



Supplementary materials

Mass transfer between serpentinites and metapelites in a paleo-subduction interface: a case study from the Yuli belt, eastern Taiwan

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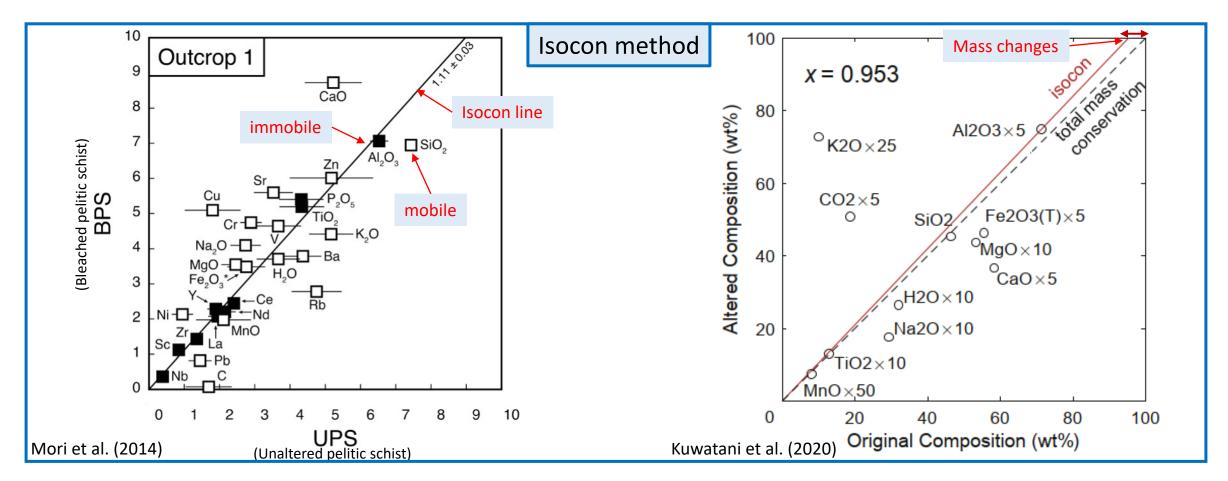


Introduction

Isocon method

- Isocon method can be used to estimate the mass, volume, or concentration changes due to mass transfers during metasomatism.
 (Grant 1986: Grant 2005: More than the mass of the m
- This method is calculated on the basis of immobile element concentrations.

(Grant, 1986; Grant, 2005; Mori et al., 2007; Kuwatani et al., 2020).



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Petrographic observations

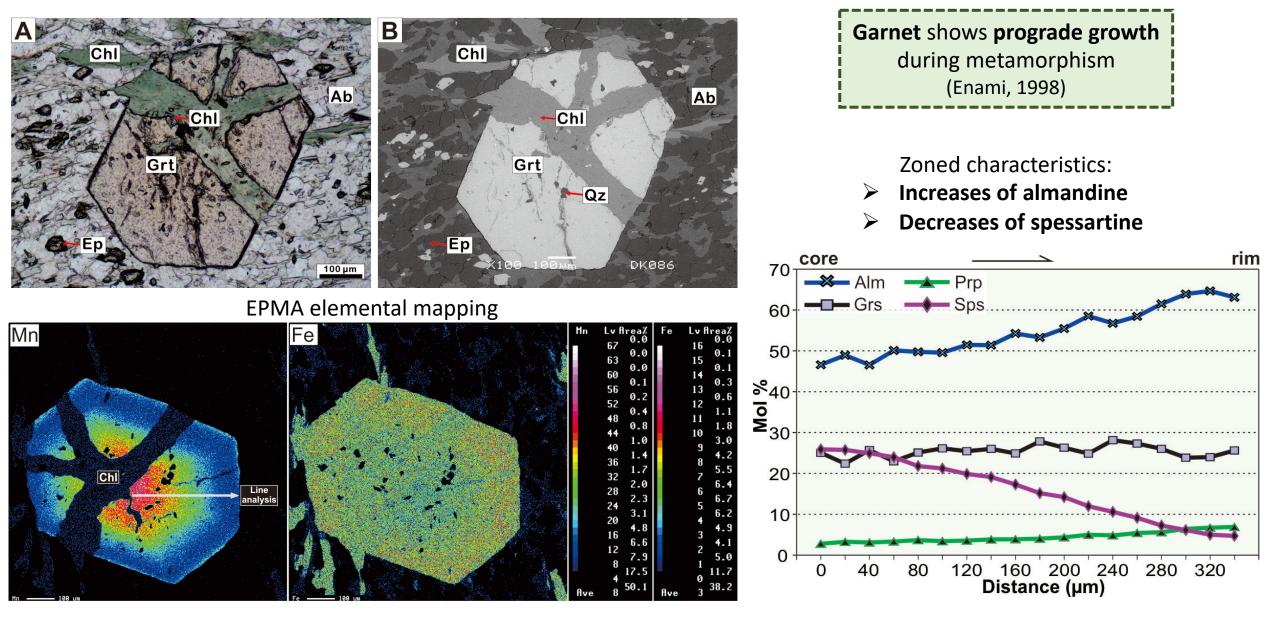
Metasomatic rocks at the Tsunkuanshan area									
Zones		Rock names	Sample code	Texture	Petrology		Coordinate		
					Major minerals (90%)	Minor (≤9%)	Assesory (≤1%)	North	East
			PS001	Schistose	Ph, qz, chl, CM	Grt, ab, ttn, aln	Ap, zr, py, tour	23°29'54.50"N	121°18'20.99"E
	PS	Pelitic schist	PS002	Schistose	Ph, ab, qz, CM	Grt, ttn, ilm	Ap, zr	-	-
			PS003	Schistose	Ph, qz, chl, CM	Grt, ab, ttn, aln	Ap, zr, py, tour	23°29'26.49"N	121°18'11.13"E
	Zone I	Chl-ab schist/rock	MZ101	Massive	Ab, chl	qz, ttn, ilm, ap, aln	Grt, py	23°29'26.49"N	121°18'11.13"E
			MZ102	Schistose	Ab, chl, ph	Grt, ttn, ilm, ep	Ap, aln, zr, py	23°29'26.49"N	121°18'11.13"E
			MZ103	Schistose	Ab, chl, ph	Ttn, ilm, ep, aln	Zr, py	23°29'27.08"N	121°18'10.75"E
	Zone II	Amp-ab rock	MZ201	Massive	Ab, chl, act, stp, qz	Ttn, rt, ilm, qz	Ap, py	23°29'26.49"N	121°18'11.13"E
			MZ202	Massive	Ab, brs, ttn	Ilm, aln, ep	Ap,zr, py	23°29'27.08"N	121°18'10.75"E
			MZ203	Massive	Ab, bio, gln	Ttn, ilm, ep	Ap, aln, zr, py	23°29'27.08"N	121°18'10.75"E
les			MZ204	Massive	Ab, brs, ttn	Ilm, aln, ep	Ap,zr, py	23°29'27.08"N	121°18'10.75"E
zones	Zone I	Chl-ab rock	MZ104	Massive	Ab, chl, ep	Ttn, rt, ilm, aln	Ap, zr,py	23°29'27.08"N	121°18'10.75"E
atic	Zone III	Ab-chl schist	MZ301	Schistose	Chl (pse. grt/amp), ph, ep, ab	Ttn, aln	Ap, py	23°29'26.49"N	121°18'11.13"E
mo			MZ302	Schistose	Chl (pse. grt/amp), ph, ep	Ttn, aln	Ap, py	23°29'26.49"N	121°18'11.13"E
Metasomatic			MZ303	Schistose	Chl, ph, ab	Ttn, grt (pse), aln	Ap, zr, py	23°29'27.08"N	121°18'10.75"E
Ŵ	Zone IV	Ep-chl schist	MZ401	Schistose	Ep, chl, ph	Ab, ttn	Ap, aln, zr, py	23°29'27.08"N	121°18'10.75"E
			MZ402	Schistose	Ep, chl, ab	Ttn, aln, ph	Ap, zr, py	23°29'27.08"N	121°18'10.75"E
			MZ403	Schistose	Ep, chl, ab	Ttn, ilm	Ap, zr	23°29'27.08"N	121°18'10.75"E
			MZ404	Schistose	Ep, chl, py	Tlc, ttn	Ilm, ap	23°29'27.08"N	121°18'10.75"E
	Zone V	Chl-tlc schist	MZ501	Schistose	Chl, tlc	Ep, ttn	Py, ap	23°29'27.08"N	121°18'10.75"E
			MZ502	Schistose	Chl, tlc	Py, ttn	Ap	23°29'27.08"N	121°18'10.75"E
			MZ503	Schistose	Tlc	Chl, mag	Ру	23°29'27.08"N	121°18'10.75"E
SP		Serpentinite	SP001	Schistose	Atg, mag	Chl, chr, car	Cr-spl, Fe-chr	23°29'26.49"N	121°18'11.13"E
			SP002	Schistose	Atg, mag	Chl, chr, car	Cr-spl, Fe-chr	23°29'26.49"N	121°18'11.13"E
			SP003	Massive	Atg, mag	Brc, chr, car	Ol, Cr-spl, Fe-chr	23°29'26.49"N	121°18'11.13"E

