

Macro-Turbulent Energy Cascades in UpperTropospheric- Lower Stratospheric Mesoscales

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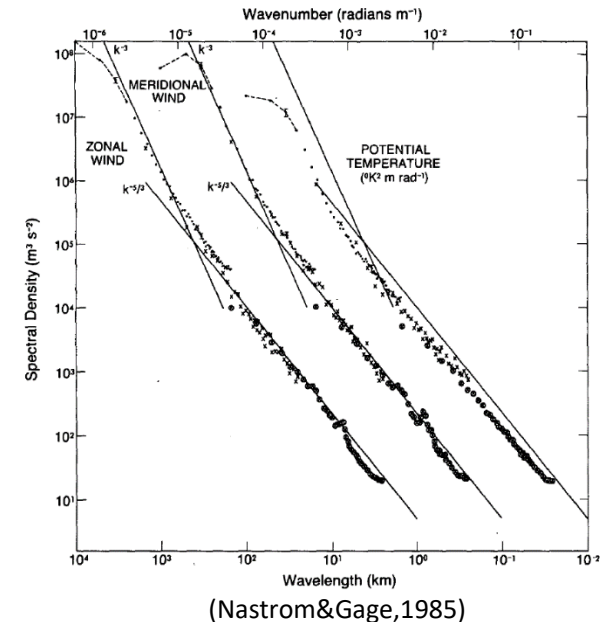


Energy transfers in
Atmosphere and Ocean



Motivation and Model Specification

Investigate the dynamics behind the simulated Nastrom & Gage Spectra, via **spectral budget analyses** of Kinetic and Available Potential Energies (**KE & APE**), **characteristic** Stratified Turbulence (**ST**) **statistics**, and Gravity Wave (**GW**) **Scale Energetics**



Kühlungsborn Mechanistic Circulation Model (KMCM):

- Spectral Dynamical Core with **T330** Triangular Truncation
- 100 Vertical Levels, extending up to $\sim 0.16 \text{ hPa}$ (beginning of Mesosphere)
- Smoothed orography with $\sim \text{T160}$ Resolution
- **Dry** Atmosphere, **no** explicit **moisture** cycle, only **large scale convective** forcing
- 16 Days of **January** Simulations
- Turbulence parameterization with a **Scale Invariant Dynamic Smagorinsky Scheme** for horizontal diffusion of **horizontal momentum** and **sensible heat**

Spectral Budget Analyses for KE & APE

Formulation of Spectral KE Budget:

$$\partial_t \vec{v} = -(\vec{v}_h \cdot \vec{\nabla}_h) \vec{v}_h + \vec{v}_h \times f \vec{e}_z - \dot{\eta} \partial_\eta \vec{v}_h - \frac{RT}{p} \vec{\nabla}_h p - \vec{\nabla} \Phi + \vec{H} + \vec{Z}$$

$$\dot{K}E_n = \frac{a_e^2}{4\pi n(n+1)} \sum_{m=-n}^n (\dot{\xi}_{mn} \xi_{mn} + \dot{D}_{mn} D_{mn})$$

$$\dot{K}E_n = \dot{K}E_n^{HA} + \dot{K}E_n^{VA} + \dot{K}E_n^{AC} + \dot{K}E_n^{MD} \quad (\text{Brune \& Becker 2013})$$

Formulation of Spectral APE Budget:

$$\partial_t(c_p T) = -\vec{v}_h \cdot \vec{\nabla}_h(c_p T) - \dot{\eta} \partial_\eta(c_p T) + \frac{RT}{p} \omega + Q \quad T = \bar{T} + \tilde{T}$$

