

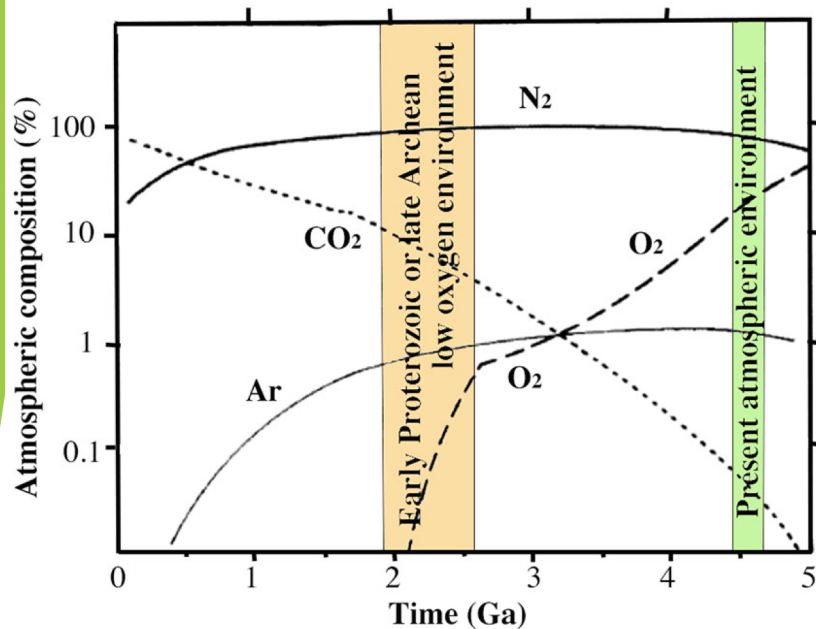
Artificial chemical weathering of basaltic rock under the earth surface conditions of the present and the Proterozoic era

S. KOBAYASHI*, Y. TAKAHASHI AND J. NAOHARA

Okayama University of Science, Okayama 700-0005, Japan. (*Correspondence: kobayashi@das.ous.ac.jp)

*https://ous.repo.nii.ac.jp/?action=pages_view_main&active_action=repository_view_main_item_detail&item_id=2467&item_no=1&page_id=13&block_id=21

In order to compare the mineral chemical effects of acid rain on surface materials in the **present oxygen level** and **the early Proterozoic or late Archean low oxygen** before Great Oxidation Event.



The probable evolution of atmospheric O₂, CO₂, N₂ and Ar over time (modified from Tajika, E. and Matsui, T., Pro. 25th ISAS Lunar Planet. Symposium, (1992), 178-183.).

Earth surface Environment of the early Proterozoic or late Archean

<Rock type on the crust: basaltic rock>

*Flood basalt studies (e.g. Pilbara) showed that basalt had been covered even on the early earth surface (e.g., Armstrong *et al.* Earth Planet. Sci. Lett., 101(1990), 90-106.).

<Atmospheric composition: high P_{CO2} with N₂>

*The major composition was CO₂ and N₂, and Cl, N₂ and S were also contained by influence of volcanic activity (L.A. Franks *et al.* Cambridge Univ. Press., (1992), 274 pp., E. Tajika & T. Matsui. Lithos 30 (1993), 267-280.).

*O₂ concentration were estimated based on studies such as banded iron formation and uranium mineralization (P. Cloud. Am. Jour. Sci. 272 (1972), 537-548).