*24 selected samples



Petrographic observations and mineral compositions

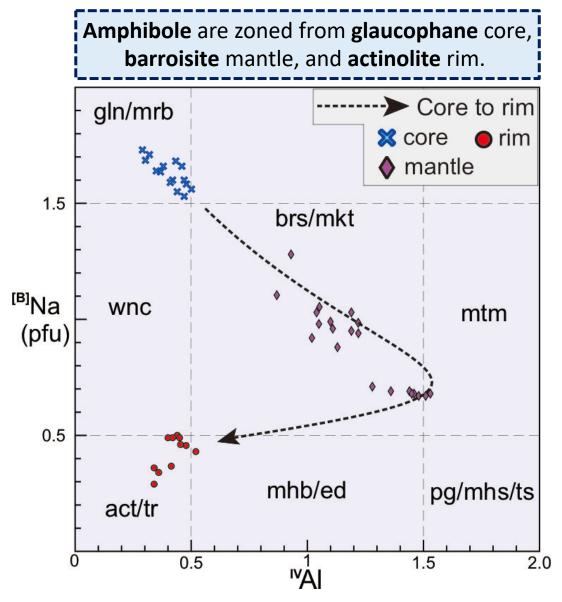
Garnet from the zone I

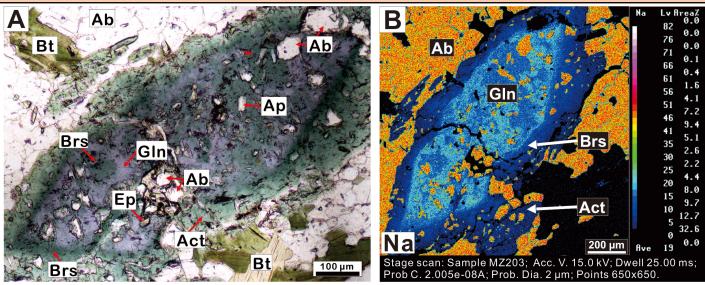




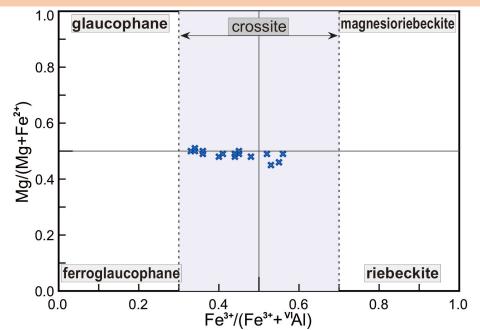
Petrographic observations and mineral compositions

