

3D Computational Modeling of Lithospheric Deformation, Asthenospheric Flow, and Deep Melt Generation with ASPECT

D. Sarah Stamps
Associate Professor
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Contributions from:

Tahiry Rajaonarison, New Mexico Tech, USA

Emmanuel Njinju, Virginia Tech, USA

John Naliboff, New Mexico Tech, USA

Asenath Kwagalakwe, Virginia Tech, USA



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*Geodynamic Processes
at Rifting and
Subducting
Margins*



EAR-1740704

EarthCube
BALTO
Brokered Alignment of Long-tail Observations



Overview

Introduction

Lithospheric Deformation

Asthenospheric Flow

Deep Melt Generation

Summary & Conclusions

Contributions from:

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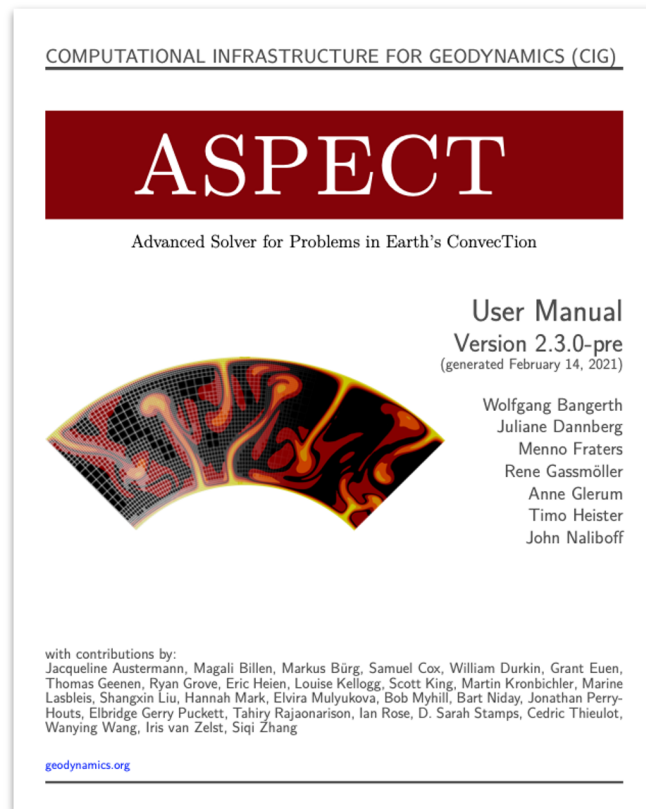


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What is ASPECT?

- Advanced Solver for Problems in Earth's ConvecTion
- NSF Computational Infrastructure for Geodynamics finite element code (EAR-1550901)
- A community code that builds on the deal.II, Trilinos, and p4est libraries
- Applications in mantle convection on Earth and other planets, lithospheric deformation, melt generation



What is ASPECT?

ASPECT solves the Stokes equations for velocity and pressure

$$-\nabla \cdot \left[2\eta \left(\varepsilon(\mathbf{u}) - \frac{1}{3}(\nabla \cdot \mathbf{u})\mathbf{1} \right) \right] + \nabla p = \rho \mathbf{g} \quad \text{Conservation of momentum}$$
$$\nabla \cdot (\rho \mathbf{u}) = 0 \quad \text{Conservation of mass}$$

η = viscosity

ρ = density

$\varepsilon(\mathbf{u})$ = strain rate

\mathbf{g} = gravitational acceleration

p = pressure

What is ASPECT?

Coupled is the **energy equation** for solving for temperature

$$\begin{aligned}
 &\left(\rho C_p - \rho T \Delta S \frac{\partial X}{\partial T} \right) \left(\frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T \right) - \nabla \cdot k \nabla T = \rho H \quad \text{Internal heat production} \\
 &\quad \text{Viscous shear heating} + 2\eta \left(\varepsilon(\mathbf{u}) - \frac{1}{3}(\nabla \cdot \mathbf{u})\mathbf{1} \right) : \left(\varepsilon(\mathbf{u}) - \frac{1}{3}(\nabla \cdot \mathbf{u})\mathbf{1} \right) \\
 &\quad \text{Adiabatic compression of material} + \alpha T (\mathbf{u} \cdot \nabla p) \\
 &\quad \text{Phase change (latent heat)} + \rho T \Delta S \frac{\partial X}{\partial p} \mathbf{u} \cdot \nabla p
 \end{aligned}$$

C_p = specific heat capacity

\mathbf{u} = velocity

T = temperature

k = thermal conductivity

S = specific entropy

α = thermal expansion coefficient

Where can I find ASPECT?

- Current release version

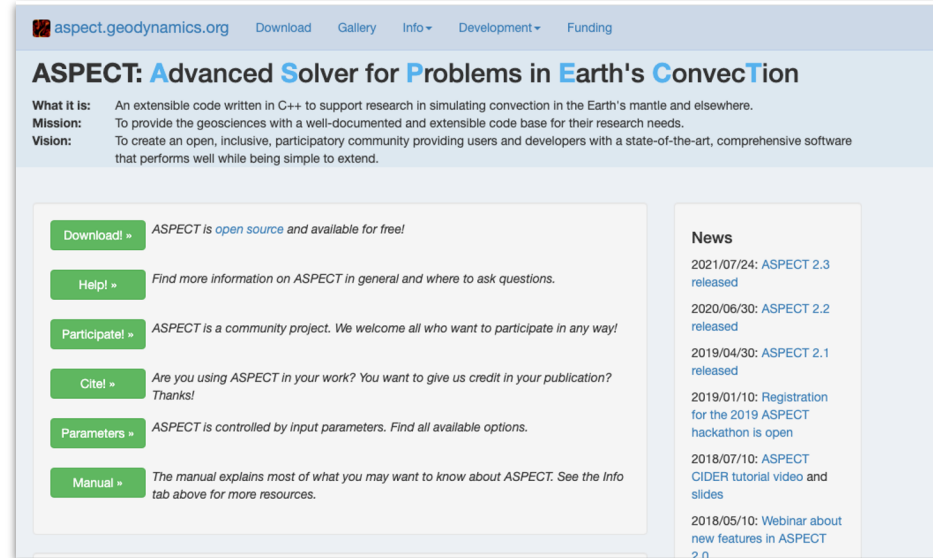
<https://aspect.geodynamics.org/>

- Current development version:

<https://github.com/geodynamics/aspect.git>

- Extensions to ASPECT from specific papers

Zenodo and/or GitHub usually



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ASPECT can be used for modeling lithospheric deformation

Geophysical Research Letters

RESEARCH LETTER

10.1029/2020GL090483

Key Points:

- 3D geodynamic models reveal that lithospheric buoyancy forces primarily drive ~E-W extension across East Africa
- Lithospheric buoyancy forces generate rigid plate motions aligned with kinematic predictions derived from GPS data
- Additional forces arising from horizontal mantle tractions may be required to explain along-rift surface motions in deforming zones

Role of Lithospheric Buoyancy Forces in Driving Deformation in East Africa From 3D Geodynamic Modeling

Tahiry A. Rajaonarison¹, D. Sarah Stamps¹, and John Naliboff^{2,3}

¹Department of Geosciences, Virginia Tech, Blacksburg, VA, USA, ²Department of Earth and Planetary Sciences, New Mexico Tech, Socorro, NM, USA, ³Department of Earth and Planetary Sciences, CA, USA

Abstract Despite decades of investigation, the origin of for highly debated. Deciphering their relative contributions is challenging due to the dependent nature of lithospheric rheology. Recent geodynamic



Rajaonarison, T. A., Stamps, D. S., & Naliboff, J. (2020b). *Trajaona/aspect: Rifting model of the East African Rift (version v1.0)*. Zenodo. <https://doi.org/10.5281/zenodo.4005094>

Rajaonarison, T. A., Stamps, D. S., & Naliboff, J. (2021). Role of lithospheric buoyancy forces in driving deformation in East Africa from 3D geodynamic modeling. *Geophysical Research Letters*, 48, e2020GL090483. <https://doi.org/10.1029/2020GL090483>



EAR-1551864

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August 28, 2020

Software Open Access

trajaona/aspect: Rifting Model of the East African Rift

Tahiry Rajaonarison; D. Sarah Stamps; John Naliboff

This software package is a modified version of ASPECT v2.2.0, which allows to reproduce model results of Rajaonarison et al. (submitted to GRL). The model setup is as follows:

- 1) In this model, the lithospheric deformation is driven by lithospheric buoyancy forces which are imposed by implementing the CRUST1.0 (density and thicknesses for lower, middle, and upper crust) and ETOPO1 data sets into the 3D model.
- 2) The lithospheric temperature structure is a steady-state conductive geothermal gradient characteristic of the continental lithosphere, which is constrained using estimates of regional lithospheric thickness and surface heat flow of the key tectonic regions (cratons, mobile belts, and rifts).
- 3) Isostatic compensation is enforced through adjustments to the mantle density (i.e. Pratt isostasy) down to 100 km, with a constant value of 3300 kg/m³ assigned from the compensation depth to 660 km.
- 4) The rheological model combines non-linear viscous flow; dislocation creep of dry quartzite with plastic failure in the crust, dislocation creep of olivine in the mantle lithosphere, and composite rheology of dry olivine in the sublithospheric mantle. In the deforming regions, a plastic strain weakening factor for cohesion and friction is applied to the lithospheric viscosity to promote strain localization.

The model can be run using the parameter file (EAR_rifting.prm) that is attached with this description. We refer the reader to the ASPECT user manual for information about how to run model in ASPECT.

This work was funded by the National Science Foundation (NSF) GeoPRISMS grant EAR-1551864.