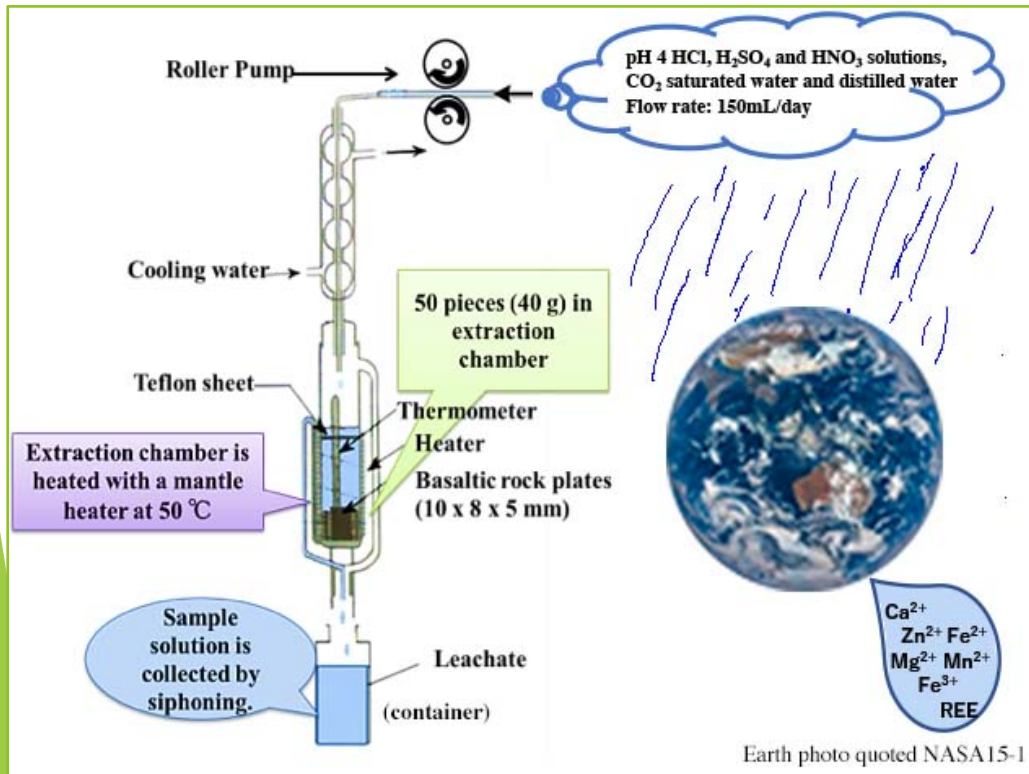
<Rain water: pH of rainwater was low >

Estimated from the atmospheric composition at the time.

Experimental methods

We have conducted experiment under **the present oxygen*** and **the low oxygen**** conditions, respectively.

Experimental image



Schematic illustration of the Soxhlet extraction apparatus, which has been improved to an open system and can be imagined as rain on-and-off.

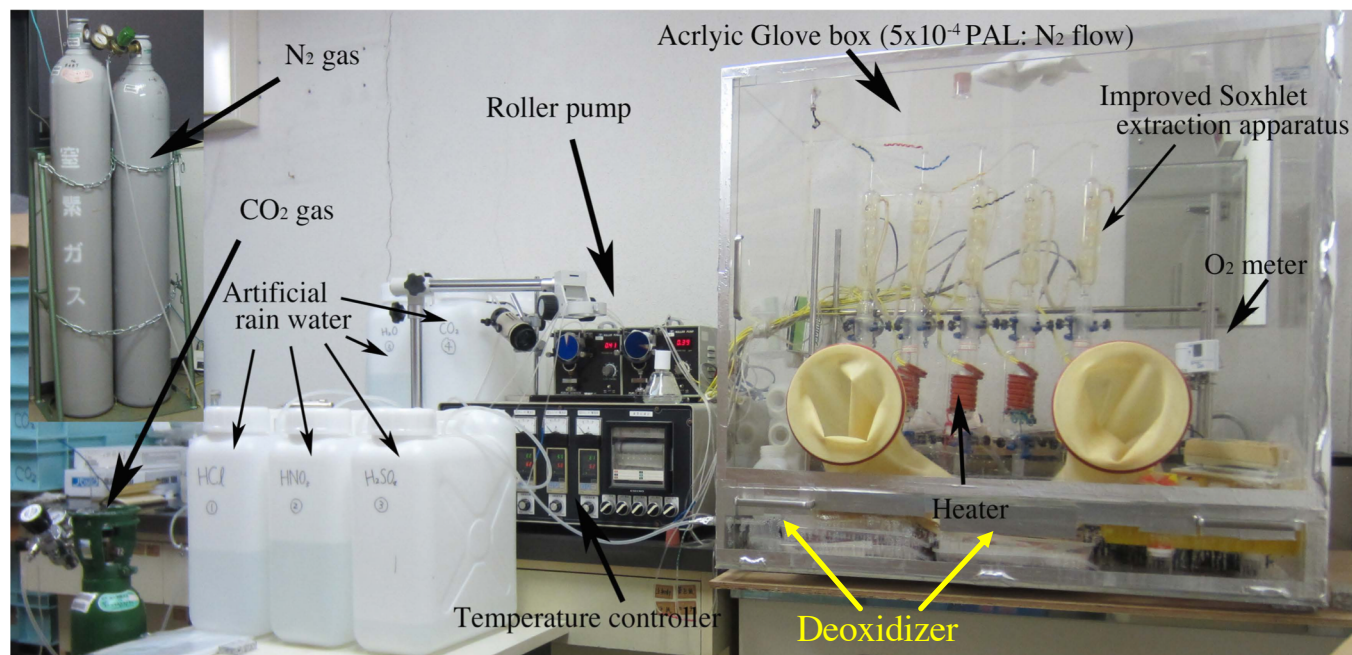
**The artificial chemical weathering experiment under the present oxygen was conducted in a room environment.*

The sample for the artificial weathering experiments is basalt in both experiments.



