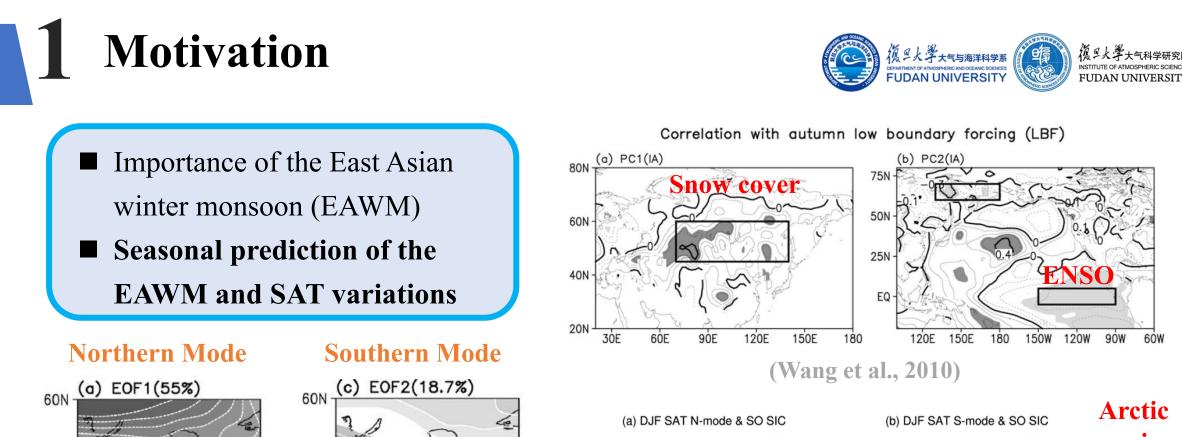




## Subseasonal strength reversal of the East Asian winter monsoon

Zhong Wogu and Wu Zhiwei

Fudan University



135E

120E

45N ·

30N -

15N

EQ

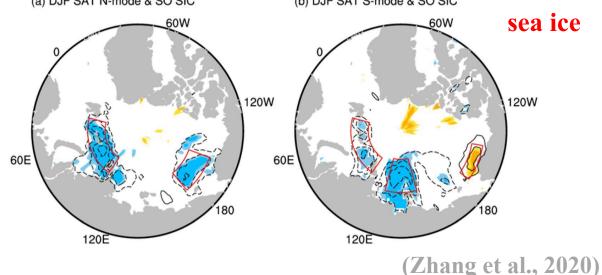
105E

45N -

30N -

15N

EQ

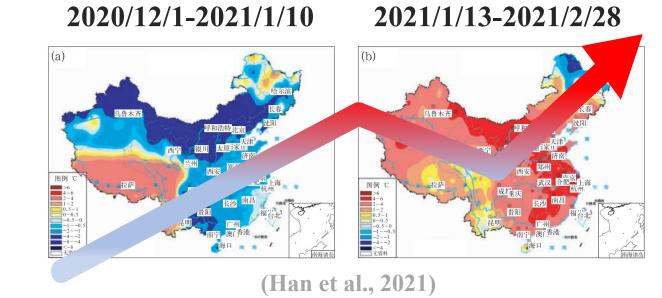


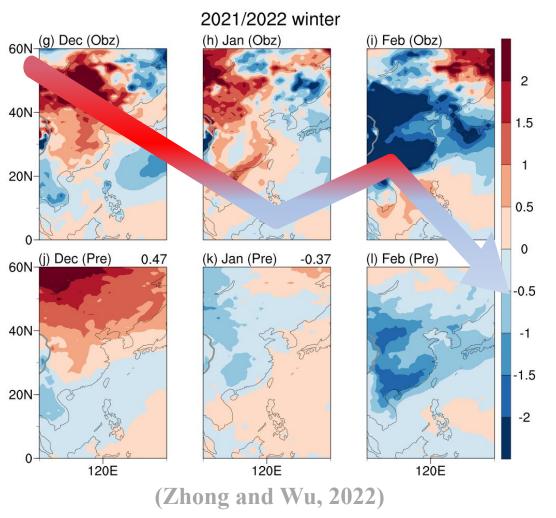
60W





A reversal of SAT anomaly (Icy weather to abnormal warm and vice versa) occurs more frequently in East Asia during the past winters.

