Impact of spatial resolution on large-scale ice cover modeling of mountainous regions







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Introduction

For reconstructing paleoclimate or studying glacial isostatic effects on the Earth's lithosphere, increasingly more studies focus on modeling the large-scale ice cover in mountainous regions over long time scales.

How well are high peaks and steep slopes resolved?



Previous studies of largescale ice cover simulation in mountain areas such as the Alps. New European the Tibetan Zealand, and Plateau, typically used kmscale resolution.

However, peaks and slopes are steep which crucial mass balance and dynamics, are poorly resolved in coarse resolution topography.

Fig. 1: A: Slope angle with respect to elevation intervals. B and C: elevation differences based on the 200 m topography in relation to elevation intervals

Methods

Experiment Setup

- Topography² of the entire **European Alps** (covering 480 240 km²)
- Present day climate with linear cooling rate until 6 °C colder
- Simulate the expanding ice cover over a time period of 160,000 years

The Instructed Glacier Model (IGM)

- Open source ice model: <u>https://github.com/jouvetg/igm</u>
- Novel 3D ice model equipped with a *Convolutional Neural Network* (CNN)
- trained from high order models to simulate ice flow
- Significant acceleration of run times
- \rightarrow Possibility of **running in higher** spatial resolution



Fig. 2: IGM model components, notations, and function of CNN, by Guillaume

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The 200 m model yields ~13.4% more total ice volume than the 2 km model.

With 1 km resolution we have ~3.4% less ice volume than with the 200 m model and ~10.4% more volume than with 2 km resolution.



Fig. 3: Ice volume of the different resolution models and temperature forcing (A). Ice volume gradient with growth phases (yellow, red and blue) (B



yellow, red, blue) for 200 m (**A**), 1 km (**B**), and 2 km (**C**)

References

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- 3)Fick, S.E. and R.J. Hijmans, 2017. WorldClim 2: new 1km spatial resolution climate surfaces for global land areas. International Journal of Climatology 37 (12): 4302-4315.

Results





Fig. 8-10: Ice volume of the different resolution models and temperature forcing (A), ice area (B), mean elevation, slope angle, surface mass balance of points that will be ice covered in the next time step (C-E), number of ice patches (F) for different catchment regions (yellow, green and magenta) as outlined in Fig. 7

This work is funded by



For **most regions**, the higher resolution model yields more ice.





EGU23-4200 Scan for abstract

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