Basalt sample after 950 days reaction with pH4H₂SO₄ in the extraction chamber.

****Artificial weathering experiment under low oxygen environment**



Photograph of the experimental equipment under 0.01% oxygen ($<5 \times 10^{-4}$ PAL) in N₂ atmosphere.

Apparatus: Soxhlet extractor, which can be *imagined as rain on-and-off*, was improved to *an open system* with N₂ flow.

O₂ concentration: 0.01% oxygen ($< 5 \times 10^{-4}$ PAL) with **deoxidizer**

Basalt sample: about 50 polished pieces of 10 x 8 x 5 mm size (40 g)

Artificial rain water: HNO₃, H₂SO₄ and HCl solutions at pH 4, and CO₂ saturated water, and distilled water after bubbling with N₂ gas.

Flow rate (drip amount) of the rain water: 150 mL/day

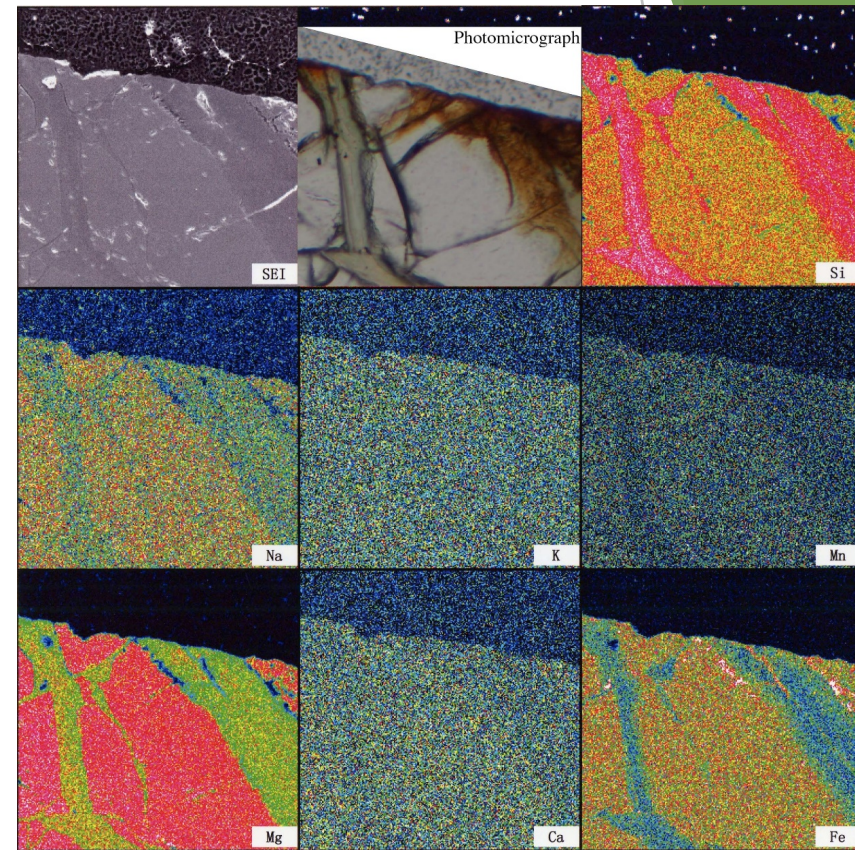
Temperature: 50 °C, **Duration:** different period of time up to 950 days

