



# Comparison between Developing and Nondeveloping Disturbances for Tropical Cyclogenesis in Different Large-Scale Flow Patterns over the Western North Pacific

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encouraged

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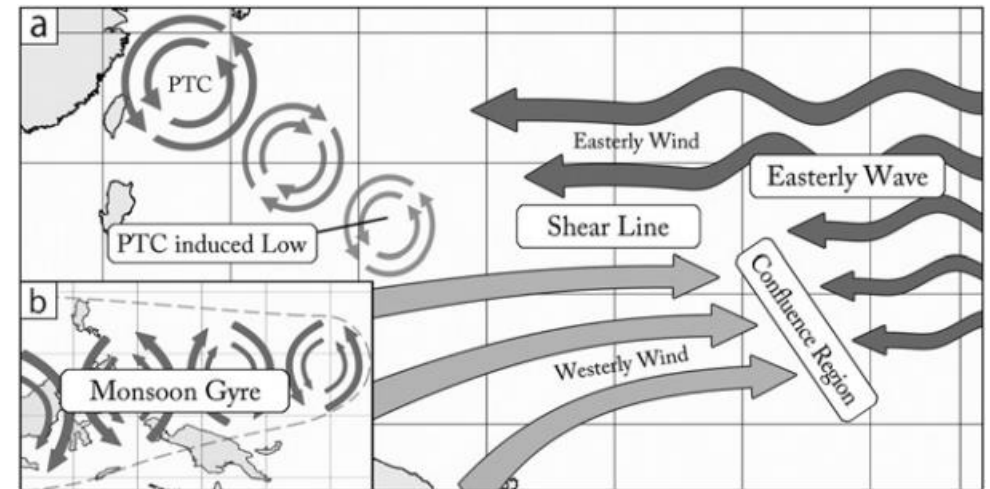
**Tropical cyclogenesis is particularly susceptible to the large-scale circulation.**

**Zehr (1992):** Monsoon-E, Monsoon-Stg,  
Monsoon-Wk, easterly pattern

**Ritchie and Holland (1999):** Rossby wave energy dispersion, Monsoon shear line (MS), Monsoon confluence (MC), Monsoon gyre (MG), Easterly wave

**Feng et al. (2014):** MS, MC, MG, reverse-oriented monsoon trough, trade easterly wind

**Most disturbances still failed to develop into a tropical depression (TD).**



A schematic image of the five flow patterns over the western North Pacific (Yoshida and Ishikawa, 2013)

**How to distinguish developing disturbances from nondeveloping ones in different large-scale circulation patterns?**



# Data



## International Best Track Archive for Climate Stewardship (IBTrACS)

- **genesis time and location for developing disturbances**
- The TC genesis time is defined as the time when the maximum sustained surface wind first reaches **11 m/s** and persists for **at least 24 hr**.

## European Centre for Medium-Range Weather Forecasts' (ECMWF's) next-generation reanalysis ERA5

- **0.25°×0.25°, 6-hourly, from June to November in 2000-2019**
- **disturbance identification, large-scale environmental parameters**

## GridSat-B1 infrared brightness temperature

- **0.07°×0.07°, 3-hourly**
- **MCS information**

A new set of criteria for five large-scale patterns

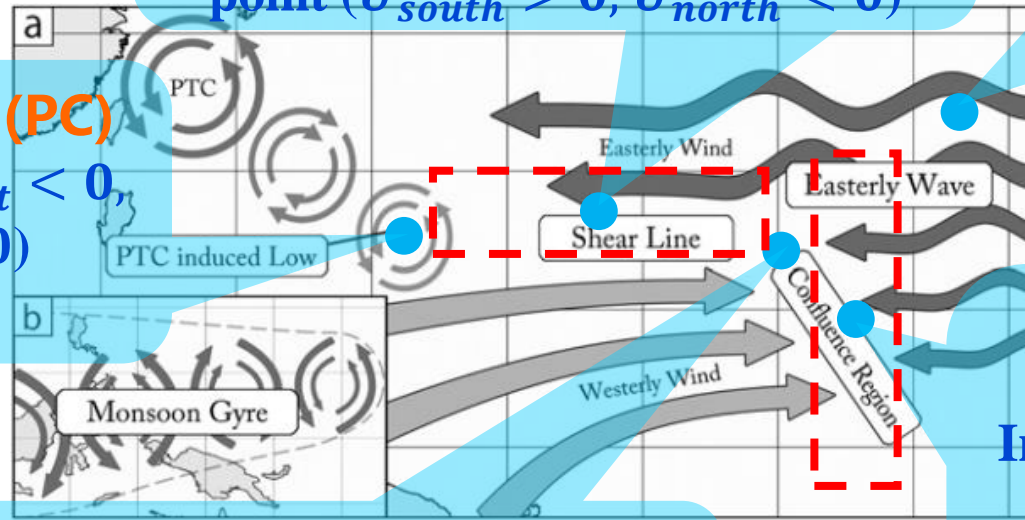
**Pre-existing Cyclone (PC)**  
 cyclone ( $U_{south} > 0, V_{west} < 0$   
 $U_{north} < 0, V_{east} > 0$ )  
 radius  $\geq 400$  km

