

# On the implications of a warm bias in modelling an eddying Southern Ocean

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## 1 Background

- High wind speeds and the zonally unbounded extent of the Southern Ocean (SO) evoke high mean and eddy kinetic energy
- SO eddies are anticipated to play a key role for the global heat and carbon budgets and the redistribution of Antarctic melt water
- Computational limits still prevent the use of global eddying ocean models for studies concerning long-term climate projections, asking for nested model setups
- Common SO biases in CMIP-type models, such as in temperature, complicate the initialisation of regional high-resolution nests
- Here, we test three different initial states in a nested climate model with regard to their effect on SO dynamics

## 2a A global climate model with a nested SO

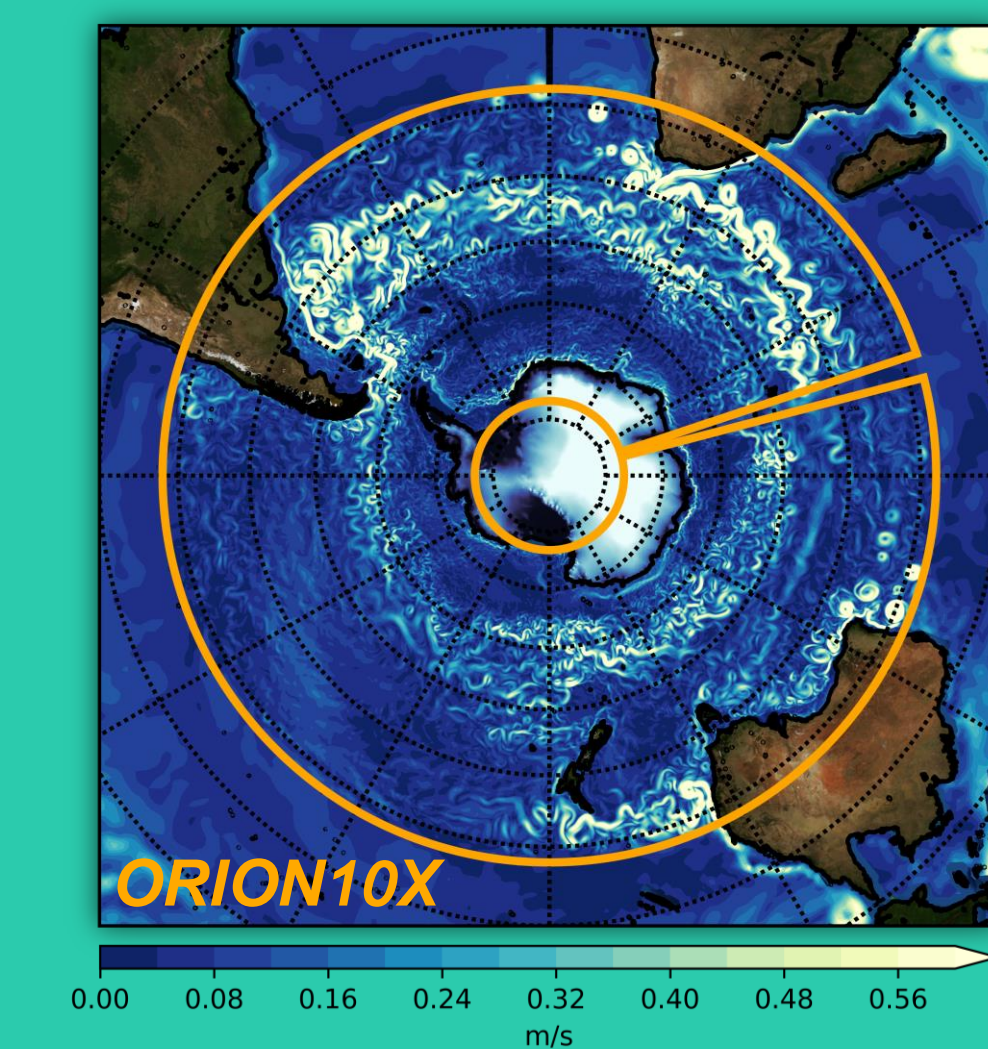


Figure 1: Snapshot of 5-daily mean surface velocity. The SO nest is outlined in orange.