98 views

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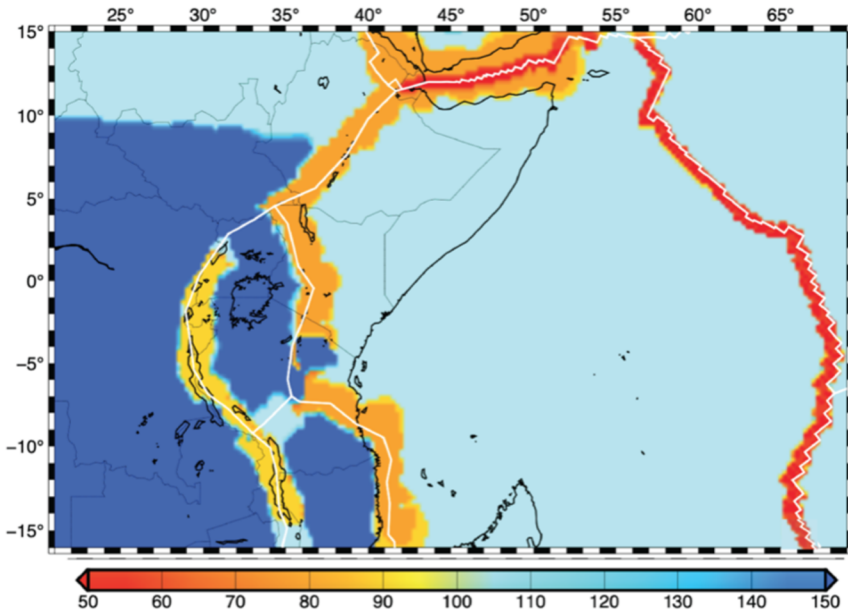
GitHub

OpenAIRE

Publication date: August 28, 2020

DOI: 10.5281/zenodo.4005094

Model Set-up: Lithospheric and Crustal Thickness



Synthetic Lithospheric Thickness (km)

Averaged from 3 seismically constrained models

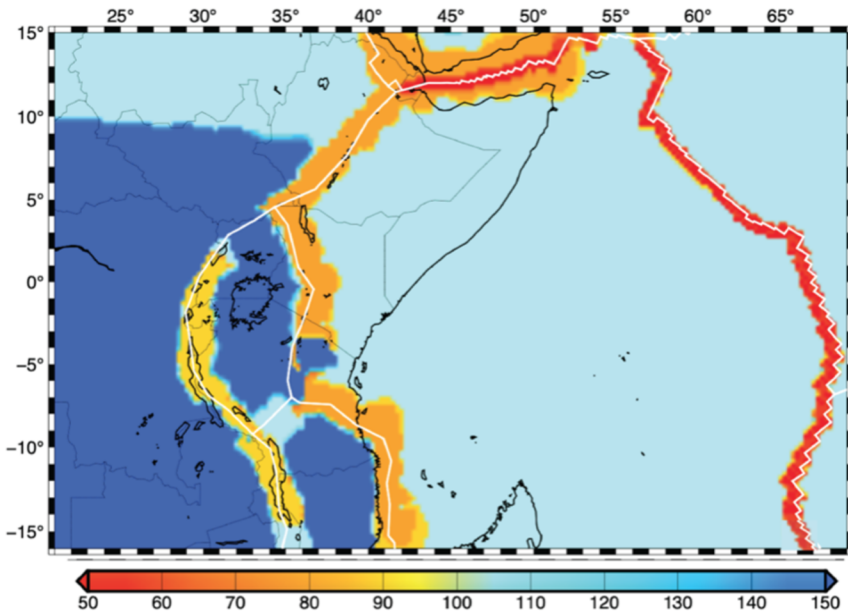
Continental rift areas = 90 km, 70 km

Cratons = 150 km

Oceanic Ridges = 50 km

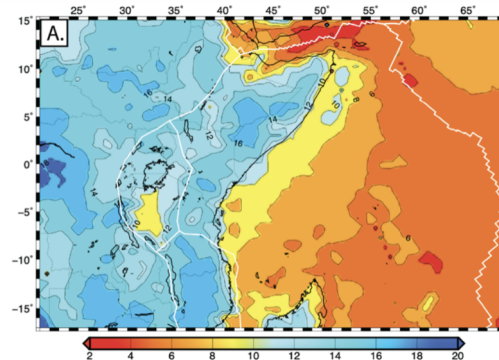
Mobile Belts+ = 100 km

Model Set-up: Lithospheric and Crustal Thickness

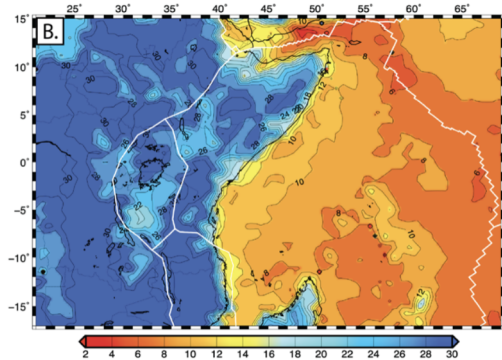


Synthetic Lithospheric Thickness (km)
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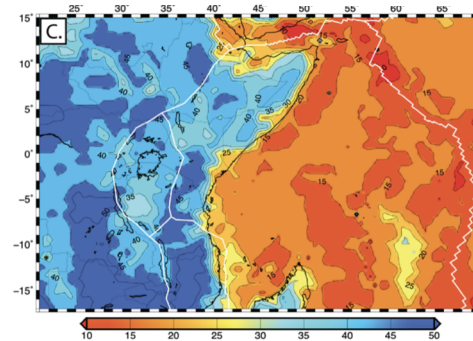
CRUST1.0 (Laske et al., 2012)



Base of Upper Crust (km)



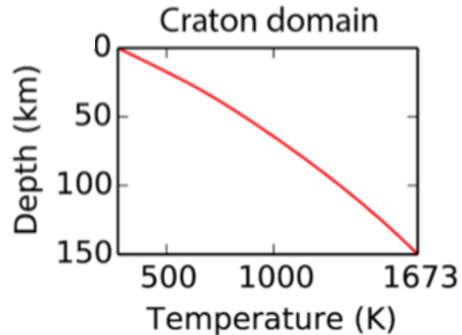
Base of Middle Crust (km)



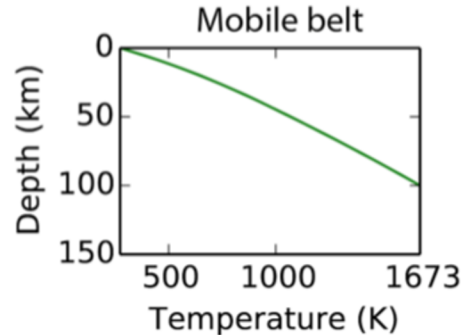
Base of Lower Crust (km)

Model Set-up: Temperature

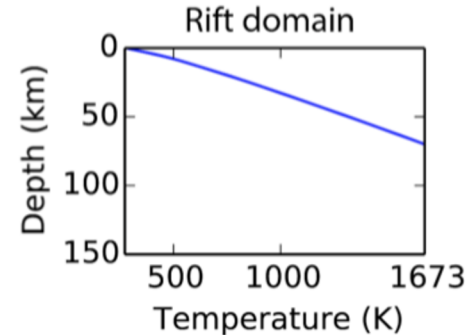
Surface heat flow: 33 mW/m²



64 mW/m²

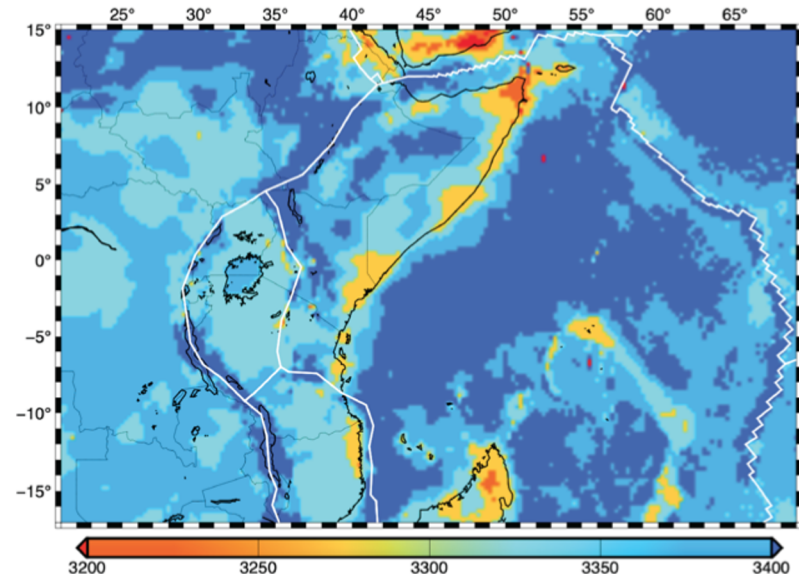


80 mW/m²



- A steady-state conductive geothermal gradient (Chapman, 1986).
- Constrained by lithospheric thickness and surface heat flow of the key tectonic regions.
- Below the lithosphere, the temperature increases approximately adiabatically.

