

A new high resolution stable isotope record from the North Atlantic Ocean: a detailed insight into the mid-Maastrichtian event

Alexa Fischer^{1,*}, Oliver Friedrich¹, André Bahr¹, Silke Voigt²

¹ Institute of Earth Sciences, Heidelberg University, Heidelberg, Germany

² Institute of Geosciences, Goethe University, Frankfurt, Germany

*Contact: alexa.fischer@geow.uni-heidelberg.de

GEOW

INSTITUT FÜR
GEOWISSENSCHAFTEN



UNIVERSITÄT
HEIDELBERG
ZUKUNFT
SEIT 1386

1. Introduction

The long-term global cooling trend during the latest Cretaceous was interrupted by an intense global warming episode at ~69 Ma known as the mid-Maastrichtian event (MME)^[1,2,3]. The MME is characterized by two positive $\delta^{13}\text{C}$ excursions with an overall magnitude of 0.6‰ to 1.5‰ separated by a negative inflection^[3]. The $\delta^{13}\text{C}$ excursions are accompanied by the extinction of inoceramid bivalves^[4] and an abrupt increase in deep-sea and sea-surface temperatures^[5,3]. Changes in ocean circulation, particularly a change in thermohaline circulation patterns, have been proposed to be one of the main drivers of the MME^[6]. Nevertheless, the driving mechanisms, timing, character, and consequences of the circulation change are still up for debate.

2. Aims

- ▶ To understand the climatic patterns causing the MME
- ▶ To unravel the drivers of the potential circulation change during the MME
- ▶ To further assess the biotic and climatic effects of the MME

3. Material and Methods

- IODP Core U1403 in the North Atlantic (J-Anomaly Ridge) (Fig. 3)
- 2 Myr-long time interval of the mid to late Maastrichtian has been analyzed at a ~2.5 to 5 kyr-resolution
- XRF core scanning
- wt% CaCO_3 analyses
- Stable oxygen and carbon isotope measurements of benthic foraminifera (*Nuttallides truempyi*)
- Bottom-water temperature reconstruction through Mg/Ca measurements of the same foraminiferal tests (*Nuttallides truempyi*)

4. Results

- MME lasted ~600 kyr
- Bottom-water temperature increase up to 4 °C (Fig. 2)
- Several CaCO_3 dissolution events (Fig. 2)
- High percentage of infaunal benthic foraminifera species
→ oxygen-poor bottom waters during the MME (Fig. 1)

