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The effect of transient lateral internal gravity wave propagation on the resolved atmosphere in ICON/MS-GWaM

Supplementary material



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- ▶ a python toy code

<https://github.com/g-voelker/python-msgwam/tree/dev>

- ▶ the ICON model

<https://gitlab.dkrz.de/icon/icon-model>

- ▶ the ICON/MS-GWaM submodule

<https://github.com/DataWaveProject/msgwam>

Mathematical model formulation

1. scale height separation

$$H_p/H_\theta = \mathcal{O}(\epsilon^\alpha)$$

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2. small Rossby number

$$1 \gg \text{Ro} = \epsilon$$

▶ quasi-geostrophic regime

▶ hydrostatic scaling

$$H_w = \epsilon L_w$$

▶ stratification and rotation

$$f = \epsilon^{5-\alpha} N$$

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3. perturbation scale separation ϵ

$$(X, Y, Z, T) = \epsilon(x, y, z, t)$$

4. perturbation expansion in small parameter (not shown)

- ▶ non-dimensional compressible Euler equations
- ▶ multiscale theory with $(\mathbf{X}, T) = \epsilon(\mathbf{x}, t)$, $\epsilon = \text{Ro}$

$$\mathbf{v} = \sum_{j=0}^{\infty} \epsilon^j \mathbf{v}_0^{(j)} + \Re \sum_{\beta} \sum_{n=1}^{\infty} \epsilon^n \mathbf{V}_{\beta}^{(n)} e^{\phi_{\beta}/\epsilon} \quad (1)$$

$$\theta = \sum_{j=0}^{\alpha} \epsilon^j \bar{\Theta}^{(j)}(Z) + \epsilon^{\alpha+1} \sum_{j=0}^{\infty} \epsilon^j \Theta_0^{(j)} + \epsilon^{\alpha+1} \Re \sum_{\beta} \sum_{n=1}^{\infty} \epsilon^n \Theta_{\beta}^{(n)} e^{\phi_{\beta}/\epsilon} \quad (2)$$

$$\pi = \sum_{j=0}^{\alpha} \epsilon^j \bar{\Pi}^{(j)}(Z) + \epsilon^{\alpha+1} \sum_{j=0}^{\infty} \epsilon^j \Pi_0^{(j)} + \epsilon^{\alpha+2} \Re \sum_{\beta} \sum_{n=1}^{\infty} \epsilon^n \Pi_{\beta}^{(n)} e^{\phi_{\beta}/\epsilon} \quad (3)$$

- ▶ field = reference + mean flow + superimposing waves β
- ▶ wave properties $\omega_{\beta} = -\partial_T \phi_{\beta}$, $\mathbf{k}_{\beta} = \nabla_{\mathbf{x}} \phi_{\beta}$

- ▶ dispersion relation

$$\dot{\xi} = (\partial_t + \mathbf{c}_g \cdot \nabla_r) \xi$$

$$\omega = \Omega(\mathbf{X}, T, \mathbf{k}) = \mathbf{k} \cdot \mathbf{U}_0^{(0)}(\mathbf{X}, T) \pm \sqrt{\frac{f^2 m^2 + N^2(k^2 + l^2)}{k^2 + l^2 + m^2}} \quad (4)$$

- ▶ eikonal equations

$$(\partial_T + \mathbf{c}_g \cdot \nabla_{\mathbf{X}}) \mathbf{k} = \dot{\mathbf{k}} = -\nabla_{\mathbf{X}} \Omega \quad \mathbf{c}_g = \nabla_{\mathbf{k}} \Omega \quad (5)$$

- ▶ polarization relations (omitted)
- ▶ wave action conservation (no TRI)

$$\mathcal{N} = \mathcal{N}(\mathbf{X}, T, \mathbf{k})$$

$$0 = \partial_T \mathcal{N} + \nabla_{\mathbf{X}} \cdot (\mathbf{c}_g \mathcal{N}) + \nabla_{\mathbf{k}} \cdot (\dot{\mathbf{k}} \mathcal{N}) + S \quad (6)$$

$$0 = \nabla_{\mathbf{X}} \cdot \mathbf{c}_g + \nabla_{\mathbf{k}} \cdot \dot{\mathbf{k}} \quad (7)$$

- ▶ changes in position

$$\dot{\lambda} = \frac{c_{g\lambda}}{r \cos \phi} \quad \dot{\phi} = \frac{c_{g\phi}}{r}$$

$$\dot{\xi} = (\partial_t + \mathbf{c}_g \cdot \nabla_r) \xi$$

$$\dot{r} = c_{gr}$$

- ▶ changes in wavenumber

$$\begin{aligned} \dot{k}_\lambda &= -\frac{1}{r \cos \phi} \left(\mathbf{k} \cdot \partial_\lambda \mathbf{U} + \frac{|\mathbf{k}_h|^2 \partial_\lambda N^2}{2\hat{\omega} |\mathbf{k}|^2} \right) && -\frac{k_\lambda}{r} c_{gr} + \frac{k_\lambda \tan \phi}{r} c_{g\phi} \\ \dot{k}_\phi &= -\frac{1}{r} \left(\mathbf{k} \cdot \partial_\phi \mathbf{U} + \frac{|\mathbf{k}_h|^2 \partial_\phi N^2 + k_r^2 \partial_\phi f^2}{2\hat{\omega} |\mathbf{k}|^2} \right) && -\frac{k_\phi}{r} c_{gr} - \frac{k_\lambda \tan \phi}{r} c_{g\lambda} \\ \dot{k}_r &= -\left(\mathbf{k} \cdot \partial_r \mathbf{U} + \frac{|\mathbf{k}_h|^2 \partial_r N^2}{2\hat{\omega} |\mathbf{k}|^2} \right) && +\frac{k_\lambda}{r} c_{g\lambda} + \frac{k_\phi}{r} c_{g\phi} \end{aligned}$$

- ▶ wave action conservation

$$0 = \partial_t \mathcal{N} + \mathbf{c}_g \cdot \nabla_r \mathcal{N} + \dot{\mathbf{k}} \cdot \nabla_{\mathbf{k}} \mathcal{N} + S$$