**Zonal Wind Shear Line (SL)**  
 In a  $30^\circ \times 7.5^\circ$  west-east band, at least 50% longitudes have a shear point ( $U_{south} > 0, U_{north} < 0$ )

**Easterly Wave (EW)**  
 easterly wind ( $U < 0$ )  
 area radius  $\geq 500$  km

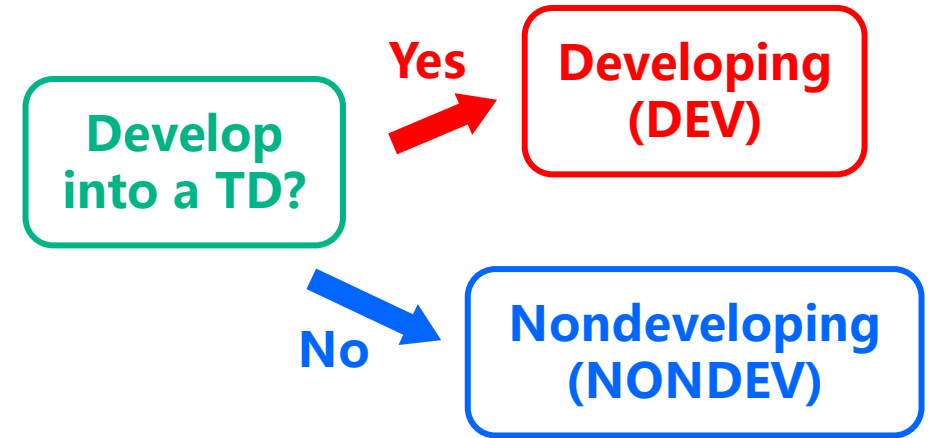
**Zonal Wind Convergence (CON)**  
 In a  $7.5^\circ \times 30^\circ$  north-south band, at least 12.5% latitudes have a convergence point ( $U_{west} > 0, U_{east} < 0$ )

**Mixed Zonal Wind Convergence and Shear Line (CON-SL)**



## Disturbance detection

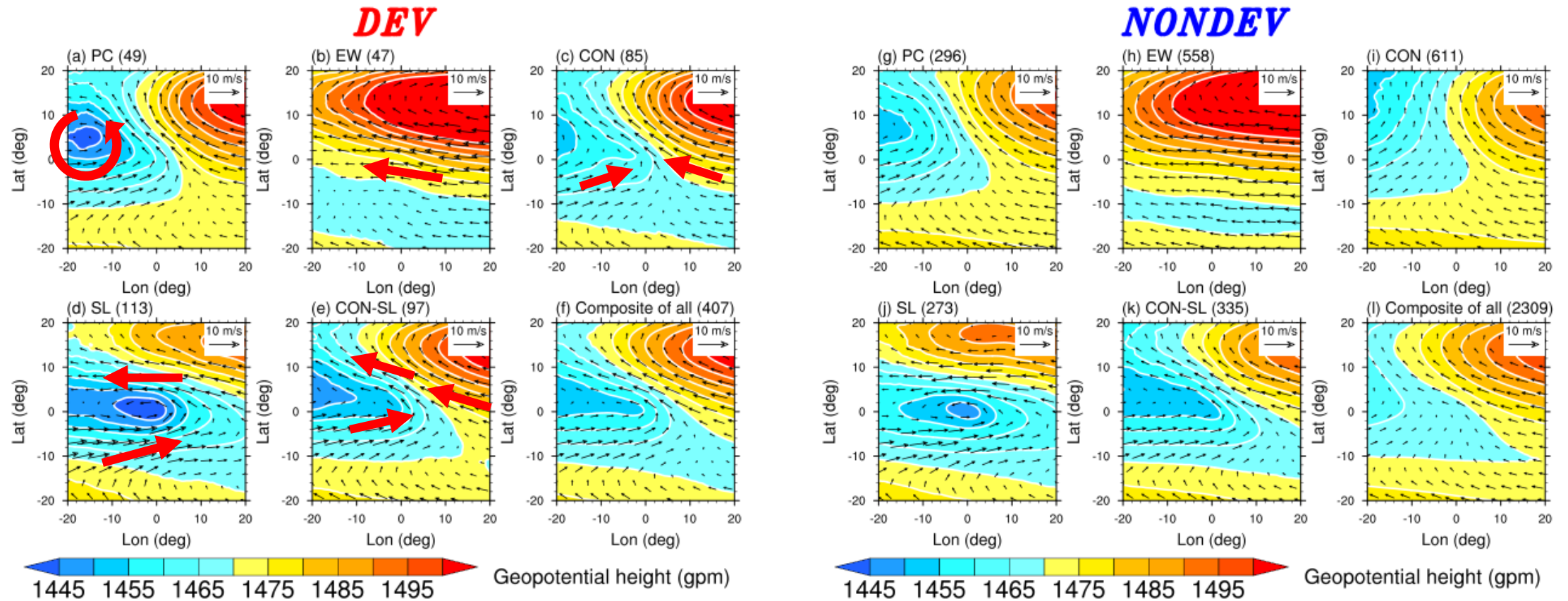
- cyclonic circulation radius  $\geq 400$  km
  - maximum relative vorticity  $\geq 3 \times 10^{-6} s^{-1}$
  - displacement velocity  $\leq 1000$  km/day
- (e.g., Kerns et al., 2008; Peng et al., 2012; Gao et al., 2019)