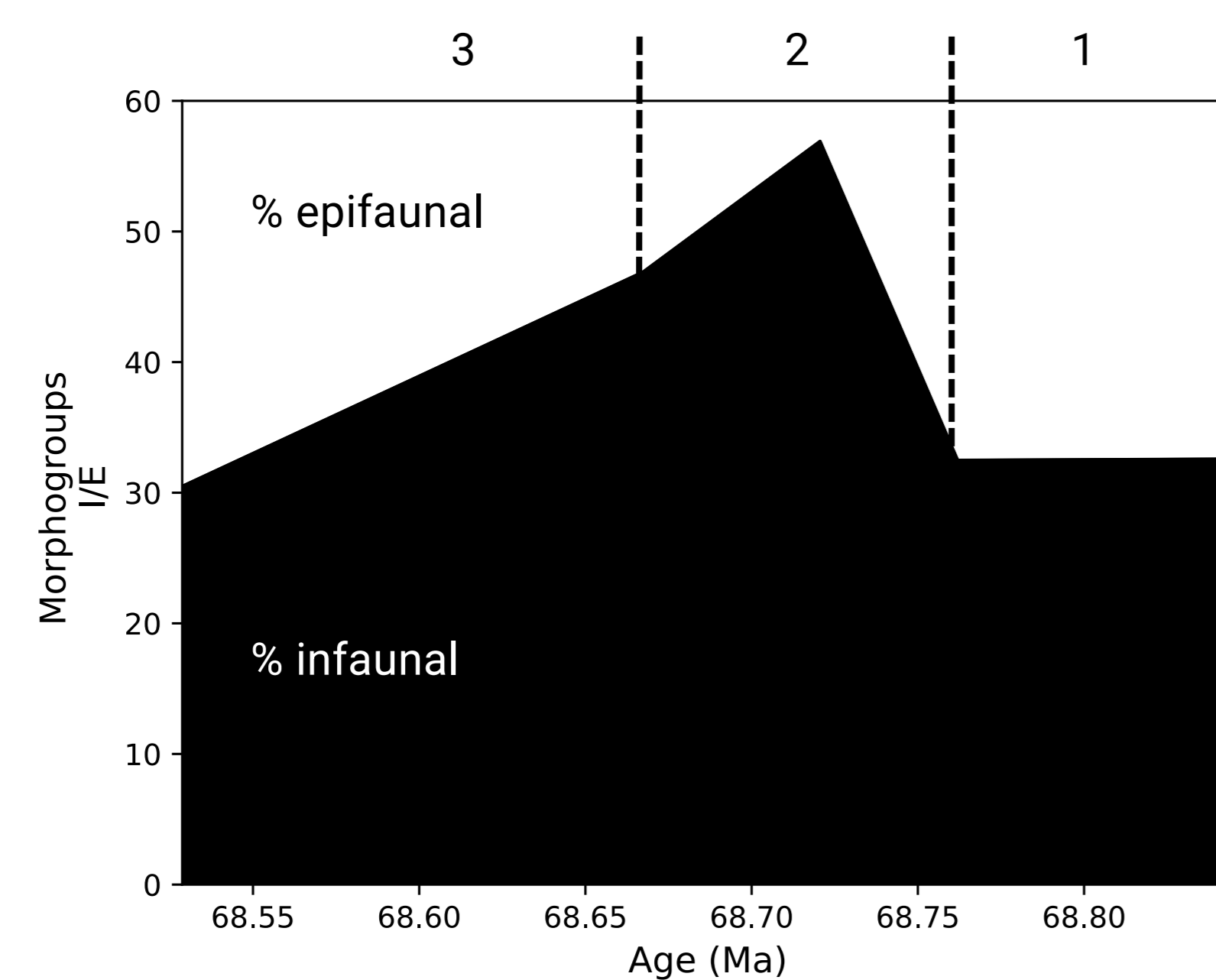


Fig. 1: Preliminary benthic foraminiferal infaunal-epifaunal ratios across the MME. Numbers indicate MME Peaks 1-3.