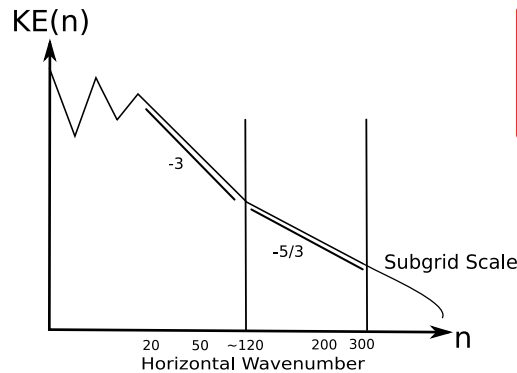
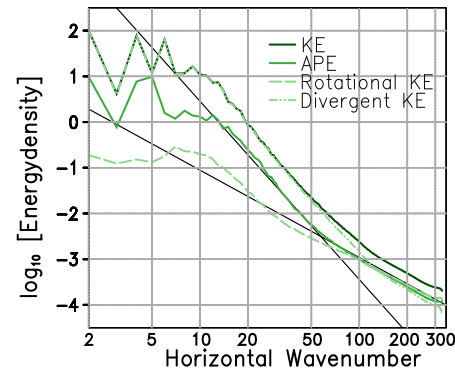
$$\dot{\tilde{T}} = \partial_t \tilde{T} = (-\vec{v}_h \cdot \vec{\nabla}_h \tilde{T} - \omega \partial_p \tilde{T} - \omega \partial_p \bar{T} + \frac{R}{c_p p} \tilde{T} \omega + \frac{R}{c_p p} \bar{T} \omega + \frac{1}{c_p} Q)$$

$$\text{Local APE} \approx \frac{g^2}{2\bar{T}^2 \bar{N}^2} \tilde{T} \cdot \tilde{T} = \frac{1}{2} C \tilde{T}^2 \quad (\text{Lorenz, 1955})$$

$$\dot{A}PE_n = \frac{1}{4\pi} C \sum_{m=-n}^n (\dot{\tilde{T}}_{mn} \tilde{T}_{mn})$$

$$\dot{A}PE_n = \dot{A}PE_n^{HA} + \dot{A}PE_n^{VA} + \dot{A}PE_n^{NBP} + \dot{A}PE_n^{DH}$$

Integral Scale Velocity and Length Scales



$$U_{h(n)} = \sqrt{2KE(n)} = \sqrt{2 \int_n^\infty e_{KE(n')} dn'}$$

$$KE_{(n>120)} = KE_{meso} + KE_{sgs}$$

(Becker, 2009)

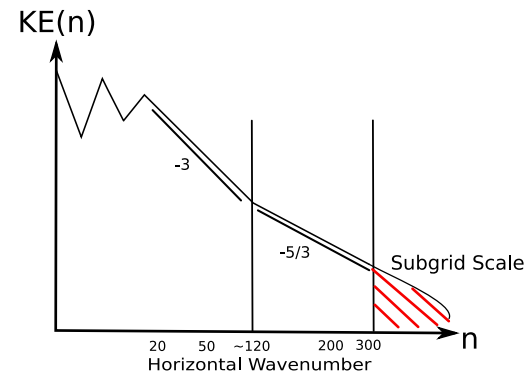
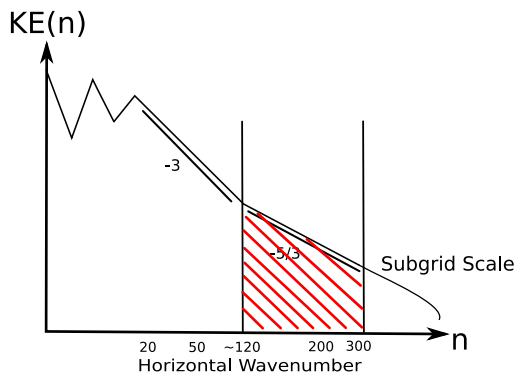
(Avsarkisov, Becker & Renkowitz, 2022)