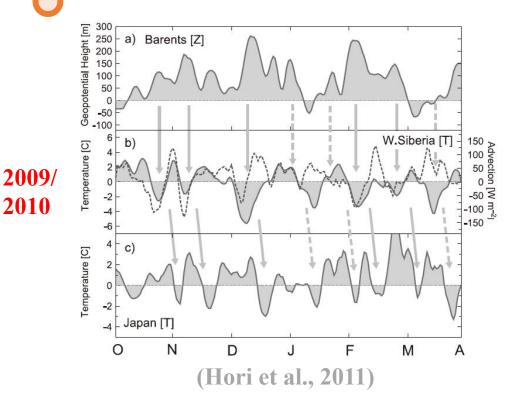


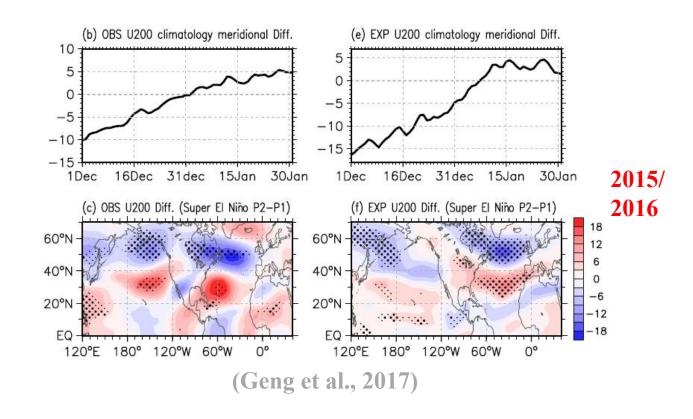






Some case studies suggested that the subseasonal SAT reversal phenomenon might be related to changes in the atmospheric condition over the Barents-Kara Sea, super El Niño, the modulation of the Pacific Decadal Oscillation on El Niño and Arctic sea ice. 2014/2015











- > The **leading modes** of subseasonal variations of the EAWM?
- > Roles of atmospheric variability and surface boundary conditions?
- ➢ How to predict subseasonal variations of the EAWM?





#### [Data]

Atmospheric circulation | ERA5 reanalysis data **(1979/1980-2021/2022)** Sea surface temperature | ERSST version 5 Sea ice concentration | HadISST ice

#### [Method]

Season-reliant Empirical Orthogonal Function (S-EOF) analysis (Wang and An,

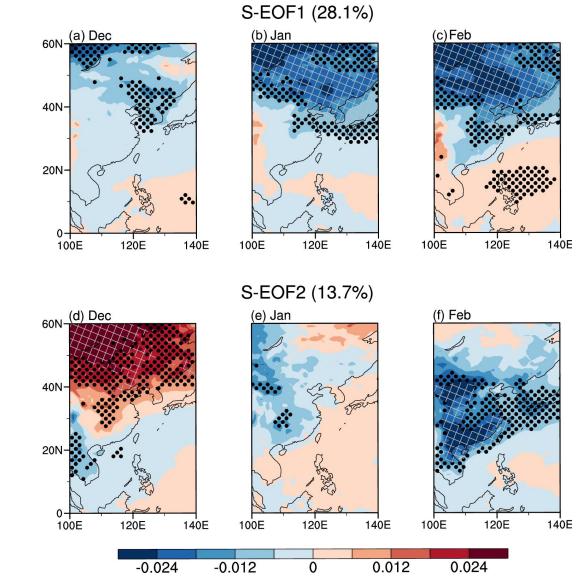
2005): to illustrate the **interannual variations** of wintertime SAT with emphasis on the monthly evolving pattern. Gridded SAT anomalies are **detrended**.

#### [Model]

Linear baroclinic model (LBM): Dry version; applied to solve steady linear equations with a given basic state and a thermal forcing; resolution: T42L20; basic state: **1979–2020** climatology based on ERA5 reanalysis.