- ▶ relax assumption of geostrophic, horizontal, and hydrostatic leading order mean-flow
- ▶ direct flux approach

$$D_t \theta + N^2 w = -\nabla_h \cdot \overline{\mathbf{u}'\theta'} \quad (8)$$

$$D_t \mathbf{u} + f \mathbf{e}_z \times \mathbf{u} = -c_p \bar{\theta} \nabla_h \pi - \frac{1}{\bar{\rho}} \nabla \cdot (\bar{\rho} \overline{\mathbf{u}'\mathbf{v}'}) + \frac{f}{g} \mathbf{e}_z \times \overline{\mathbf{u}'\mathbf{b}'} \quad (9)$$

entropy-flux convergence, momentum-flux convergence, elastic term

- ▶ for details see Achatz, Ribstein, et al. 2017; Wei et al. 2019

$$\bar{\rho} \overline{u'u'} = \mathcal{A} \left(kc_{gx} - f^2 \frac{kc_{gx} + lc_{gy}}{f^2 - \hat{\omega}^2} \right) \quad \bar{\rho} \overline{v'u'} = lc_{gx} \mathcal{A} \quad (10)$$

$$\bar{\rho} \overline{u'v'} = kc_{gy} \mathcal{A} \quad \bar{\rho} \overline{v'v'} = \mathcal{A} \left(lc_{gy} - f^2 \frac{kc_{gx} + lc_{gy}}{f^2 - \hat{\omega}^2} \right) \quad (11)$$

$$\bar{\rho} \overline{u'w'} = kc_{gz} \mathcal{A} \left(\frac{f^2}{f^2 - \hat{\omega}^2} \right) \quad \bar{\rho} \overline{v'w'} = lc_{gz} \mathcal{A} \left(\frac{f^2}{f^2 - \hat{\omega}^2} \right) \quad (12)$$

$$\bar{\rho} \overline{u'\theta} = \frac{\bar{\theta} f N^4 k_h^2}{g \hat{\omega} K^2 m (\hat{\omega}^2 - f^2)} l \mathcal{A} \quad \bar{\rho} \overline{v'\theta} = - \frac{\bar{\theta} f N^4 k_h^2}{g \hat{\omega} K^2 m (\hat{\omega}^2 - f^2)} k \mathcal{A} \quad (13)$$

$$\frac{f}{g} \overline{u'b'} = \frac{f^2 N^4 k_h^2}{\bar{\rho} g \hat{\omega} k^2 m (\hat{\omega}^2 - f^2)} l \mathcal{A} \quad \frac{f}{g} \overline{v'b'} = - \frac{f^2 N^4 k_h^2}{\bar{\rho} g \hat{\omega} k^2 m (\hat{\omega}^2 - f^2)} k \mathcal{A} \quad (14)$$

- ▶ wave action conservation (no TRI)

$$\dot{\xi} = (\partial_t + \mathbf{c}_g \cdot \nabla_r) \xi$$

$$0 = \partial_T \mathcal{A}_\beta + \nabla_{\mathbf{x}}(\mathbf{c}_{g,\beta} \mathcal{A}_\beta) \quad (14)$$

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- ▶ wave action in *phase space*

$$\mathcal{A}_\beta = \int_{\mathbb{R}^3} \mathcal{N}(\mathbf{X}, T, \mathbf{k}) \delta(\mathbf{k} - \mathbf{k}_\beta) d\mathbf{k} \quad \mathbf{k}_\beta = \mathbf{k}_\beta(\mathbf{X}, T) \quad (15)$$

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- ▶ we then find for $\mathcal{N}(\mathbf{X}, T, \mathbf{k})$

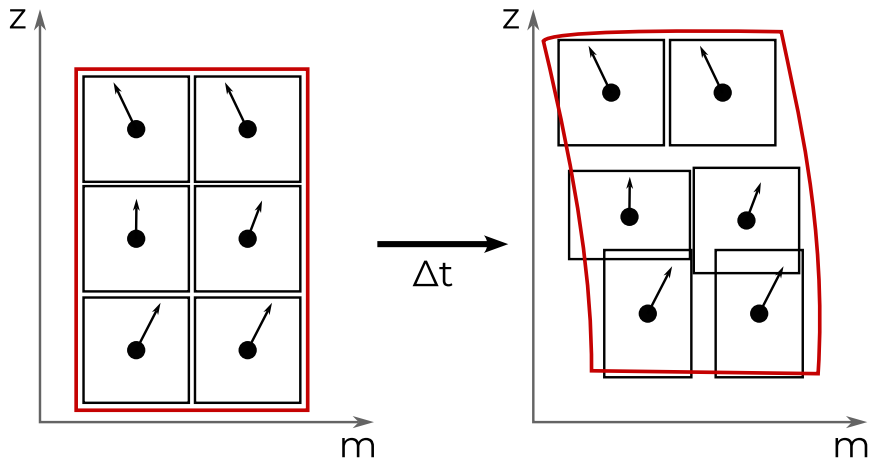
$$0 = \partial_T \mathcal{N} + \nabla_{\mathbf{x}} \cdot (\mathbf{c}_g, \mathcal{N}) + \nabla_{\mathbf{k}} \cdot (\dot{\mathbf{k}} \mathcal{N}) \quad (16)$$

$$0 = \nabla_{\mathbf{x}} \cdot \mathbf{c}_g + \nabla_{\mathbf{k}} \cdot \dot{\mathbf{k}} \quad (17)$$

- ▶ phase space volume is conserved

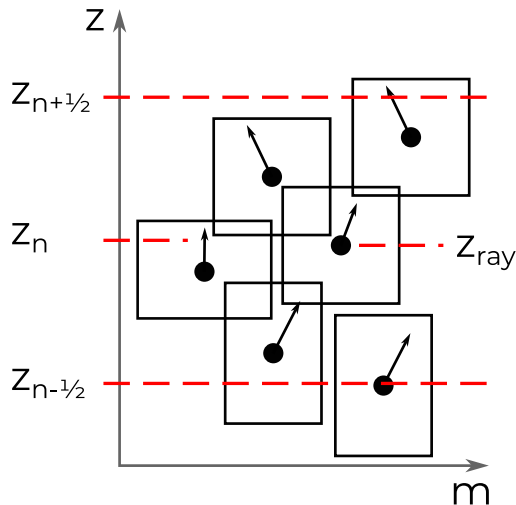
Implementation of the coupled Lagrangian model

$$0 = \nabla_{\mathbf{x}} \cdot \mathbf{c}_{g,\beta} + \nabla_{\mathbf{k}_\beta} \cdot \dot{\mathbf{k}}_\beta \quad (18)$$