Amphibole in the zone II





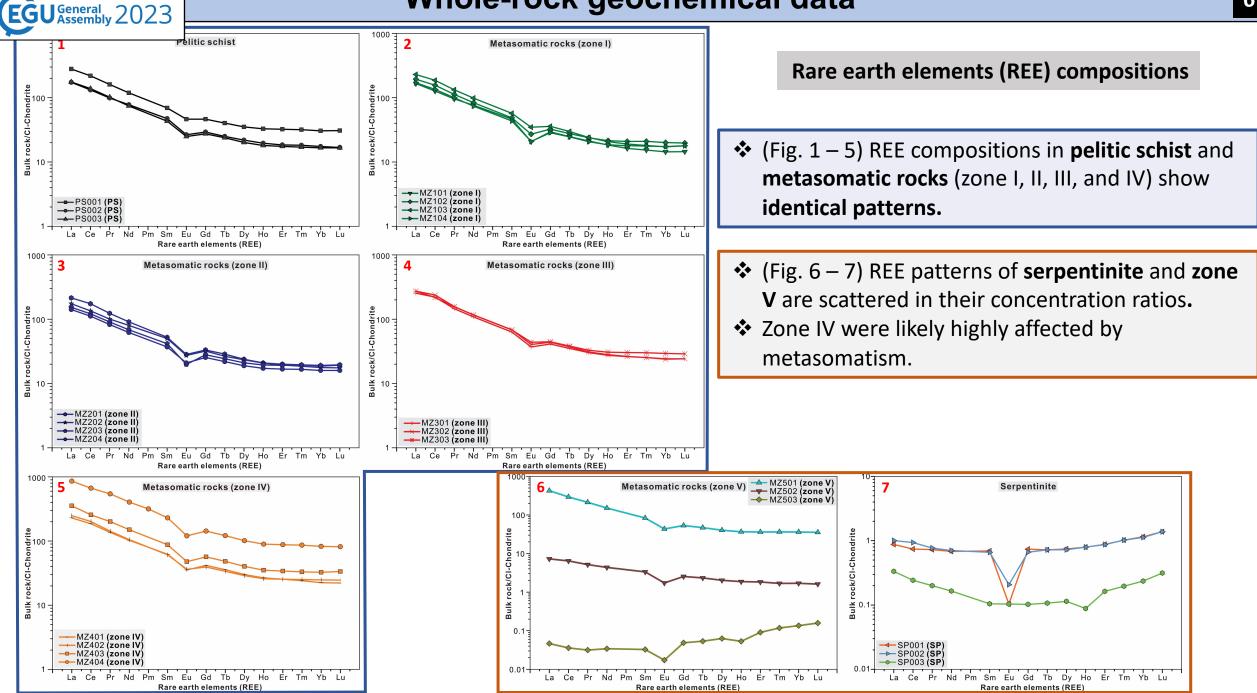
Glaucophane is a **high-pressure index mineral** which is an **indication** of **subduction zone metamorphism**





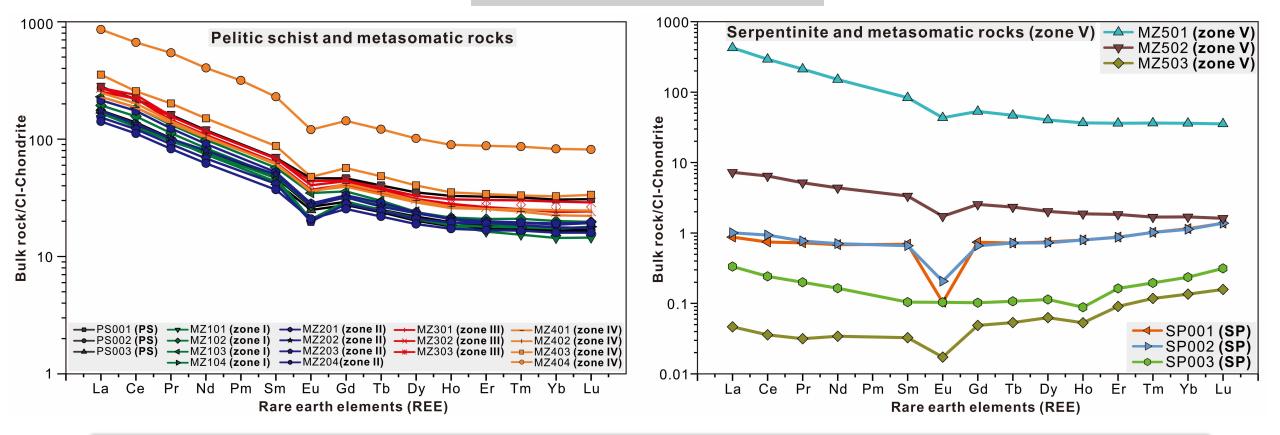
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Comparison of REE compositions

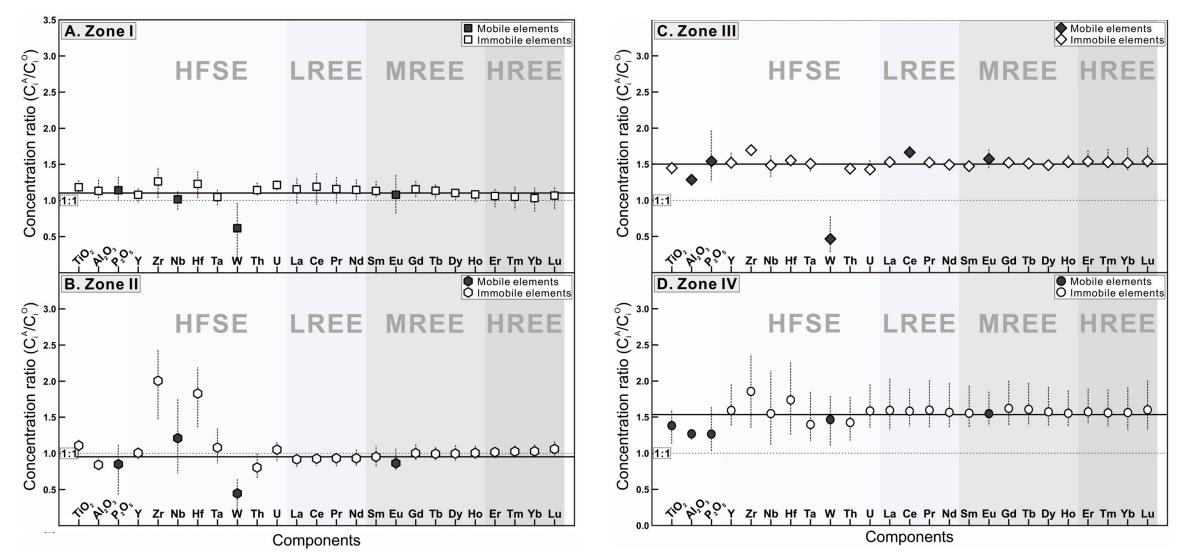


Metasomatic rocks (zone I, II, III, and IV) and pelitic schist were likely from the same origin

Isocon method can be used to estimate the **mass or concentration changed** within the metasomatic rocks.



