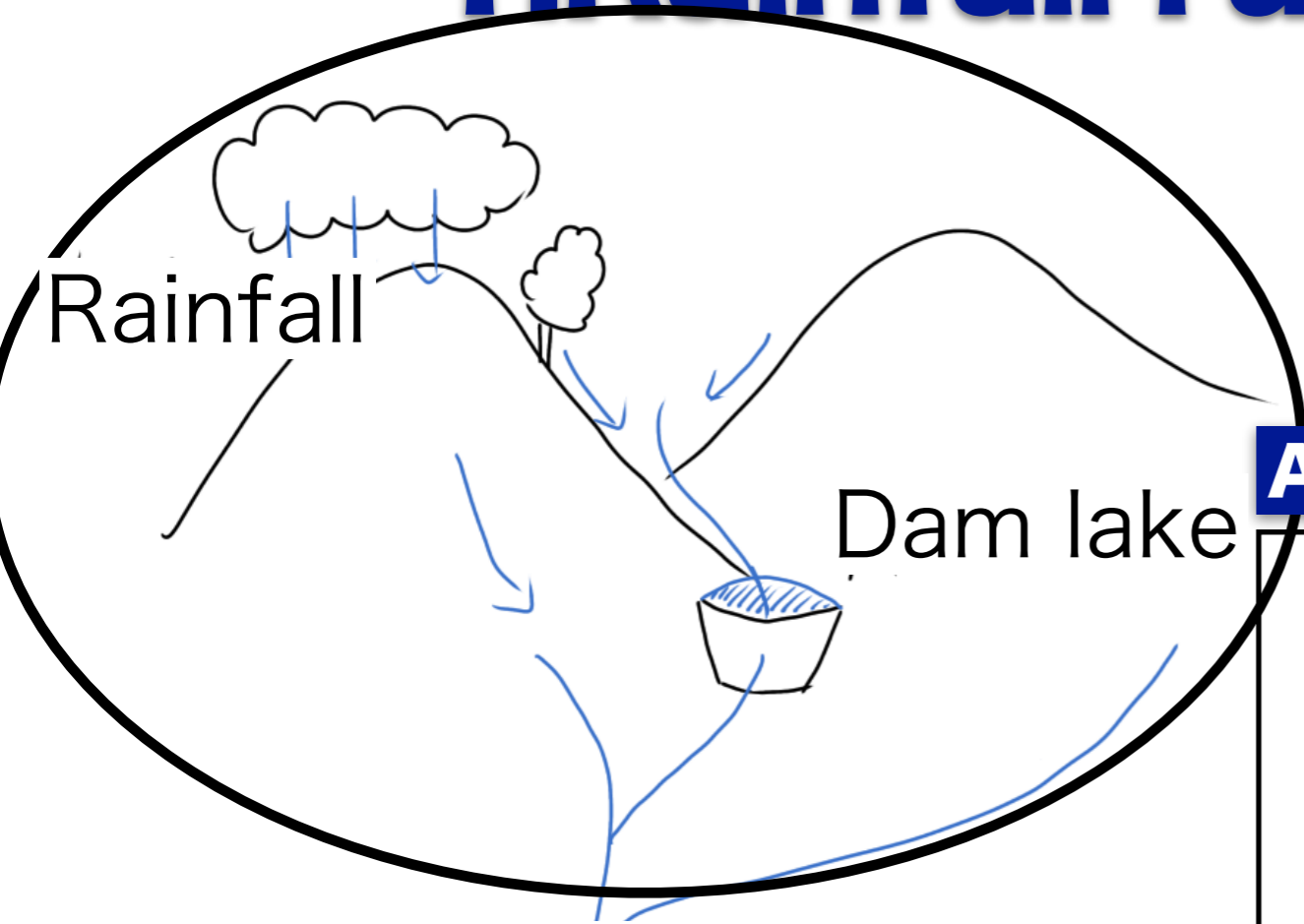


# **A study of Rainfall-Runoff Process considering two uncertainties in Basin with multiple dams**

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Prof Tomohito Yamada

# 1. Rainfall runoff process



Runoff  $\left\{ \begin{array}{l} \text{Direct runoff} \\ \text{Base flow} \end{array} \right.$

## Analysis of Runoff, Physical Model

Basic equations are derived from the Kinematic wave equation.

$$\begin{cases} \frac{\partial q}{\partial x} + \frac{\partial h}{\partial t} = r_e - \textcircled{1} \\ q = \alpha h^m - \textcircled{2} \end{cases}$$

$q$ : discharge per unit width,  $h$ : depth of flooding,  $\theta$ : gradient of steep,  $\tau_b$ : frictional stress

The observation results show that runoff coming from the vicinity of the river contribute direct runoff.

Yamada(2003)  $\downarrow$   $q(x, t) \simeq lq_*(t) - \textcircled{3}$

$$\frac{dq_*}{dt} = \alpha q_*^\beta (r_e - q_*) - \textcircled{4}$$

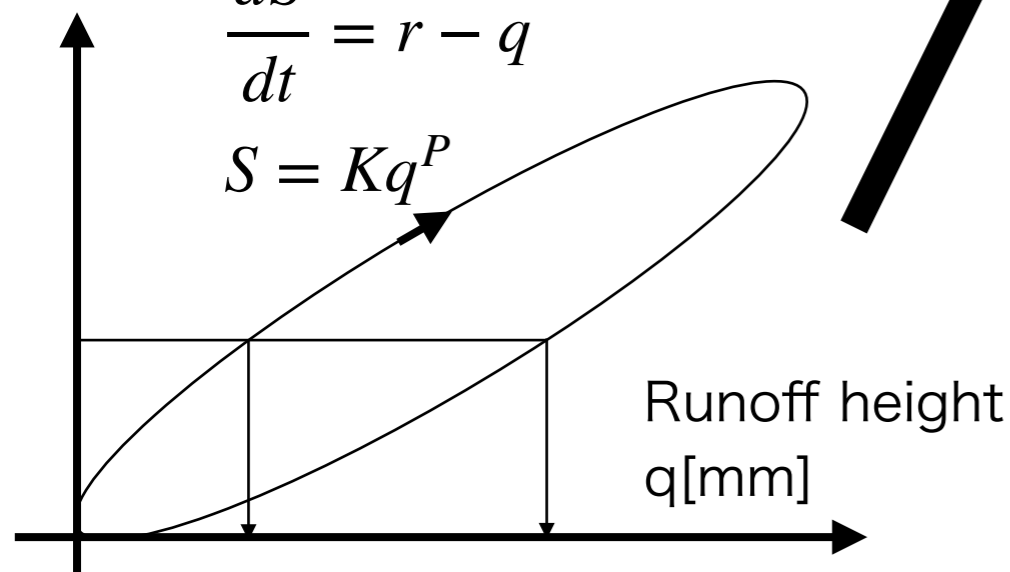
$q_*$ : runoff height,  $l$ : length of the contributing regions,  $q_*$ : runoff height,  $r_e$ : effective rainfall,  $\alpha, \beta$ : parameters

### 貯留関数法

Storage Height  $s$ [mm]

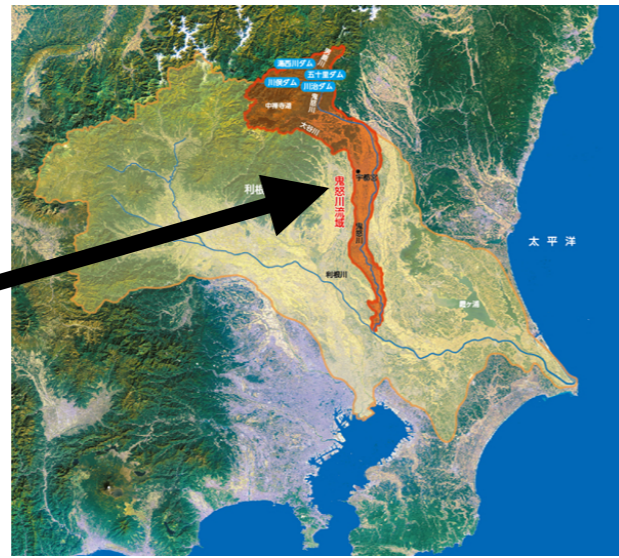
$$\frac{dS}{dt} = r - q$$

$$S = Kq^P$$



Accumulation of observation data of heavy rainfall cases in recent decades  
**Object of the study: To propose an equation form that takes into account the uncertainty of the variables (rainfall information)**