- mean \rightarrow wave:
interpolation

$$N_{ray} = N_n + \frac{N_{n+1} - N_n}{z_{n+1} - z_n} (z_{ray} - z_n)$$

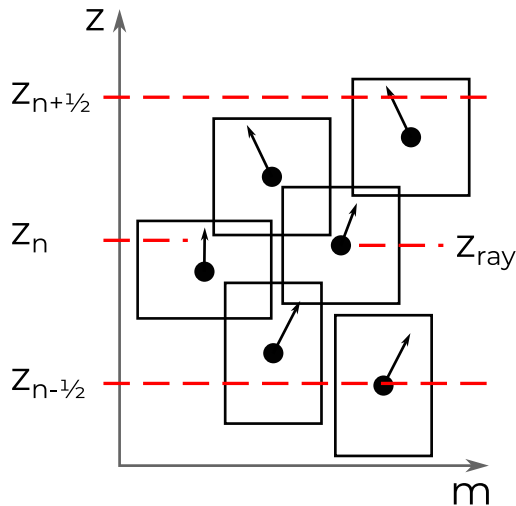


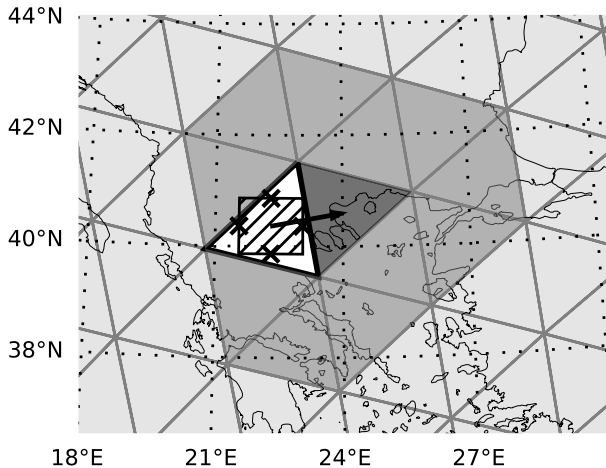
- ▶ mean \rightarrow wave:
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$$N_{ray} = N_n + \frac{N_{n+1} - N_n}{z_{n+1} - z_n} (z_{ray} - z_n)$$

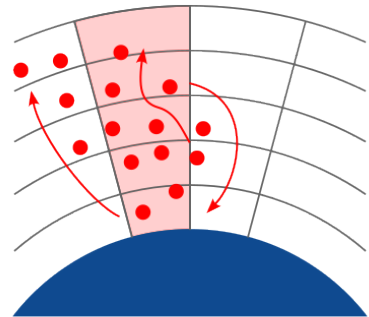
- ▶ wave \rightarrow mean:
integration and
projection

$$D_t \mathbf{U} \propto -\frac{1}{\bar{\rho}} \nabla \cdot \int_{\mathbb{R}^3} (\hat{\mathbf{c}}_g \mathbf{k} \mathcal{N}) d\mathbf{k}$$



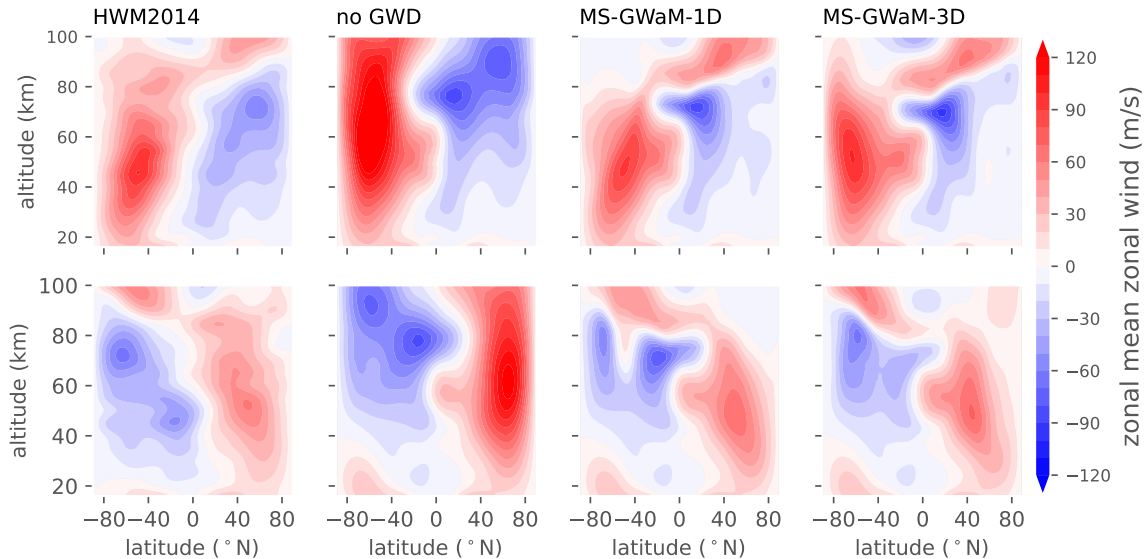


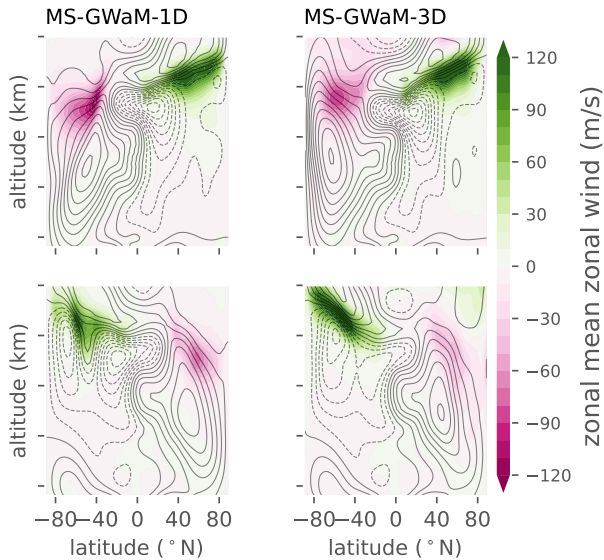
- ▶ Eulerian dynamical core
- ▶ Lagrangian wave model
- ▶ two-way coupling
- ▶ massively parallelized

