

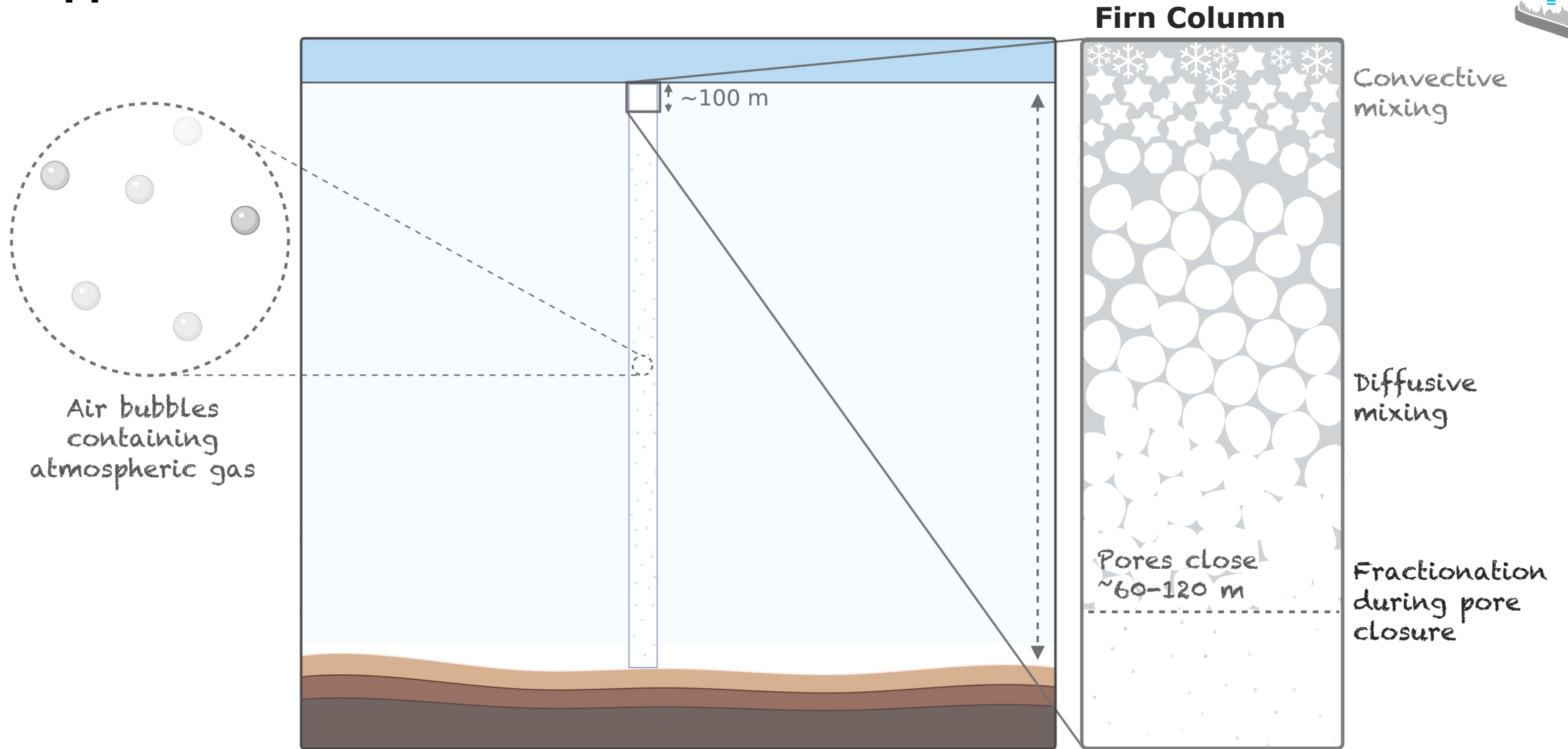
# Investigating the mechanisms controlling elemental fractionation during pore closure in polar firn

