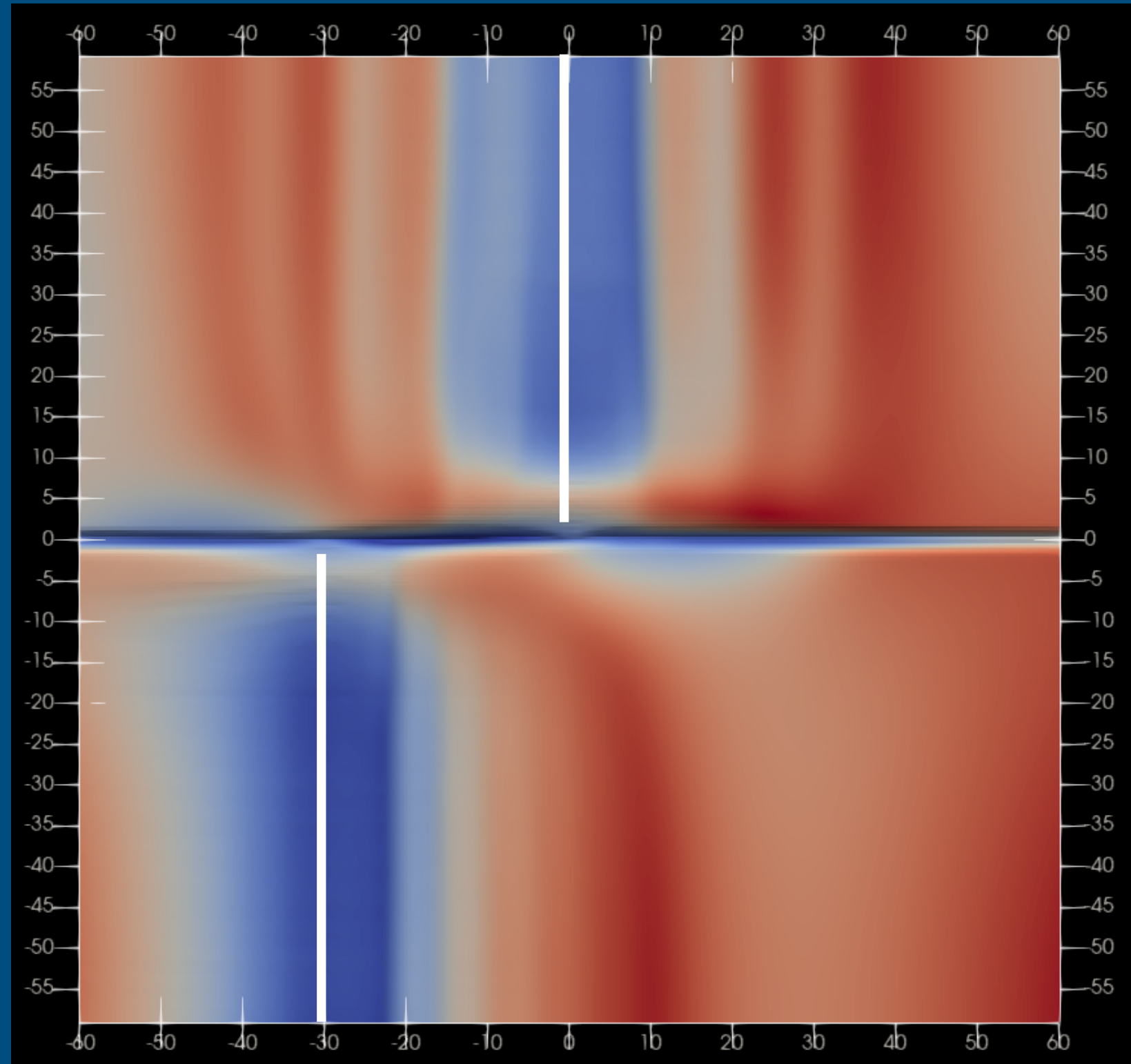
Results with different imposed weakening