Model Set-up: Density

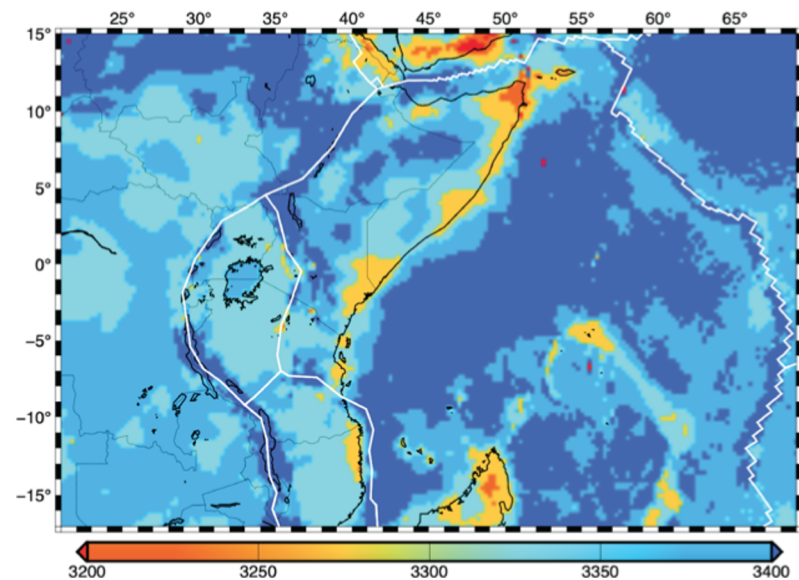


Vertically Averaged Mantle Lithosphere Density (kg/m³)

- 100 km compensation depth
- Laterally varying density in lithospheric mantle for isostatic compensation

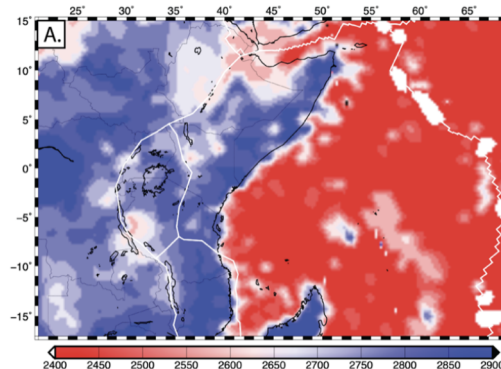
Model Set-up: Density

CRUST1.0 (Laske et al., 2012)

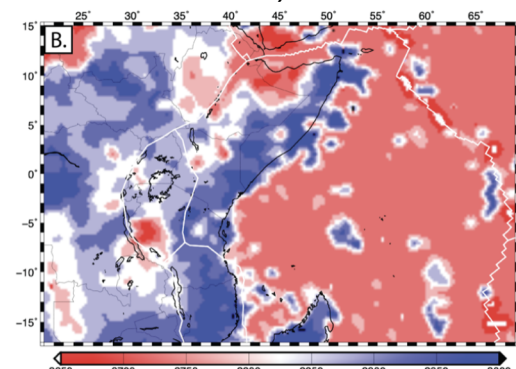


Vertically Averaged Mantle Lithosphere Density (kg/m³)

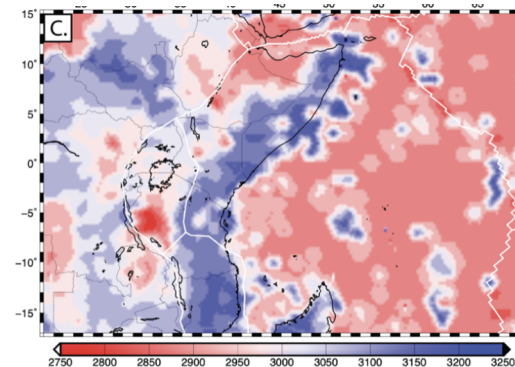
- 100 km compensation depth
- Laterally varying density in lithospheric mantle for isostatic compensation



Density of Upper Crust (km)



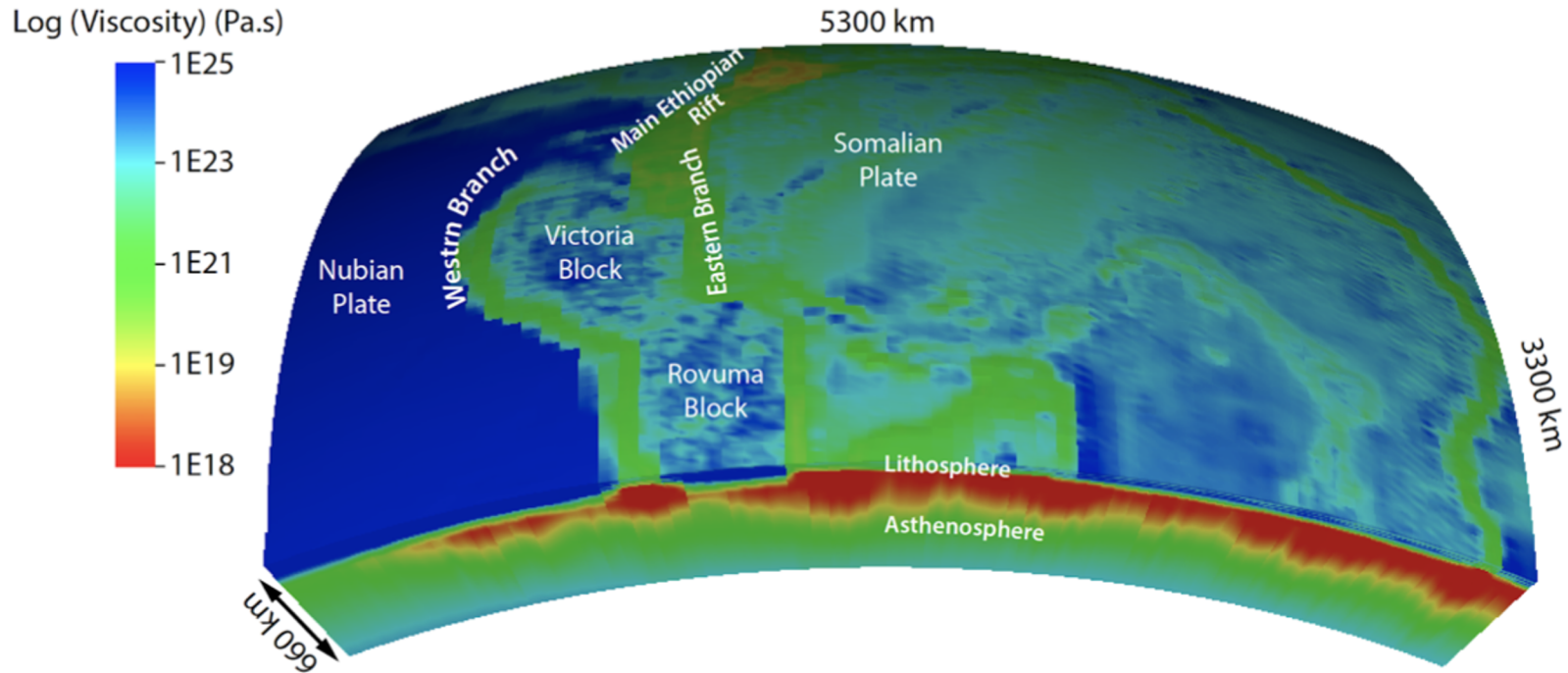
Density of Middle Crust (km)



Density of Lower Crust (km)

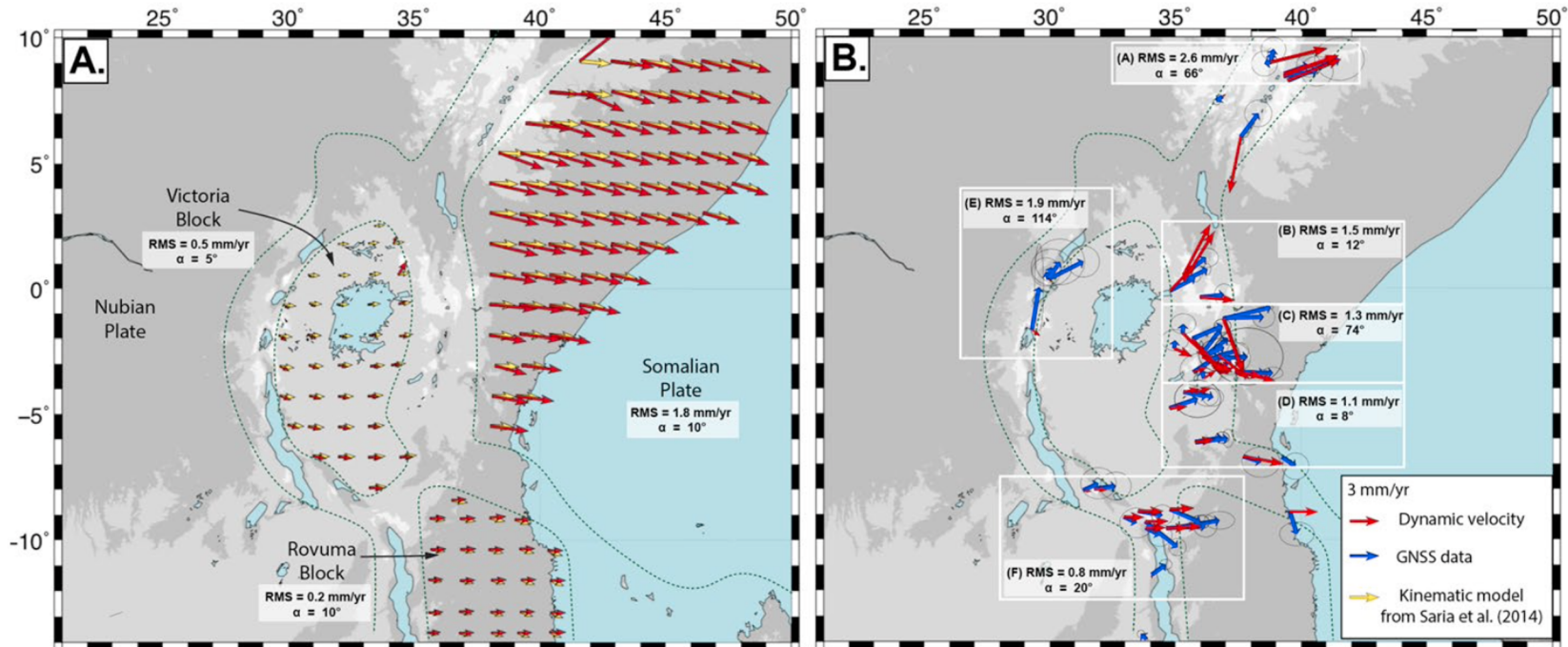
Model Set-up: Viscosity

- Crust: Combines nonlinear quartzite dislocation creep with plastic failure (Gleason et al., 1995)
- Mantle lithosphere: Combines olivine dislocation creep (Hirth & Kohlstedt, 2003) with plastic failure
- Sublithospheric mantle: Composite dry olivine diffusion and dislocation creep (Hirth & Kohlstedt, 2003)