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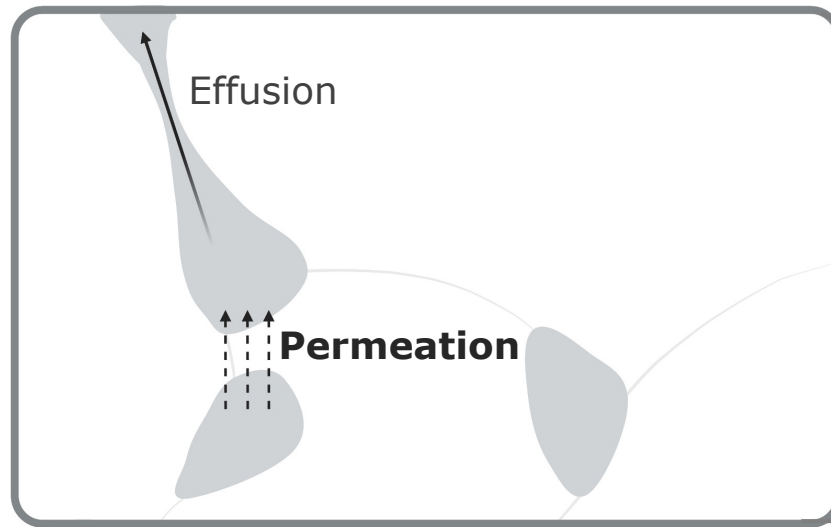
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# Physical processes in the firn column affect the composition of the air trapped in air bubbles in ice cores



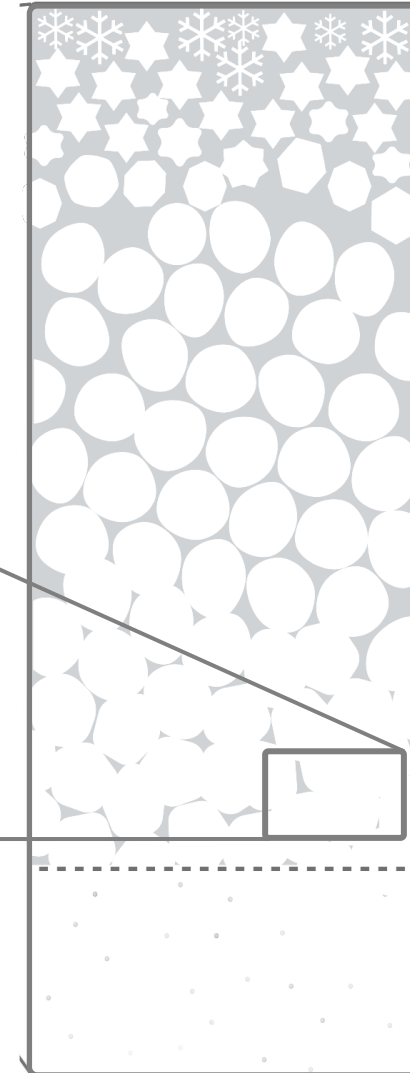
# $O_2$ is preferentially excluded during pore closure via permeation through ice matrix

- Permeability is function of grain size
- $O_2$  fractionation depends on **firn microstructure** determined near surface
- Seasonal stratification also vital (Fujita et al., 2009)