$$KE_{meso} = \frac{\vec{v}_{meso}^2}{2}$$

$$\vec{v}_{meso} = \vec{v}_{full} - \vec{v}_{(n<120)} = \vec{v}_{((n>120))}$$

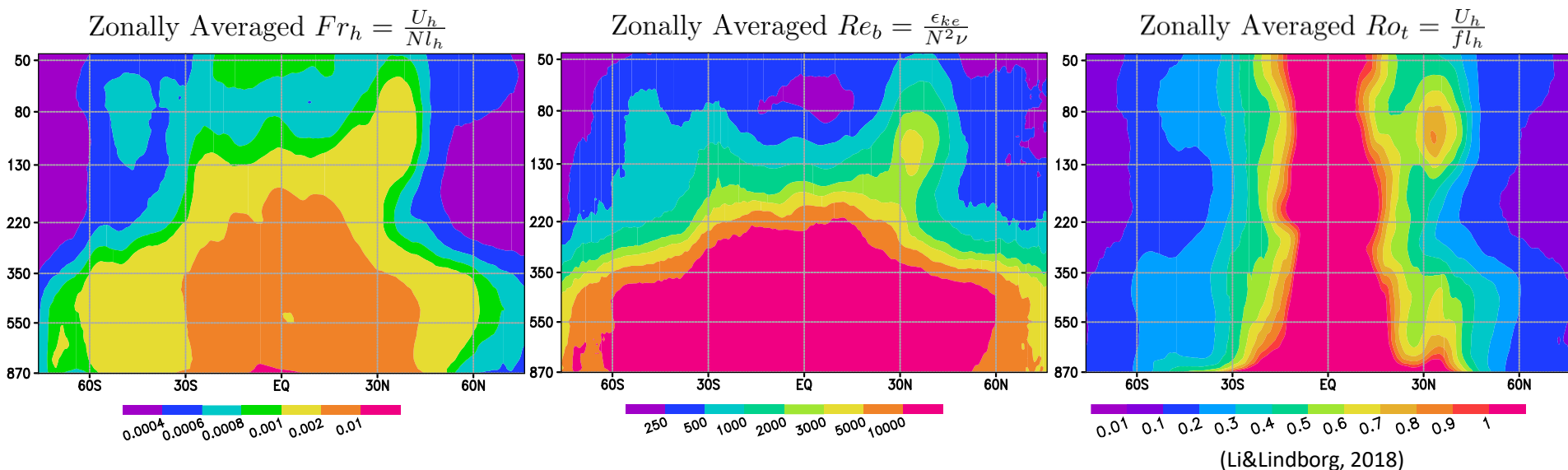
$$KE_{sgs} = \sqrt{\frac{\epsilon_{ke} K^{tot}}{C}}$$

(Launder et al., 1973)



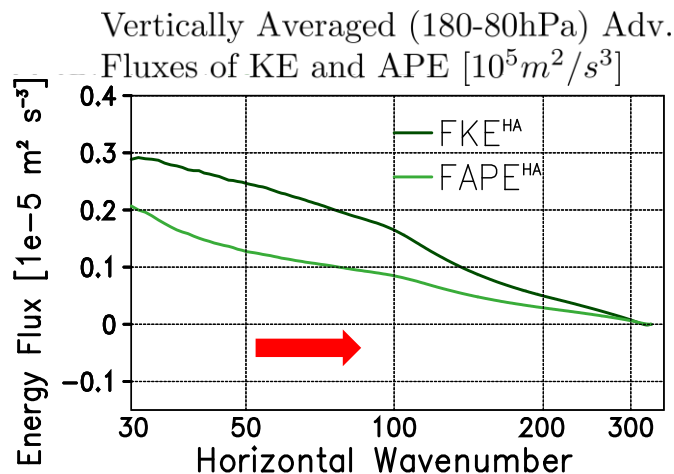
$$l_h = \frac{U_h^3}{\epsilon_{ke}} \quad (\text{Taylor, 1935})$$