# 2.Uncertainty in rainfall observations

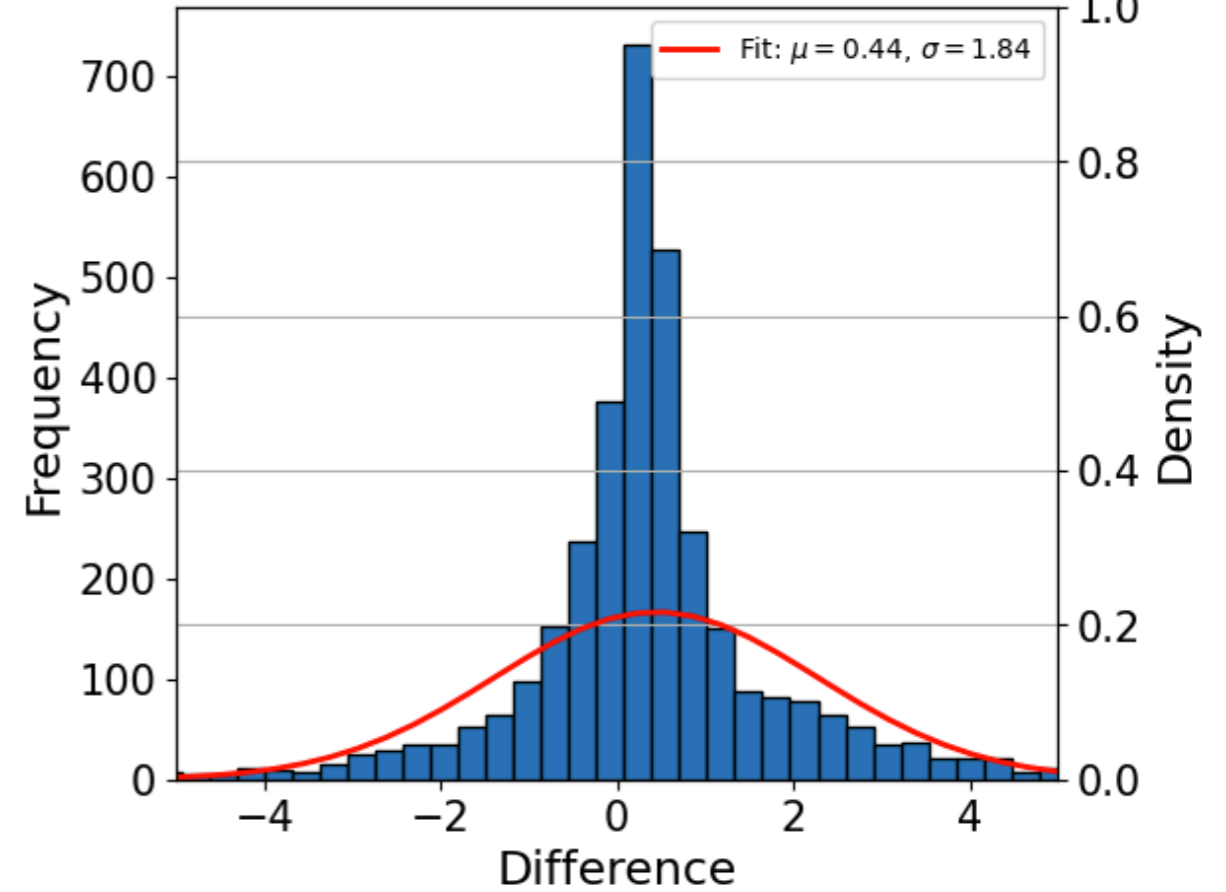


Japan, Kinugawa basin, MLIT



the 2015 Kanto Tohoku Heavy Rainfall, Nikkei x tech

Distribution of Difference between Ground Rainfall and Xrain



Histograms of Differences between XRAIN Rain Gauge and AMeDAS Rain Gauge in the Kinugawa River Basin of the 2015 Kanto Tohoku Heavy Rainfall and the 2019 East Japan Typhoon

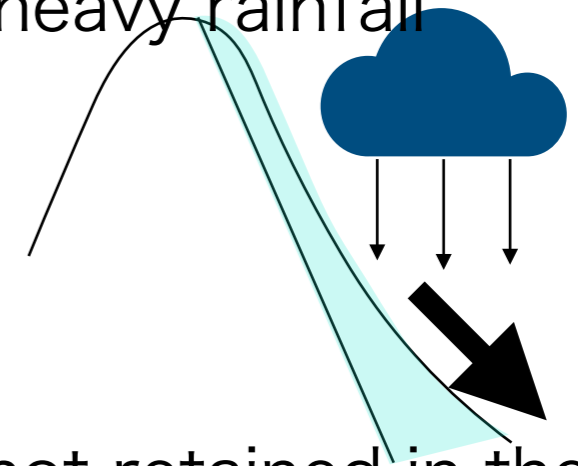
**The accumulation of observational data has made it possible to quantify the range of uncertainty.**

**Consider uncertainties due to limitations of observations in the rainfall information to be input**

# 3. Uncertainty in soil wetness 4

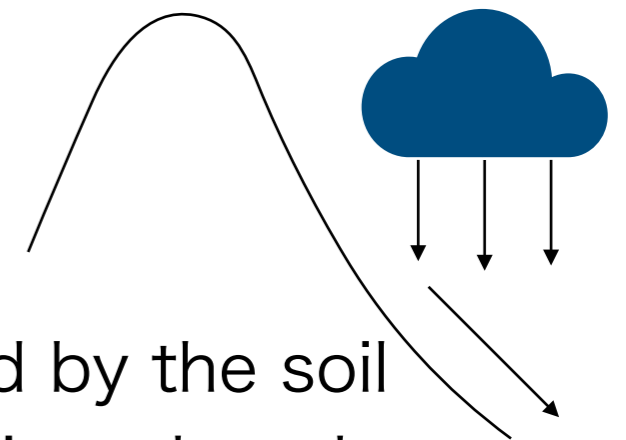
Total rainfall loss relative to total rainfall at Kawamata Dam

① If the soil is saturated just before a heavy rainfall



Rainfall is not retained in the soil and direct runoff is greater.