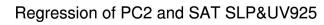


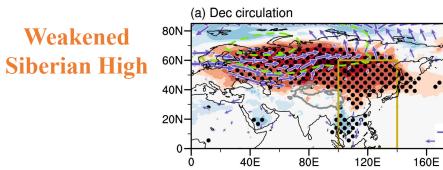


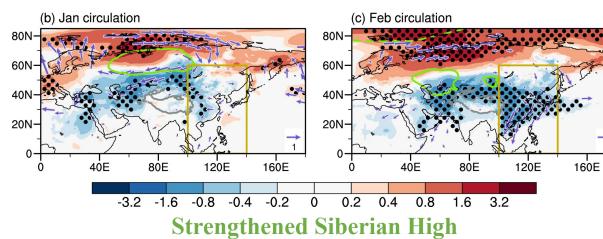
- SAT anomalies have been used as the key variable to measure the strength of the EAWM (Wu et al. 2009; Wang et al. 2010; Sun et al. 2016).
- S-EOF1 is characterized by a persistent stronger/weaker EAWM throughout the whole winter.
- S-EOF2 is featured by a reversal of the
  EAWM between early winter (December)
  and late winter (January–February).

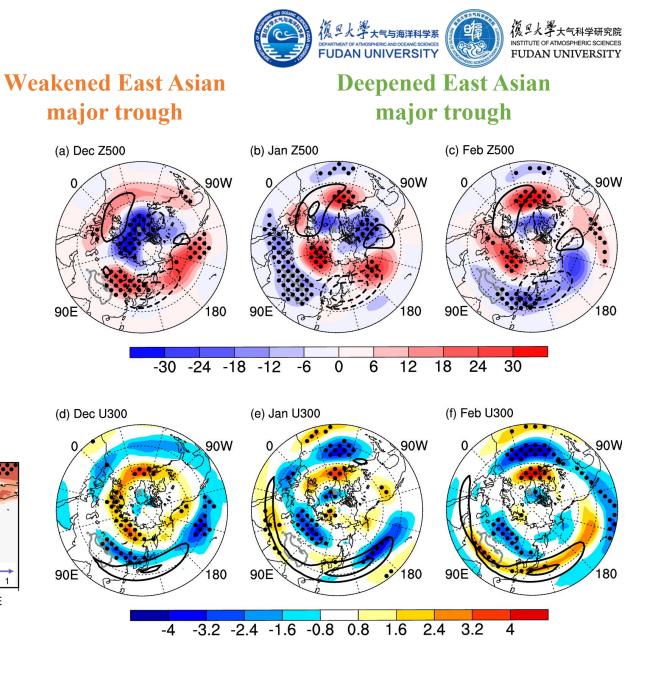