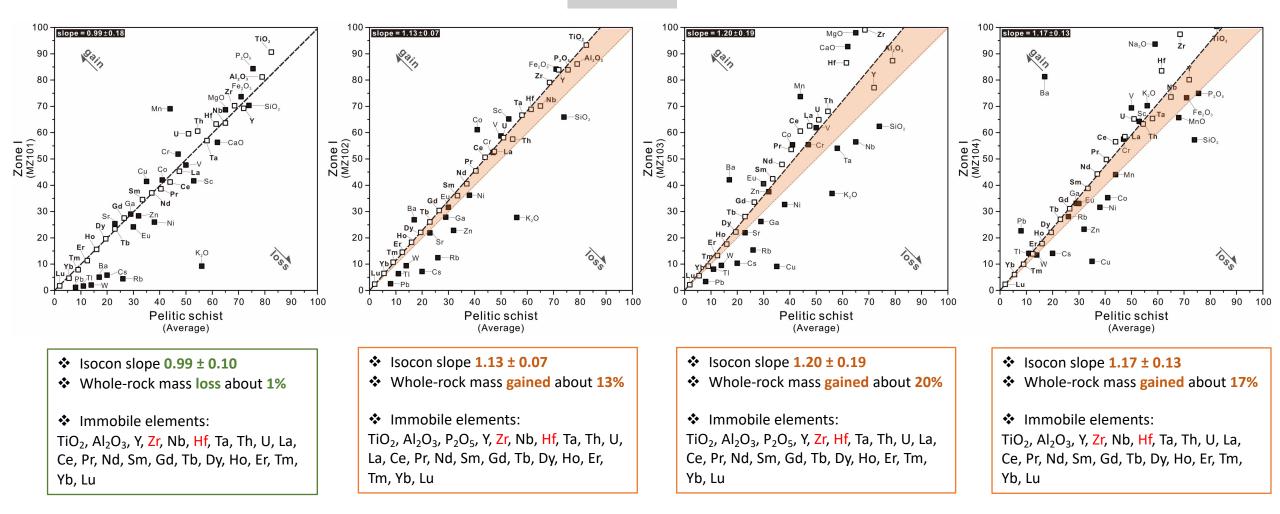
- Mobilities of components in the metasomatic zones over pelitic schist (average)
- ✤ C_i^a: component concentration of altered/metasomatic rocks
- ✤ C_i^o: component concentration of original rocks/pelitic schist





Isocon method (1)

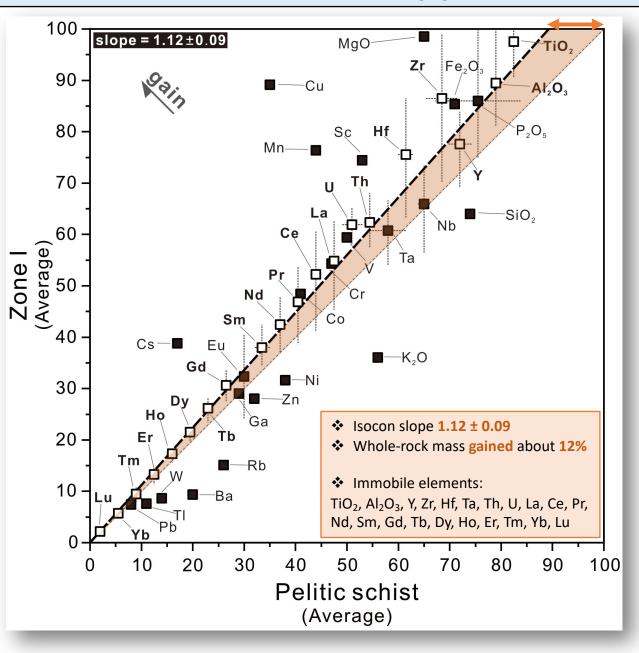






Isocon method (1)

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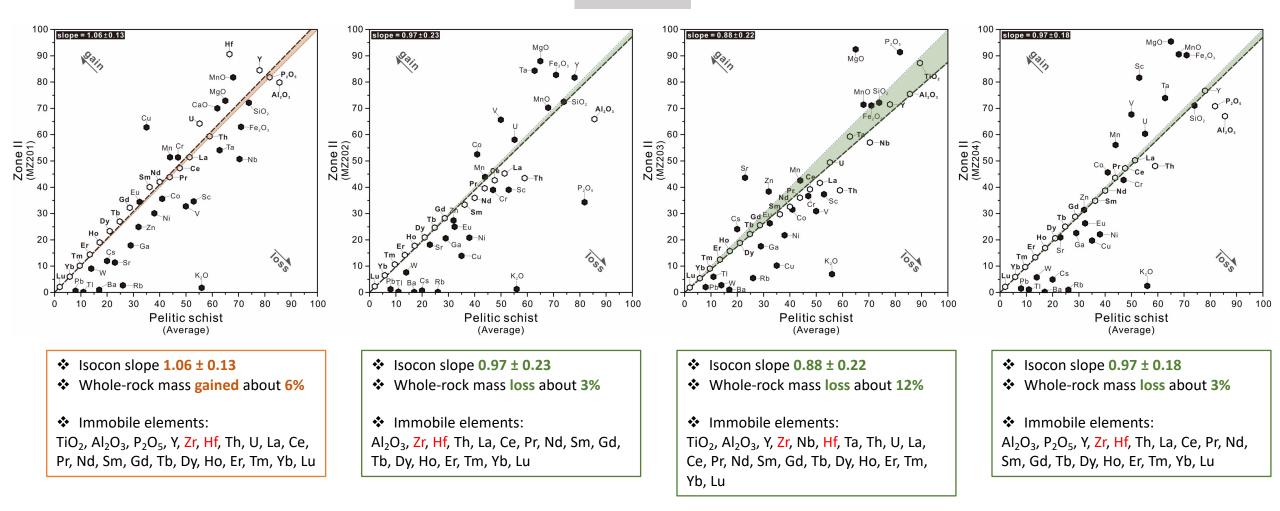


Zone I (Average)



Isocon method (2)

Zone II

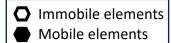




Isocon method (2)

100slope = 0.96±0.16 TiO₂ Hf 90 · 211 MgO— 80 · MnO AI_2O_3 Ta SiO 70· U $P_2 O_5$ 60 Zone II (Average) Ce Mn 50 Th Sm 40 Cr La Zn Gd Pr 30 · Nd Ho ✤ Isocon slope 0.96 ± 0.16 20 -Whole-rock mass loss about 4% Ga Tm Ъy Yb Immobile elements: 10 -Cs TiO₂, Al₂O₃, Y, Zr, Hf, Ta, Th, U, La, Ce, Pr, Nd, Sm, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu -Rb Ba 100 20 30 40 50 60 70 80 90 10 0 Pelitic schist (Average)

Zone II (average)