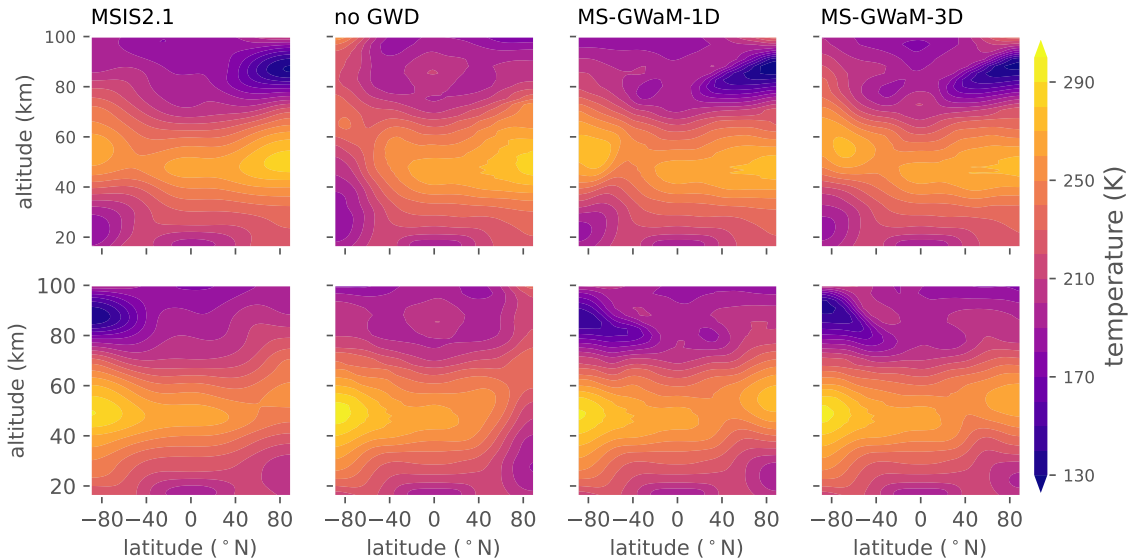


Comparing momentum flux distribution: 1D vs. 3D

Zonal mean zonal winds and temperatures



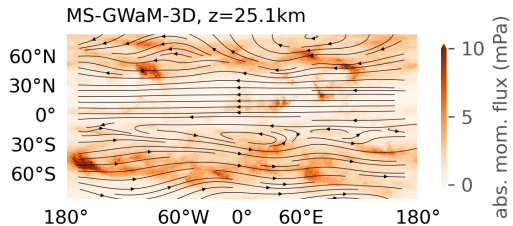
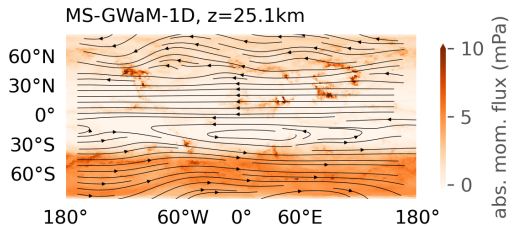




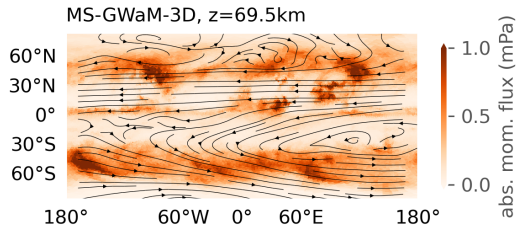
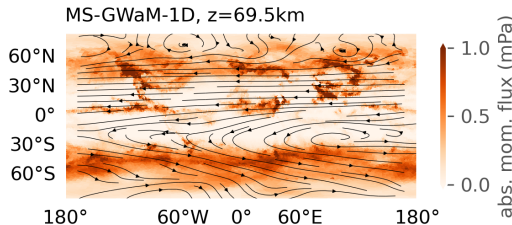
1D

3D

25km



70km

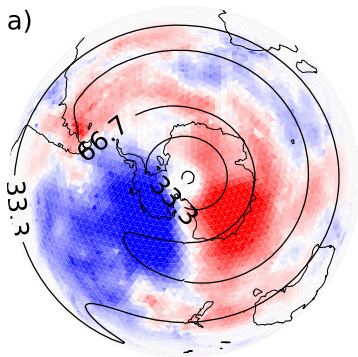


$$\mathcal{F}_{abs} = \bar{\rho}(\overline{u'w'^2} + \overline{v'w'^2})^{\frac{1}{2}}$$

June (mean of 91-98)

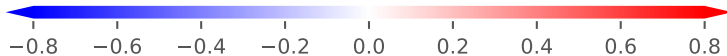
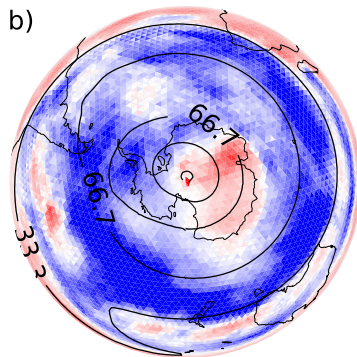
MS-GWaM 1D

a)

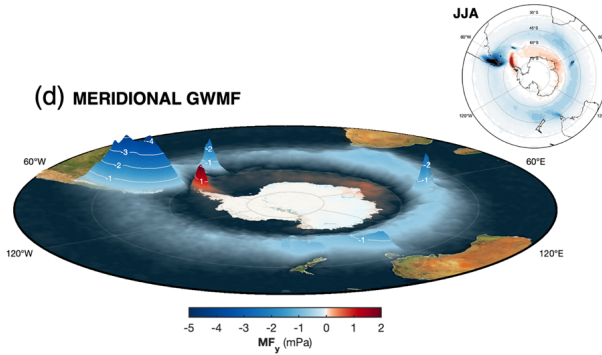


MS-GWaM 3D

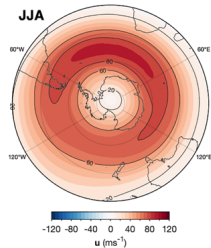
b)



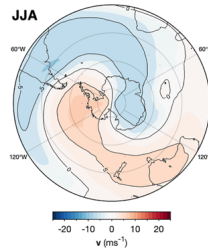
meridional momentum flux (mPa)