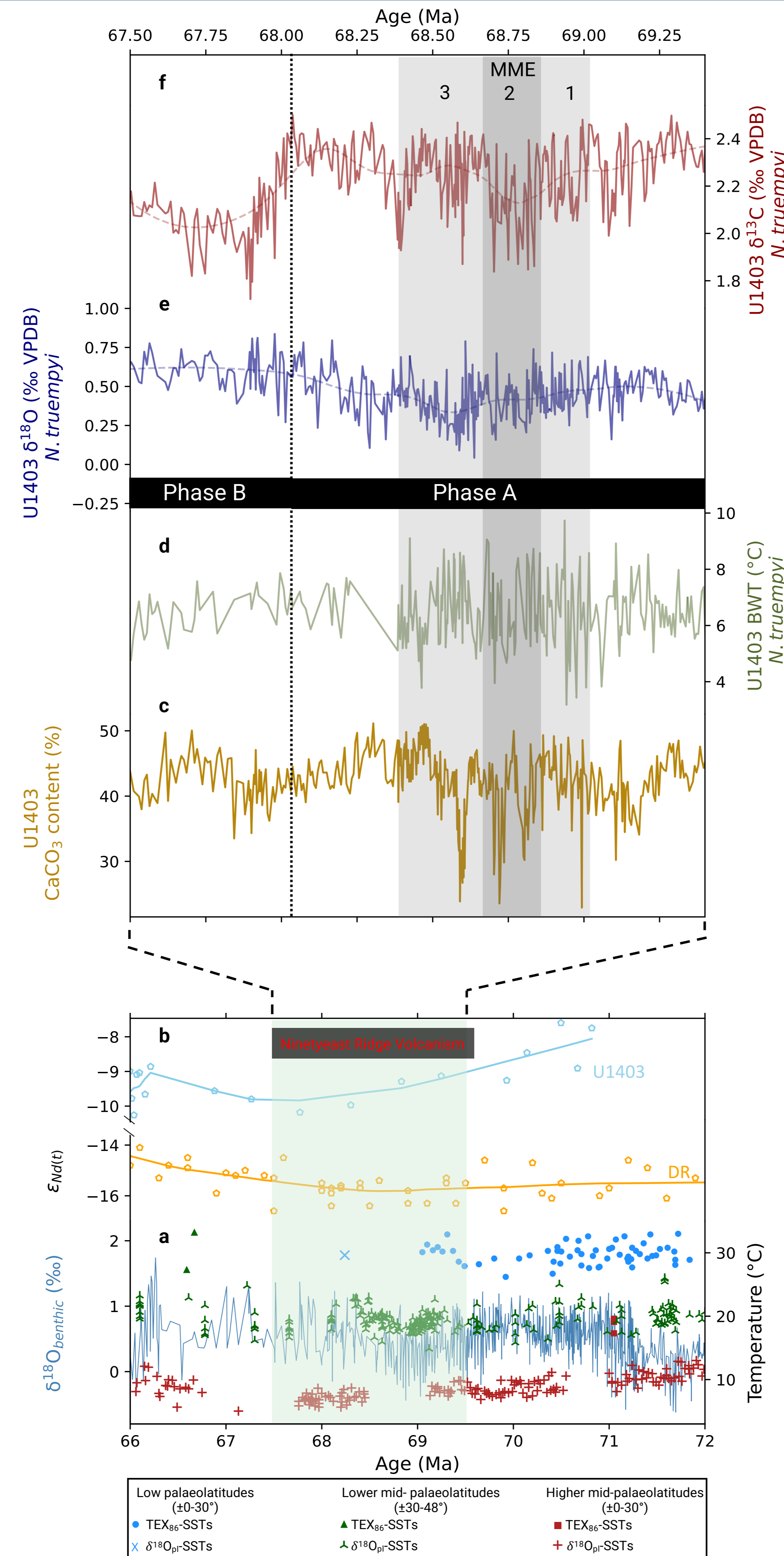


Fig. 2: High-resolution records across the latest Cretaceous of U1403. Light green shaded box, study interval for this work. a, Deep sea benthic foraminifer composite $\delta^{18}\text{O}$ data for the Maastrichtian from multiple sites^[7] and compiled $\text{TEX}_{86}^{\text{H}}$ and $\delta^{18}\text{O}_{\text{p}}$ SST^[8]. b, Nd-isotope data, in yellow: Demera Rise (DR)^[9] and in blue: U1403^[10]. c, CaCO_3 wt(%) content. d, Mg/Ca-based bottom-water temperatures. e, Benthic foraminifera (*N. truempyi*) $\delta^{18}\text{O}$ (‰, VPDB). f, Benthic foraminifera (*N. truempyi*) $\delta^{13}\text{C}$ (‰, VPDB). Ninetyeast hotspot volcanism adapted from [12]. MME: mid-Maastrichtian event.