Adapted from Severinghaus et al., 2006

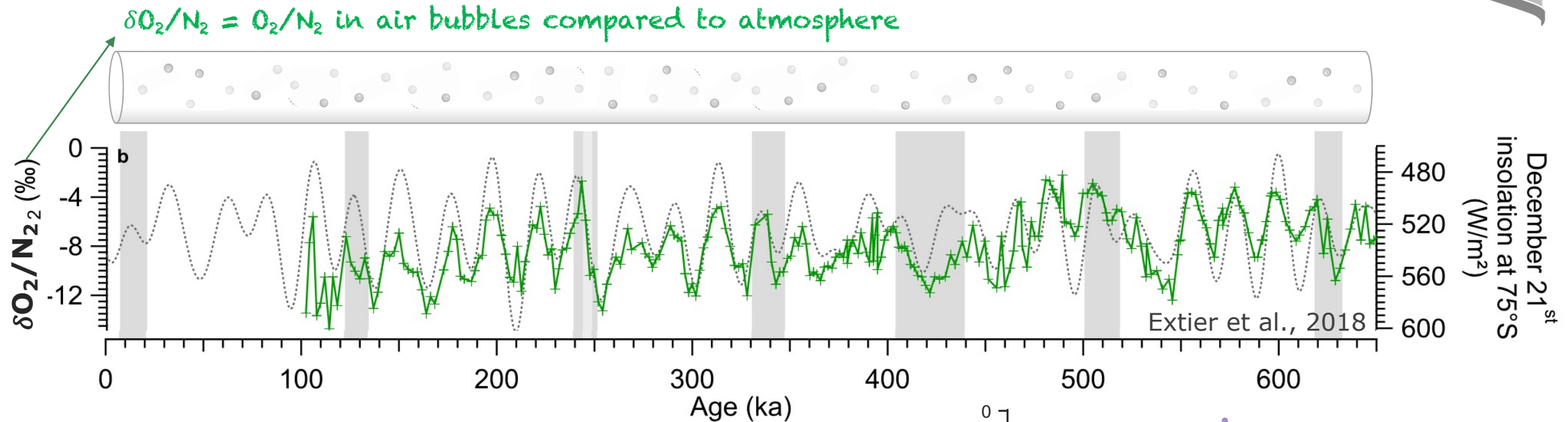
Firn Column



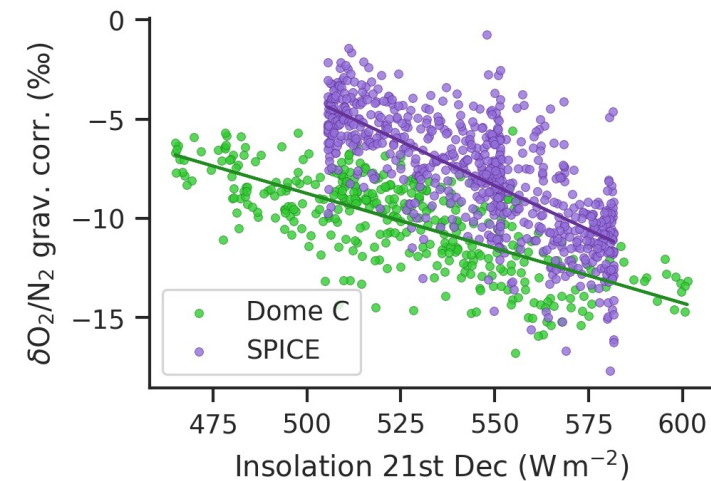
Fractionation during pore closure

Adapted from Blunier and Schwander, 2000

# Insolation paced variability in $\delta\text{O}_2/\text{N}_2$ records linked to **near-surface snow metamorphism** but exact mechanisms not quantified



- Strong fractionation of  $\text{O}_2$  with insolation
- Difference in variability between records
- Insolation cannot explain all of the variability
- Higher mean  $\delta\text{O}_2/\text{N}_2$  at South Pole suggests dependence on an additional mechanism...



# Insolation paced variability in $\delta\text{O}_2/\text{N}_2$ records linked to **near-surface snow metamorphism** but exact mechanisms not quantified

$\delta\text{O}_2/\text{N}_2 = \text{O}_2/\text{N}_2$  in air bubbles compared to atmosphere

## Insolation hypothesis:

High summer solstice insolation increases temperature gradients causing **increased grain size and seasonal layering** (Bender, 2002; Fujita et al., 2009)

## Accumulation rate hypothesis:

Increased accumulation rate will result in **decreased grain size** at a given depth given the more rapid burial of snow having less exposure to strong temperature gradients.

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- so
- Difference in variability between records