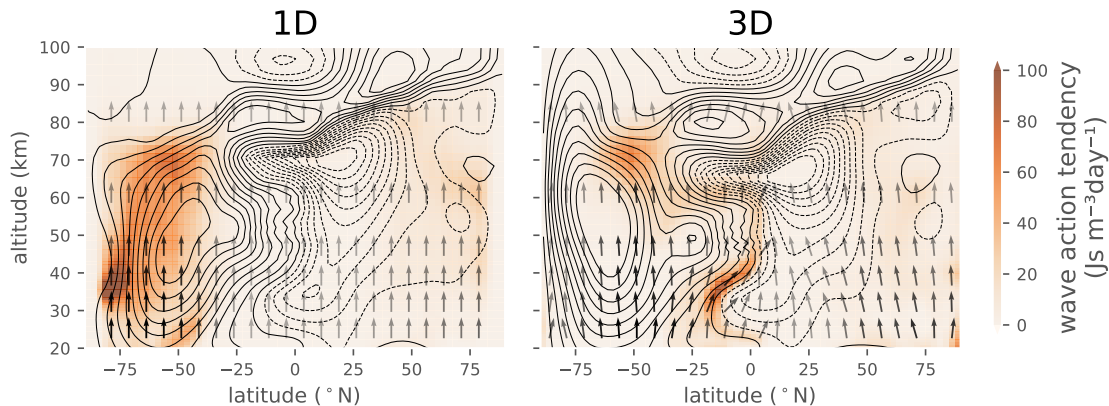


(g) ZONAL WIND

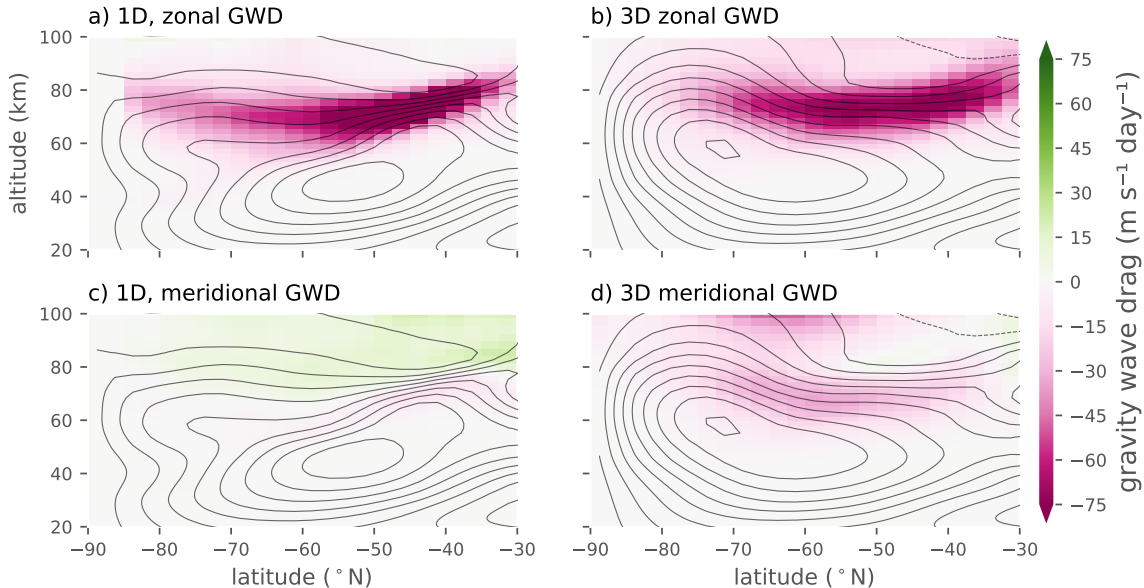


(h) MERIDIONAL WIND

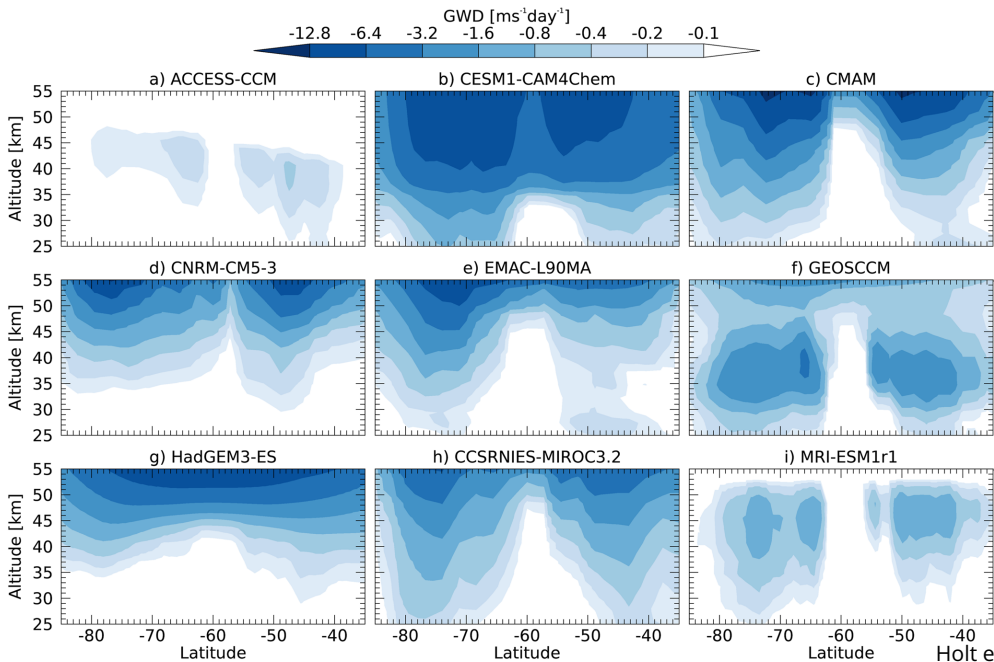


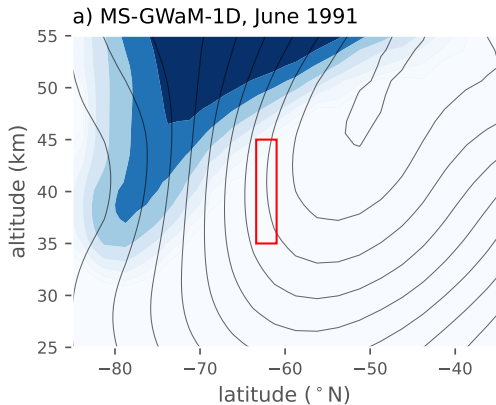
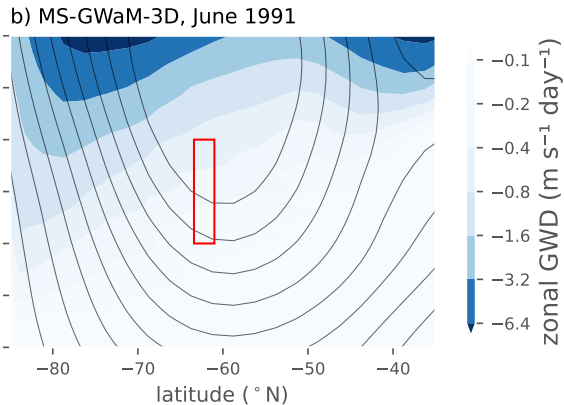


- refraction redistributes wave action dissipation

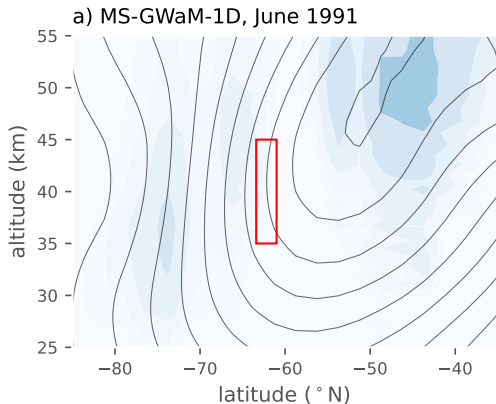
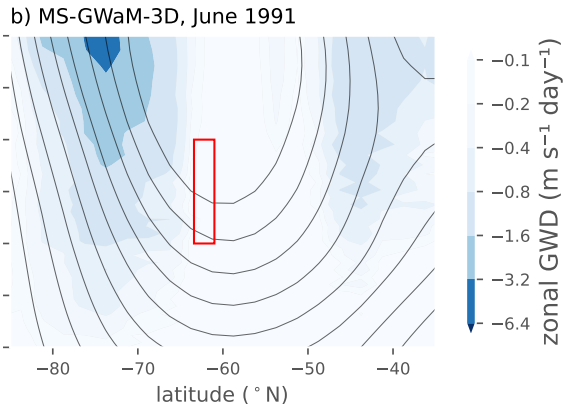