The Flexible Ocean and Climate Infrastructure (FOCI) model:

- NEMO3.6-L46 ocean + LIM2 sea ice
- ECHAM6.3-T63 atmosphere + JSBACH land
- OASIS3-MCT coupler

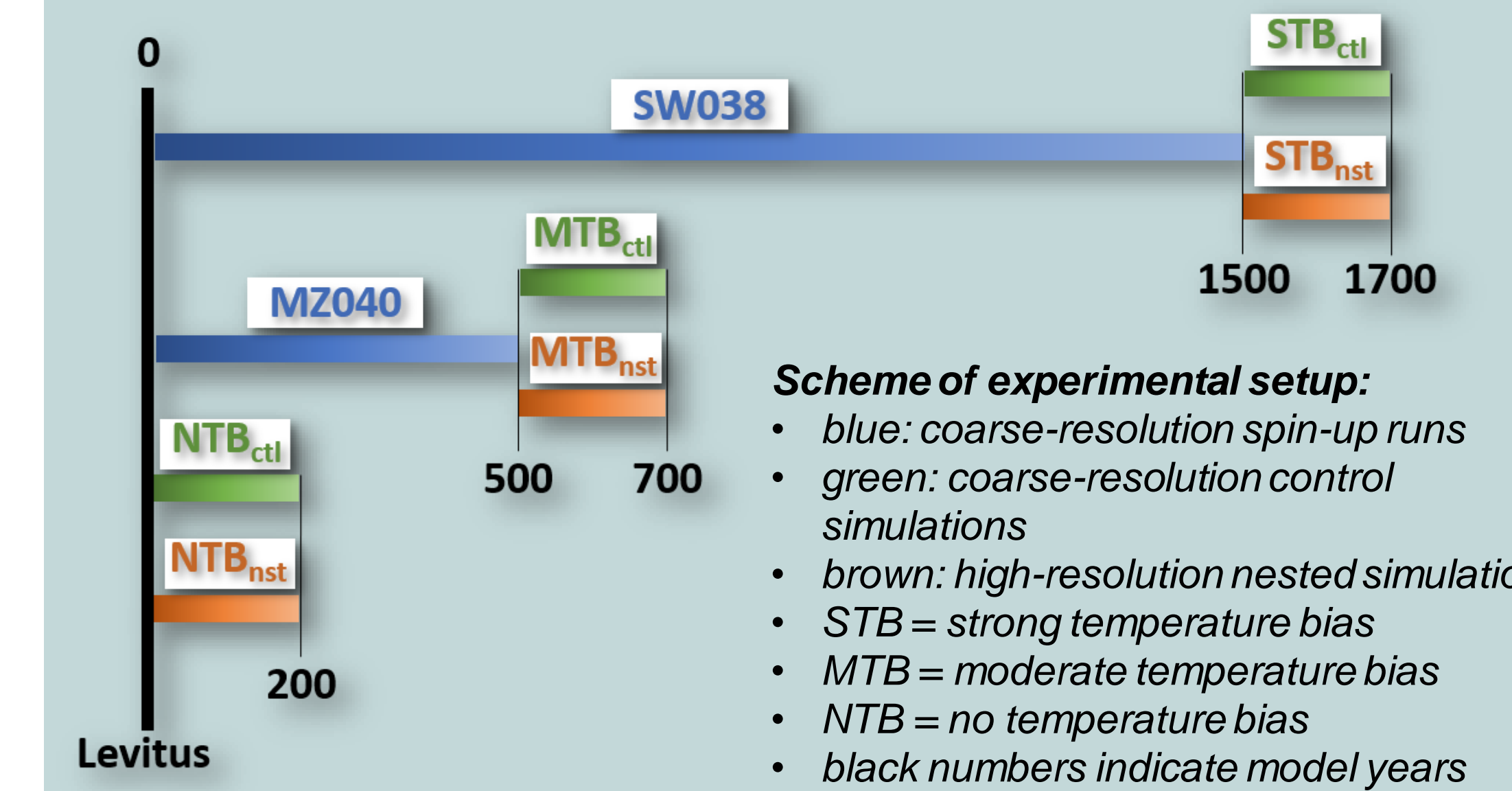
Ocean component (NEMO):