② If the soil was dry just before a heavy rainfall

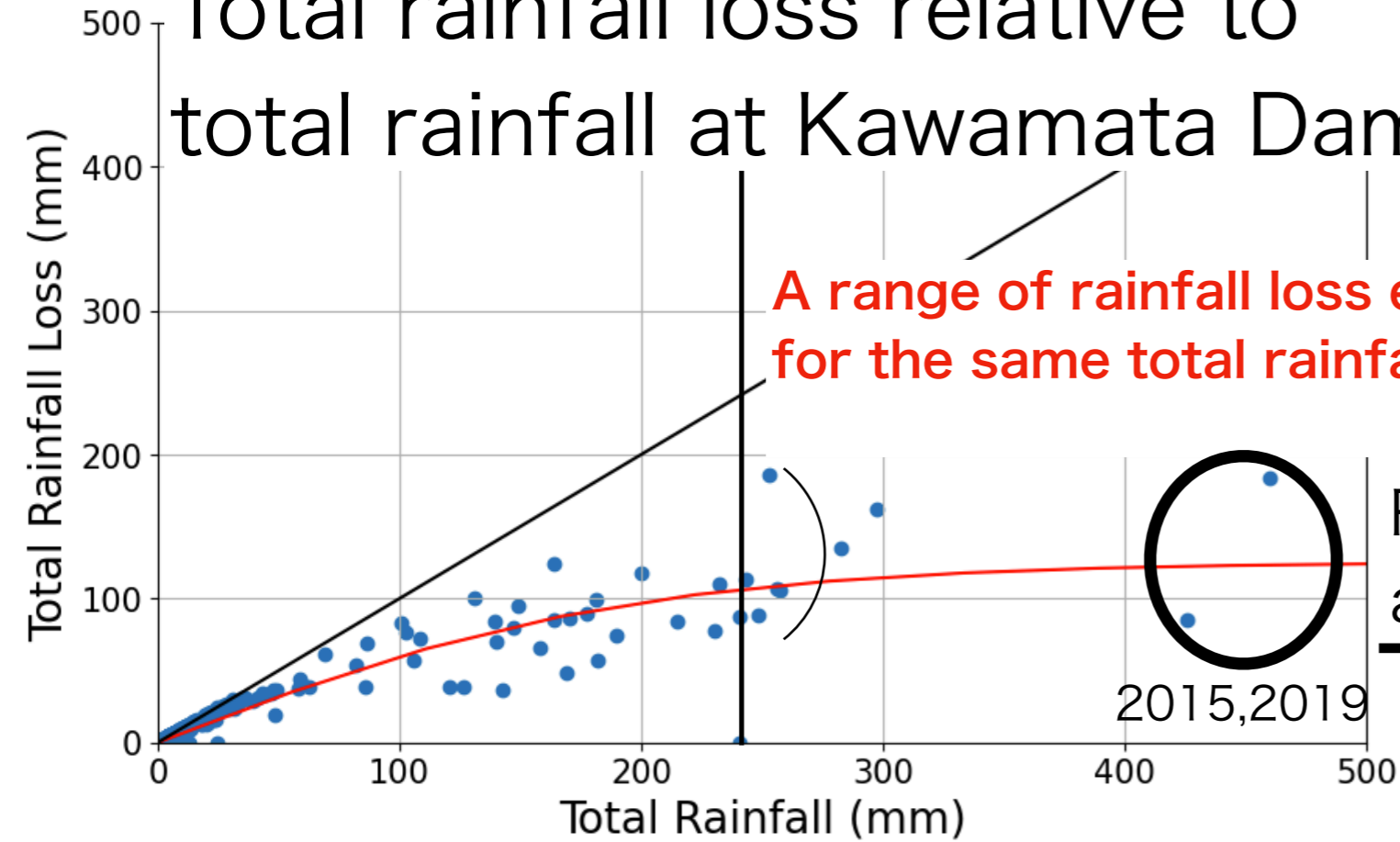


Rainfall is retained by the soil and direct runoff is reduced.

**There is uncertainty in direct runoff not only from observational uncertainty, but also from the wetness of the soil immediately prior to heavy rainfall**

A range of rainfall loss exists for the same total rainfall

2015, 2019

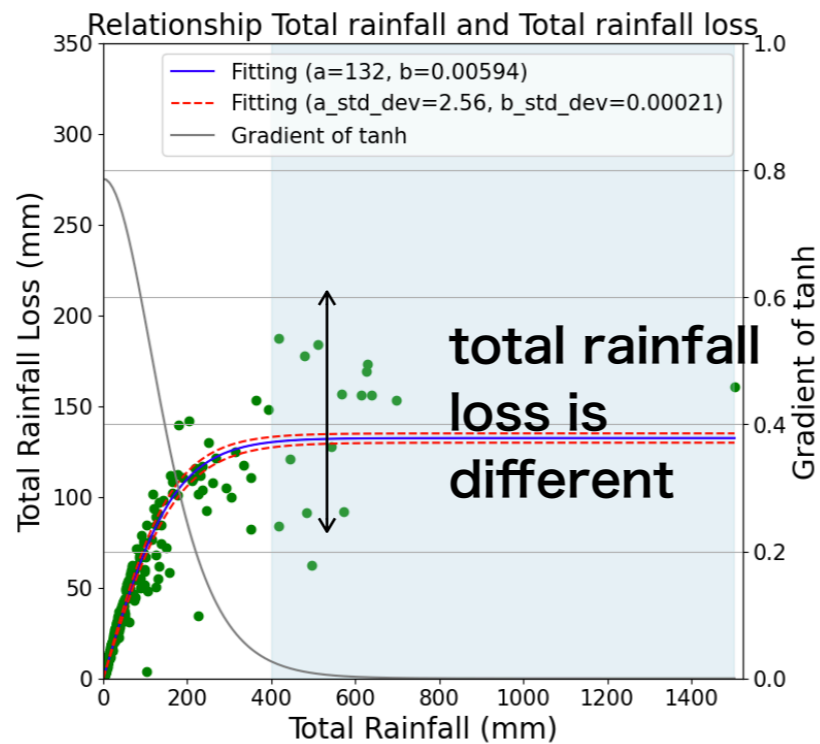
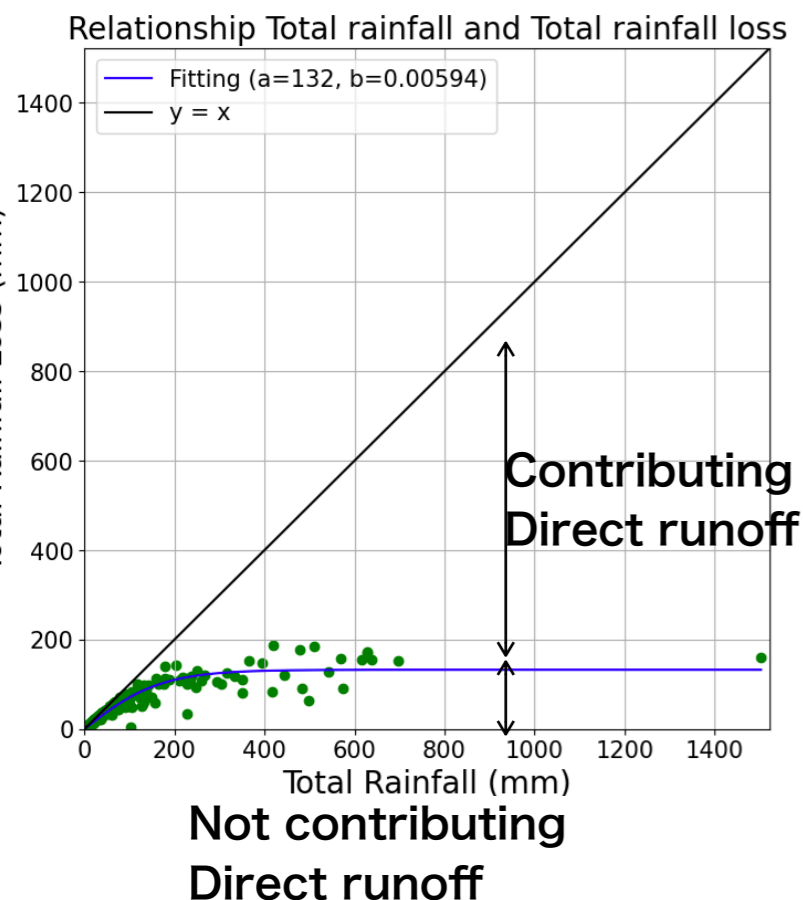
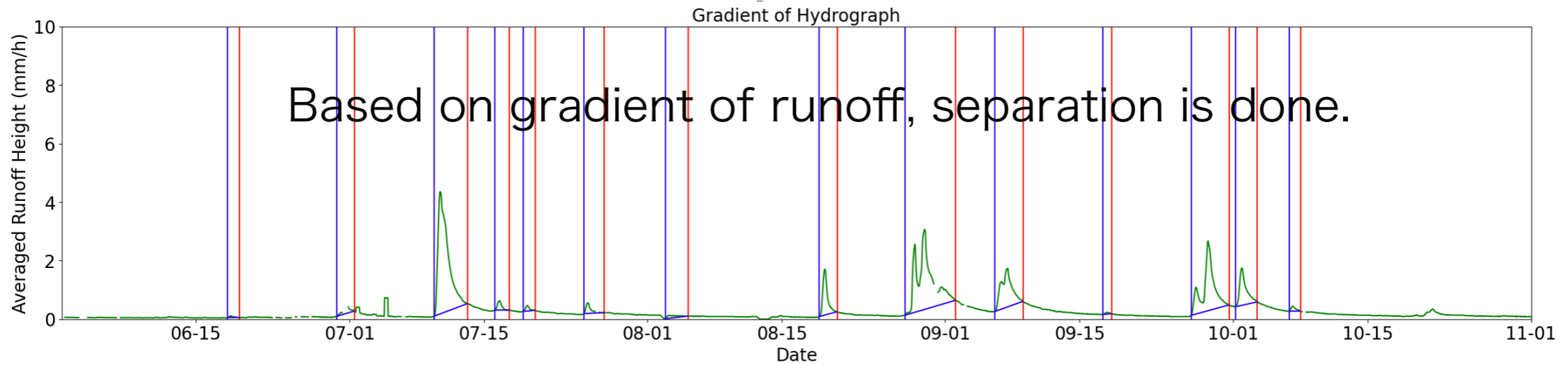
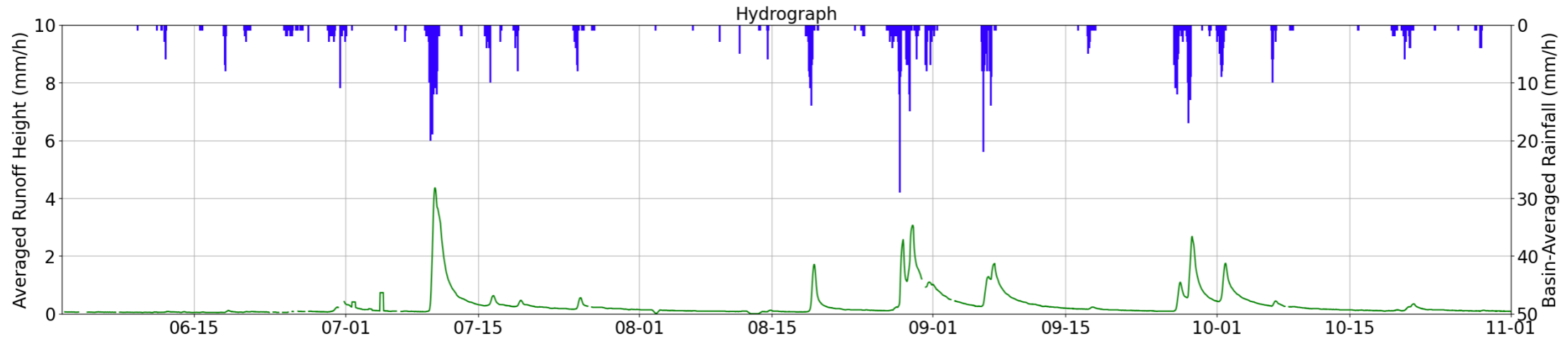


