

Comparison of historical and speleothem (paleo)climate records from northwest Yucatán (Mexico)

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1 Introduction: The Postclassic Maya



- Fig. 1** Mayapán, capital of the Postclassic Maya, located in the northwest Yucatán.
- Fig. 2** Cenote Ch'en Mul is located at the edge of the Temple of Ch'en Mul and next to the Pyramid of Kukulcan.
- Fig. 3** Stalagmite MAYA-22-7 in situ next to S.B.
- The Postclassic Maya city of Mayapán became a regional capital between 1000 and 1100 CE. At its peak, it housed 15-20,000 inhabitants and was governed by factions of rival noble families (the Cocom and the Xiu)
 - Recent paleoclimate and archaeological studies have linked evidence of repeated drought conditions to times of violent conflict amongst these rival families. Between 1350-1450, an extended dry period culminated in site abandonment just prior to the Spanish conquest of the Yucatán from 1527 to 1697 CE (Kennett et al., 2022)
 - Located next to the Temple of Kukulcan, the Cenote Ch'en Mul was of great ritual and religious importance. Speleothems from this site have distinct potential to capture a localized climate signal and examine land use change (LUC) and human-climate-environment interactions

2 Methods

- Sample Collection:** MAYA-22-7 was collected in 2022 at Cenote Ch'en Mul
- U-Th Dating:** The chronology of MAYA-22-7 was constructed via 5 U-series dates (see Fig 4). Sample preparation, analytical chemistry, and mass spectrometry were conducted by S.C. at the Department of Earth Sciences, Oxford
- Stable Isotope Analysis:** 2000 powder samples were milled at 0.1 mm resolution (~3 samples/year) along 6 growth axes (see Fig 4). Stable isotopes ($\delta^{18}O$, $\delta^{13}C$) were measured via GasBench IRMS
- Trace Element Analysis:** Trace element-to-Ca ratios (Li, Mg, Al, Si, Ti, Mn, Fe, Cu, Zn, Sr, Ba, Y, La, Ce, Nd, Yb, Lu, Pb, Th, U) were measured via LA-ICP-MS along tracks parallel to stable isotope measurements

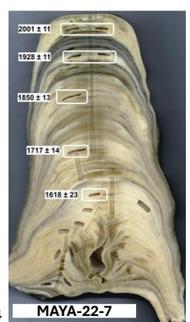
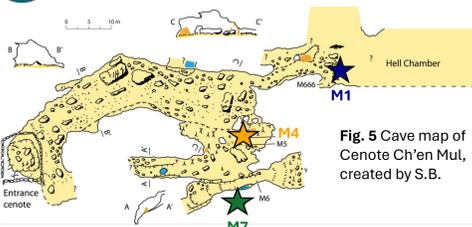


Fig. 4

3 Results: Analysis of multiple stalagmites at Mayapán



- 7 stalagmites were collected in Cenote Ch'en Mul: M1, M3, M4, M5, M6, M7, M666
- Analysis of multiple stalagmites from a single cave can help account for flow-path heterogeneity when inferring local paleoclimate signals, as climate interpretations can be distorted by karst processes (e.g., PCP, ventilation, overburden thickness)

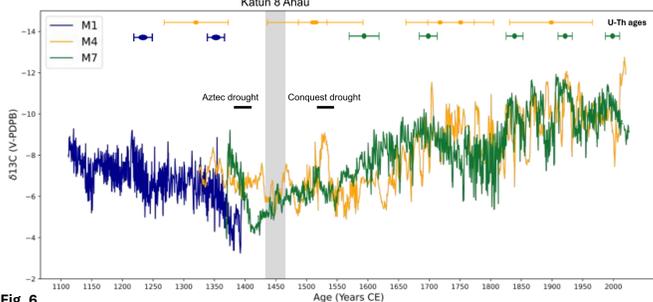


Fig. 6

- Despite variability, $d^{13}C$ shifts in M4 and M1 generally agree ($\pm 1\%$) across 1350-1980 CE
- From 1150-1400 CE, an increase from -8 to -4 ‰ is replicated across M1 and M4, suggesting a possible decrease in soil respiration
- This period terminates with depopulation, site abandonment and Katun 8 Ahau (a time of "demolition and destruction" according to Maya texts)