The (still) missing wave drag at 60°S



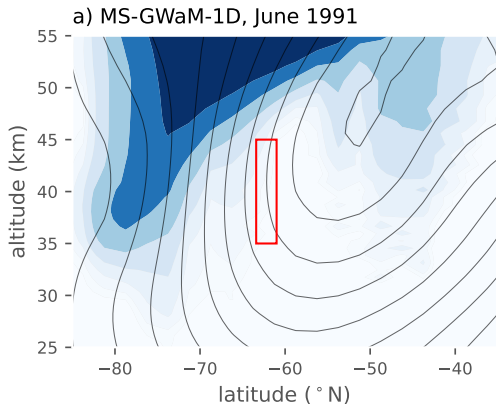
1D**3D**

► redistribution of non-orographic waves

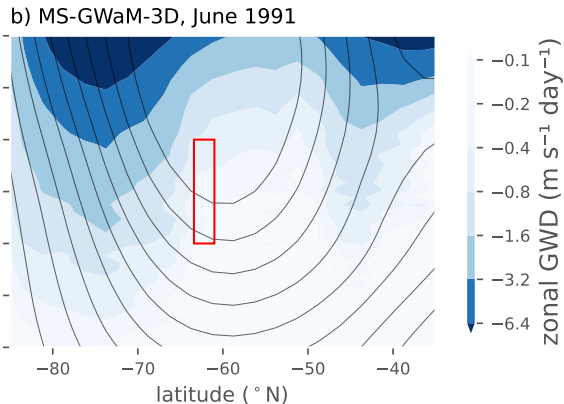
1D**3D**

► modified wind and SSO wave propagation

1D

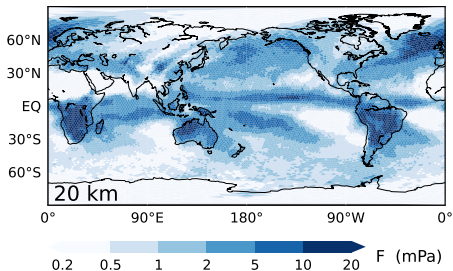
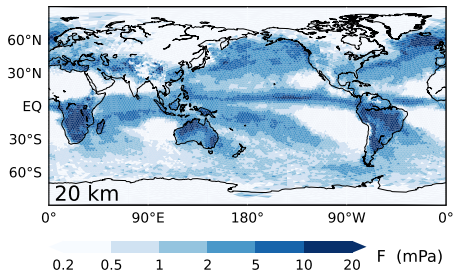


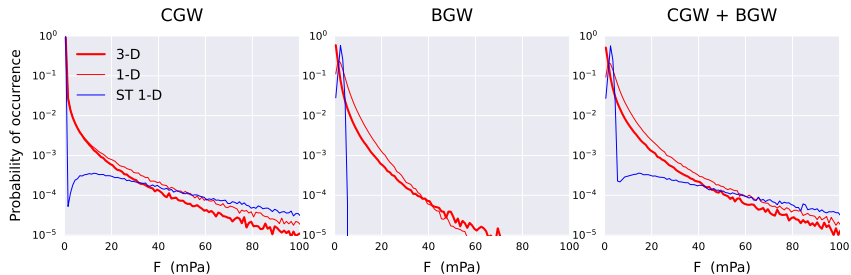
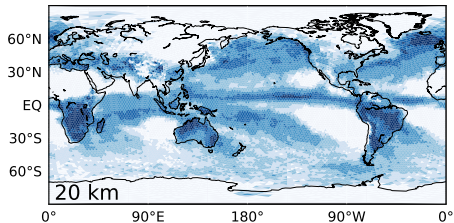
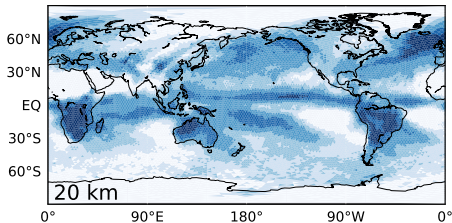
3D

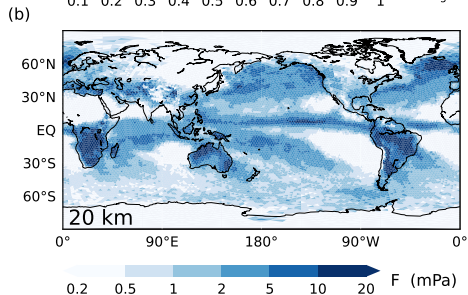
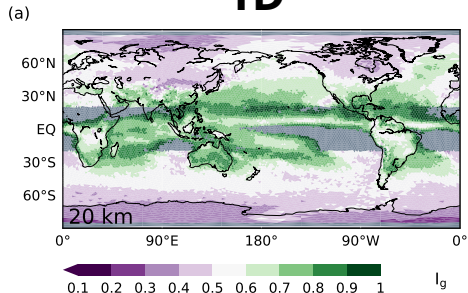
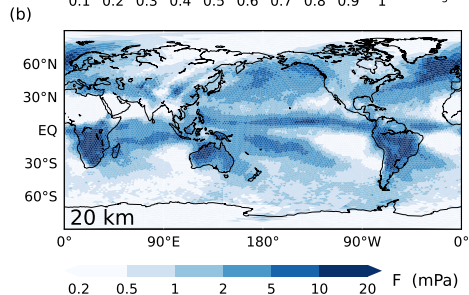
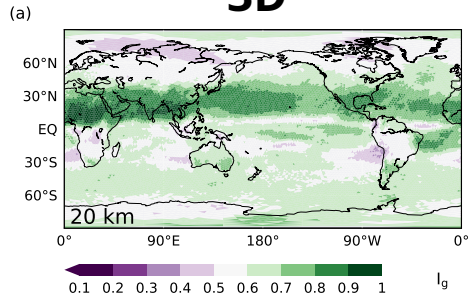


- ▶ 3D propagation of non-orographic waves alone does **not** solve it

Lateral propagation and IGW intermittency

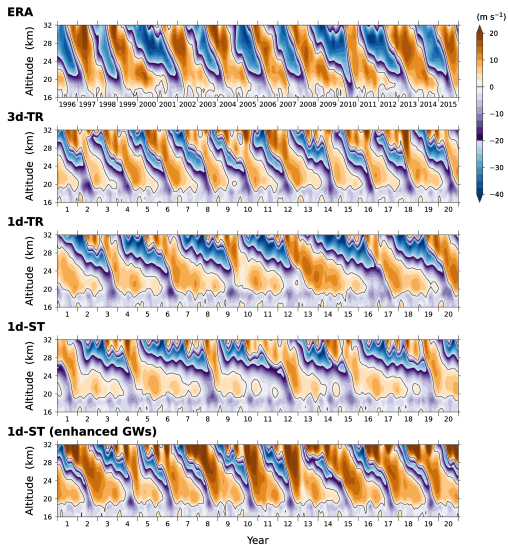


(statistics for $50\text{-}65^\circ\text{S}$)0.2 0.5 1 2 5 10 20 F (mPa)0.2 0.5 1 2 5 10 20 F (mPa)