Characteristic ST Statistics



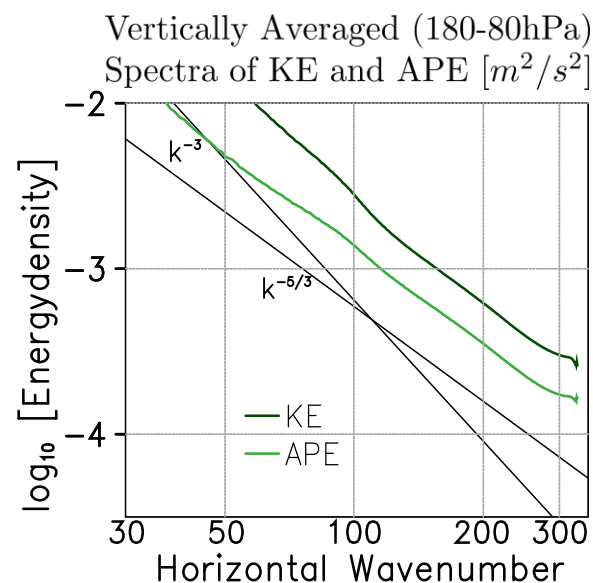
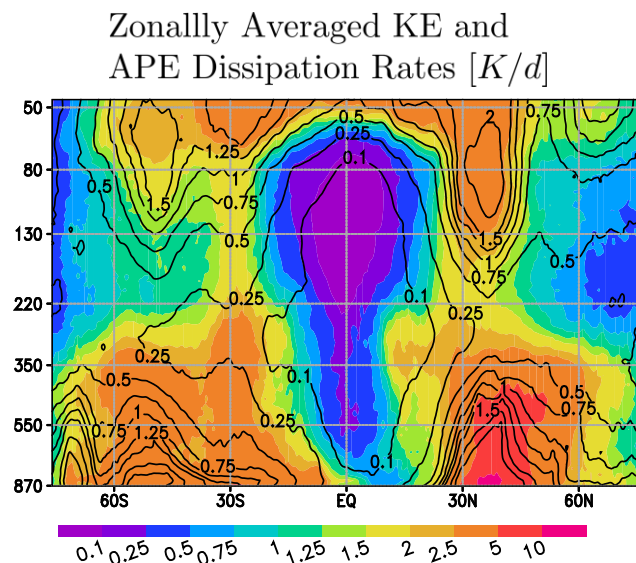
- Threshold for $Fr_h < 0.02$ (Lindborg, 2006)
- Re_b in comparison to Fr_h reasonable
 - $Re_b \sim 10^4$ at Japan on $\sim 4.5\text{km}$ ($\sim 550\text{ hPa}$) altitude & $\sim 35\text{N}$ latitude is reported from UAV measurements (Luce et al., 2019)
- Identified regime:
 - Altitude: **180<p<80hPa** - Upper Troposphere – Lower Stratosphere (UTLS)
 - Latitude: **Sub-tropical Midlatitudes**

ST Fluxes and Scaling Laws



- Forward flux $\rightarrow$ dissipation at smallest scales
- Scaling laws of ST for KE and APE (Riley&Lindborg, 2013):

$$e_{KE(n)} = C_1 \epsilon_{ke}^{2/3} n^{-5/3} \quad e_{APE(n)} = C_2 \frac{\epsilon_{ape}}{\epsilon_{ke}^{1/3}} n^{-5/3}$$



Conversion of Energy

- Adiabatic conversion couples KE and APE. **In global mean**, the conversion is represented by the **same term** in both energy reservoirs.

$$\langle \dot{K}E \rangle = C - D^{KE}$$

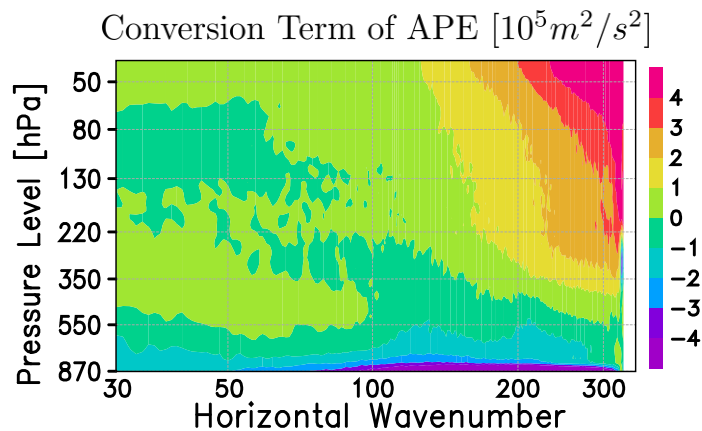
$$\langle \dot{A}PE \rangle = G - C + D^{APE}$$

- In horizontal mean**, where vertical coordinate is retained, **conversion from**