Case Study Results: East African Rift

- Lithospheric buoyancy forces primarily drive ~E-W extension across East Africa
- Lithospheric buoyancy forces generate rigid plate motions aligned with kinematic predictions derived from GPS data
- Additional forces arising from horizontal mantle tractions may be required to explain along-rift surface motions in deforming zones



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ASPECT can be used for modeling asthenospheric flow

JGR Solid Earth



RESEARCH ARTICLE

10.1029/2019JB018560

Key Points:

- Asthenospheric flow patterns beneath Madagascar are predicted from edge-driven convection and mantle wind modeling
- Comparison of predicted shear wave splitting parameters with seismic anisotropy suggests a predominantly asthenospheric source
- Dislocation creep rheology extends into the asthenosphere beneath some continental regions

Numerical Modeling of Mantle Flow Beneath Madagascar to Constrain Upper Mantle Rheology Beneath Continental Regions

T. A. Rajaonarison¹, D. S. Stamps¹, S. Fishwick², S. Brune^{3,4}, A. Glerum³, and J. Hu⁵

¹Department of Geosciences, Virginia Polytechnic Institute and State University, Blacksburg, VA, USA, ²Department of Geology, University of Leicester, Leicester, UK, ³GFZ German Research Center of Geosciences, Potsdam, Germany,

⁴Institute of Earth and Environmental Sciences, University of Potsdam, Potsdam, Germany, ⁵Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA, USA

Abstract Over the past few decades, azimuthal seismic anisotropy measurements have been widely

Rajaonarison, T. A., Stamps, D. S., Fishwick, S., Brune, S., Glerum, A., & Hu, J. (2020). Numerical modeling of mantle flow beneath Madagascar to constrain upper mantle rheology beneath continental regions. *Journal of Geophysical Research: Solid Earth*, 125, e2019JB018560. <https://doi.org/10.1029/2019JB018560>

Built into the ASPECT release code

A.74 Parameters in section Initial temperature model/Adiabatic boundary

- *Parameter name:* Adiabatic temperature gradient

Value: 0.0005

Default: 0.0005

Description: The value of the adiabatic temperature gradient. Units: K m^{-1} .

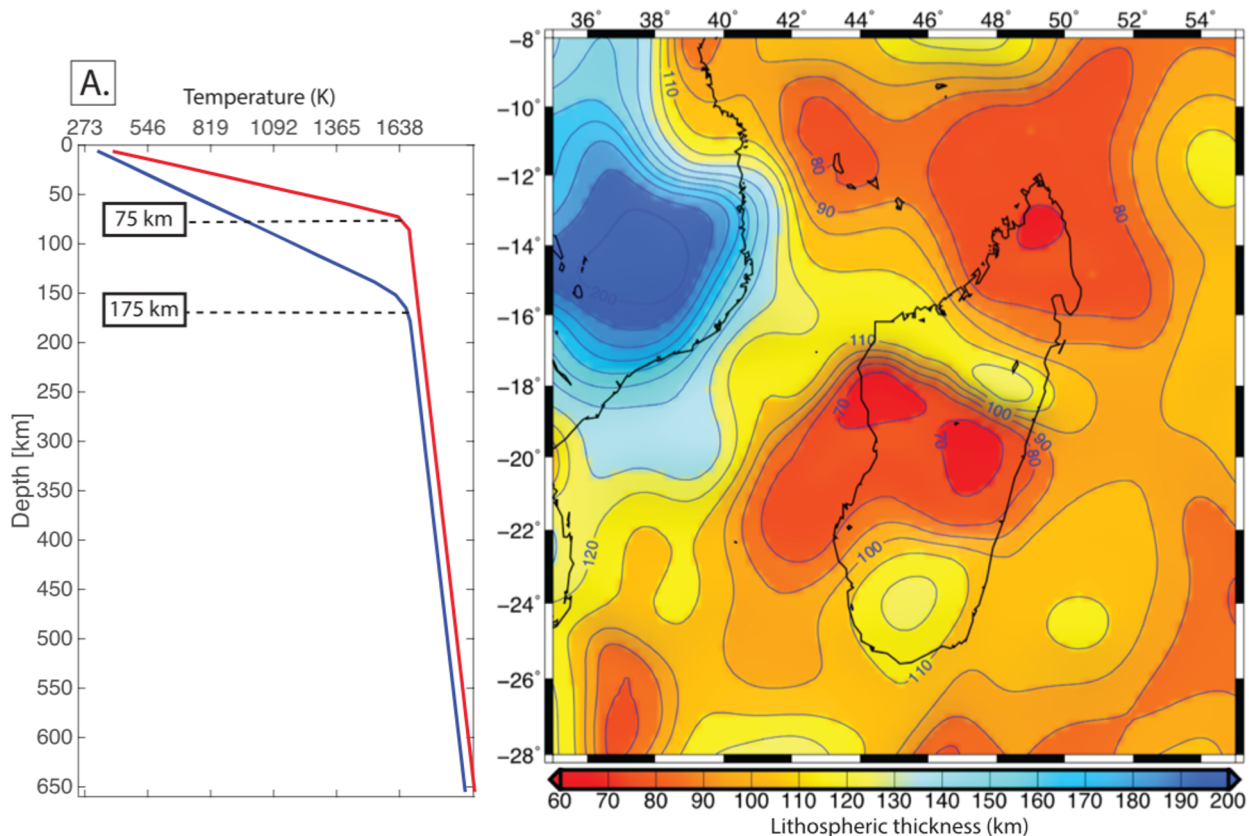
Possible values: A floating point number v such that $0 \leq v \leq \text{MAX_DOUBLE}$



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Geodynamic Processes
at Rifting and
Subducting
Margins

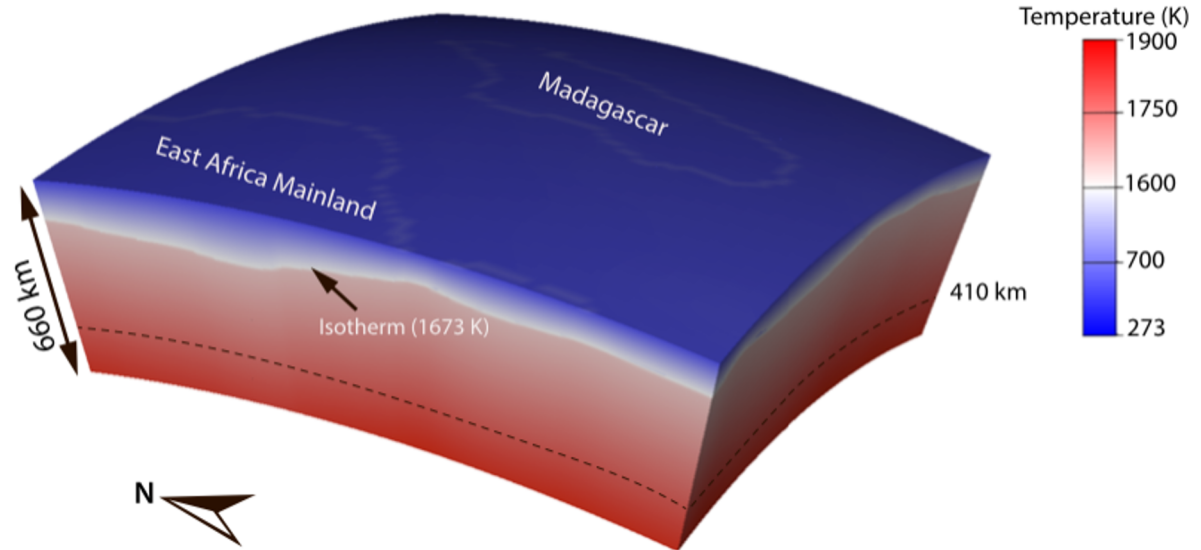
Model Setup: Initial Temperature Conditions

- Lithospheric thickness input with isotherm at lithosphere-asthenosphere boundary (LAB, i.e. Fishwick 2010, updated)
- Approximate conductive geotherm in lithosphere: linear gradient from surface to LAB isotherm (1673 K)
- Approximately adiabatic geotherm below lithosphere (i.e. 0.5 K/km)