Altered products

Altered olivine

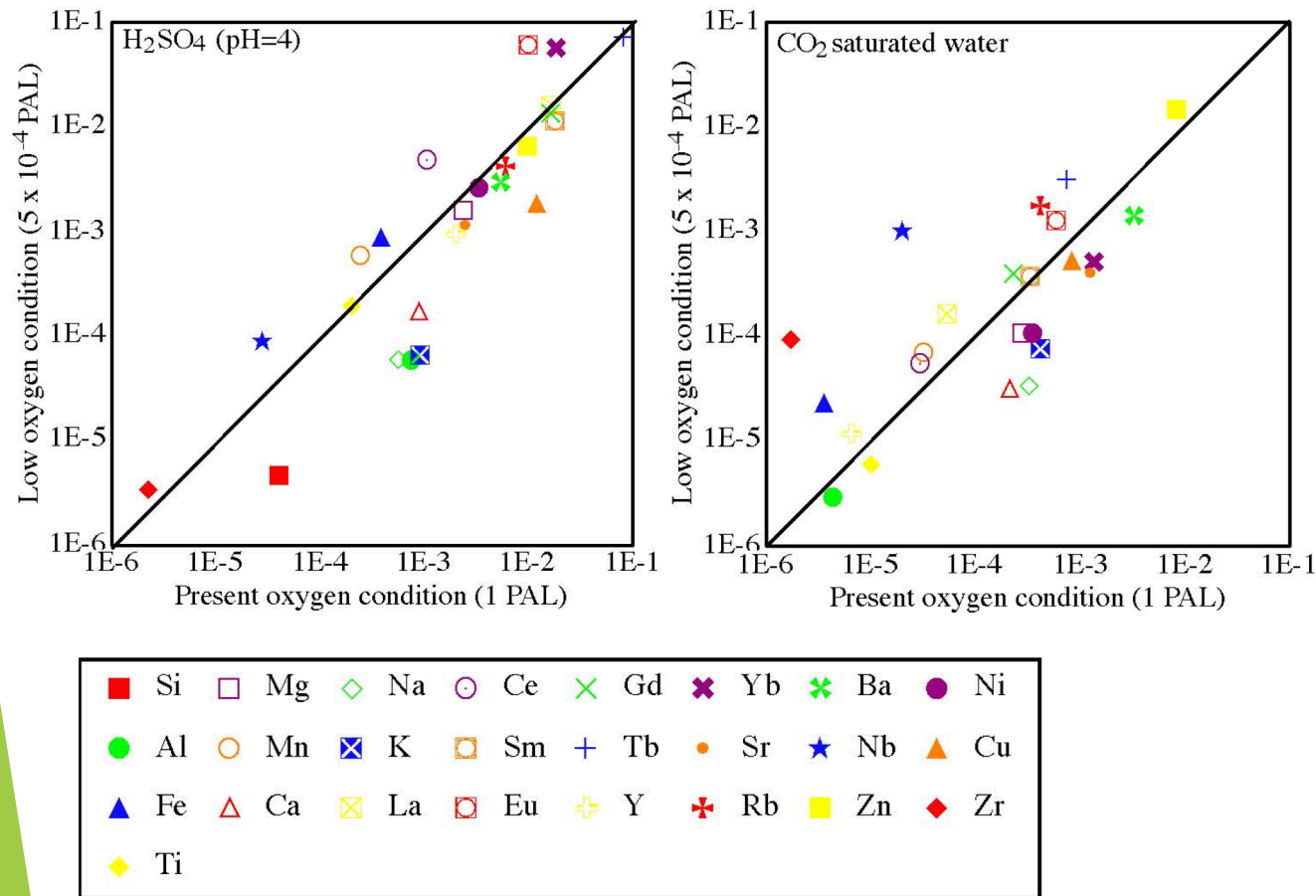


Under the present oxygen environment, A part of olivine has changed to chlorite and akaganeite ($\text{FeO}(\text{OH})$) as altered products by pH 4 H_2SO_4 treatment.



Mapping data of olivine surface after 347 days by pH4 H_2SO_4 .
The width of each field is 50 micrometers.

Comparison of the logarithmic molar ratios*** of extracted elements with rain water under the present oxygen and the low oxygen conditions.



