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POLAR LOWS IN A GLOBALLY COUPLED STORM AND EDDY RESOLVING (2.5km) CLIMATE MODEL (ICON-Sapphire) – supplementary material (case study 2)



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NOAA satellite image of comma-cloud polar low over the Barents Sea east of the Svalbard Archipelago. 1 February 2015.

Source: NERC Satellite Receiving Station, Dundee University, Scotland.



## **MOTIVATION**

#### What are polar lows?

Polar lows (PLs) are the most intense cyclones of the polar mesoscale cyclone (PMC) family, with subsynoptic scales of less than 1000 km and near-surface wind speeds of more than 15 m/s, forming over high-latitude maritime environments poleward of the polar front zone. [after Heinemann and Claude (1997)]

#### Why are polar lows important?

- Danger to shipping/air traffic, coastal communities and offshore installations
  - $\rightarrow$  forecasting with atmosphere-only models (formation processes)
- Effect on ocean/sea ice and climate?
  - $\rightarrow$  coupled global models needed
  - $\rightarrow$  Problem: kilometer-scale resolution required for atmosphere, ocean and sea ice to capture PLs



A mature polar low off the coast of Finnmark in Northern Norway from late March in 2014. Illustration: MET-Norway/NOAA (modified)



# ICON-SAPPHIRE 2.5KM GLOBALLY COUPLED CONFIGURATION (ICON2.5)

Parameter	ICON2.5
horizontal resolution	r2b10 (2.5  km)
# vertical levels (atm/oce)	90/112
$\Delta$ z-levels (oce)	6  to  532 m
$\Delta t \ (atm/oce)$	$20 \mathrm{s}/80 \mathrm{s}$
coupling frequency	$12 \min$
simulation period	2020-01-20 to 2020-03-31 (71 d)
output volume	$\sim 340 \mathrm{TB} \;(135 \mathrm{TB/month})$
output frequency	2d-atm. (30 min), 3d-atm. (1d),
	2d-oce (1h,3h), 3d-oce <200 m (3h), 3d-oce (1d)

see more details in Hohenegger et al. (2023)

#### We focus on two case studies:

- Iceland/Greenland Sea (see presentation)
- Labrador Sea (supplementary material)

**ICON2.5** snapshot (SST and sea ice concentration)





## CASE STUDY: POLAR LOW OVER LABRADOR SEA



#### PL forms at sea ice edge in Labrador Sea on 21 February 2020

- PL formed from CAO off coast from Labrador, is steered northward and rapidly intensifies at sea ice edge
- PL reaches mature stage within 24 hours (core pressure of 944 hPa) with a warm core
  - Hurricane force winds (up to 34 ms<sup>-1</sup> at sea ice edge)
  - Strong cold air outbreak (CAO) south of PL core
- PL moves east south of Cape Farewell into Irminger Sea within next 24 hour and merges with lee cyclone



## CASE STUDY: POLAR LOW OVER LABRADOR SEA



## Strong heat fluxes at sea ice edge and from leads and polynyas

- Coastal polynyas form along coat of Labrador and Baffin Island. Sea ice leads form over Baffin Bay due to northward winds
- Strong total turbulent heat flux (THF,  $> 3000 \text{ Wm}^{-2}$ ) at sea ice edge during CAO

 $> \sim 200$  to 1000 Wm<sup>-2</sup> from coastal polynyas

 $> \sim 200 \text{ Wm}^{-2}$  from sea ice leads

- Sensible heat flux (SHF) dominates latent heat flux (LHF)
- Direct cooling of boundary current



## CASE STUDY: POLAR LOW OVER LABRADOR SEA



#### Deepening of mixed layer in response to PL

- Mixed layer depth (MLD) up to 800 m along sea ice edge on 20 Feb 2020, a precursor PL was present
- Buoyant mesoscale eddies shed from West Greenland Current and inhibit deep MLDs
- During mature stage of PL (12-18UTC, 21 Feb 2020), the MLD deepened by ~50m over open ocean, more along sea ice edge
- A day later, the MLD deepened by >100m along sea ice edge and over open ocean after PL moved into Irminger Sea



## CONCLUSIONS

#### First globally coupled GCM to simulate polar lows

- ICON-Sapphire simulates all relevant processes (mesocyclones, BLFs, leads, polynyas, etc.)
- PLs induce strong heat fluxes from ocean, in particular near sea ice edge and from leads and polynyas:
  - ➤ Water mass transformation, e.g. DSOW
  - Deepening of mixed layer
  - ➢ New ice formation → brine rejection → densification of shelf water
- PLs may affect dense water formation but longer simulations needed