Total rainfall loss converges to a constant value as total precipitation increases.

Accumulation of observation data of heavy rainfall cases in recent decades

The range of uncertainty in converging values can now be quantified.

# Hachisu dam (Japan, Mie prefecture) June to November of 2002



## What do I want to do here?

I want to analyze why total rainfall loss is different even if total rainfall is the same amount.

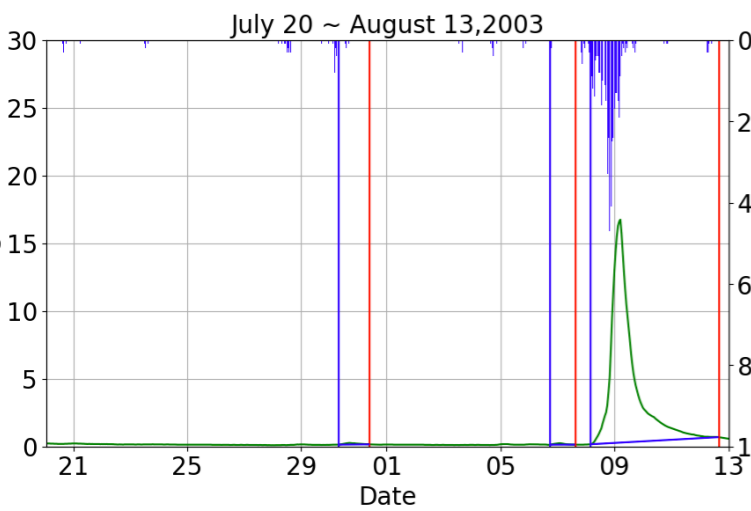
I extract rainfall event with a total rainfall amount of 400 or more and find out what happened to the previous rainfall event.

Hachisu dam (Japan, Mie prefecture)