- 0.5° horizontal resolution, 46 z-levels

ORION10X nest (Figure 1):

- high-resolution ocean nest (0.1°) embedded south of 28°S applying AGRIF
- 2-way exchange with global ocean
- fully coupled

## 2b Experimental setup



## 3 Warm bias in the Southern Ocean

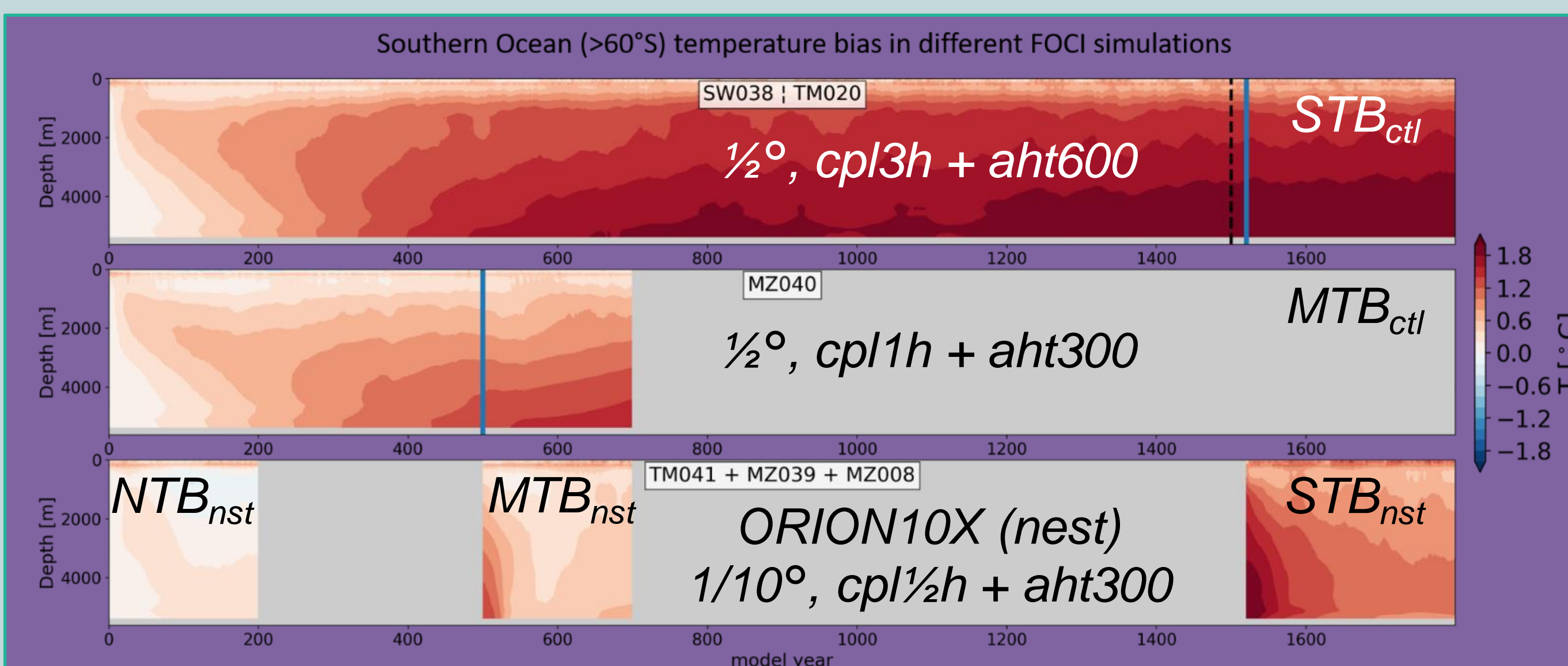


Figure 2: Temperature bias relative to World Ocean Atlas 2013 over depth and time for the standard non-nested control simulation (top panel), the non-nested control simulation with modified mixing parameters and coupling time step (middle panel), and the three nested simulations initialised from different spin-up conditions (bottom panel). The data is averaged over the Southern Ocean south of 60°S.