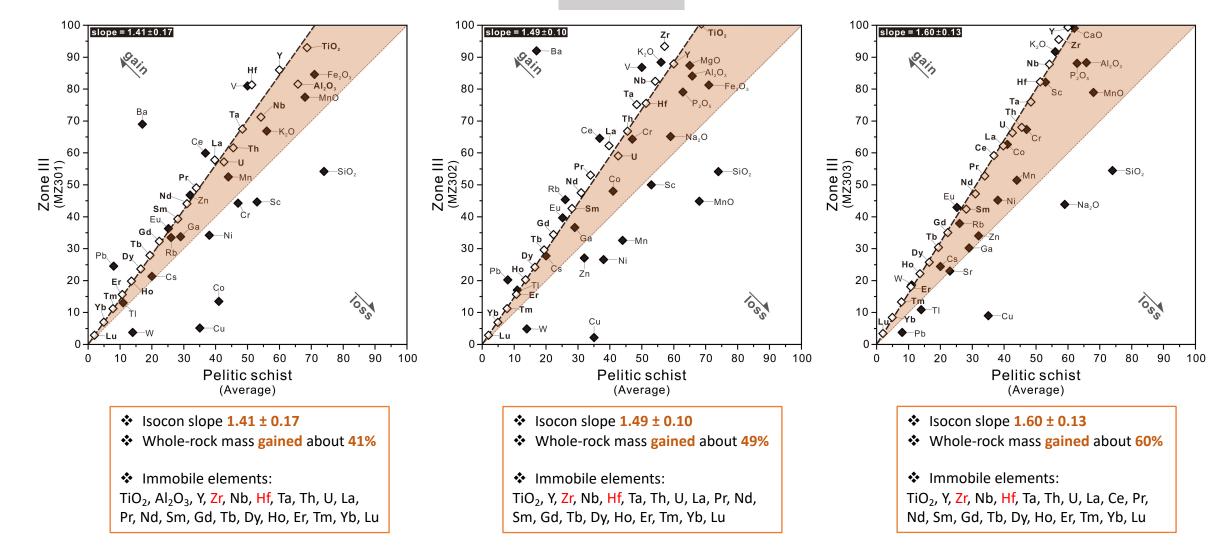




Isocon method (3)

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Zone III

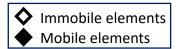




Isocon method (3)

100slope = 1.50±0.07 Zr TiO, Ba Fe₂O₃ 90 -Hf $-AI_2O_3$ 80 · Та Nb Na₂O 70 -La MnO Ce Th 60· Zone III (Average) Pr - SiO₂ Sc Sm -Mn Еu Со 40 Gd Rb Tb Ni ∕ Ga[′]Zn 30 -Dy Isocon slope 1.50 ± 0.07 Er 20 -Ph Whole-rock mass gained about 50% lo Yb Immobile elements: -Tm Cu 10 -TiO₂, Y, Zr, Nb, Hf, Ta, Th, U, La, Pr, Nd, W Sm, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu 0 20 30 80 100 0 10 40 50 60 70 90 Pelitic schist (Average)

Zone III (average)

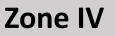


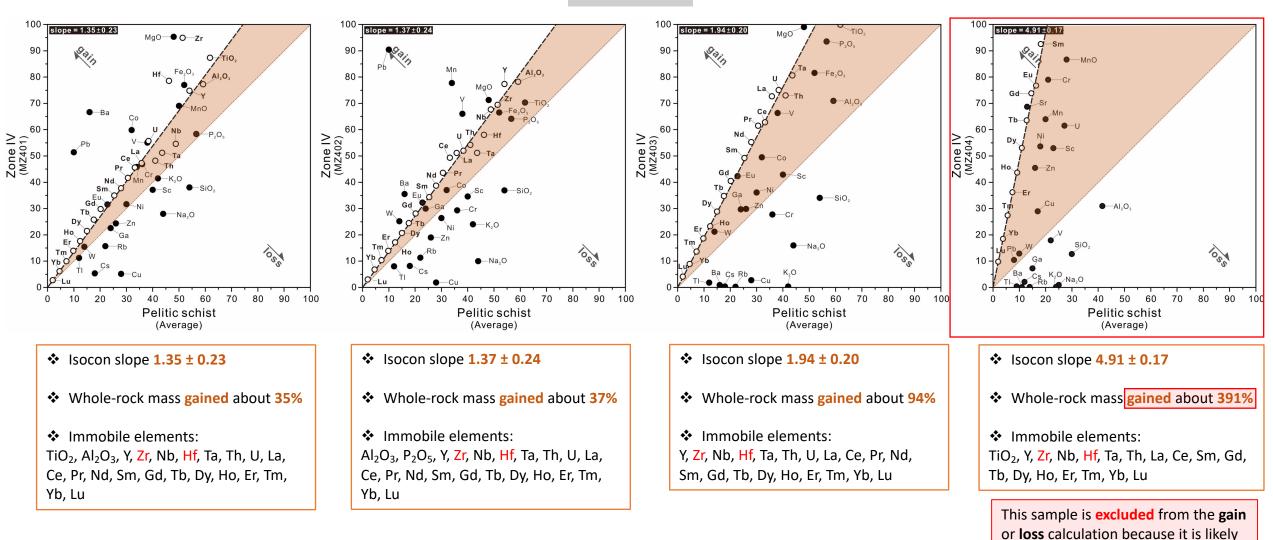


Isocon method (4)

to have been highly affected by

metasomatism than other samples.







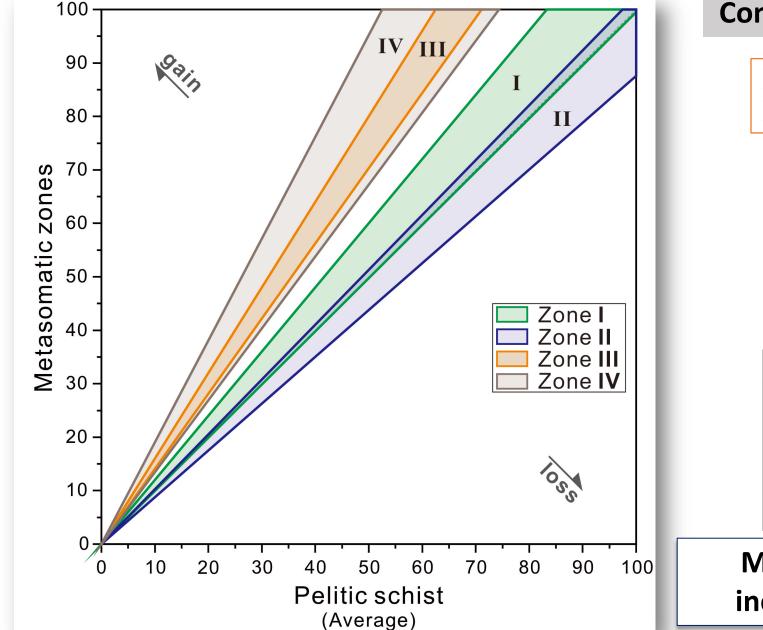
Isocon method (4)