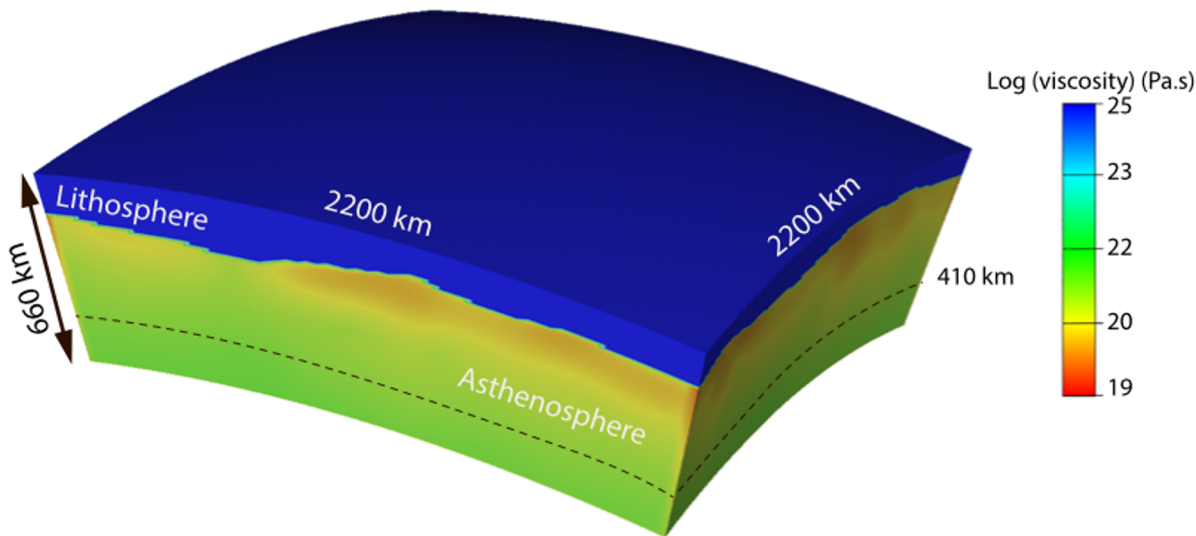
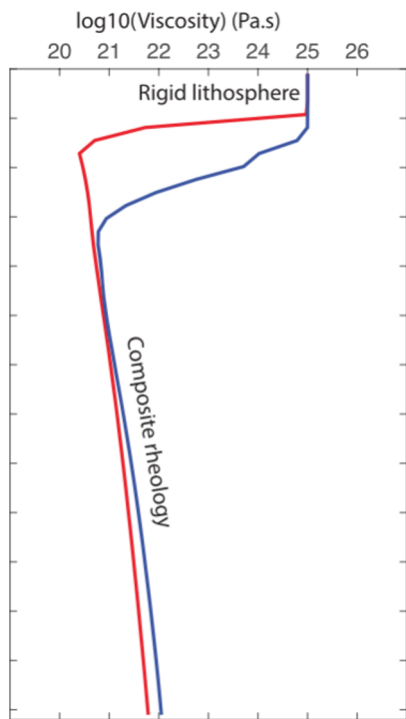
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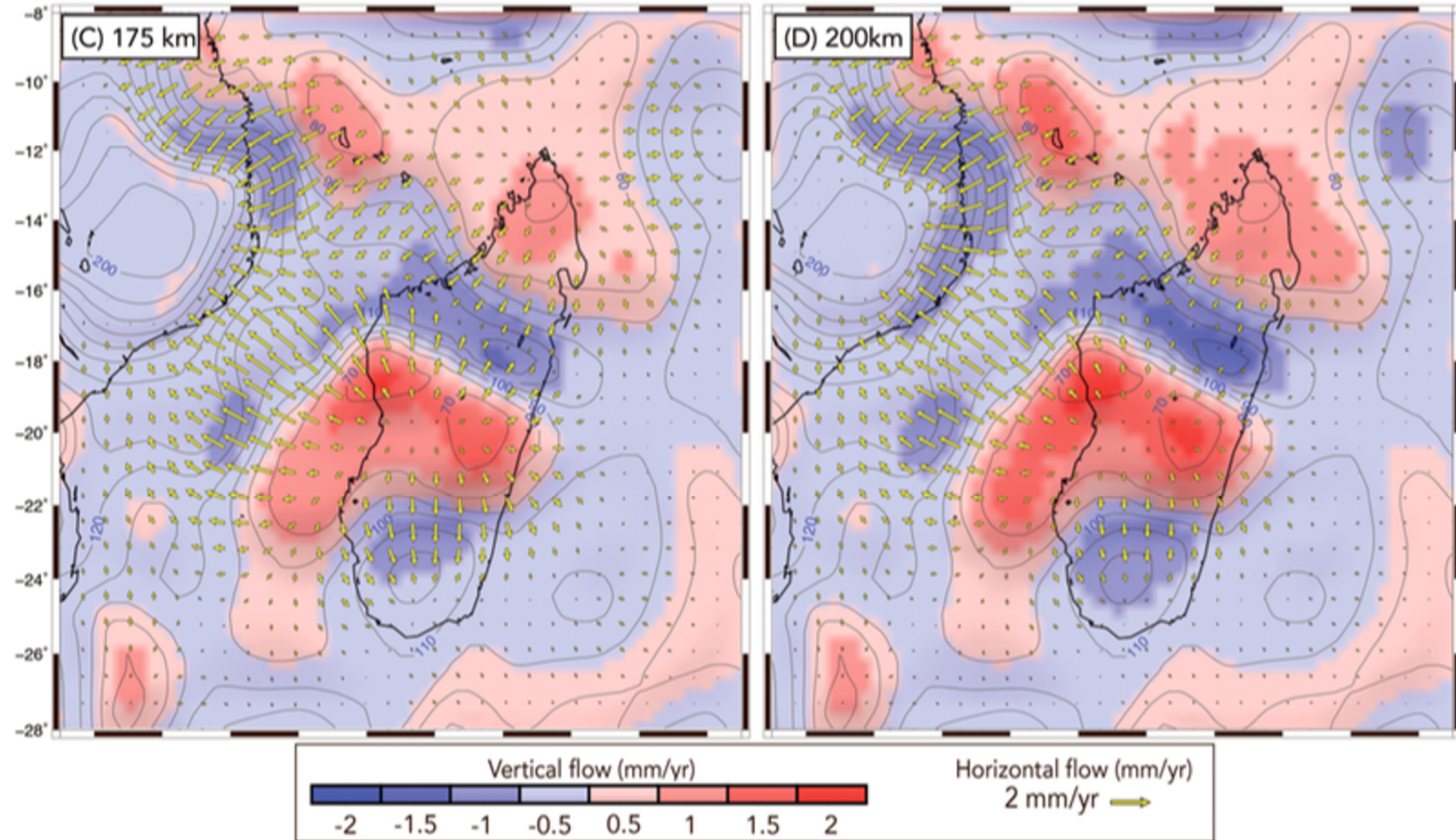


Model Setup: Viscosity and Density

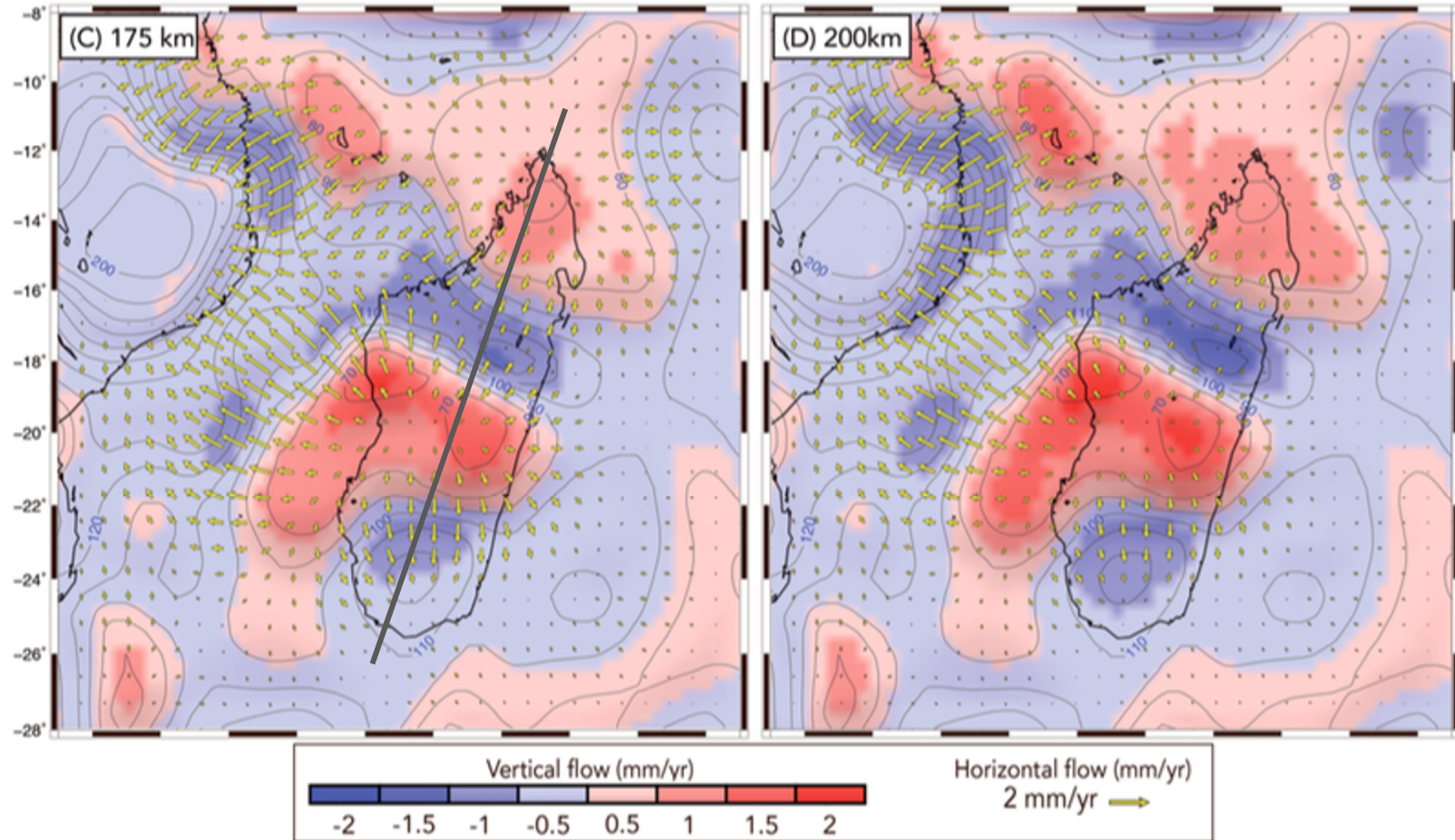
- High viscosity lithosphere ($1\text{E}25\text{ Pa.s}$)
- Composite rheology for sublithospheric mantle (Jadamec & Billen, 2010)
- Temperature dependent density below the lithosphere