### S-EOF2 can capture subseasonal changes in the EAWM subsystems and is thus labelled as the SR-EAWM.



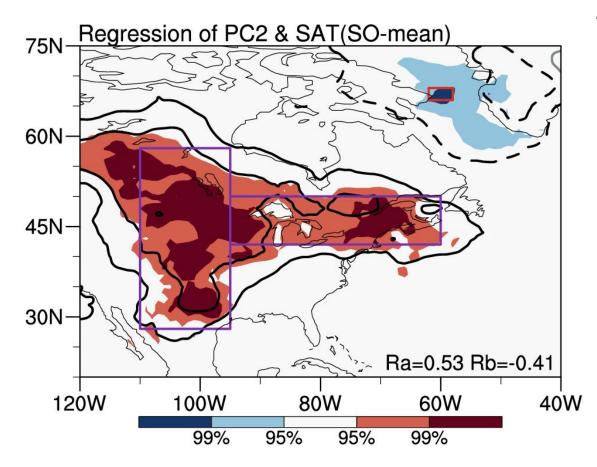




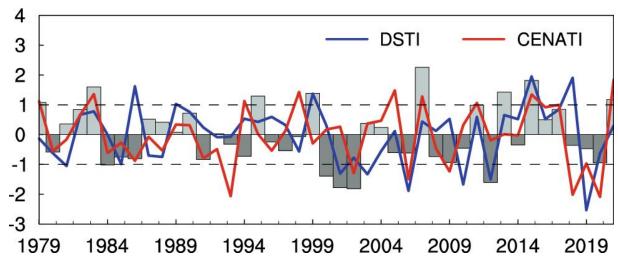




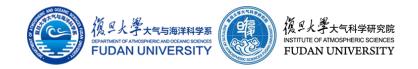




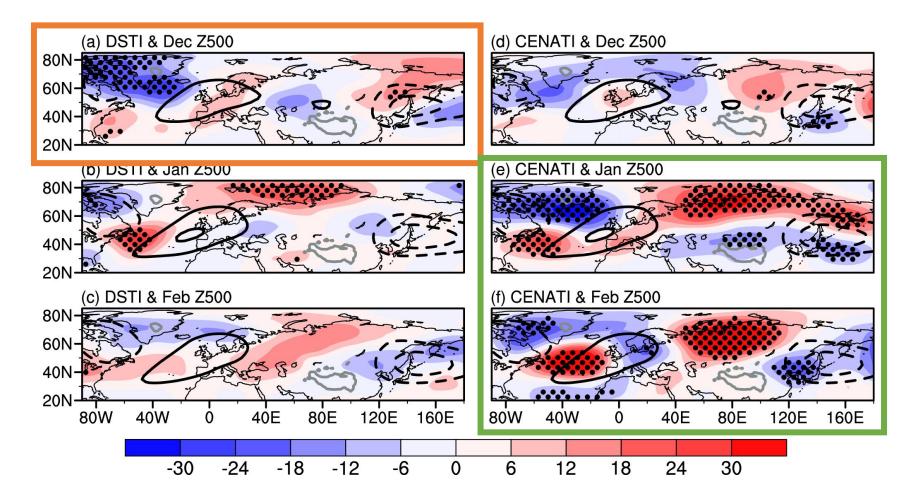
 The SR-EAWM is preceded by surface air temperature anomalies over Davis Strait (DST) and those over central-eastern North America (CENAT) in September–October.







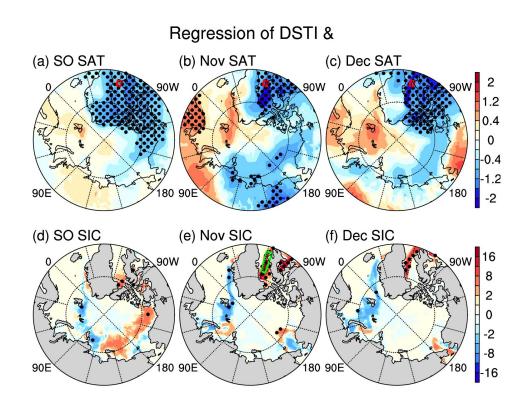
The DST mainly modulates the strength of the EAWM in early winter, while the CENAT exerts an impact on the SR-EAWM in late winter.

