





Geochemical interaction between slabderived melts and mantle at high pressure in subduction zones

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MELT-PERIDOTITE INTERACTIONS AT SUBDUCTION ZONES



Porous flow instant reaction

Focussed flow network of veins

Spandler and Pirard 2013 (Lithos)

The fate of crust-derived melts at warm subduction zones and the transport mechanism of slab-derived components to the supra-subduction mantle is still a matter of debate. Migration of crust-derived melts into the mantle by porous flow is limited by instant reaction with the peridotites and the consequent production of metasomatic orthopyroxene (±clinopyroxene) and phlogopite hybrid rocks at the slab-mantle interface.

Alternatively, the occurrence of a network of pyroxenite veins in metasomatised mantle xenoliths from arc lavas indicates that metasomatic melts may percolate the mantle by a mechanism of focused flow.

Melt-peridotite interaction via reactive porous flow has been largely studied in mantle samples from oceanic and subcontinental settings. We know little about the role of these interactions at the slab interface.



Tumiati et al. 2018 (J Pet); Pellegrino et al. 2020 (Lithos)

The Borgo outcrop of the Monte Duria area (Adula-Cima Lunga unit, Central Alps, Italy) is an ideal example of meltperidotite interaction which occurred under a deformation regime at HP. Grt-peridotites occur in direct contact with migmatised orthogneiss and eclogites that record a common HP peak at 2.8 GPa and 750 °C







We analysed the composition of 16 samples collected along a profile, from the eclogite-peridotite contact to the inner part of the peridotite lens Tremolitites occur both at the peridotite/eclogite contact and within the peridotite body and derived from the retrogression of previous Grt-websterites formed after the interaction at HP between eclogite-sourced melts and peridotites. The peridotite Grt-layering flows into the necks of the boudins, indicating that the stretching of the tremolitites (previous Grt-websterite) occurred when the peridotites were still in the Grt stability field.





Metamorphic stage	HP peak	HP hydration	HP melt/rock	Decompression	LP-(U)HT	LP-LT hydration	Retrogression
			reaction			and Chl foliation	
Chlorite peridotite							
Olivine	Ol1			Ol ₂			
Orthopyroxene	Opx ₁		Opx _{porph}	Opx ₂	*Opx _{sym}		
Clinopyroxene	Cpx ₁				*Cpx _{sym}		
Garnet							
Amphibole		Amp ₁	Amp _{porph}	Amp ₂	*Amp _{sym}	Amp₃	
Chlorite		Chl ₁				Chl ₃	
Phlogopite			Phl _{porph}			Phl ₃	
Spinel				Spl ₂	*Spl _{sym}		
Serpentine							
Tremolitite							
Orthopyroxene							
Clinopyroxene							
Garnet							
Amphibole			Hbl	Hbl		Tr ₁	Tr ₂
Phlogopite							
Talc							
Chlorite							

Opxporph 100 µm







500 µm

BULK ROCK GEOCHEMISTRY



Chl-peridotite Tremolitite Grt-peridotites (Duria) Pellegrino et al. 2020 (Lithos) Grt-orthopy. and websterite (Dabie-Shan) Malaspina et al. 2006 (EPSL) Sp-Grt-pyroxenite (External Ligurides) Montanini et al. 2012 (EPSL)

Chl-peridotites: Mg# (90), high Ni, and low $Al_2O_3 + CaO$. Tremolitites: Mg# (0.91) and Ni up to 1390 µg/g plotting into the field of the ultramafic compositions (marked difference with respect to Grt- and Spl-pyroxenites of subcontinental ophiolites). They show high SiO₂, high CaO and Al_2O_3 vs SiO₂/MgO close to the composition of metasomatic Grt-orthopyroxenites and websterites from Dabie-Shan, which were formed after the interaction of Grt- harzburgites with Si-rich crust-derived melts at UHP



BULK ROCK GEOCHEMISTRY



Peridotites: REE close to or slightly lower than the DMM, with fractionated patterns enriched in LREE relative to the MREE and HREE.

Tremolitites: REE up to 4.63 × PM with enrichments in MREE and LREE/HREE two orders of magnitude higher than subcontinental pyroxenites from External Ligurides. Both peridotites close to the contact with the migmatised eclogites, and tremolitites display a negative Eu anomaly, resembling that of eclogite leucosome produced from a PI-bearing source.



Tremolitites show a progressive increase in LREE from the contact up to about 30 m within the adjacent peridotite. Among the fluid mobile elements, both tremolitites and peridotites show a progressive depletion in Pb from the contact to 80 m within the peridotite body, whereas Sr is almost constant.



In the first step, the crustal melt derived from partial melting of eclogites reacts with the peridotite at the eclogite-peridotite interface. The liquids resulting from increasing interactions are assumed to form veins of websterites that are supposed to be the protoliths of tremolitites preserving their original REE composition. After several interactions, the final reacted melt crystallised at the eclogite-peridotite contact likely forms the first websterite DB177, showing a diffusive contact with retrogressed Grt-peridotite .

To test this hypothesis, we have numerically simulated the REE gradient applying the Plate Model of Vernieres et al. (1997) using the REE composition of the eclogite leucosome of Pellegrino et al. (2020) as starting melt and the DMM as peridotite matrix.





We assume an initial peridotite porosity of 20 % that reflects a high perido- tite assimilation coupled with the progressive melt consumption through melt-peridotite reaction during the percolation (i.e. a high extent of transient melt crystallisation), simulated by the progressive increase in crystallisation rate of 50 % Opx, 20 % Cpx, 20 % Grt, 10 % Phl

After the interaction at the eclogite-peridotite contact, the reacted melt infiltrates the peridotite producing the REE gradient observed in the tremolitites profile.







LaN and CeN/YbN resulting from Step 2 calculations (yellow squares) compared with those measured in our tremolitite (coloured squares) along the first 30 m from the eclogite-peridotite contact. Grey scale symbols show the sensitivity of the model to the different extents of initial and final crystallisation rate of the percolating melts in the fractionation of LREE/HREE

POSSIBLE PERSPECTIVES

The numerical simulation aims to model the effect of interaction between crust-derived melts produced by partial melting of mafic components of the slab with the supra-subduction mantle peridotite at sub-arc depths. This includes a first step of crustal magma stagnation and meltperidotite reaction at the slab-mantle interface and the following metre-scale percolation of reacted melt within the overlying peridotite that buffers the composition of the infiltrating melt. Reactive melt infiltration at HP is a plausible mechanism to modify the REE budged of mantle peridotites that lie on top of the subducting crustal slab. Samples from those settings tend to show peculiar LREE "spoon-like" fractionations.



SENSITIVITY OF THE MODEL



PM-normalised Yb vs La/Sm of melt resulting from the percolation of the liquid after Step1 at variable porosity (squares) and 30% or 50% of orthopyroxene (triangles and circles) in the crystallizing assemblage. The model shows that at high Ma/Mc ratio the addition of variable amount (30 or 50%) of orthopyroxene (having rather low Kd for the REEs) in the crystallizing assemblage does not impact significantly on the melt REE composition resulting from the reaction. We also evaluated the role of variable extent of initial porosity: Figure S-6b shows that porosity higher than 0.1 fails to reproduce the low LaN/SmN of melt that generated the retrogressed Grt-websterite (DB177) close to the contact with the peridotite-eclogite interface. On contrary, assuming porosity lower than 0.1 the model results in too fast melt LaN/SmN lowering.