Supplementary Material for

Spectral Albedo of Dusty Martian CO₂ Snow and Ice

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SNICAR-Adv4 Model

Since the model perceives snow as a mixture of independently dispersing ice grains suspended in an air medium, a layer of snow has the same bulk refractive index as air. Ice is described as independently dispersing air bubbles within a solid ice medium with spectrally variable refraction (Picard et al., 2016; Warren and Brandt, 2008).

CRISM dataset

CRISM observation ID FRT000075B1_01 (L_s = 303.4° of MY 28) of Wolf glacier (as described by Smith et al., 2022) is processed, analyzed and used for comparison with modeling results.

Modeling Results



Figure S1: Albedo variation with change in thickness of the topmost layer pure CO_2 snow (snow grain size: 5000 µm, density=327.15 kg/m³) is overlaying CO_2 ice (layer thickness= 10m, density= 1200 kg/m³)



Figure S2: Variation of pure CO_2 snow albedo with effective grain radius. Density of CO_2 snow assumed: 327.15 kg/m3)



Figure S3: Initial results on CO_2 glacier ice albedo variation with effective grain radius. Density of CO_2 ice assumed: 1000 kg/m³

The reflection by very dense ice is due to the Fresnel reflection. The ice spectral albedo has a peak at 3.5 μ m due to the reflectivity of ice at normal incidence based on the spectrally varying indices of refraction utilized in SNICAR-ADv4.



Figure S4: Initial results on variation of CO_2 glacier ice albedo with change in density of the ice for air bubble radius of 5000 μ m.

While there is no effect on spectral curve with variation in density of CO_2 snow, there is noticeable change in spectral profile of CO_2 glacier ice with change in density, keeping the air bubble size and layer thickness constant.

Note: The initial results (Figure S3 and S4) maybe not accurate as the model is still being improved for representing CO_2 glacial ice.

References

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