

# CONTROLS OF INTERMODEL UNCERTAINTY IN LAND CARBON SINK PROJECTIONS

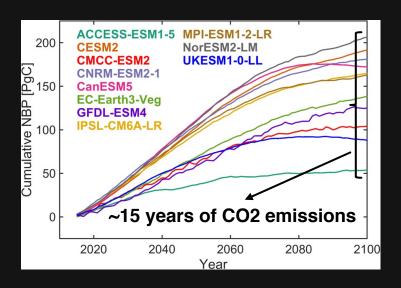
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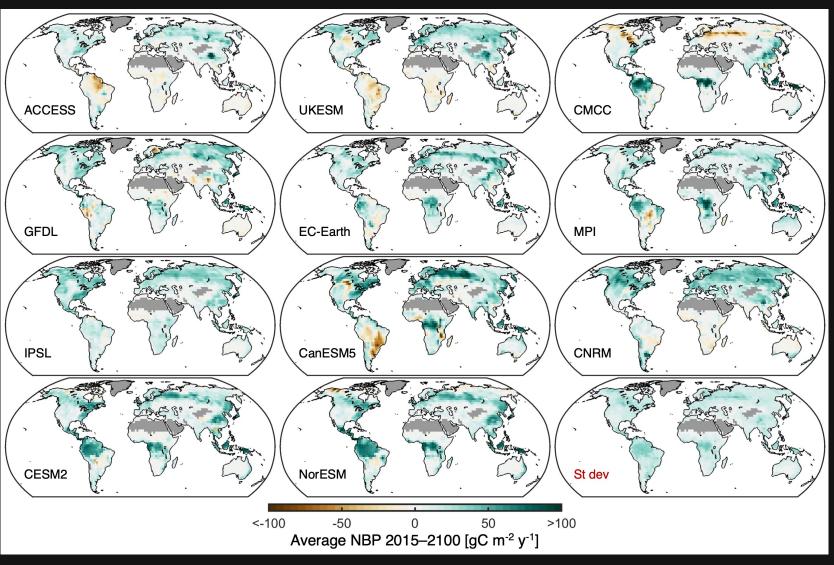


# ~150 PGC OF SPREAD IN CUMULATIVE NBP BY 2100 WITH 2C WARMING



Even models with a similar C sink can have very different regional response.

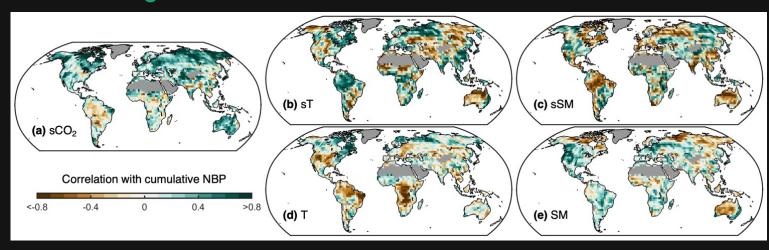
Large differences in the tropics and boreal forests.



#### CONTROLS OF INTERMODEL DIFFERENCES IN NBP PROJECTIONS

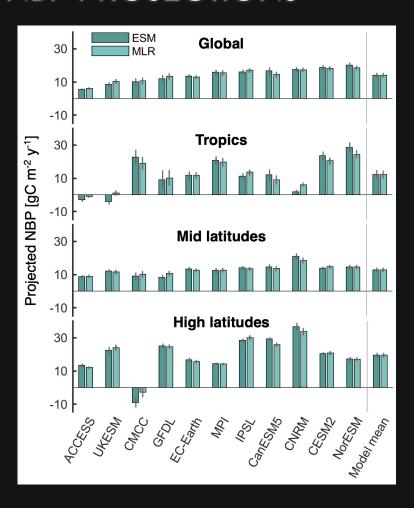
- Sensitivity of GPP to CO<sub>2</sub> (sCO<sub>2</sub>)
- Sensitivity of NBP to interannual T variability (sT)
- Sensitivity of NBP to interannual SM variability (sSM)
- Long-term average T
- Long-term average SM

**Higher NBP** is favored in models with **higher sCO<sub>2</sub>**, **higher sT**, **lower sSM**, **lower T**, and **higher SM**.