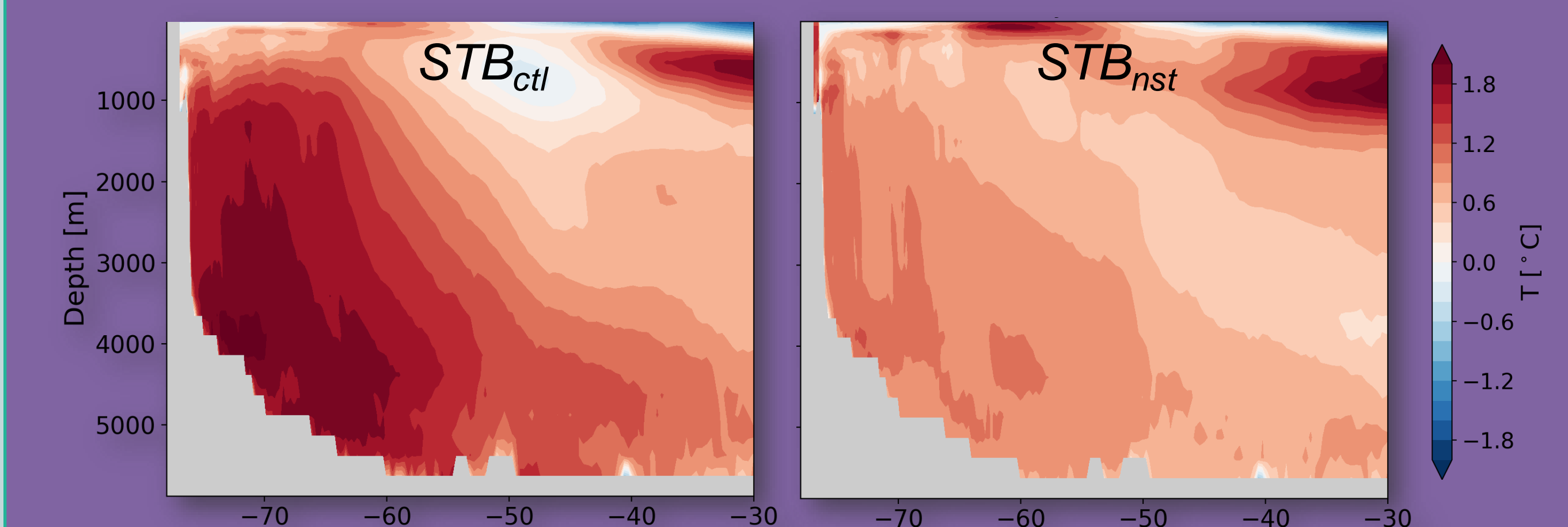


Figure 3: Zonal average temperature bias relative to Levitus for  $STB_{ctl}$  (left panel) and  $STB_{nst}$  (right panel). The data is averaged over the last 50 years of each simulation.

- build-up of warm bias over time in coarse-resolution FOCI spin-up runs (Figure 2, top)
- warm bias reduced by increasing ocean-atm. coupling frequency from 3 hours to 1 hour and reducing iso-neutral diffusion from 600 m<sup>2</sup>/s to 300 m<sup>2</sup>/s (Figure 2, middle)
- increasing horizontal resolution from 0.5° to 0.1°, thus including mesoscale dynamics, results in vast decline of warm bias within 2-3 decades (Figure 2, bottom)
- initialising nest from observed conditions does not entail a heat build-up (Figure 2, bottom)

- SO warm bias in FOCI (Figure 3, left) causes unstable conditions when branching off the nested ORION10X simulation
- ... causing massive open ocean deep convection (Figure 4), which releases the excess heat to the atmosphere
- eddying simulation yields improved temperature distribution, esp. in the deep ocean (Figure 3, right)