100slope = 1.56±0.16 90 -MgO -TiO₂ Hf Mn 80 · Рb AI_2O_3 Nb 70· P₂O₄ 60· Zone IV (Average) Sm 40 Sc Ba ●—SiO₂ Cr 30 -Cs D Isocon slope 1.56 ± 0.16 20 -Whole-rock mass gained about 56% Gа Immobile elements: 10 Y, Zr, Nb, Hf, Ta, Th, U, La, Ce, Pr, Nd, ● Rb Ni Sm, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu 0 30 60 80 90 100 0 10 20 40 50 70 Pelitic schist (Average)

Zone IV (average)







Comparison of whole-rock mass changes

Zone I, III, and IV show mass gained
Zone II displays mass loss

Mass changes	
✤ Zone I: -1% to +20%	
✤ Zone II: -12% to +6%	
✤ Zone III: +41% to +60%	
✤ Zone IV: +35% to +94%	

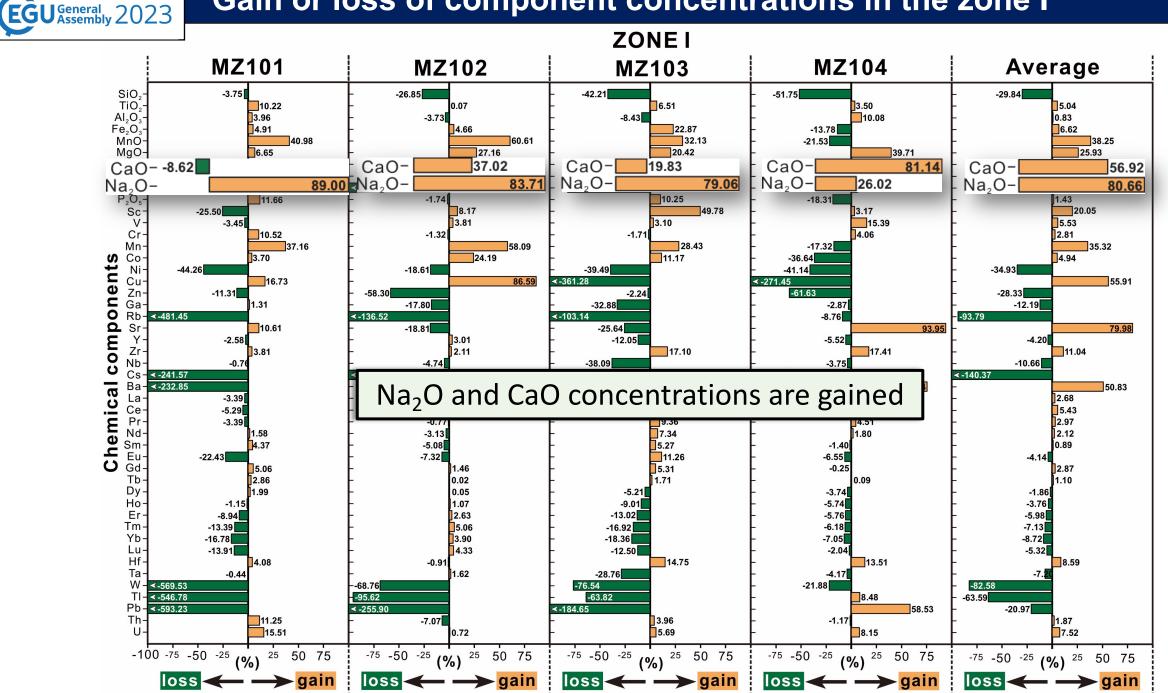
average

Zones	Isocon slope	Mass changed
	1.12 ± 0.09	~12% gained
	0.96 ± 0.16	~4% loss
	1.50 ± 0.07	~50% gained
IV	1.56 ± 0.16	~56% gained

Metasomatism intensities are likely increased from the zone I to zone IV.

Gain or loss of component concentrations in the zone I

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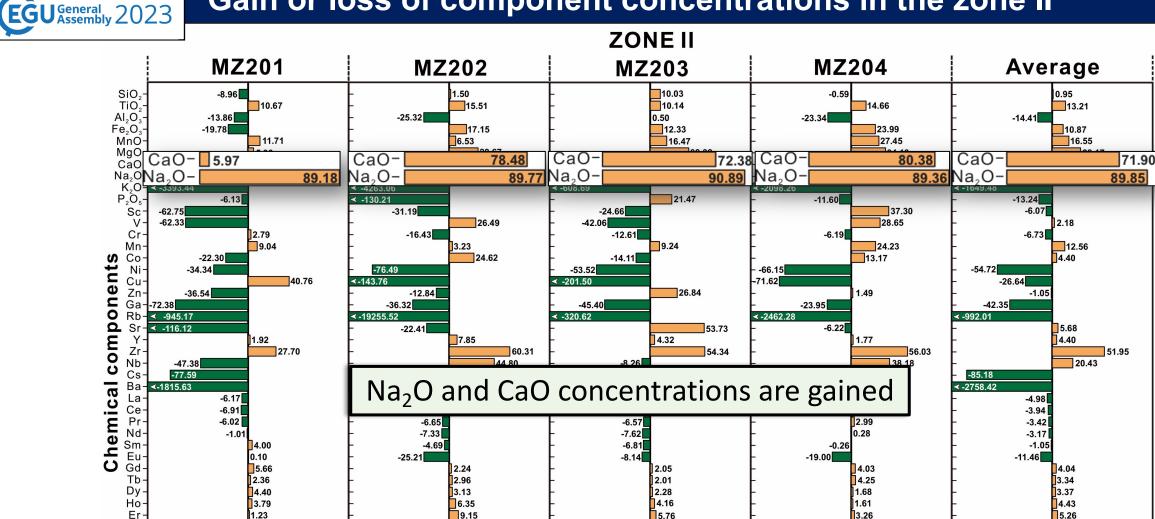
Gain or loss of component concentrations in the zone II