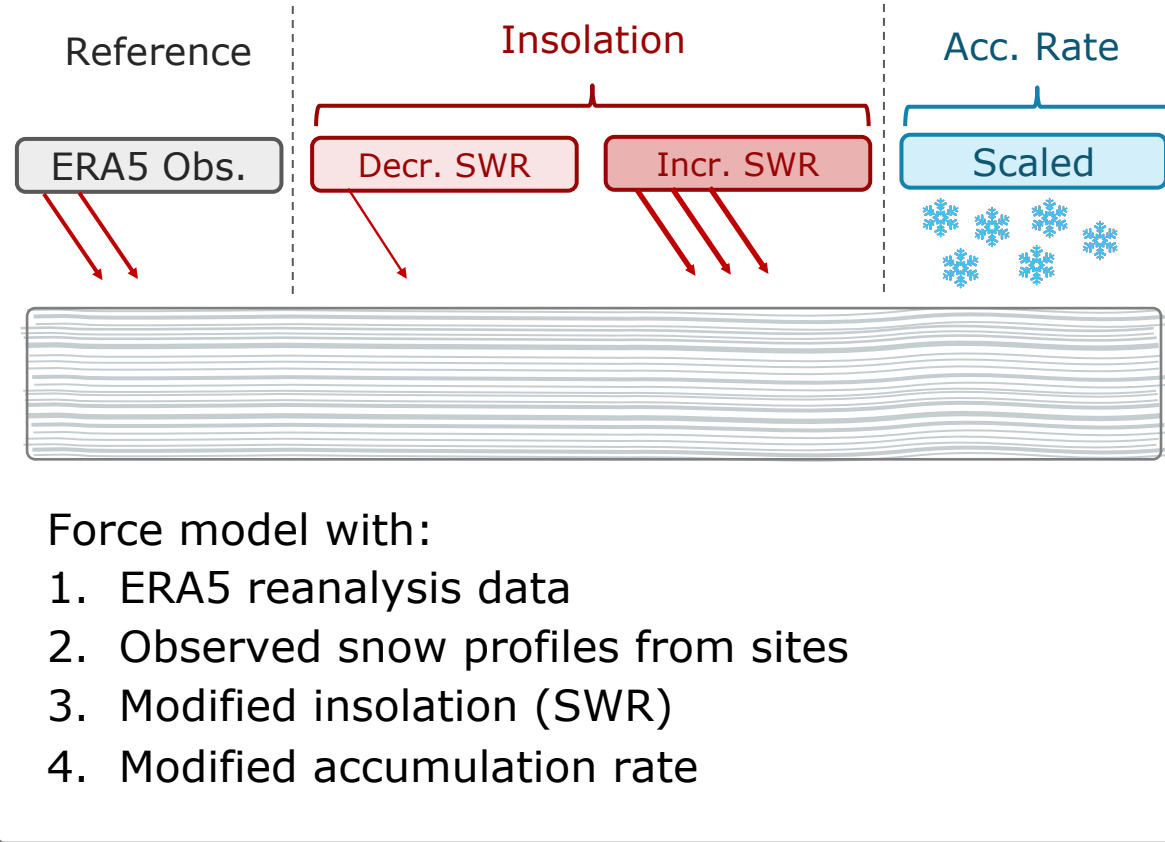
Remondino et al., 2007, 2009

December 21<sup>st</sup>  
insolation at 75°S  
(W/m<sup>2</sup>)

