



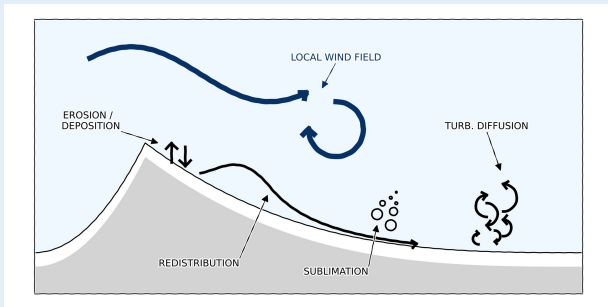
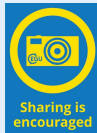
EGU General Assembly, CR6.2, 24 April 2023

Simulating snow drift in WRF

First results and future plans of a novel module

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Get the abstract:
EGU23-12384



Get the model:

Code available at : <https://github.com/manuelsaigger/WRFsnowdrift>

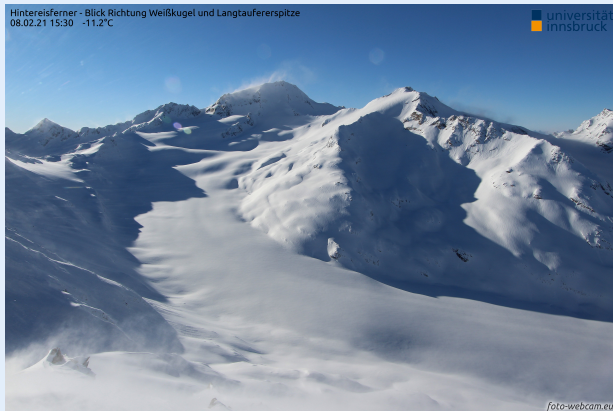
Schmid, C., 2021: Implementierung eines Schneerdriftmoduls in das Weather Research and Forecasting (WRF) Modell und eine erste Evaluierung, Doctoral thesis, Friedrich-Alexander-Universität Erlangen-Nürnberg.



Friedrich-Alexander-Universität
Erlangen-Nürnberg

WHY STUDY SNOW DRIFT?

- ▶ Redistribution of snow → influence on glacier mass balance
- ▶ Interaction of drifting snow with ambient flow field.



Model Components

ATMOSPHERE

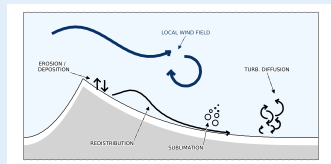
SNOW

EROSION

DEPOSITION

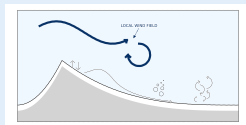
SUBLIMATION

ADVECTION + TURB. DIFFUSION



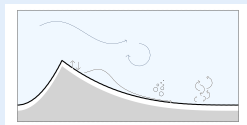
Atmospheric Model

- ▶ Snow drift module included as a physical parameterization into the Weather Research and Forecasting Model (WRF, Skamarock et al. 2019).
- ▶ Tested for WRF versions 4.1 to 4.3, idealized and real-case applications, meso-scale and LES simulations.



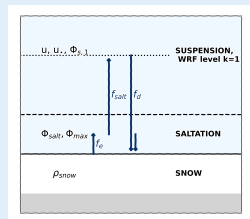
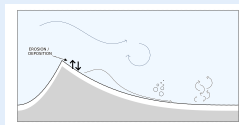
Snow Module

- ▶ Use snow model of land-surface model Noah-MP (Niu et al. 2011).
- ▶ Three-layer snow
- ▶ Solves energy and mass balance of snow (processes: precipitation, heat conduction, melting, liquid-water storage and transport, refreezing of meltwater, compaction, interception)
- ▶ Interaction with drifting snow as a function of snow density.



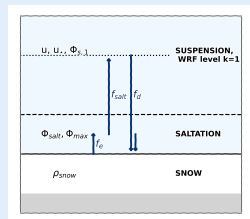
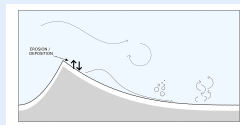
Erosion Module

- ▶ Divide drifting snow into saltation layer and suspension
- ▶ Erosion depending on excess shear stress
→ flux into saltation layer f_e where $u_* > u_{th}$ (Gallée et al. 2001; Sauter et al. 2013)
- ▶ Threshold friction velocity u_{th} following Walter et al. (2004) as function of snow density
- ▶ Bulk transfer approach for flux to suspension f_{salt} (into lowest mass level of WRF) following Gallée et al. (2001) and Sauter et al. (2013)
- ▶ Exact equations [here](#)