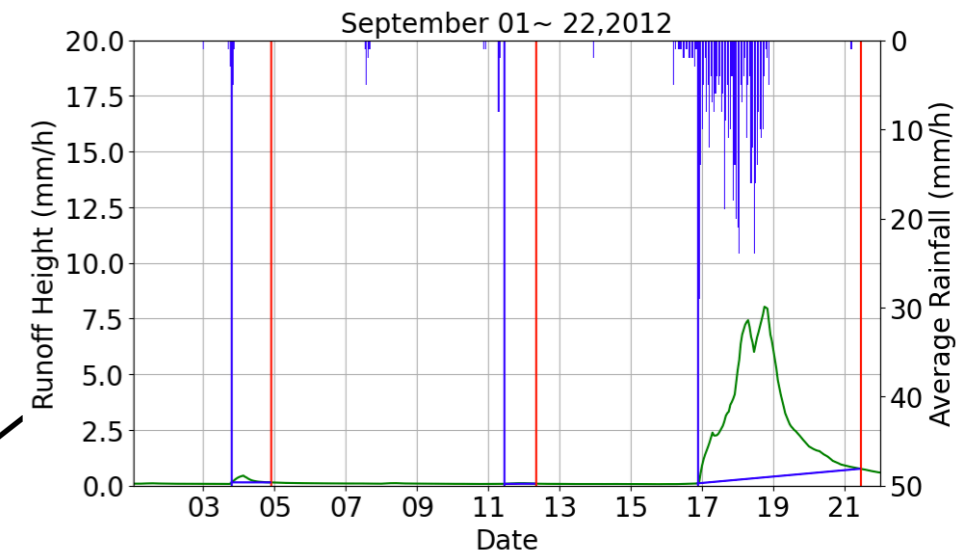
400 ≤ Total rainfall ≤ 500

# Runoff characteristics below the fitting curve have rained in previous rainfall events

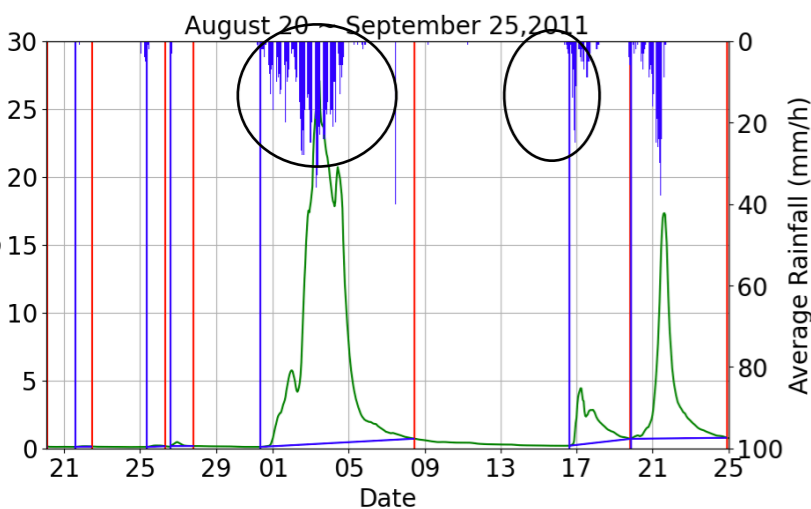
August 8~12,2003



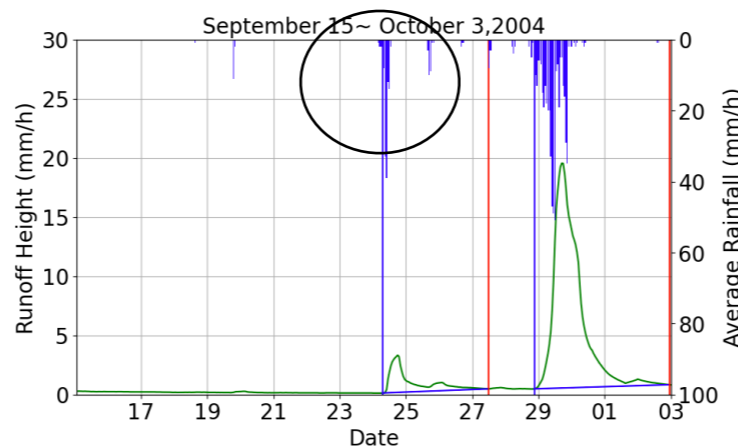
September 16~21,2012



September 19~24,2011

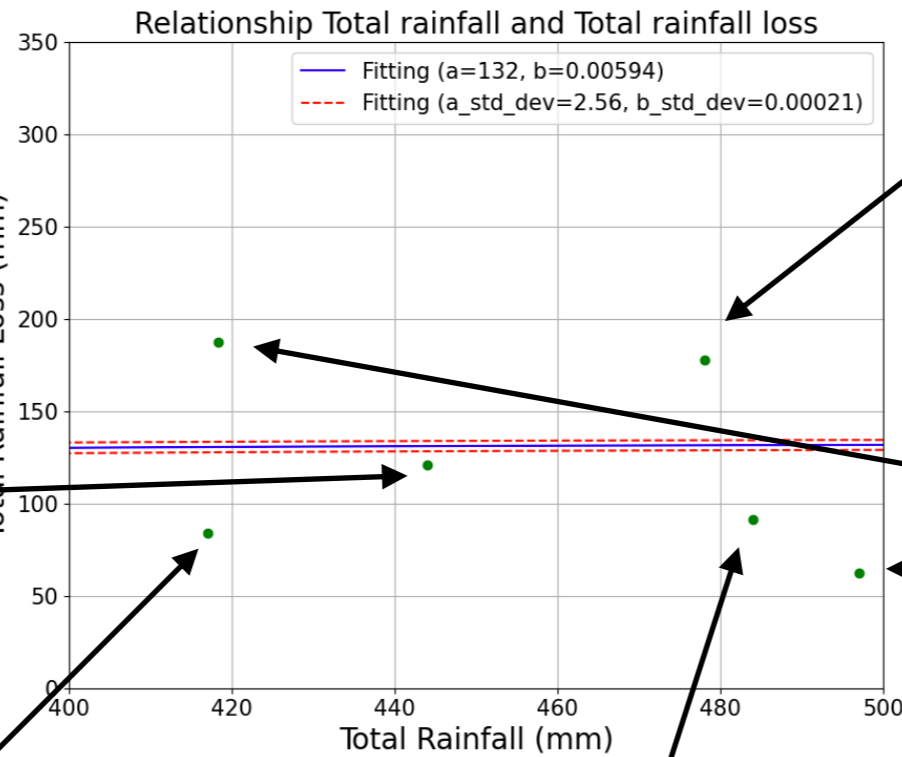
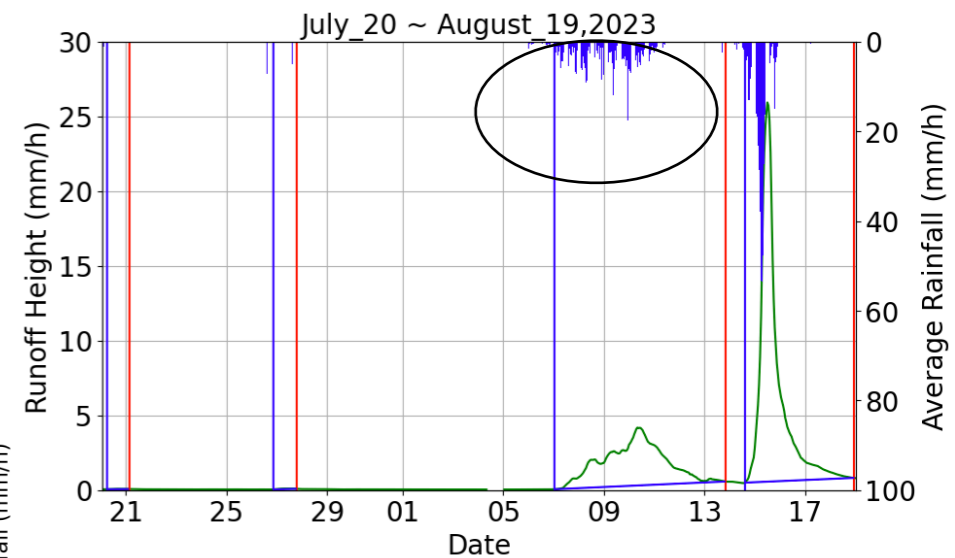