# Our approach: sensitivity tests using Crocus snowpack model

## INPUTS

### Modelling snowpack sensitivity to:



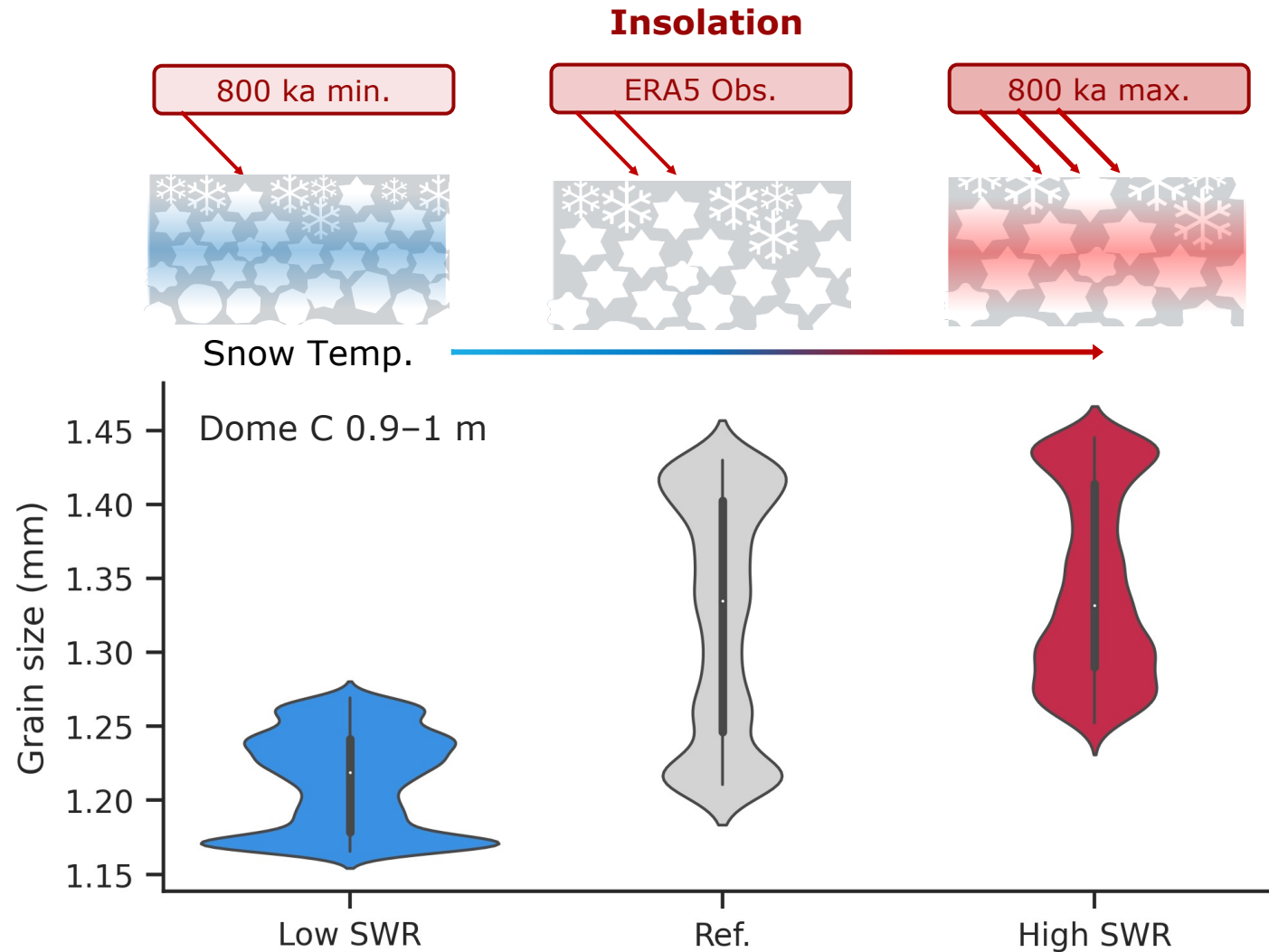
Acc. Rate =  $80 \text{ kg m}^{-1} \text{ a}^{-1}$