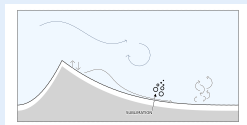
Deposition Module

- ▶ Particle fall speed similar to Vionnet et al. (2014)
→ spherical particles, balance between gravity and drag force
- ▶ Particle size depending on height and friction velocity (Pomeroy 1988; Gordon et al. 2010)
- ▶ Deposition where $u_* < u_{th}$ similar to Sauter et al. (2013)
- ▶ Optional: define constant fall speed or particle size at the surface
- ▶ Exact equations [here](#)



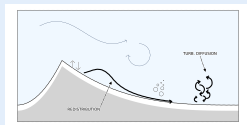
Sublimation Module

- ▶ Sublimation following Thorpe and Mason (1966)
→ solve energy and mass balance for single particle, transfer to ensemble of particles (similar to, e.g., Sauter et al. 2013; Vionnet et al. 2014)
- ▶ Feedback to temperature and moisture fields: energy required for sublimation is taken from surrounding air (→ cooling), water vapor is given to specific humidity
- ▶ optional: switch off sublimation, switch off sublimation feedback



Advection and Turbulent Diffusion

- ▶ Snow particles treated as passive tracers (subsidence treated in deposition module)
→ use WRF-internal modules for advection and diffusion.
- ▶ possible to use all available schemes and options for advection and turbulent diffusion in WRF



Equations of Erosion and Deposition

$$f_e = e_{salt} \rho_{air} \left(\left(u_* + (u_{th} - u_*) \left(\frac{\Phi_{salt}^2}{\Phi_{max}^2} \right) \right)^2 - u_{th}^2 \right)$$

$$u_{th} = 0.0195 + (0.021 \sqrt{\rho_{snow}})$$

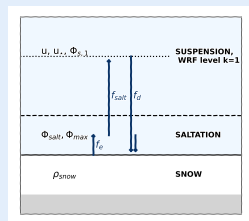
$$\Phi_{max} = \frac{\rho_{air}}{3.29 u_*} \left(1 - \frac{u_{th}^2}{u_*^2} \right)$$

$$f_{salt} = u C_d (\Phi_{salt} - \Phi_{s,1})$$

$$v_p = \frac{-A}{r} + \sqrt{\left(\frac{A}{r} \right)^2 + B r}; A = \frac{6.203 \nu_{air}}{2}; B = \frac{5.516 \rho_{ice} g}{4 \rho_{air}}$$

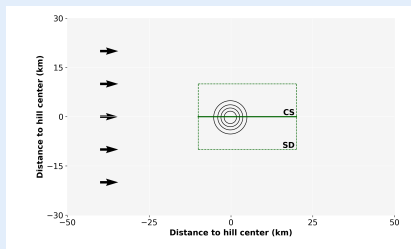
$$r(z) = r_0 z^{-0.258}; r_0 = 0.5 \left(\frac{7.8 \cdot 10^{-6} u_*}{0.036} + 3.1 \cdot 10^{-5} \right)$$

$$f_d = v_{p,1} \Phi_{s,1} \cdot \max \left(\frac{u_{th}^2 - u_*^2}{u_{th}^2}, 0 \right)$$



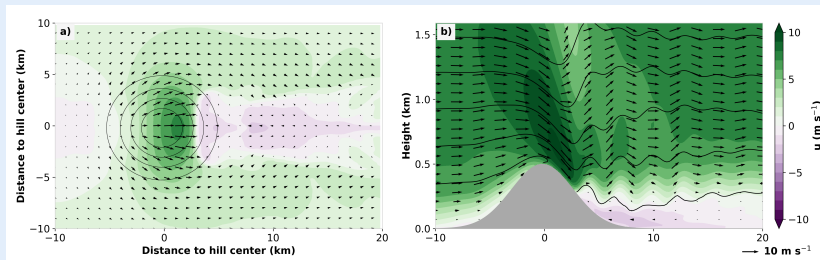
Setup Test Simulations

- ▶ WRF ARW version 4.1.2
- ▶ Domain with one bell-shaped hill in the center ($h = 500$ m)
- ▶ $\Delta x = \Delta y = 200$ m
- ▶ Lowest mass level at 5 m agl, model top at 10000 m, $\Delta z = 10$ m close to surface.
- ▶ Initialize with constant profiles of $u = 7.5 \text{ m s}^{-1}$ (log profile in lowest 100 m), and $N = 0.015 \text{ s}^{-1}$
→ non-dimensional mountain height $\epsilon = 1$
- ▶ $rh = 80$ % close to the ground, decreasing with height
- ▶ 1 m of snow with $\rho_{snow} = 100 \text{ kg m}^{-3}$ over entire domain
- ▶ total model runtime 6 h, use 3 to 6 h for evaluation (steady-state after spin-up)