1D**3D**

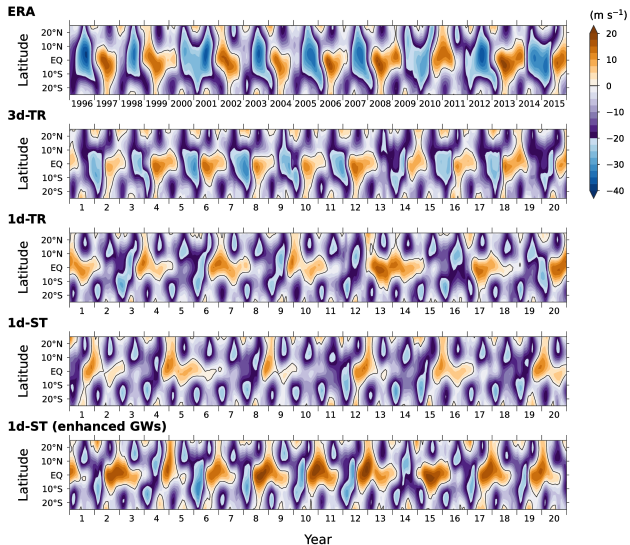
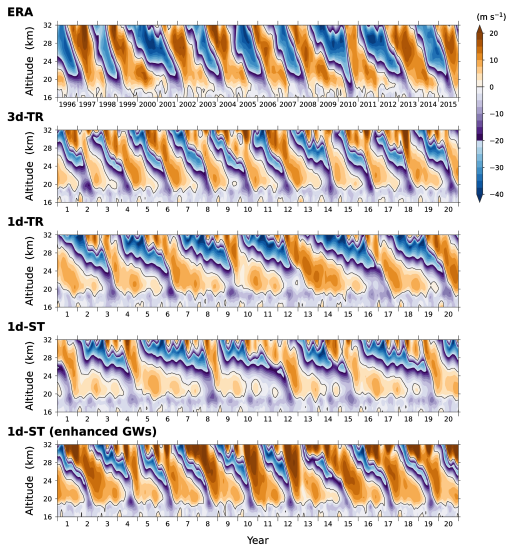
Simulating the QBO with MS-GWaM (details)

$$5^{\circ}\text{S} < \phi < 5^{\circ}\text{N}$$

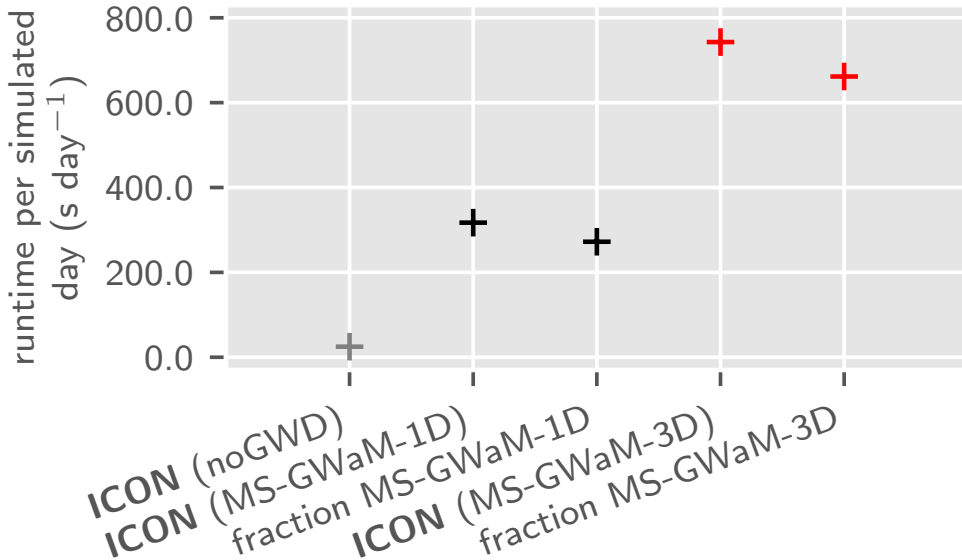


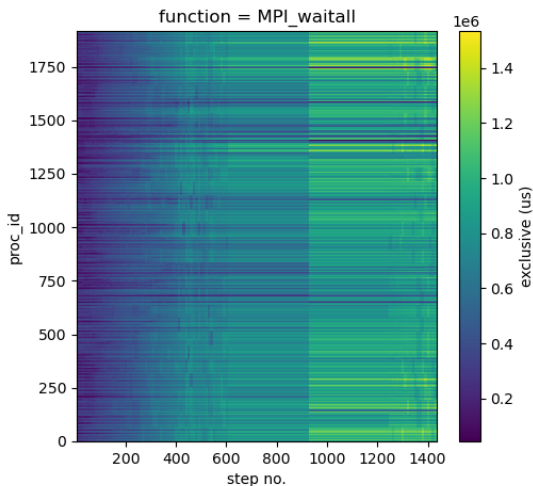
$5^{\circ}\text{S} < \phi < 5^{\circ}\text{N}$