+ South Pole

Acc. Rate =  $25 \text{ kg m}^{-1} \text{ a}^{-1}$

+ Dome C

# Grain size increases with insolation due to increased snow temperatures and macroscale temperature gradients

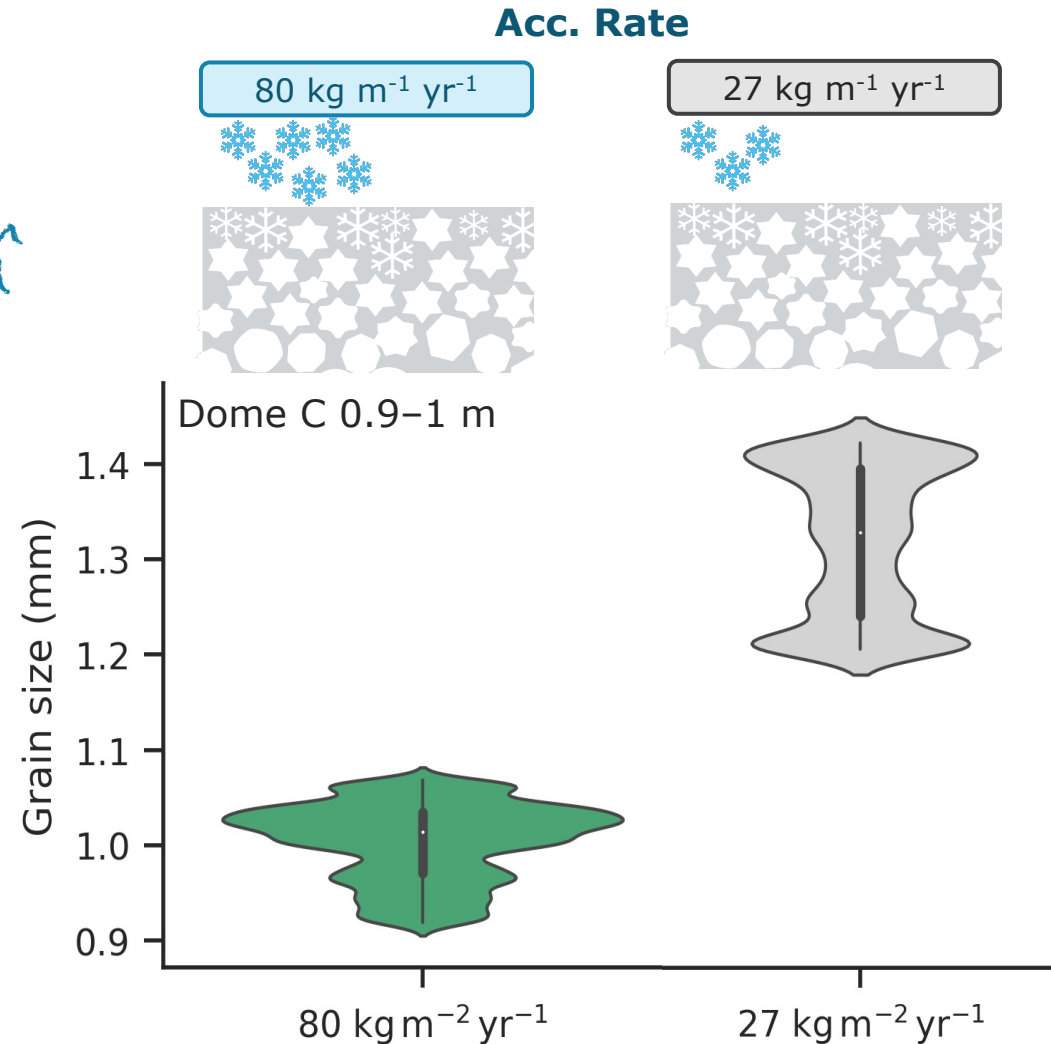


**Supports observation of increased O<sub>2</sub> fractionation during high summer solstice insolation**



# Grain size decreases with increased accumulation rate due to less time exposed to strong temperature gradients

We swap the accumulation rate between Dome C and South Pole

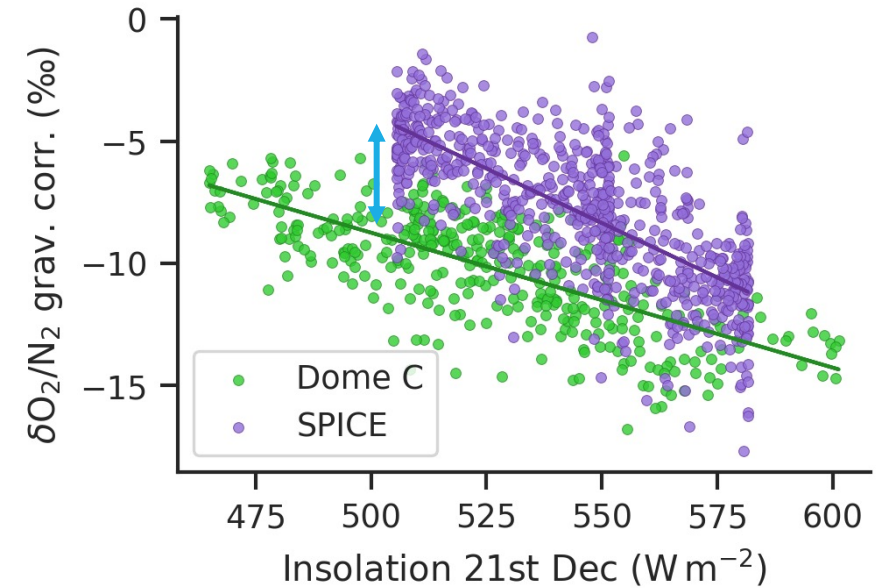


Can explain increased mean  $\delta O_2/N_2$  at high accumulation sites South Pole compared to Dome C



# Conclusions

- Grain size increases with insolation at both sites
- Supports hypothesis of increased grain size causing stronger fractionation via permeability
- Grain size is decreased with increased accumulation rate
- Potentially explains the increased mean  $\delta O_2/N_2$  at high accumulation sites South Pole compared to Dome C



## Ongoing work:

- Identify influence of stratification vs. near-surface snow microstructure
- Quantify relative influence of accumulation rate and insolation
- Implement in firn densification models