4 Results: Speleothem MAYA-22-7

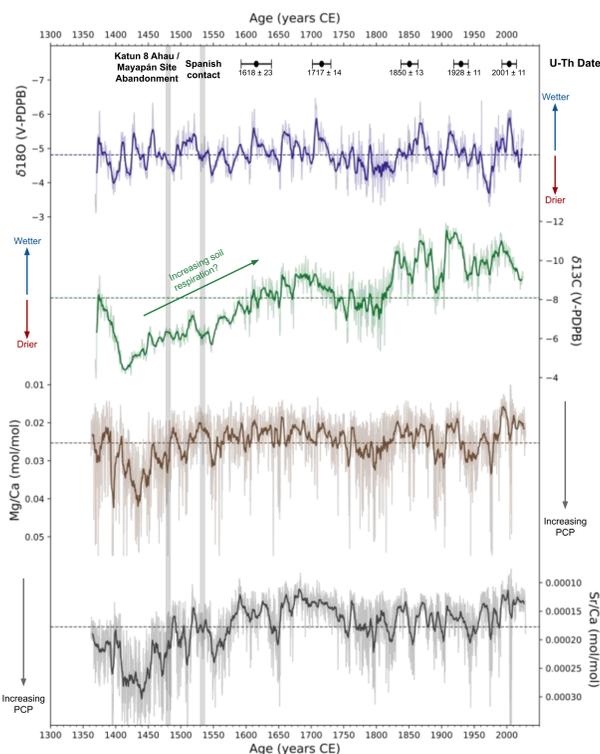


Fig. 7 MAYA-22-7 proxy records against age (age-depth model linearly interpolated; see Fig. 8).

Age-depth model

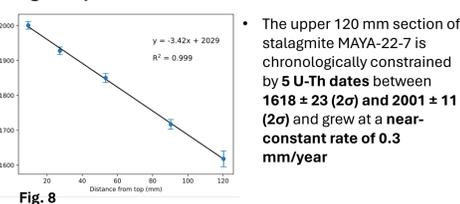


Fig. 8

- The upper 120 mm section of stalagmite MAYA-22-7 is chronologically constrained by 5 U-Th dates between 1618 ± 23 (2σ) and 2001 ± 11 (2σ) and grew at a near-constant rate of 0.3 mm/year
- $\delta^{18}O$ record capture abrupt shifts (1-2 ‰) above/below baseline, tracking changes in precipitation amount ("Amount Effect") above the cave
- 180 to 120 mm dft: $\delta^{13}C$ values steadily shift from -5 to -9 ‰, replicated by a muted shift in Mg/Ca and in Sr/Ca, indicating an increase in soil respiration (increased plant activity and/or increased soil cover), possibly related to reforestation after Maya dispersal
- 139 to 87 mm dft: $\delta^{13}C$ values demonstrate distinct 2 ‰ shifts from -10 to -8 ‰ and frequent (3-5 year) repeated oscillations, suggesting a strong, rapid response of the overlying ecosystem's carbon budget to climate variability, perhaps related to the El Niño/Southern Oscillation (ENSO)

Prior Calcite Precipitation

- The 1-year and 5-year cross-plots of trace element values (not shown) are consistent with a PCP control; the slope falls within the estimated slope related to prior calcite precipitation ($m=0.7-1$) (Sinclair et al. 2011)
- This indicates that PCP reflects the degree of water infiltration into the cave (e.g., increasing PCP = higher aridity)

5 Results: Capturing the historic record

- MAYA-22-7 captures historic droughts with ± 2 years precision, additionally capturing an extensive dry period spanning 1725-1830 CE in both the $d^{13}C$ and $d^{18}O$
- This extensive dry period, amongst earlier dry intervals, is well-replicated by YOK-1 $d^{13}C$, a speleothem from the southeast Yucatán, which provided a paleoclimate background for the relationship between civil conflict and climate change at Mayapán (Kennett et al., 2022)

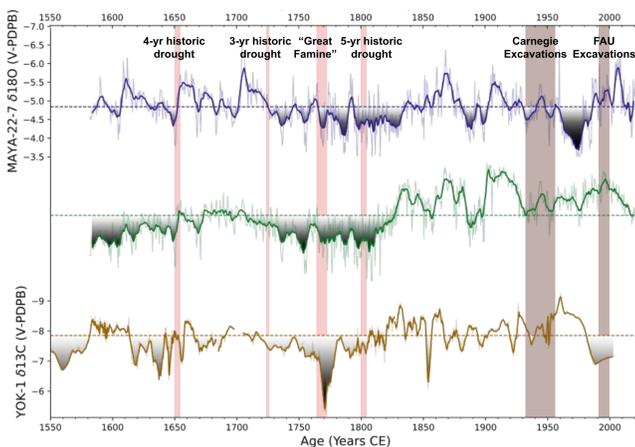


Fig. 9. Historic droughts documented via Spanish conquest records, as compiled by Hoggarth et al., 2017.