September 28~October 02,2004



August 7~13,2023

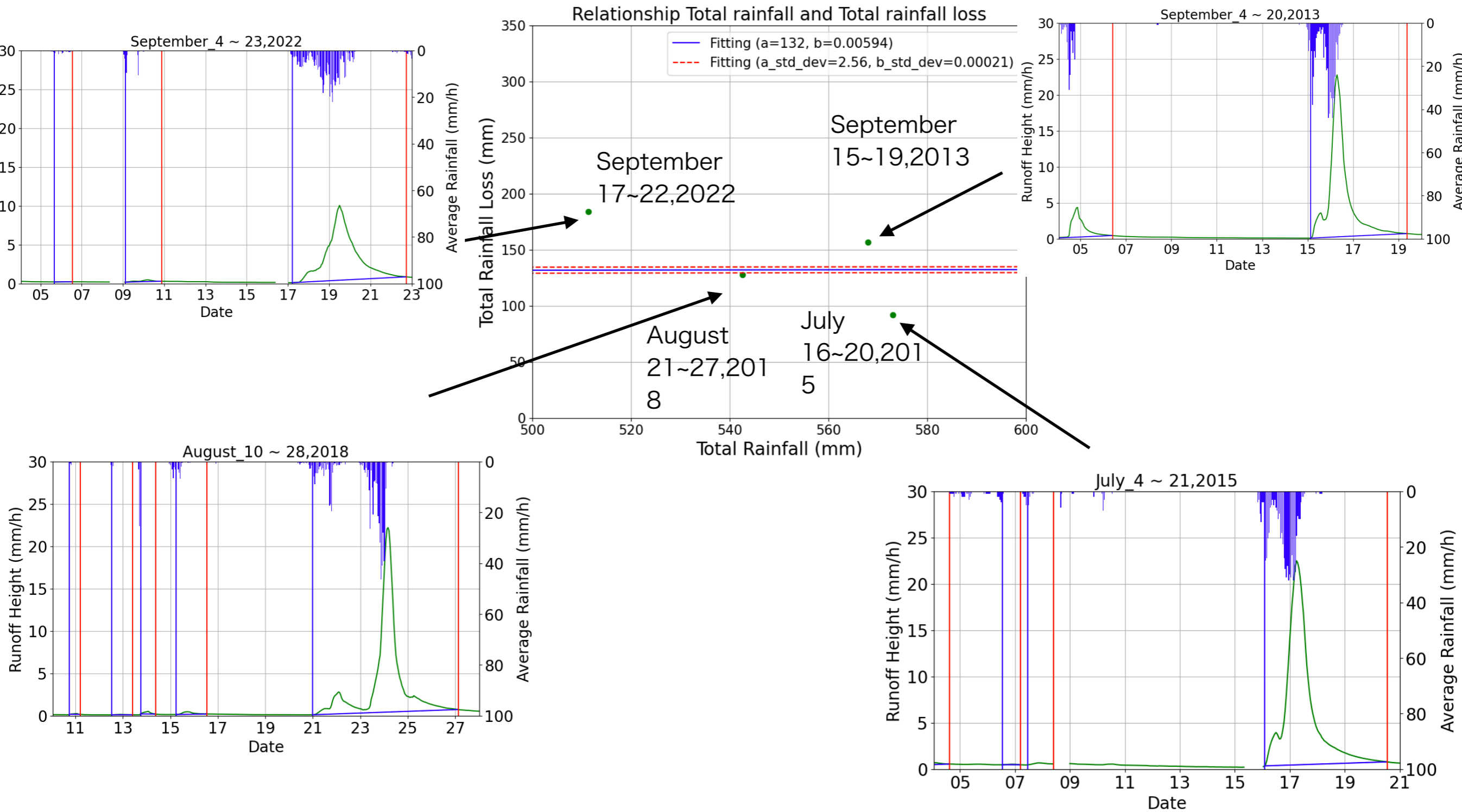
August 14~18,202



# Hachisu dam (Japan, Mie prefecture)

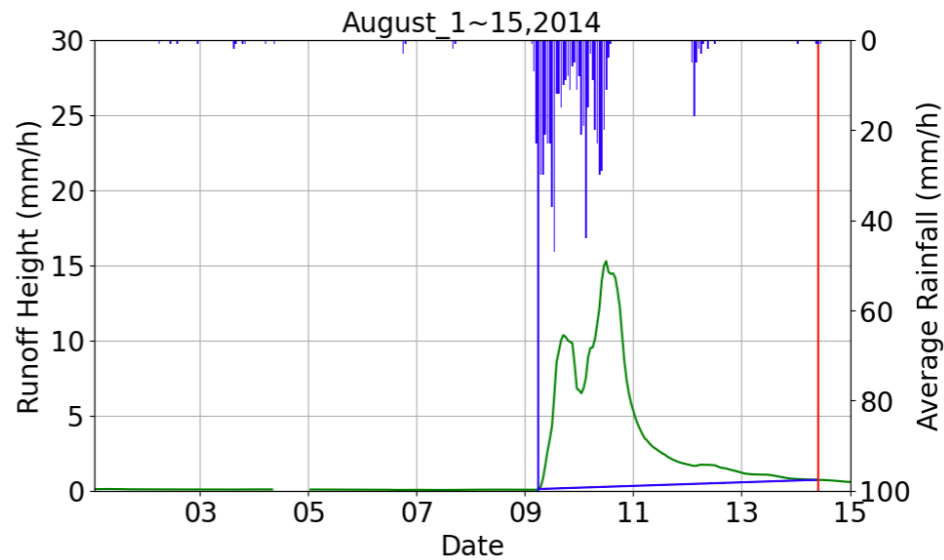
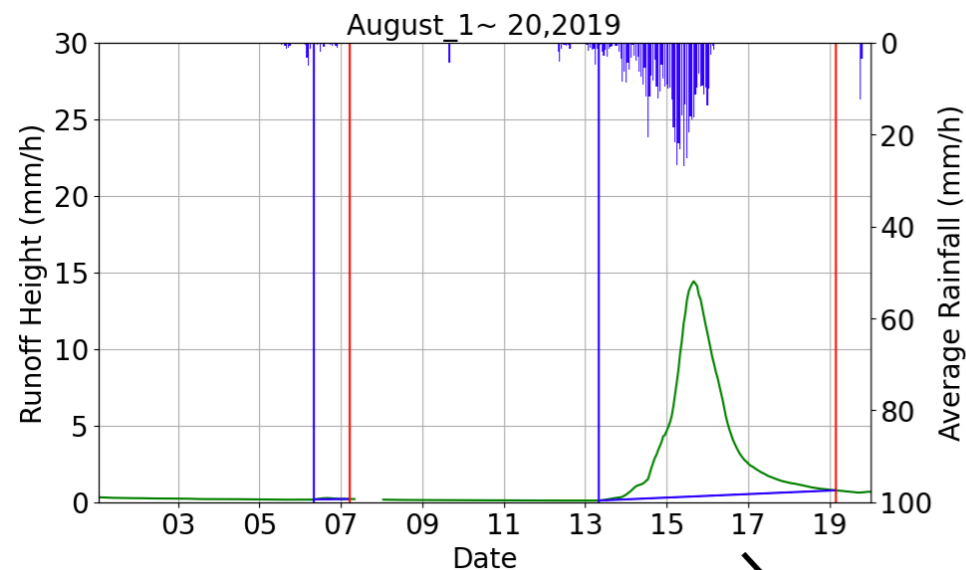
500 <= Total rainfall <= 600

**Runoff from July 4th to 21st, 2015 shows smaller values of lost rainfall than that in the fitting curve, but there was no significant rainfall in the previous rainfall event**

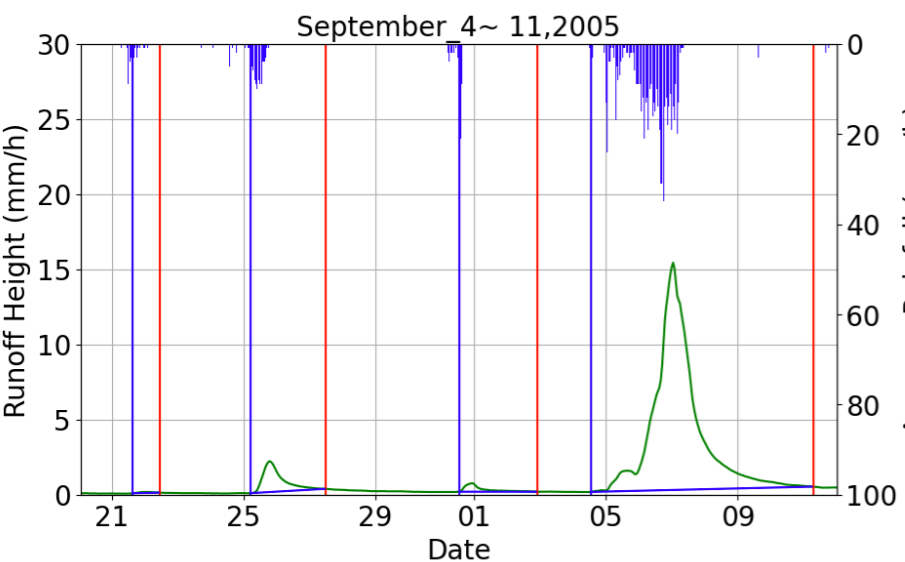
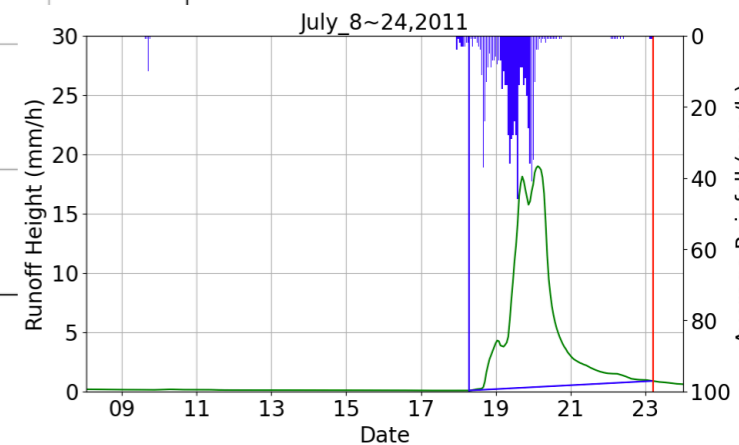
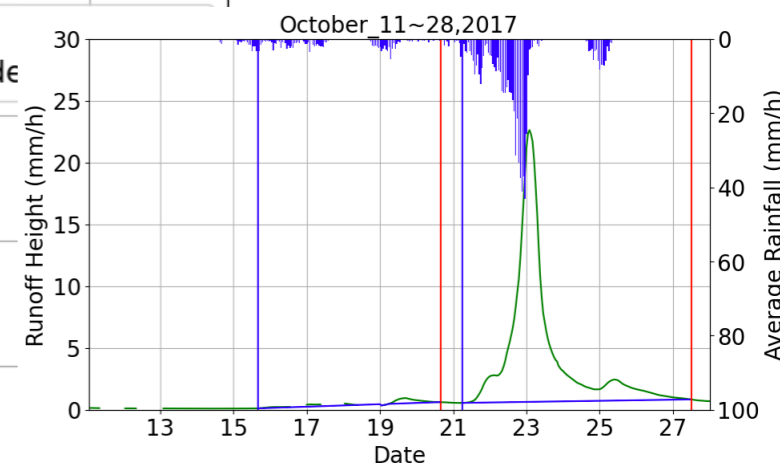
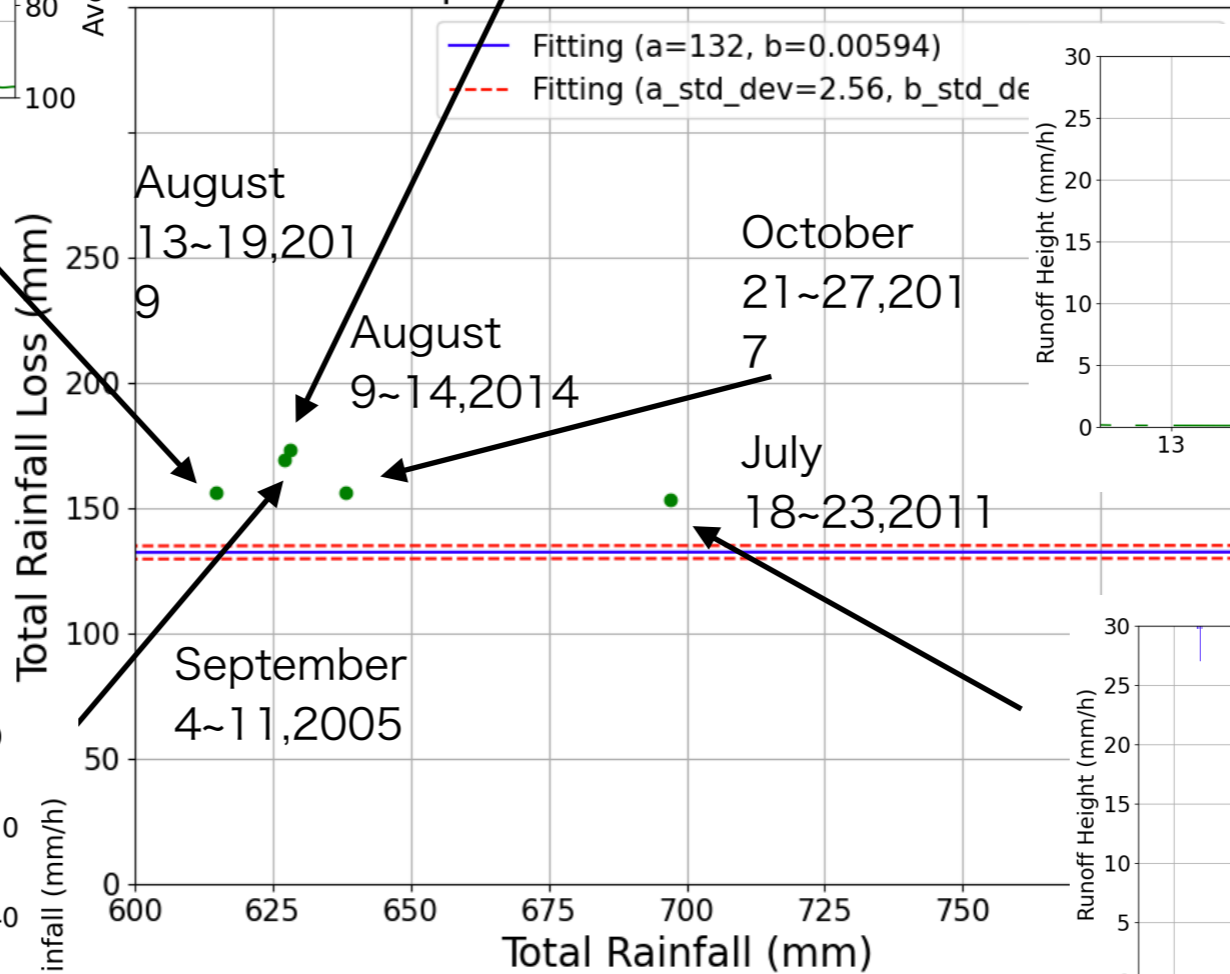