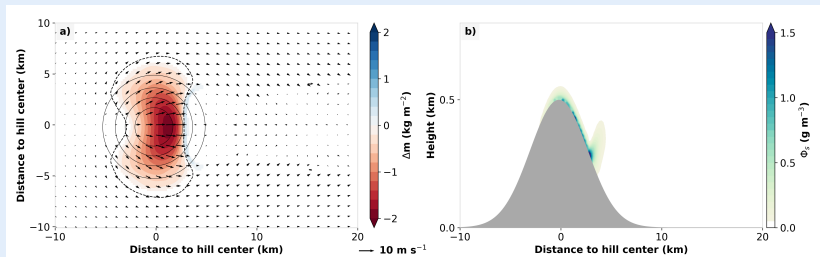
Test Simulations: Wind Fields

- ▶ mean fields of wind and potential temperature (b, contour interval 1 K)



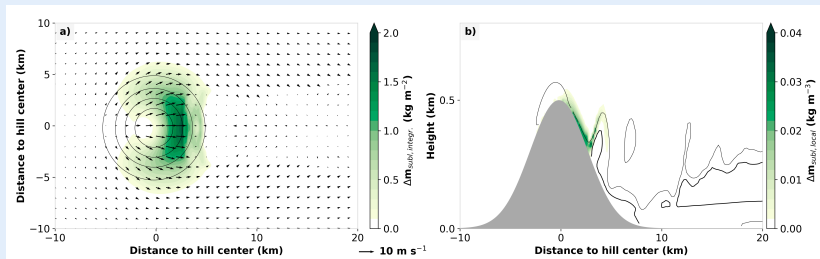
Test Simulations: Drifting Snow Fields

- ▶ Accumulated erosion and deposition (a), mean snow particle concentration (b)
- ▶ Erosion in regions of highest wind speed, where $u_* > u_{th}$ (dashed contour).
- ▶ Snow particles stay in shallow layer, deposition in region of flow separation.



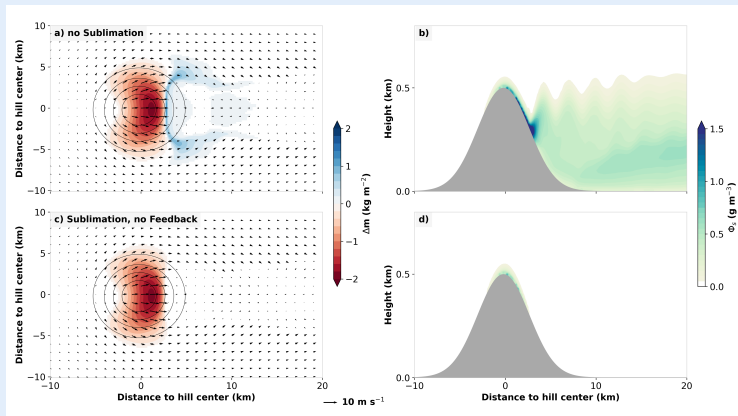
Test Simulations: Drifting Snow Sublimation

- ▶ Accumulated sublimation (a: vertically integrated, b: local), mean relative humidity (b, thin/thick contours 95/99%)
- ▶ Sublimation removes large parts of eroded snow
- ▶ Damping due to increased air moisture



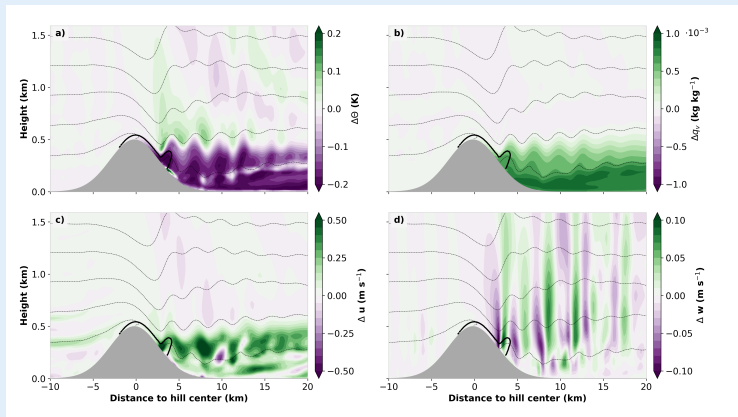
Test Simulations: Influence of Sublimation

- no sublimation: stronger deposition + lee-ward transport; no sublimation feedback: no damping of sublimation due to moisture increase → less deposition



Test Simulations: Non-local Effect of Sublimation

- ▶ Difference in mean fields for control simulation and with switched off sublimation
- ▶ local cooling and moistening due to sublimation, but also non-local effect downstream and influence on gravity wave field



Future Plans

How does drifting snow influence glacier mass balance?

- ▶ Case studies with Large-Eddy-Simulations and terrestrial laser scanning on Hintereisferner (Austria)
→ see Voordendag and Goger (PICO3a.3, EGU23-5236)
- ▶ (Multi-) seasonal perspective → incorporate machine-learning techniques

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