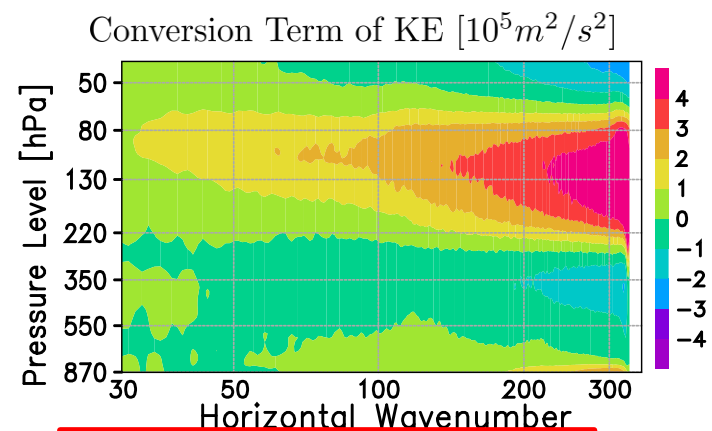
KE includes vertical geopotential fluxes as well! (Brune&Becker,2013 & Augier&Lindborg2013)

$$[\dot{K}E]_{(p)} = C_{(p)}^{KE} - D_{(p)}^{KE}$$

$$[\dot{A}PE]_{(p)} = G_{(p)} - C_{(p)}^{APE} + D_{(p)}^{APE}$$



$$C_{(p)}^{APE} = -\left[\frac{\omega}{\rho}\right]$$



$$C_{(p)}^{KE} = \left[\frac{\omega}{\rho}\right] - \partial_p [gz\omega]$$

GW Scale Energetics

Define KE and APE budget for GW scale perturbations:

(Becker, 2017)

- Anelastic
- For mid-frequency GWs
- No GW Scale (mesoscale) forcing
- Single-Column Approximation

$$\dot{K}E^{GW} = -(\overline{\vec{v}'w'}) \cdot \partial_z \vec{v} + \frac{g}{T_r} \overline{T'w'} - \frac{1}{\rho_r} \partial_z \overline{p'w'} - (K_z) \overline{(\partial_z \vec{v}')^2} - K_h \overline{|\mathbf{S}'_h|^2}$$

$$\dot{A}PE^{GW} = -\frac{g}{T_r} \overline{T'w'} - \frac{g^2}{N_r^2 T_r^2} \left[\frac{K_z}{P_r} \overline{(\partial_z T')^2} + \frac{K_h}{P_r} \overline{(\nabla T')^2} \right]$$

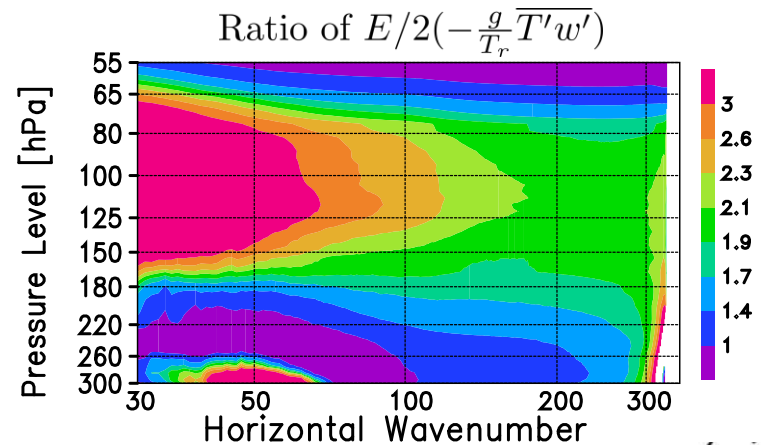
Define accordingly the Energy Deposition of nonconservative GWs:

(Hines&Reddy, 1967 & Lindzen, 1973)

$$E = -\frac{1}{\rho_r} \partial_z \overline{p'w'} - (\overline{\vec{v}'w'}) \cdot \partial_z \vec{v} = -\frac{g}{T_r} \overline{T'w'} + (K_z) \overline{(\partial_z \vec{v}')^2} + K_h \overline{|\mathbf{S}'_h|^2}$$