# Hachisu dam (Japan, Mie prefecture)

**600 ≤ Total rainfall ≤ 800**



Relationship Total rainfall and Total rainfall loss

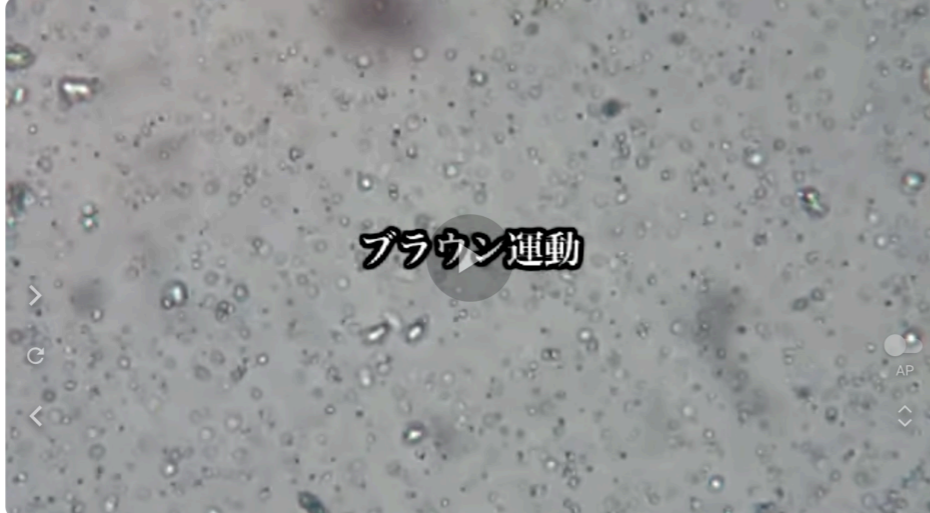
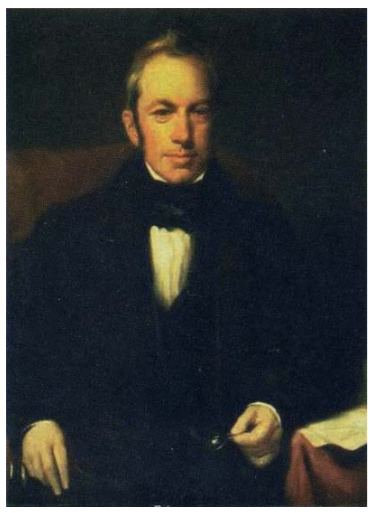




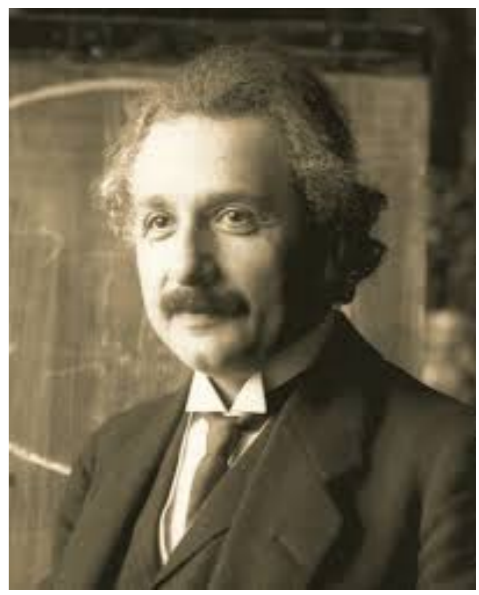
# 4.A method for mathematically evaluating the range of uncertainty

Wikipedia

Robert Brown (1773~1858)



The phenomenon of Brownian motion was theorized, and it was shown that molecules exist.



Einstein (1879~1955)

He discovered that the movement of microparticles in water, which was thought to be biological, is a physical phenomenon.

Langevin equation

$$m \frac{dv}{dt} = -\mu v(t) + f(t)$$

Average force  
Random force

$m$ :mass,  $v$ :velocity,  $\mu$ :coefficient

Ito stochastic differential equation

$$dx(t) = \mu(x, t)dt + \sum_{j=1}^n \sigma_j(x, t)dw_j$$

Fokker-Planck equation

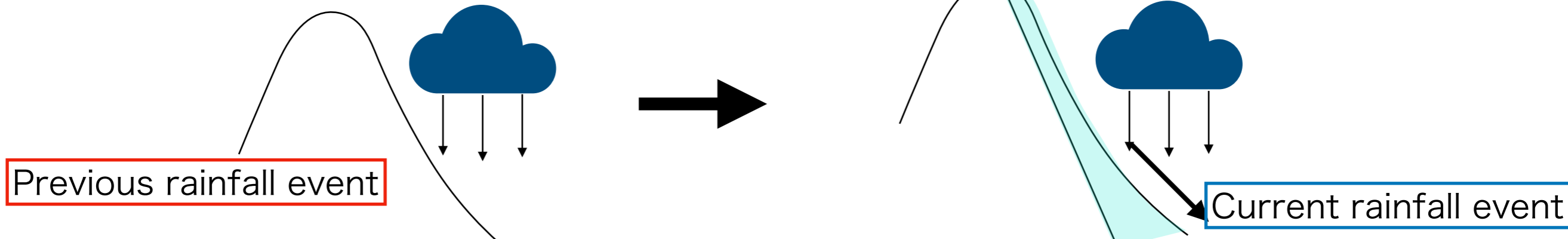
$$\frac{\partial p(x, t)}{\partial t} = -\frac{\partial (\mu(x, t)p(x, t))}{\partial t} + \frac{1}{2} \frac{\partial^2 \left( \sum_{j=1}^n \sigma_j^2(x, t)p(x, t) \right)}{\partial x^2}$$