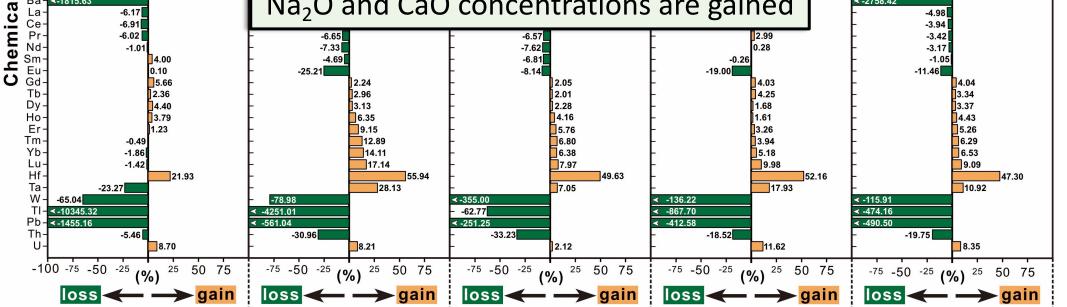
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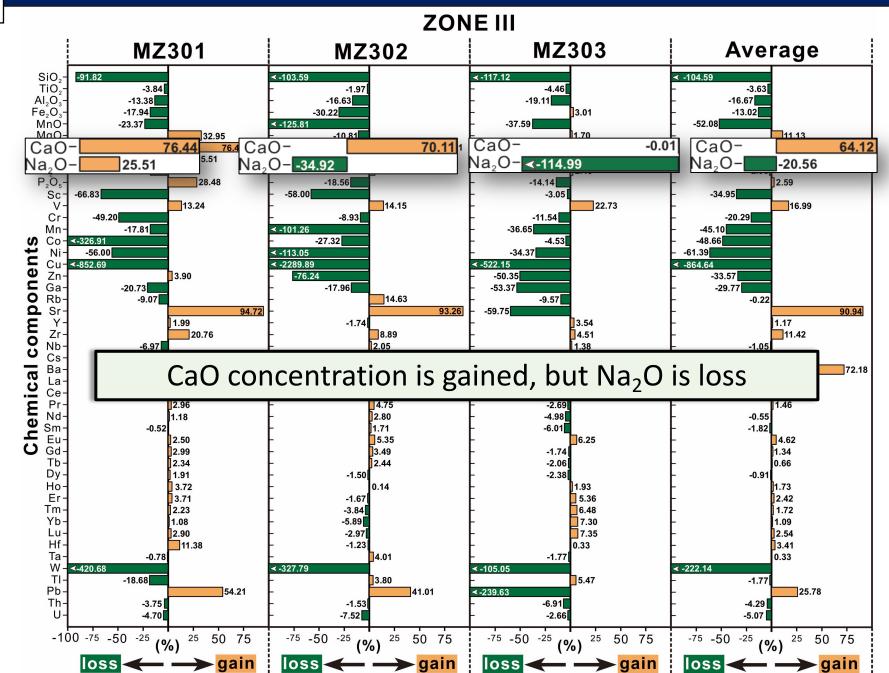


Gain or loss of component concentrations in the zone III

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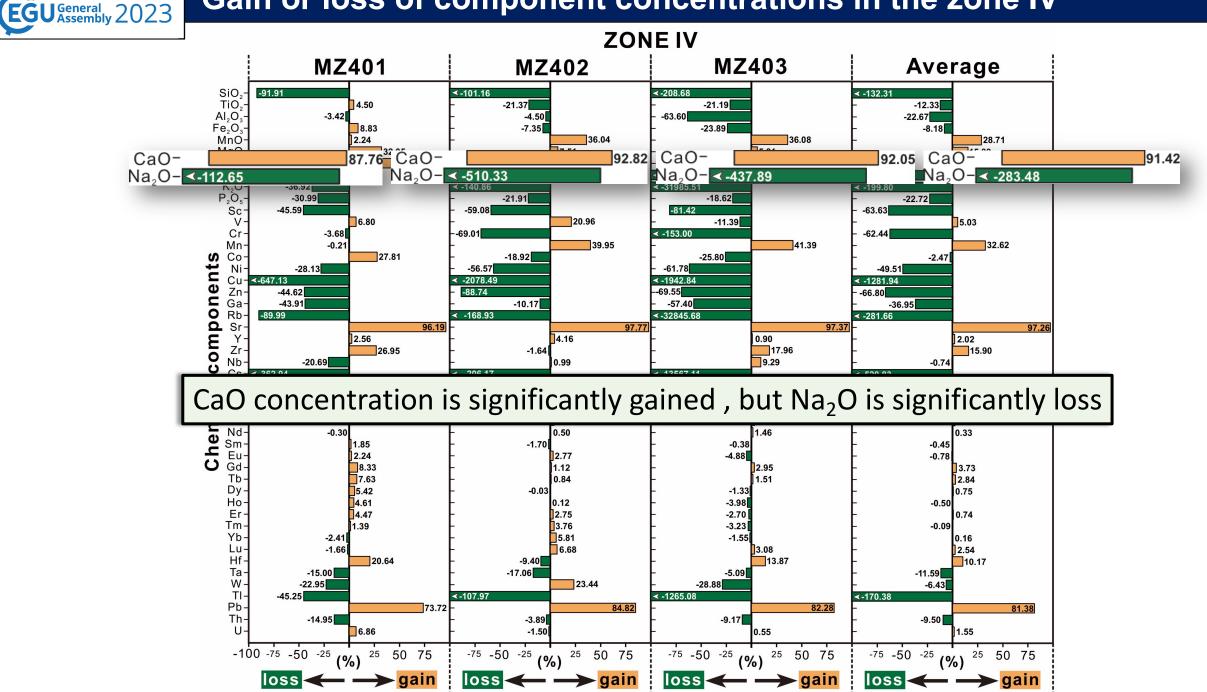