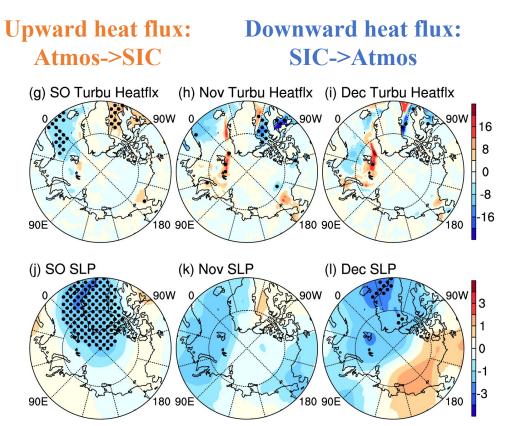






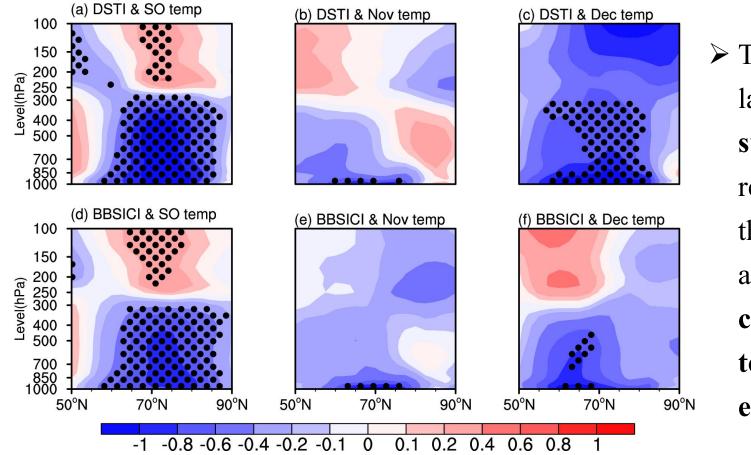
The DST mainly influences the SR-EAWM in early winter through a "sea ice bridge" of the November Baffin Bay sea ice concentration anomaly (BBSIC). The BBSIC could intensify the DST in December by altering surface heat flux.









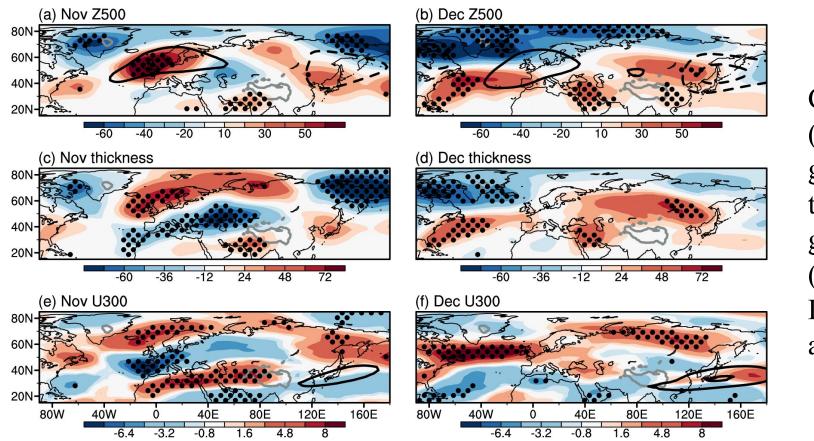


 $\blacktriangleright$  The formation of the BBSIC in late autumn greatly affects the surface energy balance and restricts the heat transfer between the warmer ocean and the colder atmosphere, thus the surface cooling intensifies and develops to the middle troposphere in early winter.





Air-sea ice interaction leads to a NAO-like circulation pattern, thus modulating the strength of the EAT in early winter.



Composite differences of (a–b) Z500 anomalies (units: gpm), (c-d) 1000–300-hPa thickness anomalies (units: gpm), (e–f) U300 anomalies (units: m/s) in November and December between heavy and light BBSIC winters.



(a) Forcing

80W

-1

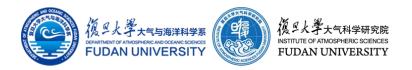
80N

70N-

60N-

50N

40N-



> The LBM experiment suggests that a cooling over Davis Strait can generate a NAO-like circulation anomaly and modulate the SR-EAWM in early winter.

Heating profile

K/day

(b)