# 5. Relationship between two uncertainties 10

Uncertainty due to limitations of observations and uncertainty in soil wetness are independent

uncertainty in soil wetness

Uncertainty due to limitations of observations



The effective rainfall term can be decomposed as a linear sum of the effective rainfall from observations, the uncertainty in soil wetness, and the uncertainty in observations

$$r_e = \bar{r}_e + r' + r''$$

( $r'$ : the uncertainty in observations,  $r''$ : the uncertainty in soil wetness)

( $\sigma_i, T_i, dw$  is standard

$$\frac{dq_*}{dt} = \alpha q_*^\beta (r_e - q_*) \longrightarrow dq_* = \alpha q_*^\beta (\bar{r}_e - q_*) dt + \alpha q_*^\beta \sigma_1 \sqrt{T_w} dw + \alpha q_*^\beta \sigma_2 \sqrt{T_w} dw$$

deviation, time constant (0, dt) normal distribution)

**Two uncertainties are introduced into the traditional deterministic basic equation**

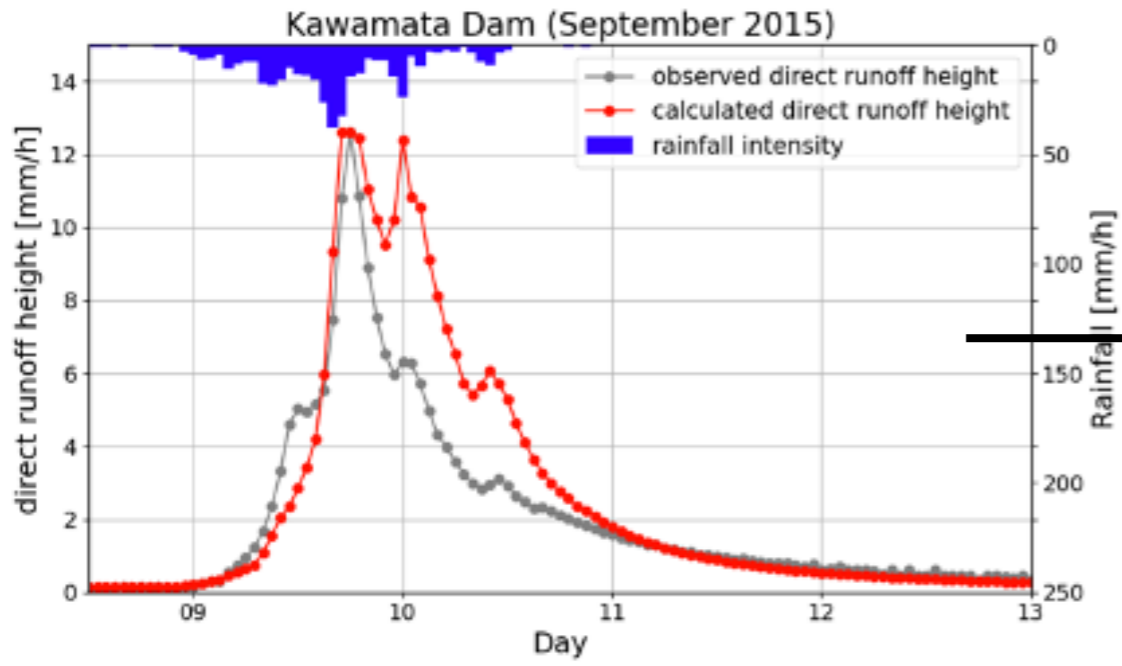
# 6. Stochastic rainfall runoff process

Deterministic method  $\longrightarrow$  Method considering uncertainty

$$\frac{dq_*}{dt} = \alpha q_*^\beta (r_e - q_*)$$

$\alpha, \beta$ : parameter  $q_*$ : runoff

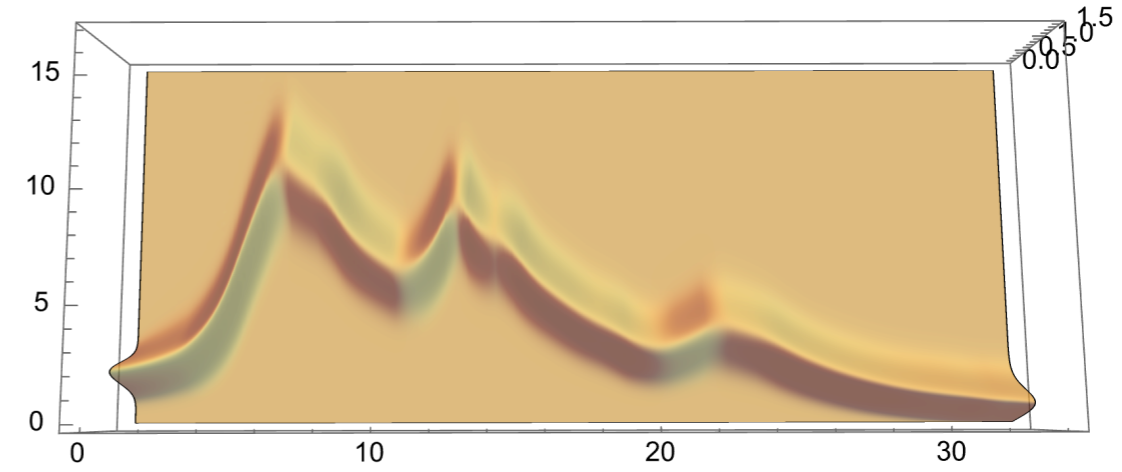
$r_e$ : effective rainfall



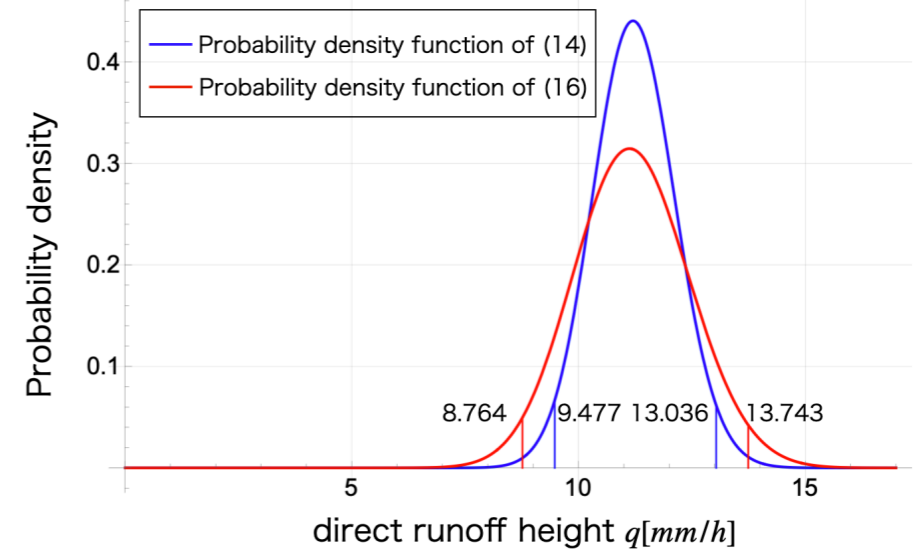
Conventional deterministic models settle on one outflow

$$\frac{\partial p(q_*, t)}{\partial t} = - \frac{\partial (a_0 q_*^\beta (r_e - q_*) p(q_*, t))}{\partial t} + \frac{1}{2} \frac{\partial^2 ((a_0 q_*^\beta \sigma_1(q_*, t))^2 p(q_*, t))}{\partial q_*^2} + \frac{1}{2} \frac{\partial^2 ((a_0 q_*^\beta \sigma_2(q_*, t))^2 p(q_*, t))}{\partial q_*^2}$$

Probability Density Function of Direct Runoff Height  $p(q, t)$



Probability density function of peak direct runoff height



Now that the uncertainty of rainfall information become known, it is possible to evaluate runoff heights taking the uncertainty into account.