Gain or loss of component concentrations in the zone IV

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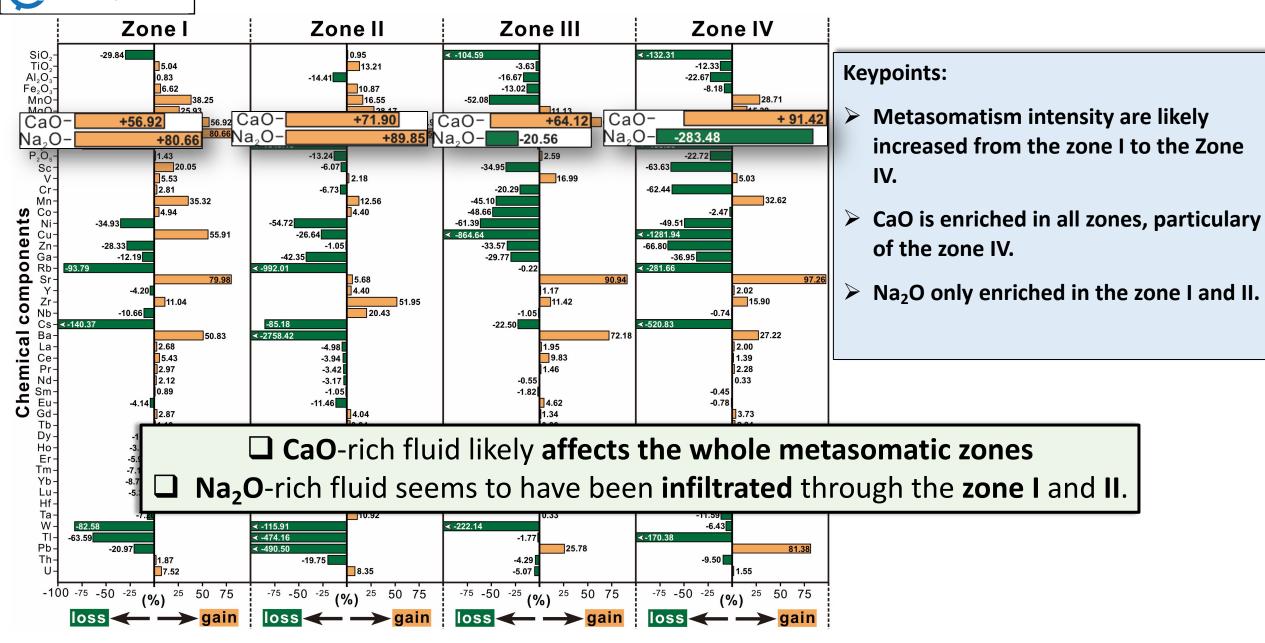
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Comparison gain or loss of chemical components (average)

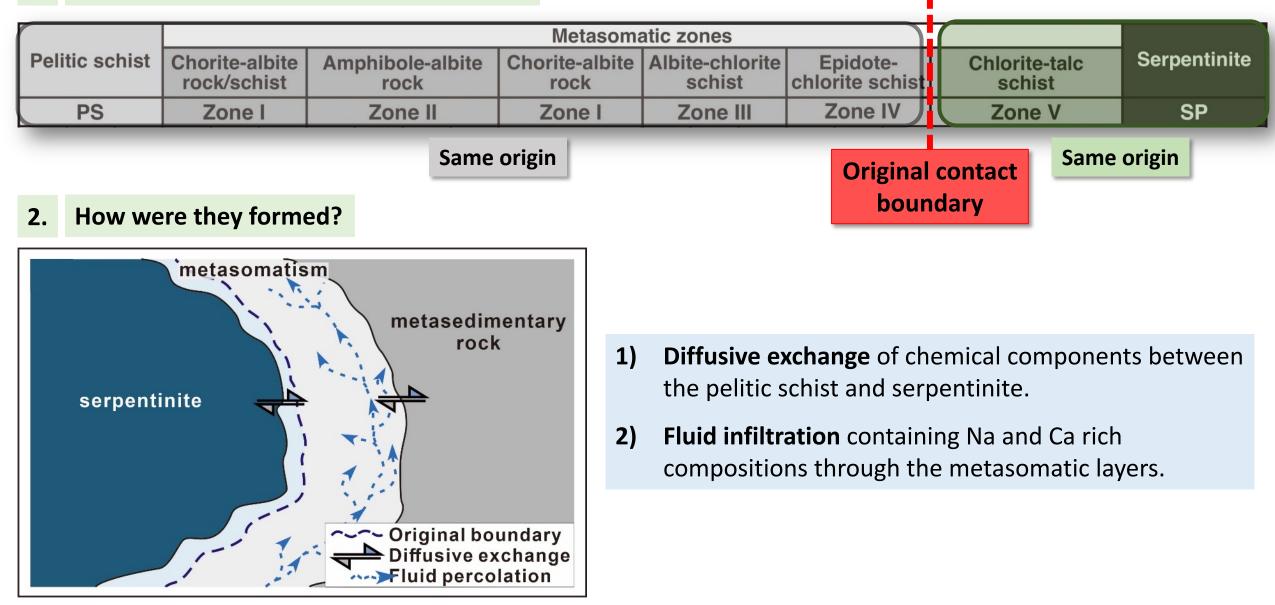






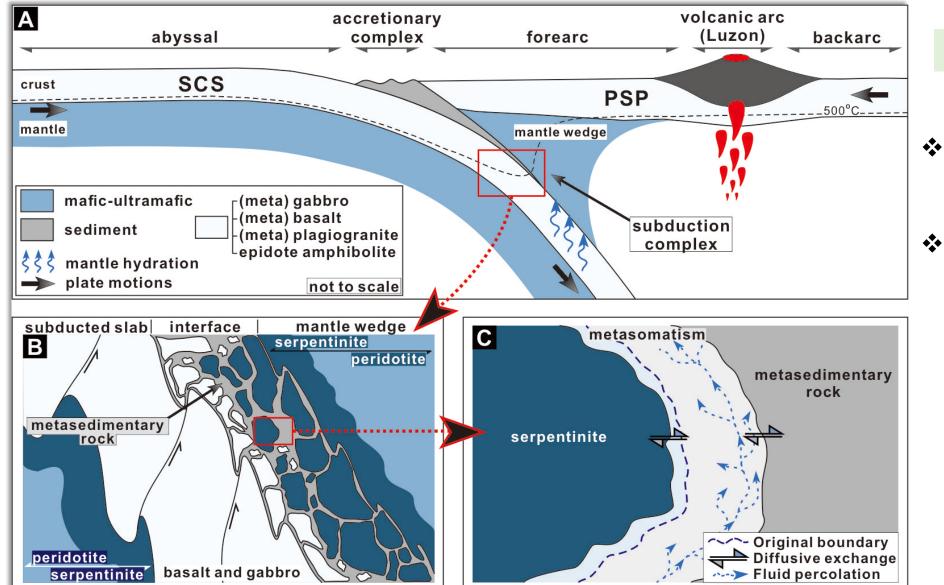
Discussion (1)

1. What is the origin of metasomatic rocks?



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Discussion (1)



3. Tectonic implications ?

- Metasomatic rocks indicate materials from the slabmantle interface.
- Metasomatic rocks record indications of fluid-rock interactions during subduction metamorphism.



Conclusions

- Pelitic schist and metasomatic zones I IV are from the same origin, whereas zone V is from the serpentinite origin. The original contact boundary is between zone IV and V.
- **2. Diffusive exchange** and **fluid infiltration** (Na and Ca rich) are responsible for the formation of the metasomatic zones.
- 3. Metasomatic rocks record evidence of **fluid-rock interactions** at the **slabmantle interface** during **subduction metamorphism**.

Reference list