5. Conclusion

Ocean circulation phases

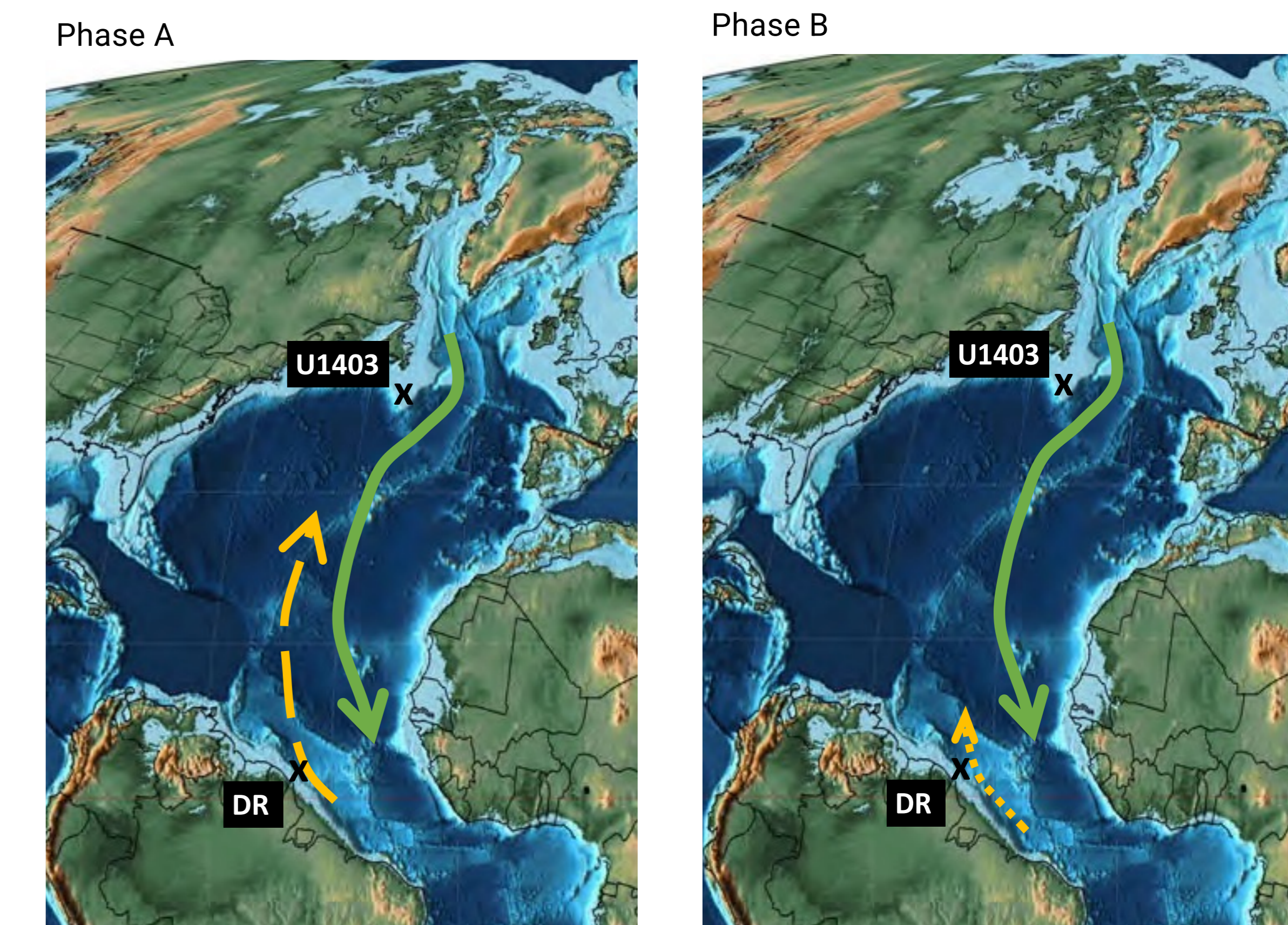


Fig. 3: Schematic representation of the proposed North Atlantic circulation patterns during the late Maastrichtian (68 Ma). Green arrow: waters from the high-latitudes. Yellow dashed arrow: waters from the low-latitudes DR: Demera Rise. Paleomap via [13].

- Interplay between waters from the high- and low-latitudes between 69.25 and 68 Ma (Fig. 2, Fig 3 Phase A)
 - ▶ Alternating warm and cold bottom-water temperatures
 - ▶ $\delta^{13}\text{C}$ fluctuations
 - ▶ CaCO_3 dissolution events
- Northern-sourced water mass as main deep-water source starting around ~68 Ma (Fig. 2, Fig. 3 Phase B)
 - ▶ Stable bottom-water temperatures
 - ▶ Stable and lighter $\delta^{13}\text{C}$ values
 - ▶ No CaCO_3 dissolution events

References:

- Dvorik, S. I., Nordt, L., & Atchley, S. (2005). Determining terrestrial paleotemperatures using the oxygen isotopic composition of pedogenic carbonate. *Earth and Planetary Science Letters*, 237(1-2), 56-68. <https://doi.org/10.1016/j.epsl.2005.06.054>
- Nordt, L., Atchley, S., & Dvorik, S. (2003). Terrestrial Evidence for Tertiary Greenhouse Events in the Latest Cretaceous. *GSA Today*, 13(12), 4. [https://doi.org/10.1130/1052-5173\(2003\)013<4:TEFGES>2.0.CO;2](https://doi.org/10.1130/1052-5173(2003)013<4:TEFGES>2.0.CO;2)
- Voigt, S., Gals, A., & Jung, C. (2012). Global correlation of upper Cretaceous-Maastrichtian successions using carbon-isotope stratigraphy: Development of a new Maastrichtian timescale. *Newsletters on Stratigraphy*, 45(1), 25-53. <https://doi.org/10.1127/0078-0421/2012/0016>
- MacLeod, K. G. (1994). Bioturbation, inoceramid extinction, and mid-Maastrichtian ecological changes. *Geology*, 22(2), 139. [https://doi.org/10.1130/0091-7613\(1994\)022<0139:BIEMM>2.3.CO;2](https://doi.org/10.1130/0091-7613(1994)022<0139:BIEMM>2.3.CO;2)
- Keller, G. (2001). The end-Cretaceous mass extinction in the marine realm: year 2000 assessment. *Planetary and Space Science*, 49(8), 817-830. [https://doi.org/10.1016/S0032-0633\(01\)00032-0](https://doi.org/10.1016/S0032-0633(01)00032-0)
- MacLeod, K. G., & Huber, B. T. (2001). The Maastrichtian record at Blake Nose (western North Atlantic) and implications for global palaeoceanographic and biotic changes. *Geological Society, London, Special Publications*, 189(1), 111-130.
- Friedrich, O., Norris, R. D., & Eisecher, J. (2017). Evolution of middle to late Cretaceous oceanic $\delta^{18}\text{O}$ SSTs: Record of Earth's temperature and carbon cycle. *Geology*, 45(2), 107-110. <https://doi.org/10.1130/G42701.1>
- O'Brien, C. L., Robinson, S. A., Pancost, R. D., Stinnette Damsted, J. S., Schouten, S., Lunt, D. J., Aksen, H., Bornemann, A., Böttini, C., Brassell, S. C., Farnsworth, A., Foster, A., Huber, B. T., Inglis, G. N., Jenkyns, H. C., Linnert, C., Littler, K., Markwick, P., McAnena, A., ... Vribe, N. E. (2017). Cretaceous sea-surface temperature evolution: Constraints from TEX_{86} and planktonic foraminiferal oxygen isotopes. In *Earth-Science Reviews* (Vol. 172, pp. 224-247). Elsevier B.V. <https://doi.org/10.1016/j.earscrv.2017.07.012>
- MacLeod, K. G., Izaola Londoño, C., Martin, E. E., Jiménez Berrocoso, A., & Basak, C. (2011). Changes in North Atlantic circulation at the end of the Cretaceous greenhouse interval. *Nature Geoscience*, 4(11), 779-782. <https://doi.org/10.1038/ngo1284>
- Balenburg, S. J., Friedrich, O., Moriya, K., Voigt, S., Coumède, C., Blum, P., Bornemann, A., Fiebig, J., Hasegawa, T., Hull, P. M., Norris, R. D., Röhl, U., Sexton, P. F., Vesterhold, T., & Wilson, P. A. (2018). Late Maastrichtian carbon isotope stratigraphy and cyclostratigraphy of the Newfoundland Margin (Site U1403, IODP Leg 342). *Newsletters on Stratigraphy*, 51(2), 245-260. <https://doi.org/10.1127/nos/2017/0398>
- Balenburg, S. J., Voigt, S., Friedrich, O., Obovne, A. H., Bornemann, A., Klein, T., Peter-Delz, L., & Frank, M. (2018). Major intensification of Atlantic overturning circulation at the onset of Paleogene greenhouse warmth. *Nature Communications*, 9(1). <https://doi.org/10.1038/s41467-018-07457-7>
- Keller, G., Purnet, J., & Haseo, P. (2016). Upheavals during the Late Maastrichtian: Volcanism, climate and faunal events preceding the end-Cretaceous mass extinction. *Paleogeography, Paleoclimatology, Paleocology*, 441, 137-151. <https://doi.org/10.1016/j.palaeo.2015.06.034>
- Scotese, C. R. (2014). Atlas of Late Cretaceous Maps, PALEOMAP Atlas for ArcGIS, volume 2, The Cretaceous, Maps 16 - 22, Mollweide Projection, PALEOMAP Project, Evanston, IL.