## 4 Implications on dynamics

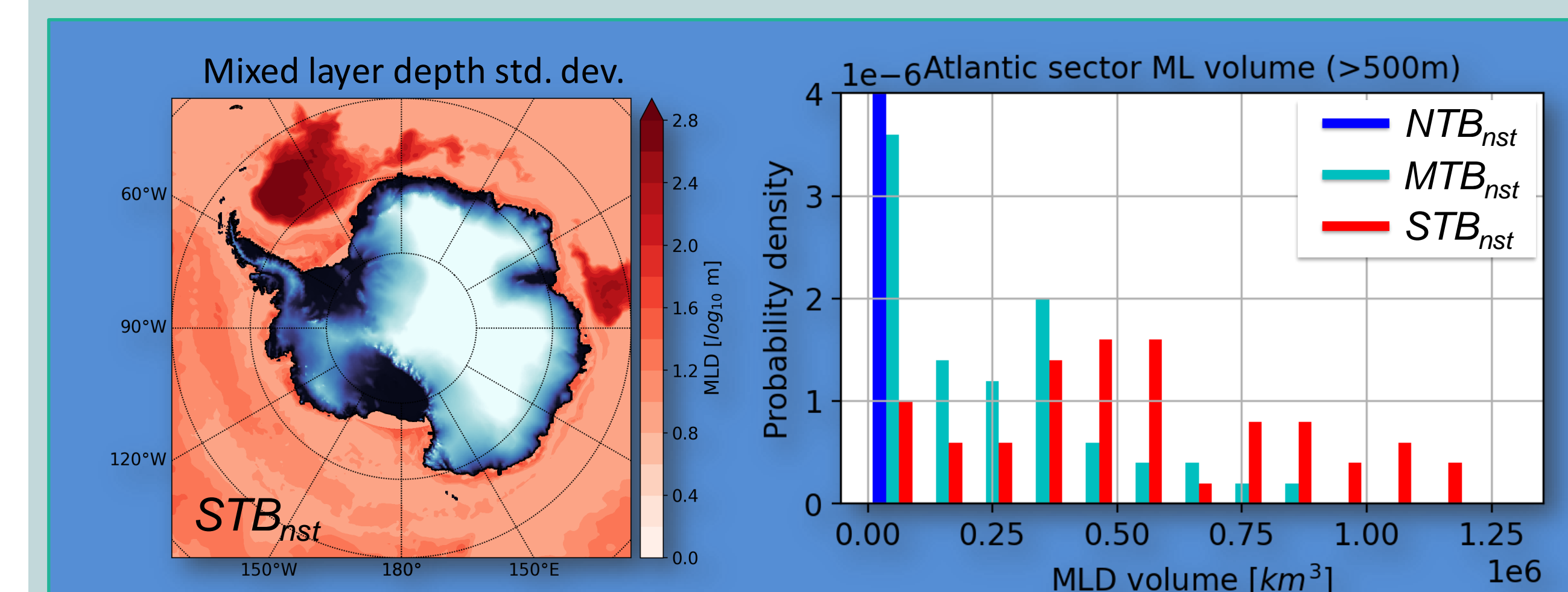


Figure 4: Left: Annual standard deviation of mixed layer depth for  $STB_{nst}$  simulation. Note that the colour bar is on log scale. Right: Histogram of annual Atlantic sector mixed layer volume (>500m) for the three high-resolution nested simulations.