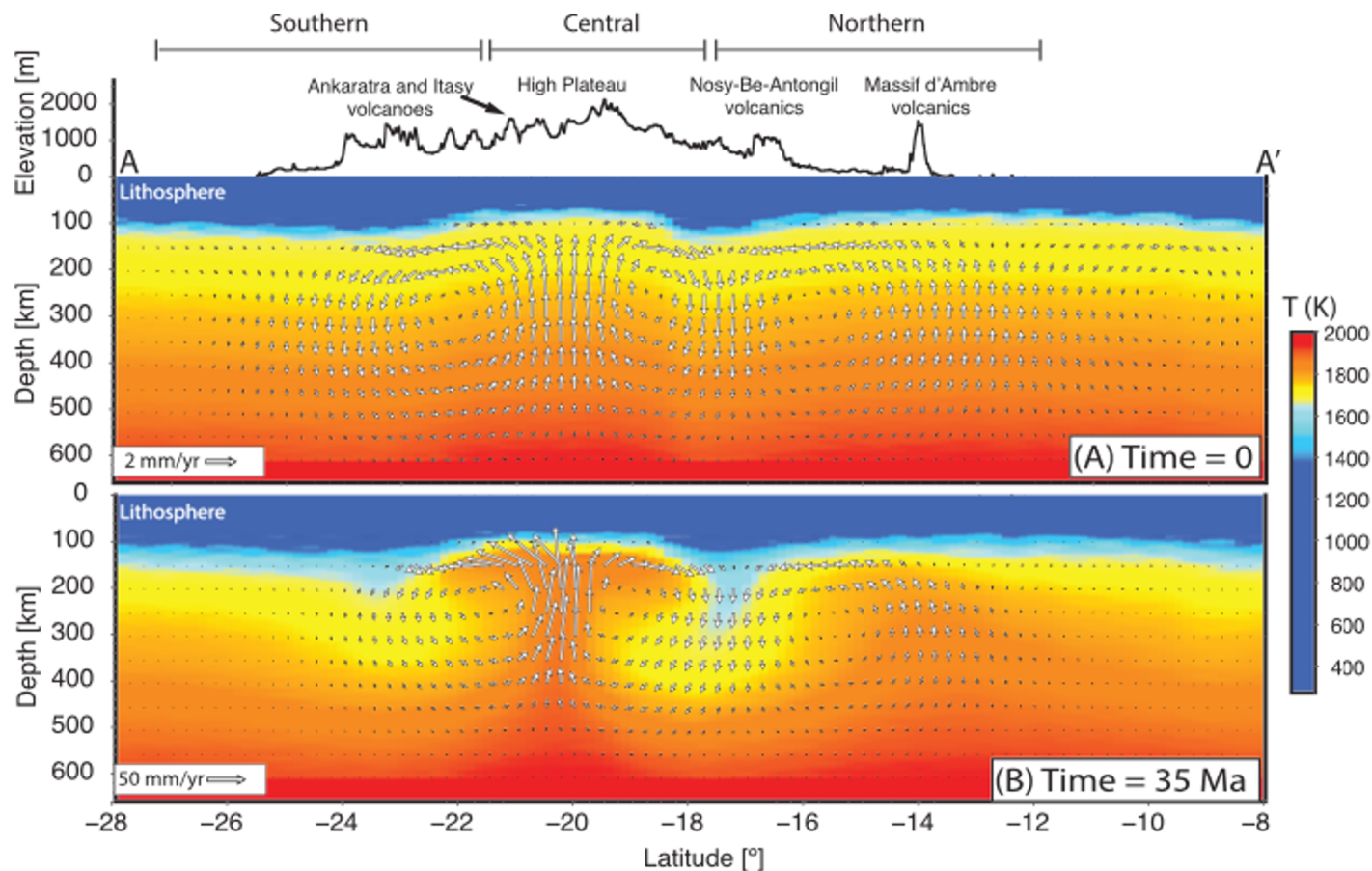
Case Study: Madagascar



Case Study: Madagascar

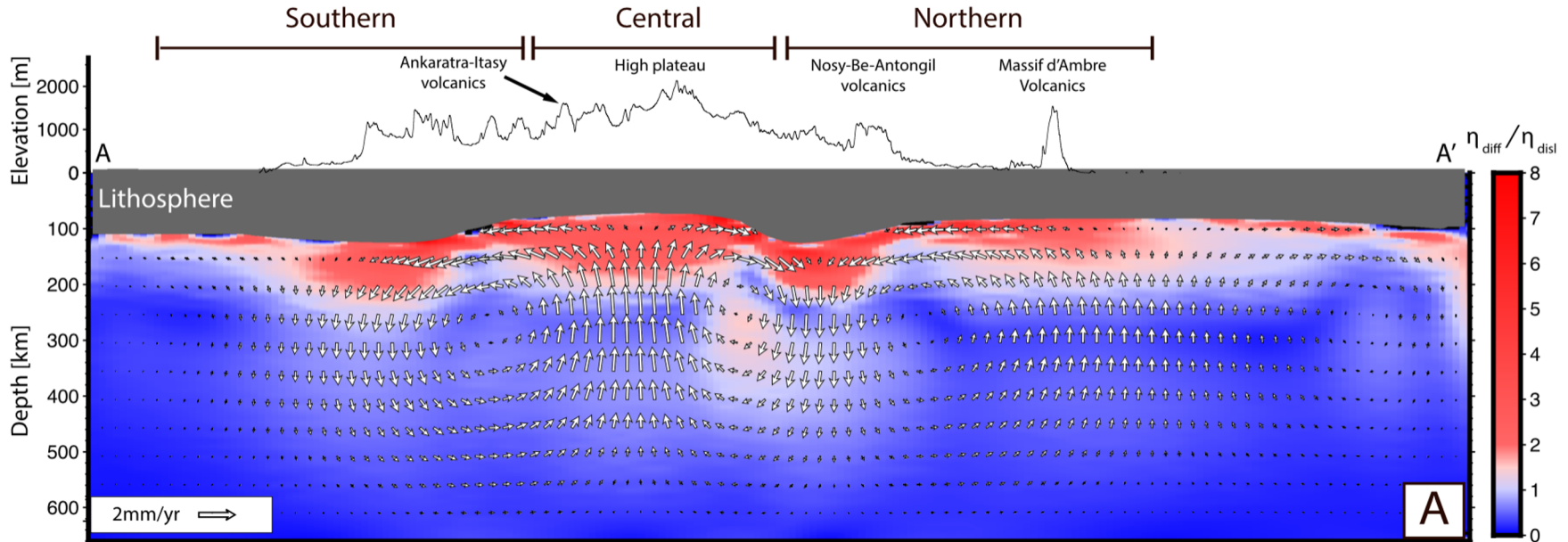


Case Study: Madagascar



Case Study: Madagascar

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ASPECT can be used for modeling deep melt generation

JGR Solid Earth

RESEARCH ARTICLE

10.1029/2020JB020728

Key Points:

- Lithospheric modulated convection beneath the Rungwe Volcanic Province generates melt when mantle potential temperatures are elevated
- Plume material beneath the Rungwe Volcanic Province is required to explain geochemical and geophysical observations
- Lithospheric modulated convection enables the entrainment of plume material beneath the Rungwe Volcanic Province

Lithospheric Control of Melt Generation Beneath the Rungwe Volcanic Province, East Africa: Implications for a Plume Source

Emmanuel A. Njinju¹, D. Sarah Stamps¹, Kodi Neumiller², and James Gallagher²

¹Department of Geosciences, Virginia Tech, Blacksburg, VA, USA, ²OPeNDAP, Narragansett, RI, USA

Abstract The Rungwe Volcanic Province (RVP) is a volcanic center in an anomalous region of magma-assisted rifting positioned within the magma-poor Western Branch of the East African Rift (EAR). The source of sublithospheric melt for the RVP is enigmatic, localized, unlike the Eastern Branch of the EAR. Some studies suggest that the melt beneath the RVP arises from thermal perturbations in the upper mantle beneath the African superplume flowing from the SW, while others suggest that the melt is generated by a localized plume source.