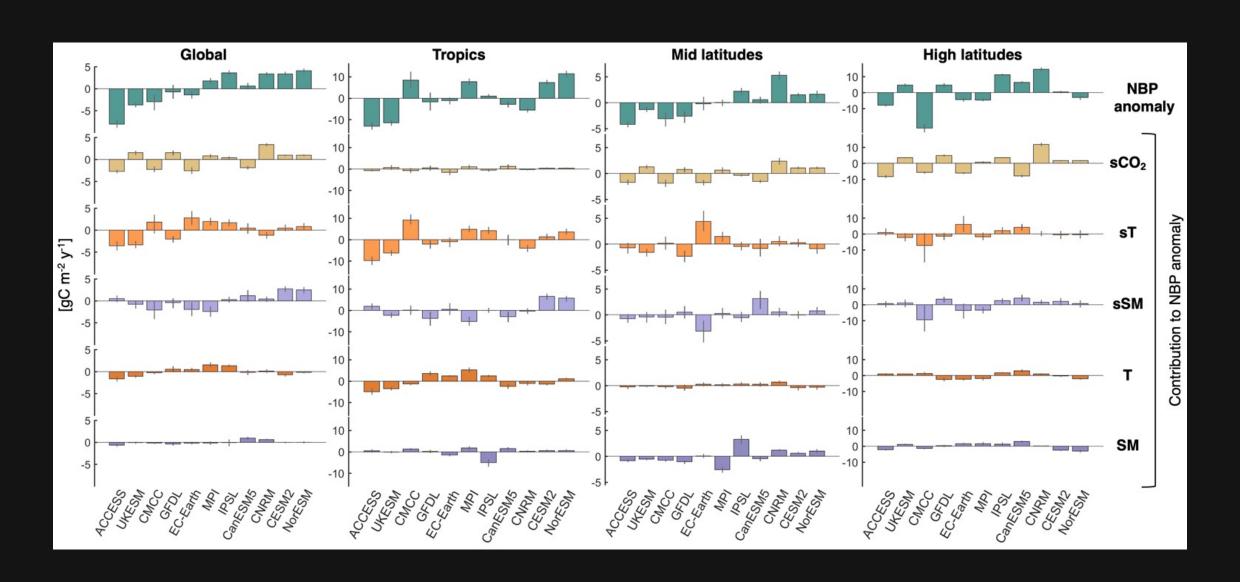
At every grid cell: multiple linear regression (MLR) to estimate cumulative NBP of every ESM m.

$$NBP_m = b_0 + b_1 *sCO2_m + b_2 *sT_m + b_3 *sSM_m + b_4 *T_m + b_5 *SM_m$$



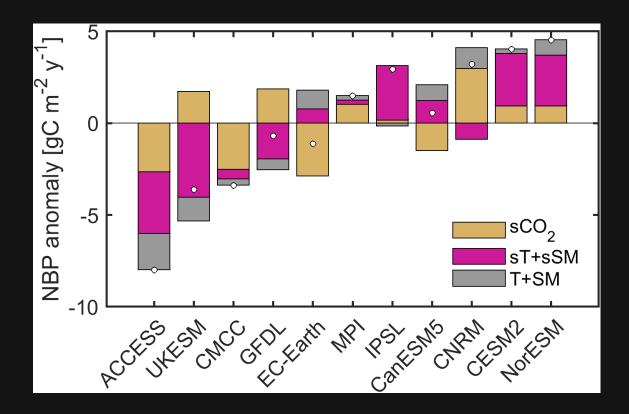
Global and regional intermodel differences in NBP are well represented by the regression estimate

### CONTRIBUTIONS OF THE CONTROLS TO EXPLAIN DIFFERENCES IN NBP



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Dominant role of the response of NBP to interannual temperature and soil moisture variability over that of the CO2 fertilization effect and average climate conditions.

# SOIL MOISTURE IS THE MAIN CONTROL OF INTERANNUAL VARIABILITY OF NBP