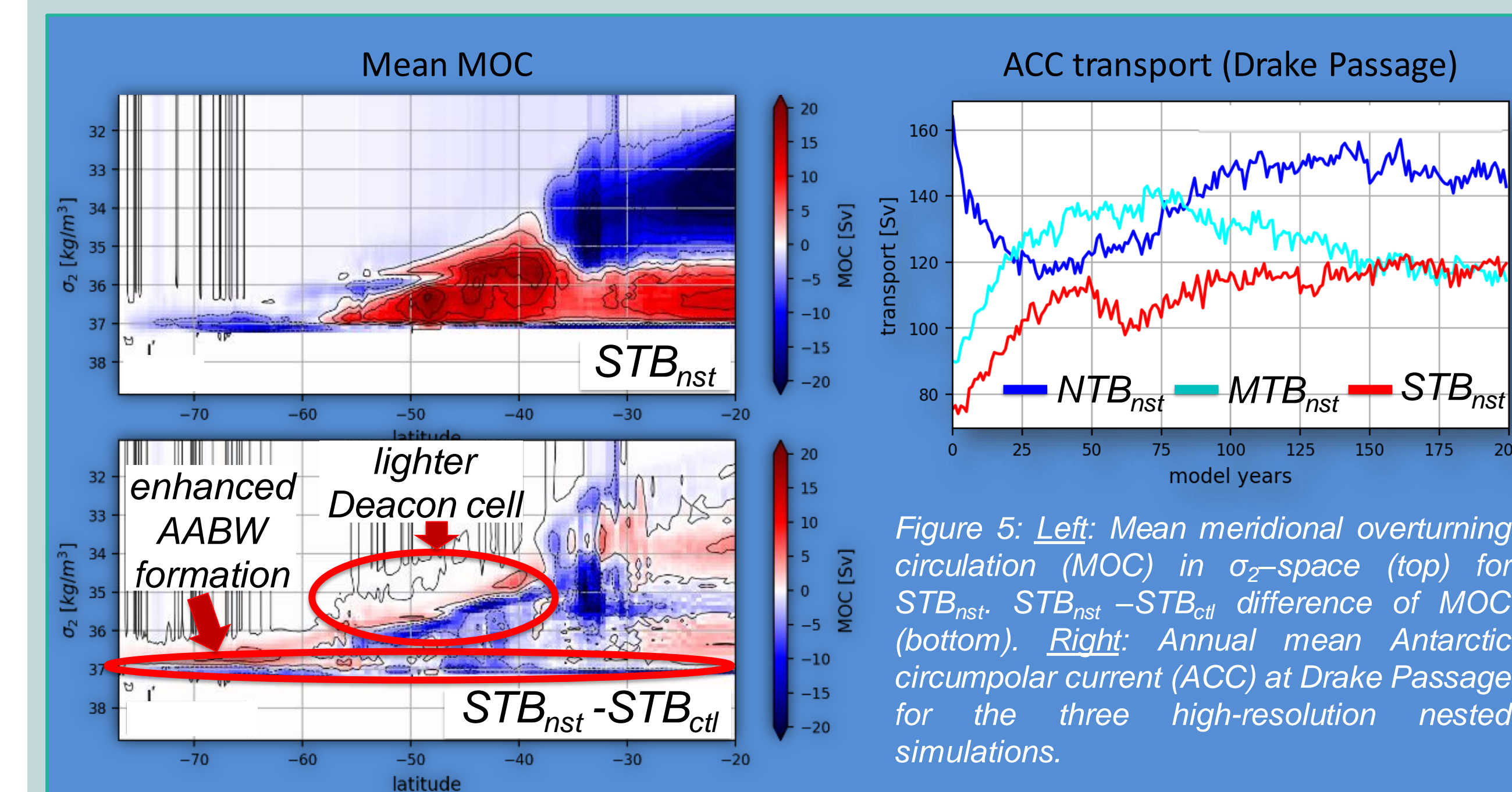


Figure 5: Left: Mean meridional overturning circulation (MOC) in  $\sigma_t$ -space (top) for  $STB_{nst}$ ,  $STB_{nst} - STB_{ctl}$  difference of MOC (bottom). Right: Annual mean Antarctic circumpolar current (ACC) at Drake Passage for the three high-resolution nested simulations.

## 5 Conclusions

- nesting reduces warm bias in the Southern Ocean
- initialising a high-resolution nest from a biased SO leads to baroclinic instabilities due to different dynamics (resolved vs. parameterized processes)
- these instabilities initiate massive deep convection, in turn inducing spurious anomalies in the global circulation
- best practice is to initialise a SO nest from observed conditions

- Deep convection in the nest mainly occurs in Weddell Gyre & Amery basin (Fig. 4, left)
- Enhanced spurious deep convection with increasing warm bias (Figure 4, right)
- Deep convection leads to modifications of the horizontal (ACC, Figure 5, right) and overturning (MOC, Fig. 5, left) circulations → high production rate of Antarctic bottom water (AABW,  $STB_{nst}$ : 13 Sv  $\pm$  5 Sv) → upper branch of Deacon cell becomes lighter in nested simulations