No initial imposed weakening



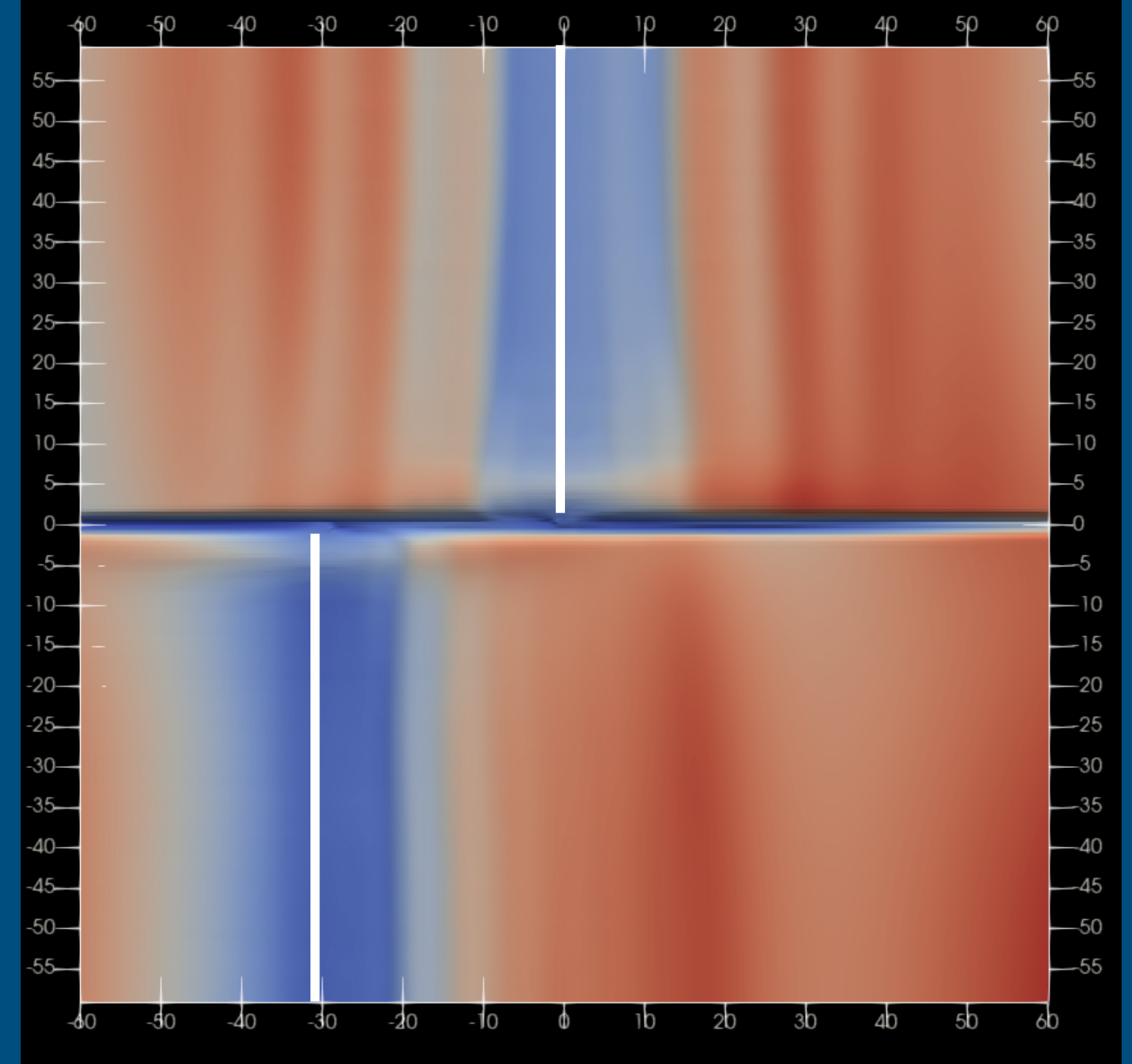
Only imposed on Fracture zone:

initial friction angle: 10°
initial cohesion: 10 Pa



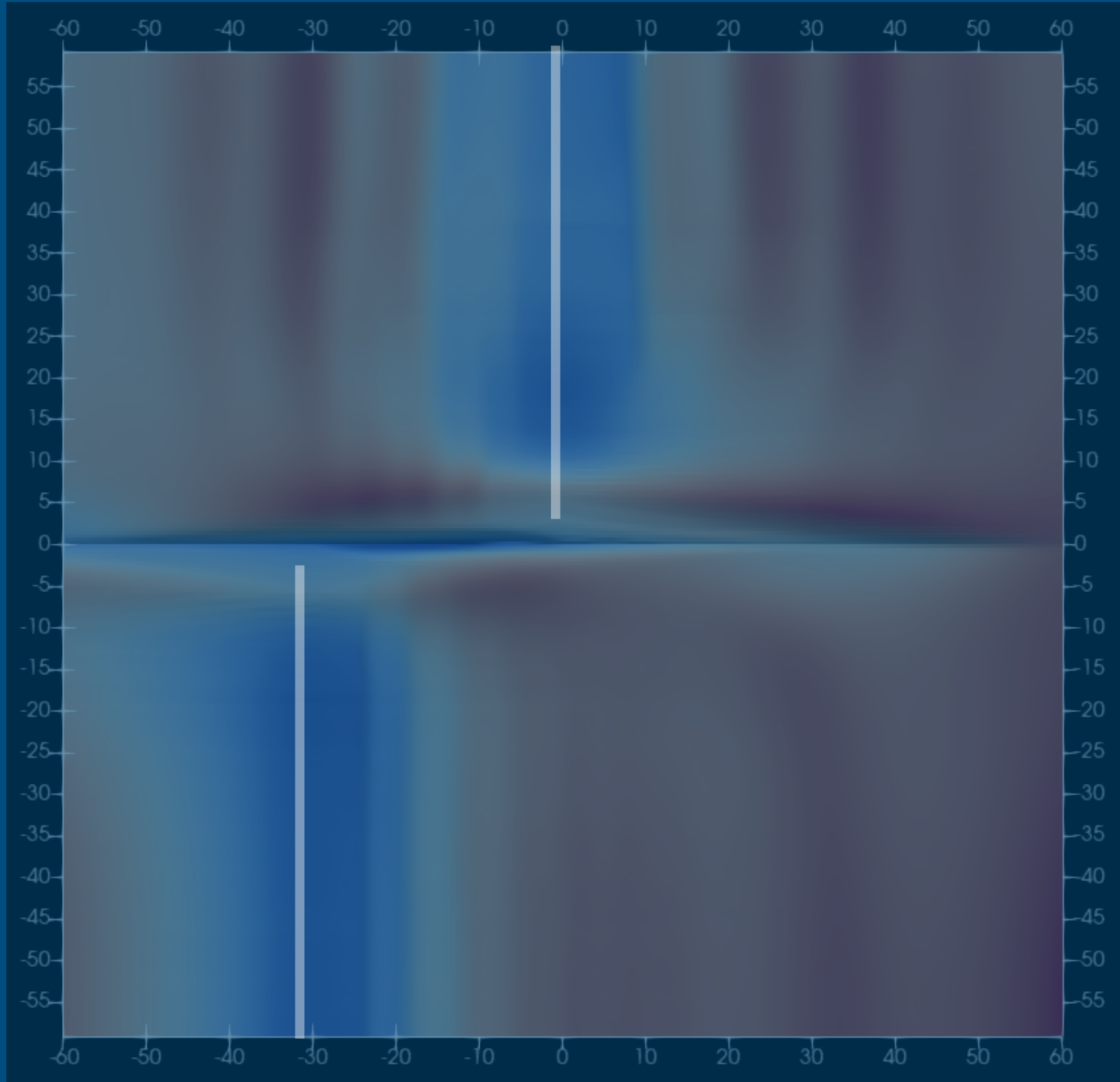
Imposed on FZ and TF:

initial friction angle: 10°
initial cohesion: 10 Pa



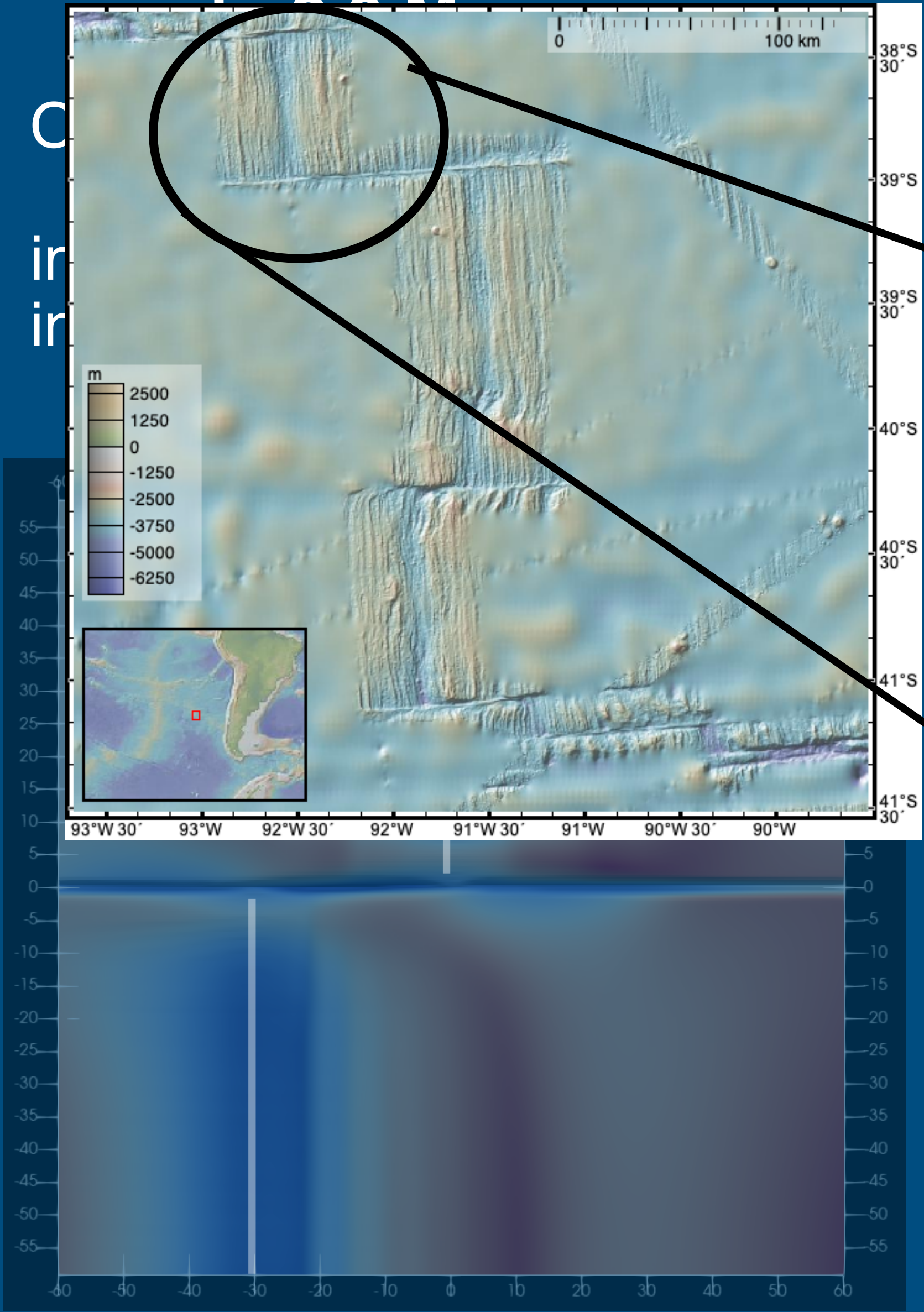
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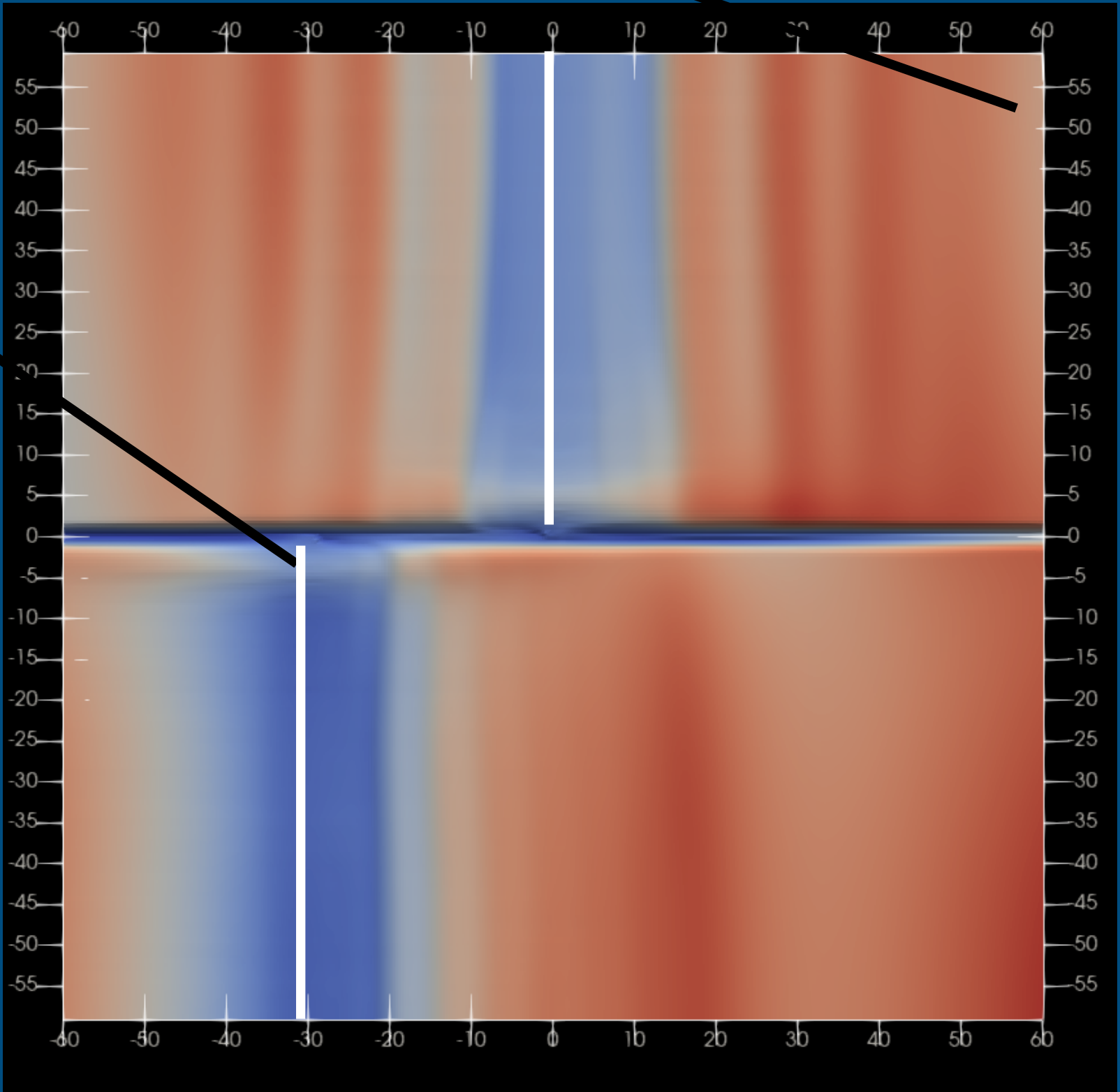
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Imposed on FZ and TF:

initial friction angle: 10°
initial cohesion: 10 Pa



Take home messages & Outlook

- *NTOs, TFs and ITSC* develop depending on the *size of offset* between ridges and the *distance over which magmatic extension decreases*
- a relatively *weak fracture zone* is necessary to *promote faulting on the inside and outside* of MORs (in our models) and lead to a more natural-looking TF

Next steps:

- define a clear boundary between TF and NTO (new?)
- relate relative weakness of FZ vs TF to fault spacing and occurrence in nature