4-year historic drought of 1650-1654

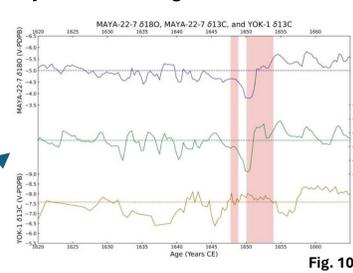


Fig. 10

"Great Famine" of 1765-1773 and 5-year historic drought of 1800-1805

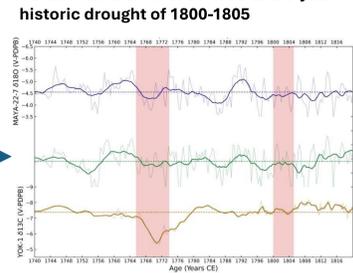


Fig. 11

6 Results: Instrumental validation

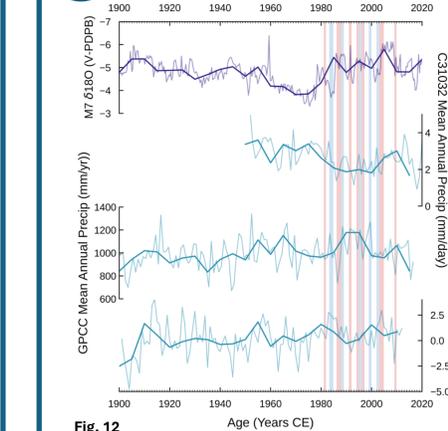


Fig. 12

Fig. 13 (left) MAYA-22-7 against instrumental records: C31032 (local meteorological station), GPCP (NOAA meteorological station compilation), and MXDA PDSI (GPCC-derived Palmer Drought Severity Index, where $>+4$ represents very wet conditions and <-4 represents an extreme drought).

- Under the existing age-depth model, MAYA-22-7 performs poorly in capturing the instrumental record
- MAYA-22-7 5-year mean is offset from 5-year meteorological station data (15-yr offset) and 5-year NOAA GPCP (5-year offset)
- From 1980-2011, MAYA-22-7 does not consistently capture annual variations in precipitation (De la Barreda 2020; horizontal red shading represents a dry period and blue represents wet)

7 Results: Regional proxy comparison

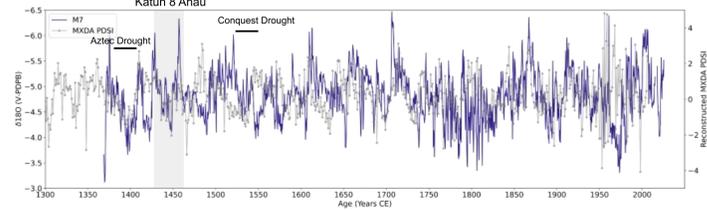


Fig. 13. MXDA PDSI reconstructed from tree-ring-record compilation across Central Mexico (Stahle et al., 2011) compared against MAYA-22-7 $d^{18}O$. Despite a lack of tree-ring records in the Yucatán, MAYA-22-7 $d^{18}O$ displays good agreement.

8 Next steps

- Employ Hierarchical Dynamic Time Warping to match proxy records of disparate resolution
- Introduce wavelet analysis to identify periodicity and forcings
- Compare greyscale variations to dO_{18} (capture archaeological excavations, flood layers)
- Verify age-depth model through the analysis of 8 additional U-Th dates
- Analyze lime plaster used as Maya paving material to decouple impact on speleothem trace element ratios

9 Conclusions

- We present MAYA-22-7, a new speleothem record from Mayapán, northwest Yucatán spanning 1318-2022 CE (linear age-depth model)
- Despite poor coherence with instrumental records, speleothem MAYA-22-7 exhibits good agreement with historic compilations of drought (see Fig 9) and the tree-ring-derived Palmer Drought Severity Index (see Fig 13)
- We provide initial insights into LUC (soil disturbance and reforestation) preceding and following Katun 8 Ahau and site abandonment at the terminal Postclassic

Corresponding historic records of drought

"In the year [16]50 at harvest time it was not understood that the harvest would be extremely brief, at the start of the year [16]51 word began to spread of the very great shortfall of maize for that year's sustenance"
-López de Cogolludo, 1957

Table 1. Population change in Yucatán from 1528 to 1781

Year CE	P	r	Per capita (%)	Population change (%)	Expected change
1609	176,320	30	6	18	+
1629	208,742	4	1	21	+
1643	209,180	23	-3.9	-48	-
1666	193,000	21	-6	-18	-
1684	99,042	12	-3.7	-56	+
1700	136,270	10	<1	-41	+
1710	156,788	26	<1	-19	+
1720 ^a	175,100	6	1.5	-5	+
1750 ^b	184,500	8	6.1	39	+
1772/1773 ^c	113,000	8	5.5	55	+

Population estimates via Hoggarth et al., 2017

"The pastures and cornfields that were not eaten by locusts dried or did not germinate because in 1769 there was a shortage and lack of rain. But the drought worsened in 1770. The population of the peninsula suffered imponderable havoc from the defect of water"
- María Isabel Campos Goenaga

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