Njinju, E. A., Stamps, D. S., Neumiller, K., & Gallagher, J. (2021). Lithospheric control of melt generation beneath the Rungwe Volcanic Province, East Africa: Implications for a plume source. *Journal of Geophysical Research: Solid Earth*, 126, e2020JB020728. <https://doi.org/10.1029/2020JB020728>

njinju. (2022). njinju/aspect-2.2.0_njinju_et_al_2021_JGR: Lithospheric Control of Melt Generation Beneath the Rungwe Volcanic Province, East Africa: Implications for a Plume Source (v2.2.0). Zenodo. <https://doi.org/10.5281/zenodo.6533974>

zenodo

May 9, 2022

njinju/aspect-2.2.0_njinju_et_al_2021_JGR: Lithospheric Control of Melt Generation Beneath the Rungwe Volcanic Province, East Africa: Implications for a Plume Source

njinju

ASPECT code for lithospheric modulated convection (LMC) and melt generation by Njinju et al., 2021-JGR (<https://doi.org/10.1029/2020JB020728>)

Preview

aspect-2.2.0_njinju_et_al_2021_JGR.zip

The previewer is not showing all the files

aspect-2.2.0_njinju_et_al_2021_JGR

- clang-ldy 1.0 kB
- AUTHORS 639 Bytes
- CITATION 113 Bytes
- CMakeLists.txt 22.6 kB

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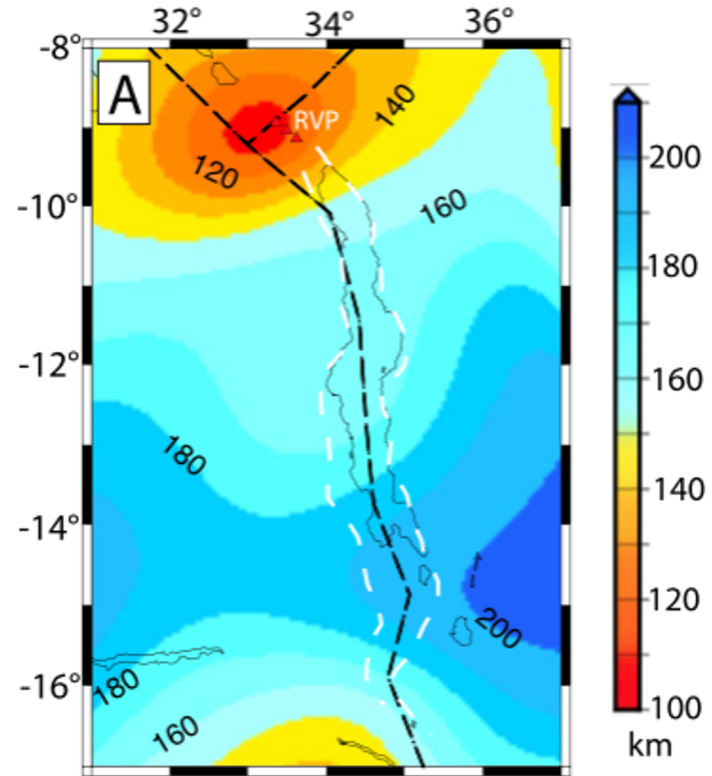
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Model Setup: Initial Temperature Conditions

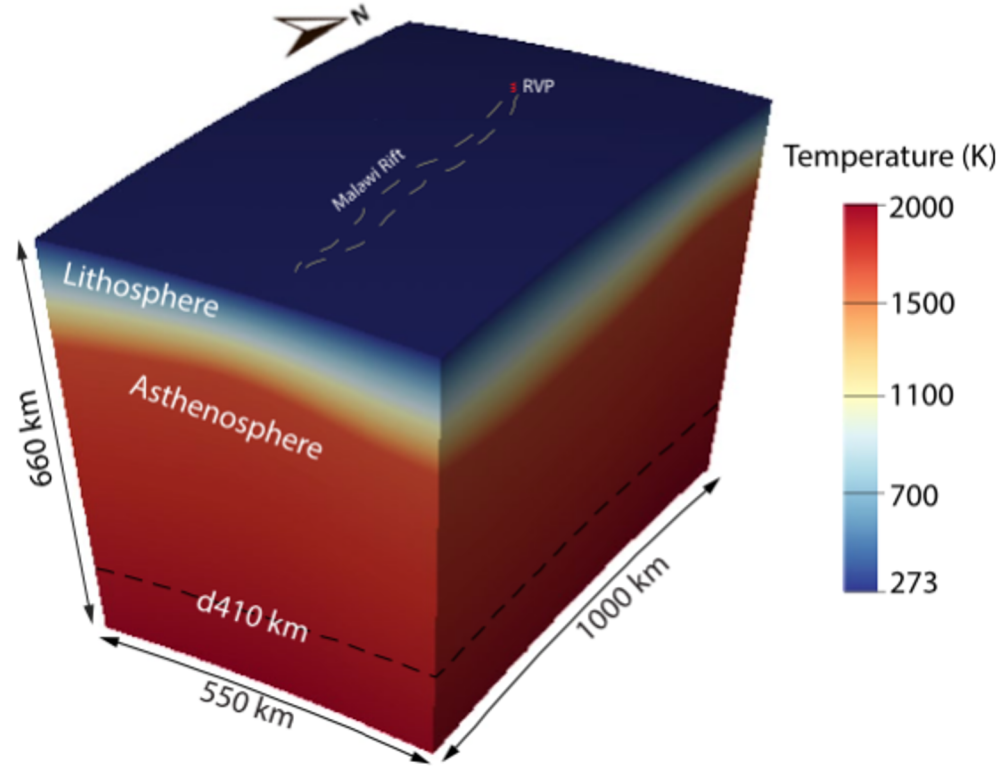
- Lithospheric thickness input with isotherm at lithosphere-asthenosphere boundary (LAB, i.e. Fishwick 2010, updated)
- Approximate conductive geotherm in lithosphere: linear gradient from surface to LAB isotherm (1673 K)
- Approximately adiabatic geotherm below lithosphere (i.e. 0.4 K/km)



Lithospheric Thickness

Model Setup: Initial Temperature Conditions

- Lithospheric thickness input with isotherm at lithosphere-asthenosphere boundary (LAB, i.e. Fishwick 2010, updated)
- Approximate conductive geotherm in lithosphere: linear gradient from surface to LAB isotherm (1673 K)
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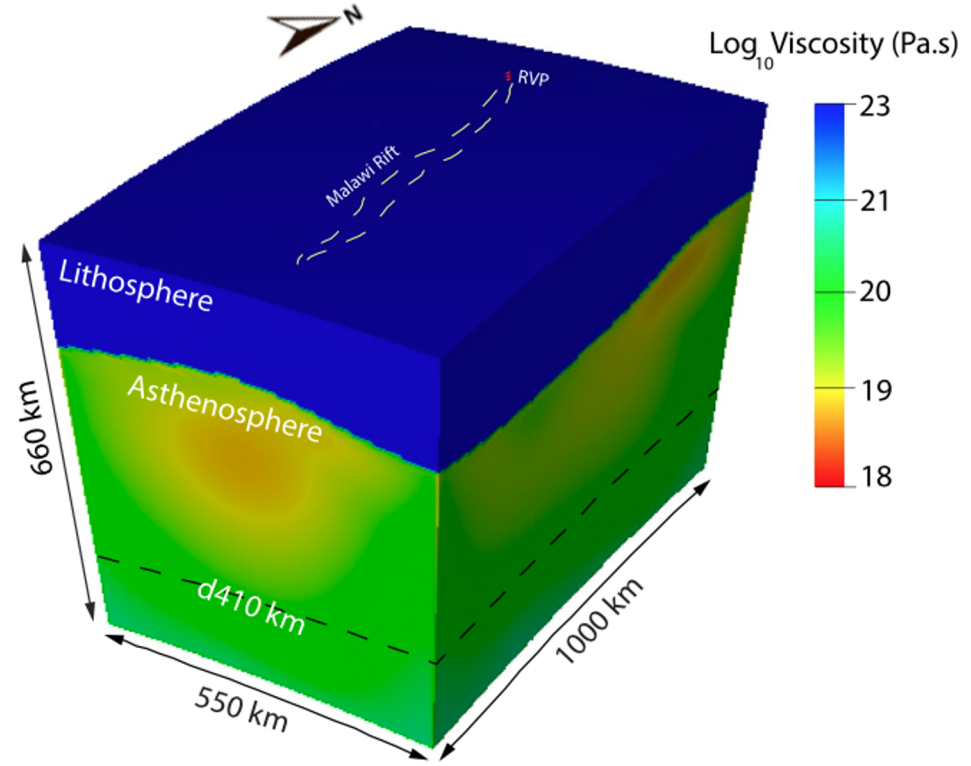
Model Setup: Viscosity

- Rigid lid lithospheric viscosity
- Composite rheology below the LAB (Jadamec & Billen, 2010)

$$\eta_{comp} = \frac{\eta_{diff} \times \eta_{disl}}{\eta_{diff} + \eta_{disl}}$$

η_{diff} = viscosity for diffusion creep

η_{disl} = viscosity for dislocation creep



Model Setup: Density and Melt Generation

- Crust = 2700 kg/m³, Mantle_ref = 3300 kg/m³
- Density of solid rock solid and the density of melt melt vary with both temperature and pressure below the LAB as follows:

$$\rho_{solid,melt} = \rho(T, p) = \rho_0 \left[1 - \alpha(T - T_0) \right] e^{\left[\beta(p - p_0) \right]}$$