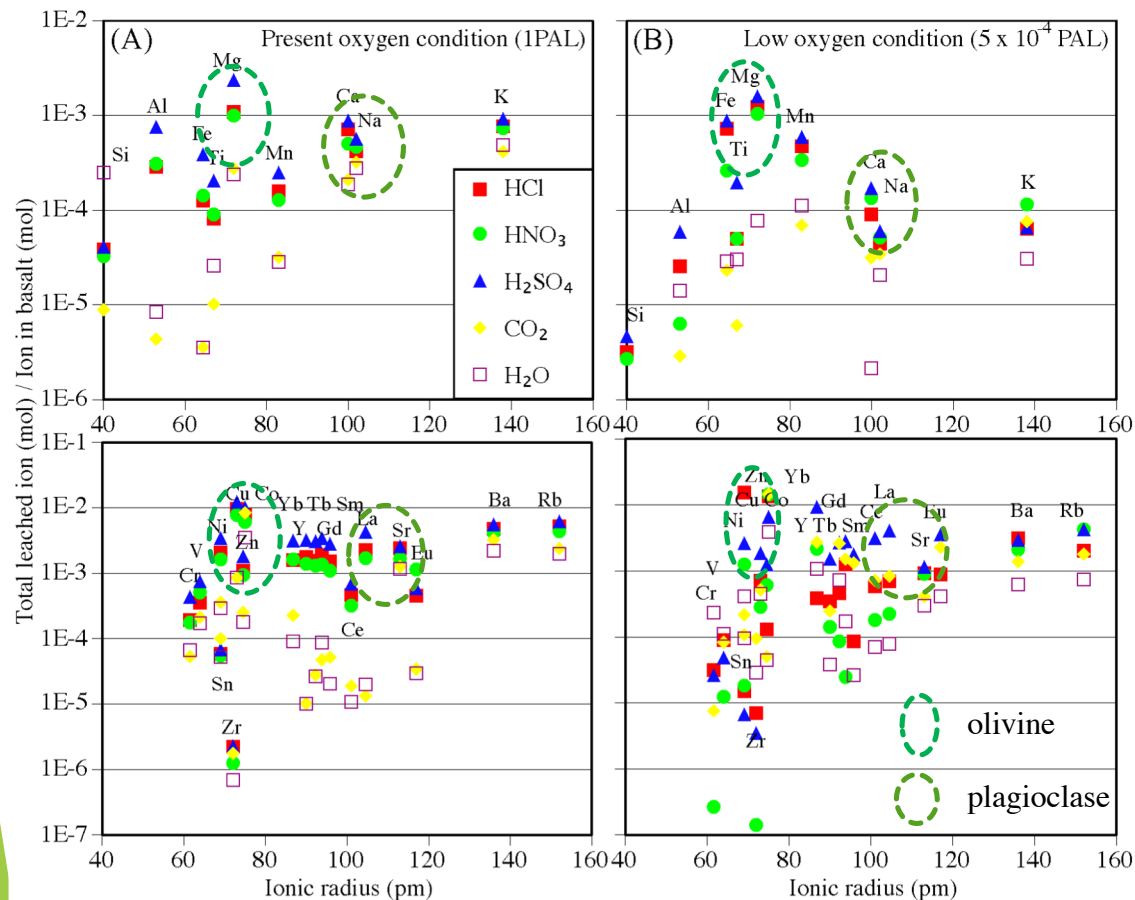
Logarithmic molar ratios of Fe and Mn are high (easy to extract) relatively under the low oxygen rather than present oxygen environment.

High field strength (HFS) elements tend to be easily extracted in any solution in the low oxygen environment. This tendency is remarkable when treated with CO₂ saturated and distilled water.

Under a low oxygen environment, Ce and Eu in REE tend to be easily extracted.

*** Molar ratio is calculated by dividing the cumulative total mole of each extracted element by the mole of individual element in the unaltered basaltic rock.

The relation between the logarithmic molar ratios and ionic radius under present (A) and low (B) oxygen environments.



In both environments, the ratios of Fe, Mg, Ni, Zn and Co near 70 pm in ionic radius are high, and reflect the dissolution from the octahedral coordination of olivine. On the other hand, mineral surface and chemical changes indicate that the extraction from CPX is low.

The high ratio of Ca, Na, Ce, La, Eu and Sr near 100-120 pm, which are occupied in the cavities within the framework, reflect the dissolution of plagioclase.

These results indicate that the first step in the weathering of basalt by acid solutions is the dissolution of olivine and plagioclase. The molar ratios of many elements indicate that the reactivity of the solution to the basalt is approximately pH 4 H₂SO₄, HCl and HNO₃, and CO₂ saturated water and distilled water.

Elements abundant in seafloor mineral resources.