0.0

0.2

0.4

0.6

0.8

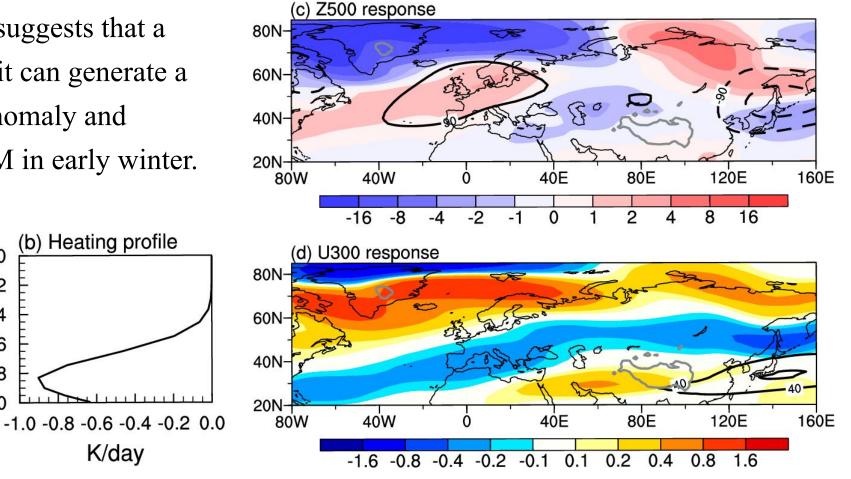
1.0

40W

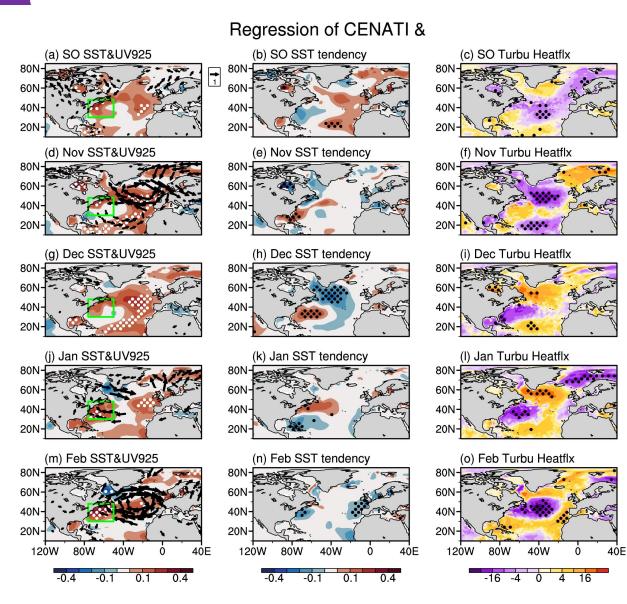
-0.8 -0.6 -0.4 -0.2

20W

60W



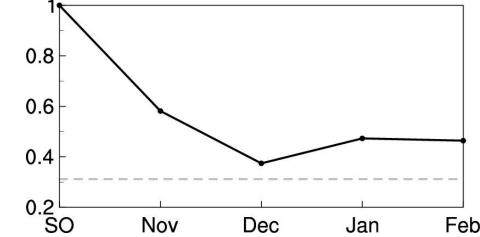




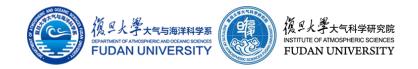


 The CENAT affects the SR-EAWM in late winter by inducing an "ocean bridge" of the western North Atlantic sea surface temperature anomaly (WNASST). The WNASST can persist into late winter.

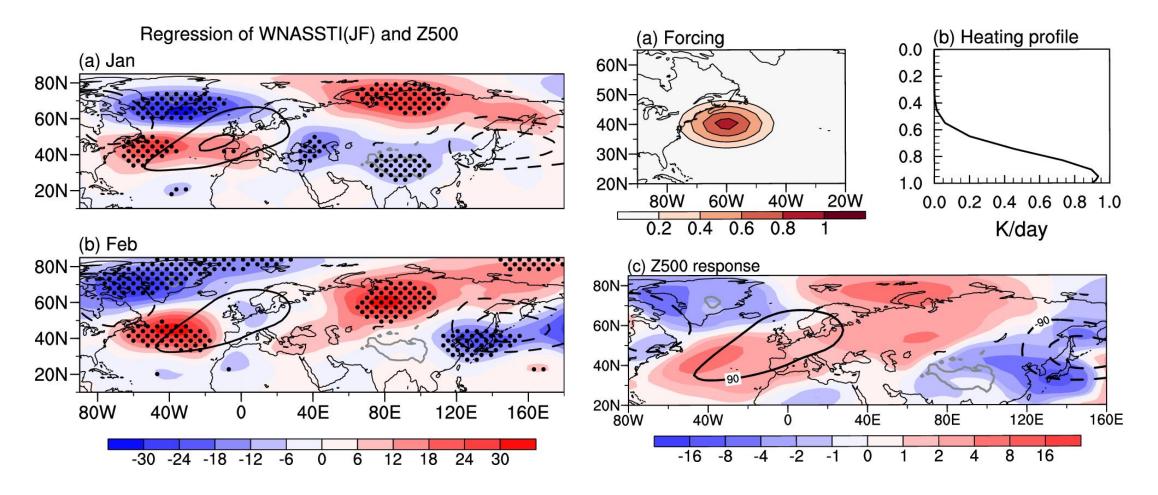
Auto CC of WNASST(SO-mean)