ρ_0 = reference density

β = compressibility coefficient

α = thermal expansivity

p = pressure

T_0 = reference temperature

p_0 = reference pressure

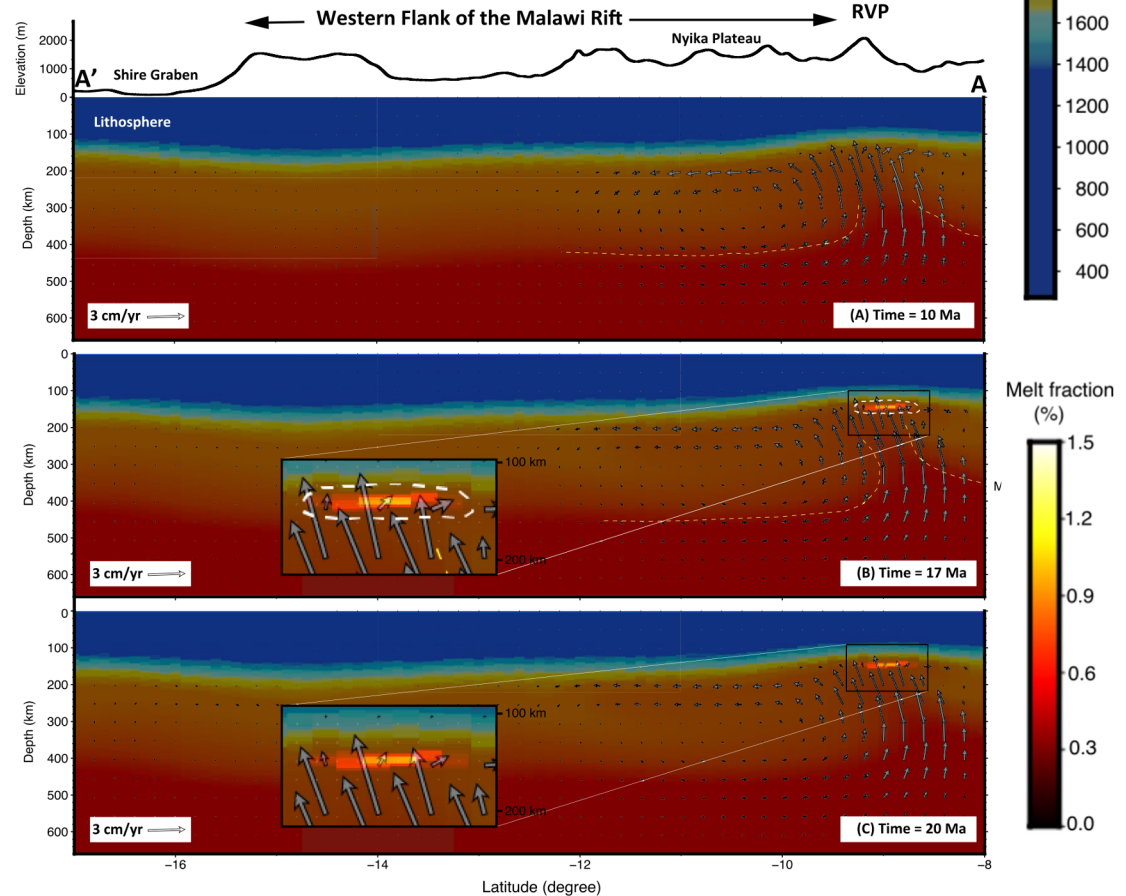
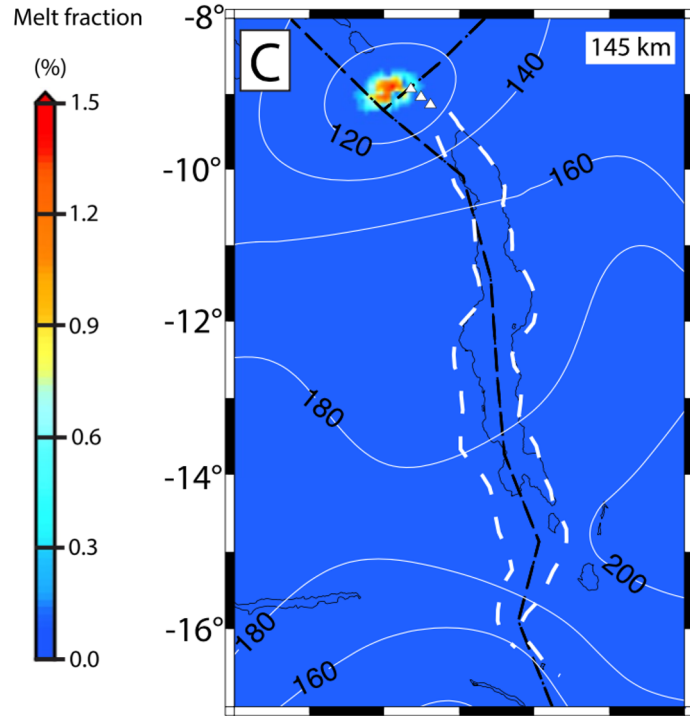
T = temperature

- Melt fraction:

$$F(p, T) = \left(\frac{T - T_{solidus}}{T_{liquidus} - T_{solidus}} \right)^{1.5}$$

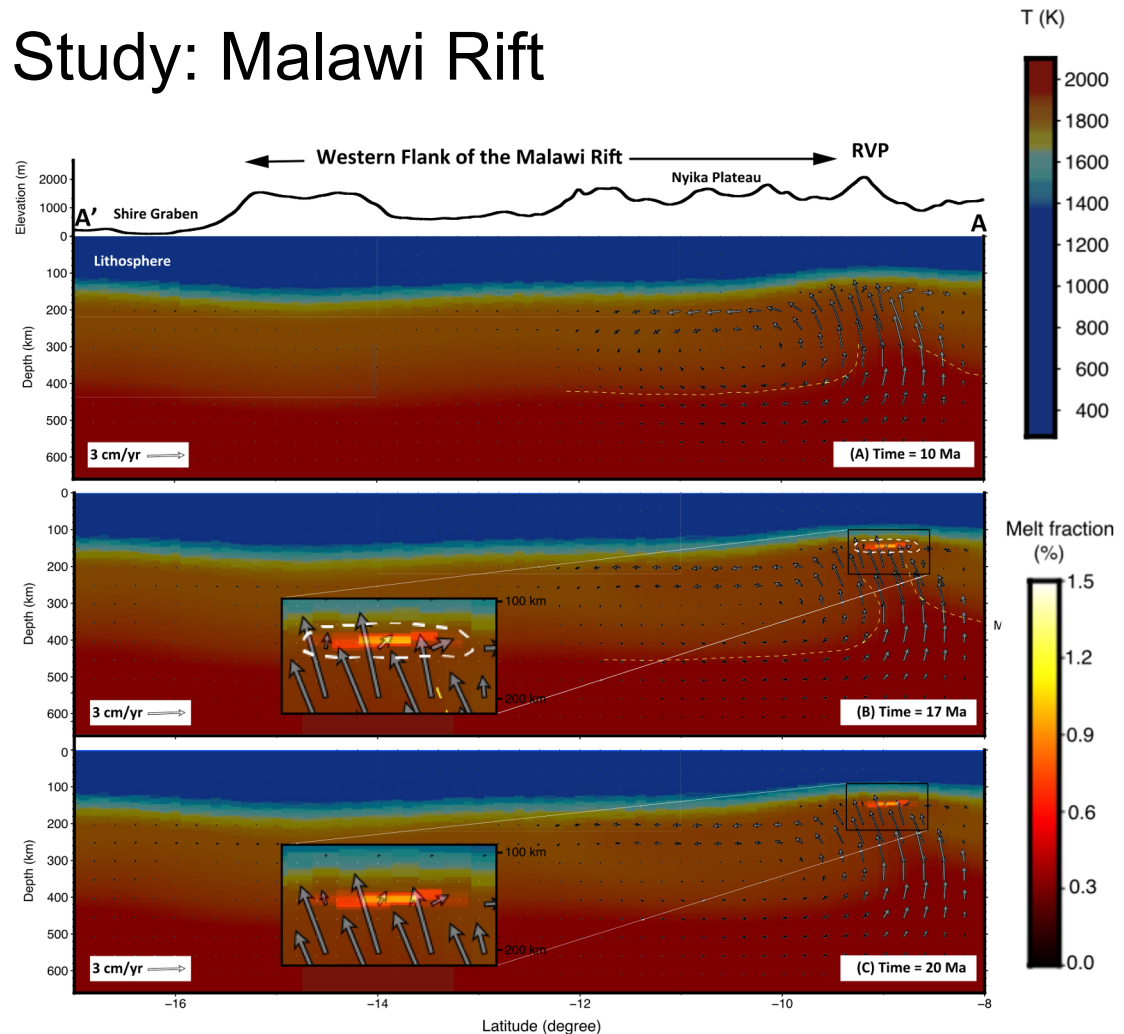
for $T_{solidus} \leq T \leq T_{liquidus}$

Case Study: Malawi Rift



Case Study: Malawi Rift

- Lithospheric modulated convection beneath the Rungwe Volcanic Province generates melt when mantle potential temperatures are elevated
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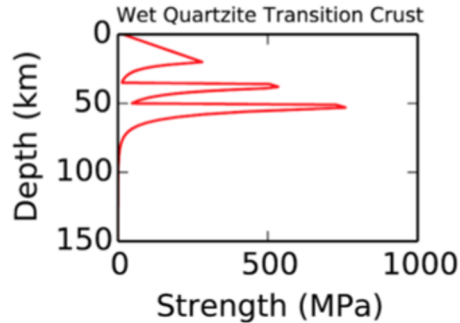


Summary and Conclusions

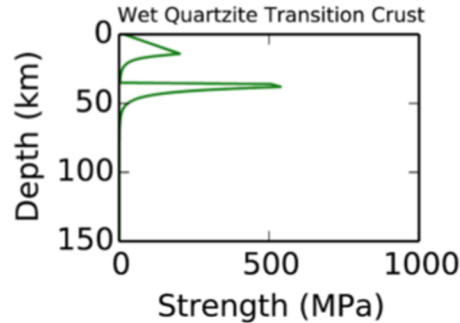
- ASPECT can be used for:
 - Lithospheric deformation
 - Asthenospheric flow
 - Deep melt generation
- Applications to:
 - East African Rift
 - Rigid plate motions explained by lithospheric buoyancy forces
 - Madagascar
 - Dislocation creep extends into the upper asthenosphere
 - Malawi Rift
 - Lithospheric control on the location of melt generation
- All developments are open access and available via Zenodo or the ASPECT release code

Model Set-up: Rheology

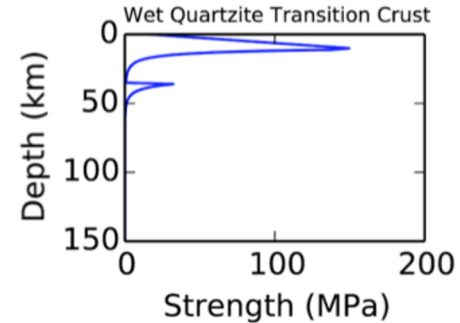
Surface heat flow: 33 mW/m²



64 mW/m²



80 mW/m²



- Crust: combines nonlinear dislocation creep (dry quartzite) with plastic failure
- Mantle lithosphere: combines olivine dislocation creep with plastic failure
- Sublithospheric mantle: composite rheology