Manganese nodule

期	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
1	H<1																	He
2	Li10	Be2											B25	C	N	O	F	Ne
3	Na1	Mg1											Al<1	Si<1	P1	S<1	Cl<1	Ar
4	K1	Ca<1	Sc<1	Ti1	V3	Cr<1	Mn161	Fe1	Co90	Ni217	Cu378	Zn19	Ga1	Ge<1	As64	Se	Br<1	Kr
5	Rb<1	Sr1	Y7	Zr3	Nb9	Mo520	Tc	Ru10	Rh43	Pd2	Ag2	Cd150	In<1	Sn<1	Sb185	Te4200	I<1	Xe
6	Cs79	Ba6	*1	Hf2	Ta<1		Re	Os37	Ir19	Pt97	Au<1	Hg<1	Tl338	Pb41	Bi177	Po	At	Rn
7	Fr<1	Ra<1	Th5	U8														
*1 rantanoid	La2	Ce12	Pr32	Nd2	Pm	Sm2	Eu29	Gd1	Tb52	Dy1	Ho23	Er1	Tm71	Yb<1	Lu<1			

Manganese crust

期	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
1	H<1																	He
2	Li<1	Be2											B13	C	N	O	F	Ne
3	Na<1	Mg<1											Al<1	Si<1	P8	S5	Cl49	Ar
4	K1	Ca1	Sc1	Ti2	V5	Cr<1	Mn163	Fe2	Co283	Ni91	Cu40	Zn9	Ga1	Ge1	As72	Se(3)	Br28	Kr
5	Rb<1	Sr3	Y7	Zr3	Nb6	Mo374	Tc	Ru12	Rh115	Pd1	Ag(4)	Cd49	In(10)	Sn1	Sb161	Te8000	I<1	Xe
6	Cs(1)	Ba4	*1	Hf3	Ta1	W107	Re	Os24	Ir50	Pt305	Au1	Hg12	Tl378	Pb67	Bi177	Po	At	Rn
7	Fr<1	Ra<1	Th<1	U8														
*1 rantanoid	La10	Ce25	Pr22	Nd8	Pm	Sm11	Eu9	Gd7	Tb13	Dy16	Ho9	Er15	Tm<1	Yb9	Lu11			

Hydrothermal sulfide ore

期	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
1	H																	He
2	Li<1	Be<1											B<1	C	N	O	F	Ne
3	Na<1	Mg<1											Al<1	Si1	P<1	S879	Cl<1	Ar
4	K<1	Ca3	Sc<1	Ti<1	V<1	Cr<1	Mn1	Fe5	Co36	Ni1	Cu3370	Zn4361	Ga<1	Ge<1	As180	Se<1	Br<1	Kr
5	Rb<1	Sr3	Y<1	Zr<1	Nb<1	Mo118	Tc	Ru<1	Rh<1	Pd<1	Ag<1	Cd11075	In<1	Sn<1	Sb1160	Te<1	I<1	Xe
6	Cs<1	Ba200	*1	Hf<1	Ta<1	W<1	Re	Os<1	Ir<1	Pt<1	Au3169	Hg<1	Tl<1	Pb300	Bi<1	Po	At	Rn
7	Fr<1	Ra<1	Th<1	U<1														
*1 rantanoid	La<1	Ce<1	59Pr<1	Nd<1	Pm	Sm<1	Eu<1	Gd<1	Tb<1	Dy<1	Ho<1	Er<1	Tm<1	Yb<1	Lu<1			

<Numerical values show the ratio with rock on the surface.>

200	Very high
50	High
10	Slightly high

海底鉱物資源。オーム社/
Cronan(2000)/Hein(1997), Hannington(1995)/Usui,Someya(1997)

One consideration of the relationship between the extracted elements and the sea water

Material supply to the oceans is primarily thought to come from the continental and oceanic crust, the Earth's interior, and the atmosphere. Based on the results of this study, one consideration of supply of elements is shown below, using Mn deposits on the seafloor as an example.

Mn, Mg, Cu, Zn, Ni, Co and some REE elements are shown relatively easily extracted from basalt in this research. Among them, Mn, Ni, Co and REE are more abundant in manganese nodule and crust than in hydrothermal deposits. This study suggests that it is important to consider the weathering of materials near the earth surface for the supply of resources. Moreover, Fe, Mn, and Zn, which are easily extracted from basalt under the low oxygen condition.

The results of this study are expected to provide basic data for research on material supply to the ocean from the early Earth to the present.

Usui, A. (2010) Kaitei_koubutsu_shigen, Ohmsha Press./
Croman, D.S. (2000) Marine Mineral Deposits, CRC Press.
Hein, J.R. et al. (1997) Geochemistry and Mineralogy of Terrestrial and Marine Deposits, 119, 123-138.
Hannington, M.D. & Jonasson, I.R. (1995) Seafloor Hydrothermal Systems, 115-157.
Usui, A. & Someya, M. (1997) Geol. Soc. London Special Publ., 119, 177 – 198.