$z=24\text{km}$

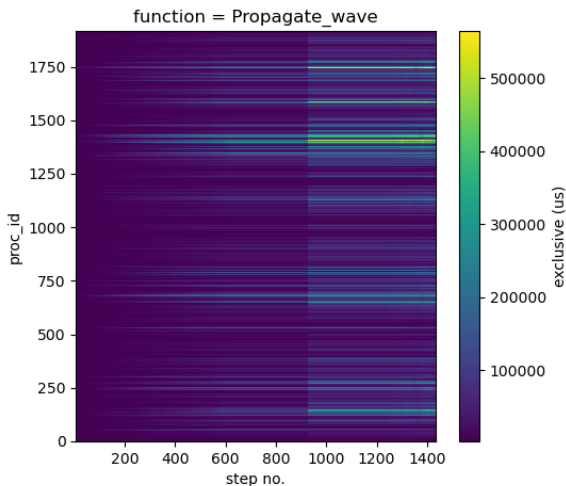


Notes on the model performance

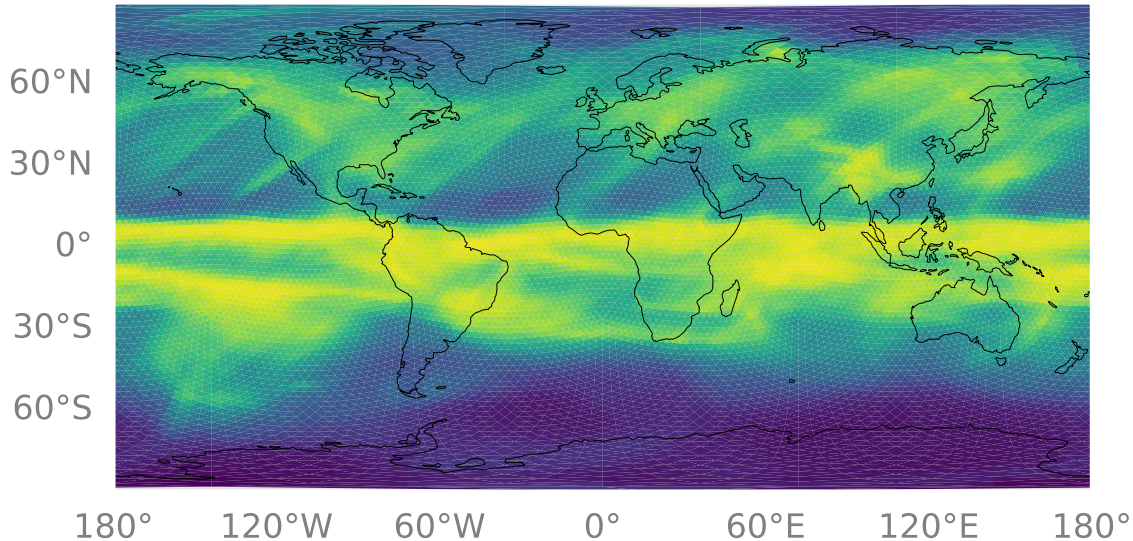


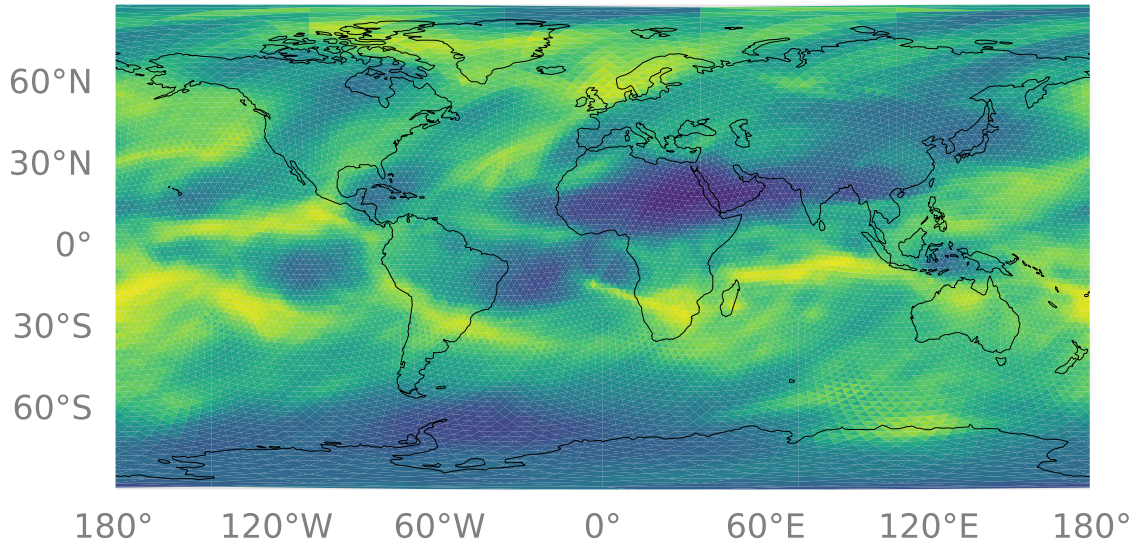


MPI barrier



computation





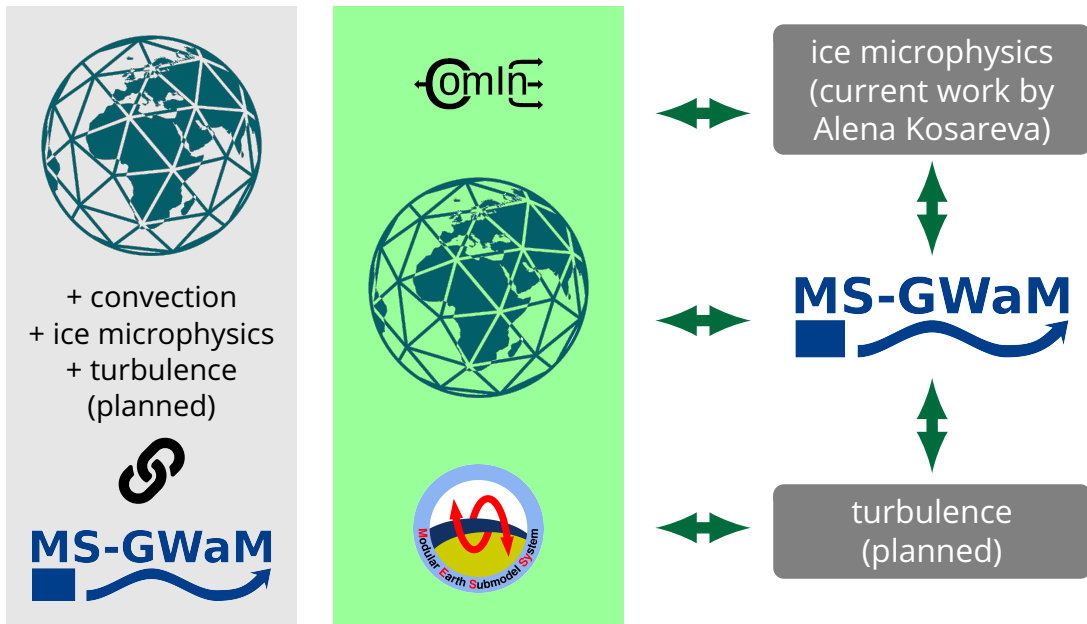
Towards a community-available model



+ convection
+ ice microphysics
+ turbulence
(planned)



MS-GWaM

Important references

- Achatz, U., B. Ribstein, F. Senf, and R. Klein (2017). "The Interaction between Synoptic-Scale Balanced Flow and a Finite-Amplitude Mesoscale Wave Field throughout All Atmospheric Layers: Weak and Moderately Strong Stratification". *Quarterly Journal of the Royal Meteorological Society* 143.702, pp. 342–361. DOI: [10.1002/qj.2926](https://doi.org/10.1002/qj.2926).
- Achatz, U., Y.-H. Kim, and G. S. Voelker (2023). *Multi-Scale Dynamics of the Interaction Between Waves and Mean Flows: From Nonlinear WKB Theory to Gravity-Wave Parameterizations in Weather and Climate Models*. DOI: [10.48550/arXiv.2310.07334](https://doi.org/10.48550/arXiv.2310.07334). *Journal of Mathematical Physics*, accepted: [arXiv:2310.07334](https://arxiv.org/abs/2310.07334).
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