Impose ST argument for similar amounts of KE and APE dissipation rates:

$$E = 2\left(-\frac{g}{T_r} \overline{T'w'}\right)$$

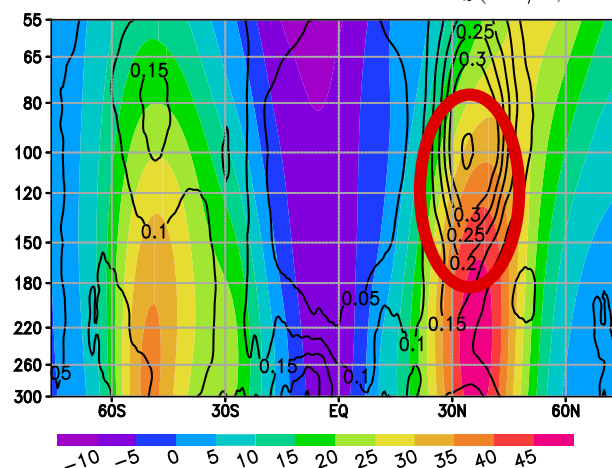


Resolved Saturated Westward GWs

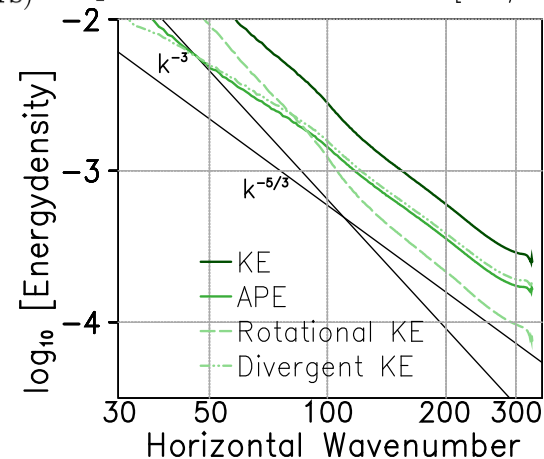
Further predictions from the theory of Superposition of **Saturated GWs** hold as well

(Lindzen, 1981 & Becker, 2004):

Zonally Averaged Zonal Wind u (m/s, colors)
and Vertical Diff. Coeff. of Mom. K_z (m^2/s , contours)

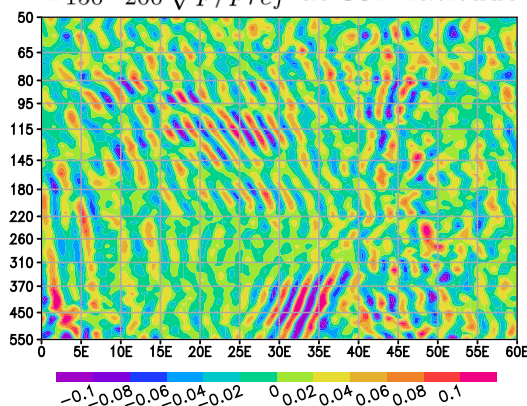


Vertically Averaged (180-80hPa)
Spectra of KE and APE [m^2/s^2]

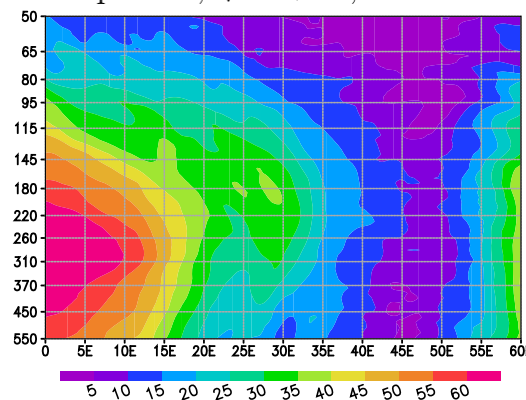


Instantaneous snapshots of Temperature perturbations and wind amplitude further supports the **existence of Westward GWs** (Viudez, 2007 & Plougonven & Zhang, 2014 & Yasuda et al., 2015):

Temperature perturbation, i.e.,
 $T_{150-200} \sqrt{p/p_{ref}}$ at 30N latitude



Wind Amplitude, $\sqrt{u^2 + v^2}$, at 30N latitude



Conclusions & References

1. Mesoscale shallowing at UTLS altitudes are **primarily governed by resolved westward mid-frequency GWs**, for which each GW is at the **saturation level due to the direction** of their propagation.
2. This saturation is the cause of energy deposited at UTLS altitudes, **which can be explicitly formulated** via the anelastic, single column approximated GW scale energetics.
3. **Resolved saturated WGWs induce the forward cascades** of both APE and KE, which are **governed by strongly nonlinear ST dynamics**.

Thank you for your attention! Any questions?

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