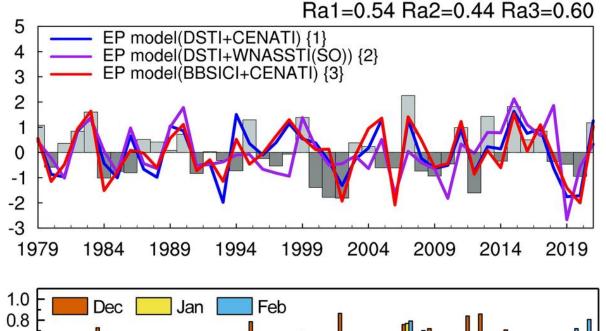


The WNASST significantly affects the SR-EAWM by regulating Eurasian circulation anomalies and the downstream EAT, which can be further verified by the LBM.







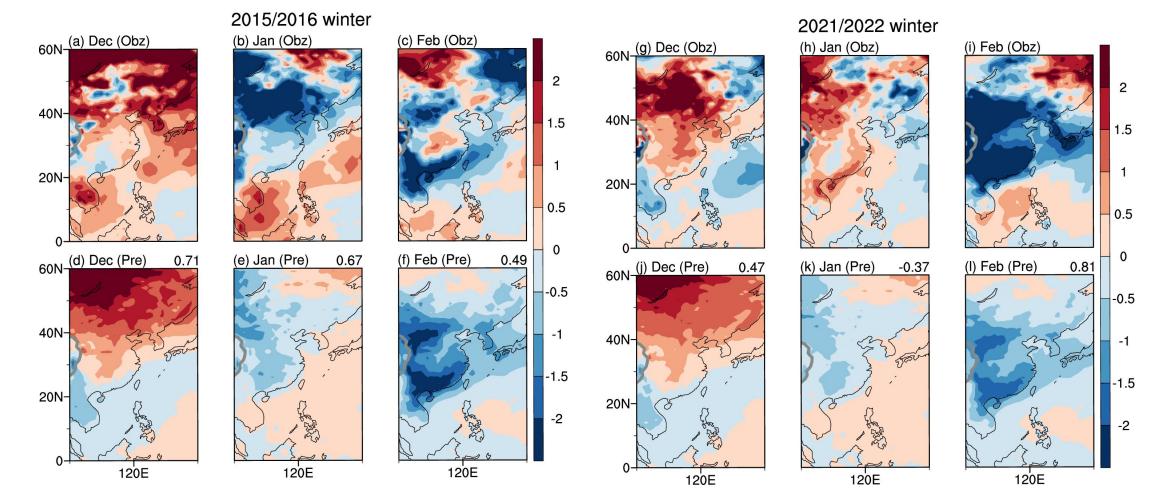


- 0.6 0.4 0.2 0.0 -0.2 -0.4 -0.6 -0.8 -1.0 1979 1984 1989 1994 1999 2004 2009 2014 2019
- Two physical-empirical models are established using the DST/BBSIC and the CENAT indices, respectively. Both exhibit promising prediction skills.
- The results highlight that the DST,
  BBSIC, and CENAT are crucial
  predictability sources for the
  SR-EAWM.





The physical-empirical models show good performances in predicting the subseasonal reversals of the EAWM in the winter of 2015/2016 and 2021/2022.









The leading modes of subseasonal variations of the EAWM?
 S-EOF1 A consistent in-phase EAWM
 S-EOF2 A reversal of the EAWM between Dec and JF

- Roles of atmospheric variability and surface boundary conditions? Interactions between the preceding atmospheric variability and surface boundary conditions
- How to predict subseasonal variations of the EAWM?

Physical-empirical models based on the DST/BBSIC and the CENAT indices

# Thank You For Your Attention



Reference: Zhong Wogu, Wu Zhiwei\* (2022). Subseasonal strength reversal of the East Asian winter monsoon. *Climate Dynamics*, DOI: 10.1007/s00382-022-06610-9.