Flow pattern	Number of disturbances		TC yield
	DEV	NONDEV	
PC	49 (12.0%)	296 (12.8%)	14.2%
EW	47 (11.5%)	558 (24.2%)	7.8%
CON	85 (20.9%)	611 (26.5%)	12.2%
SL	113 (27.8%)	273 (11.8%)	29.3%
CON-SL	97 (23.8%)	335 (14.5%)	22.5%
Unclear	16 (3.9%)	236 (10.2%)	6.3%
<b>Total</b>	<b>407</b>	<b>2309</b>	<b>15.0%</b>

least favorable pattern

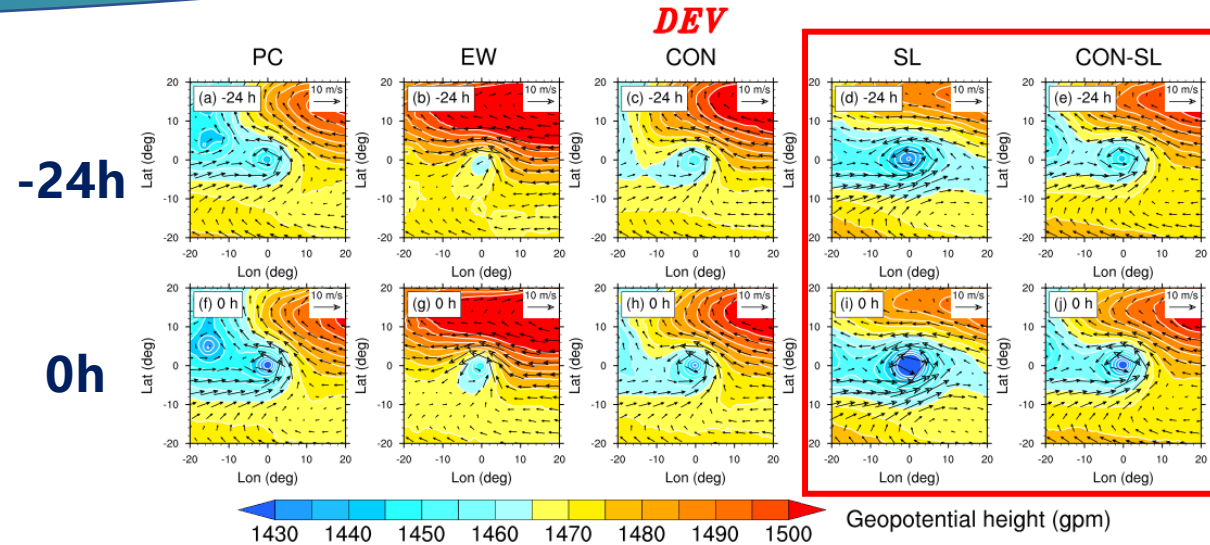
most favorable patterns



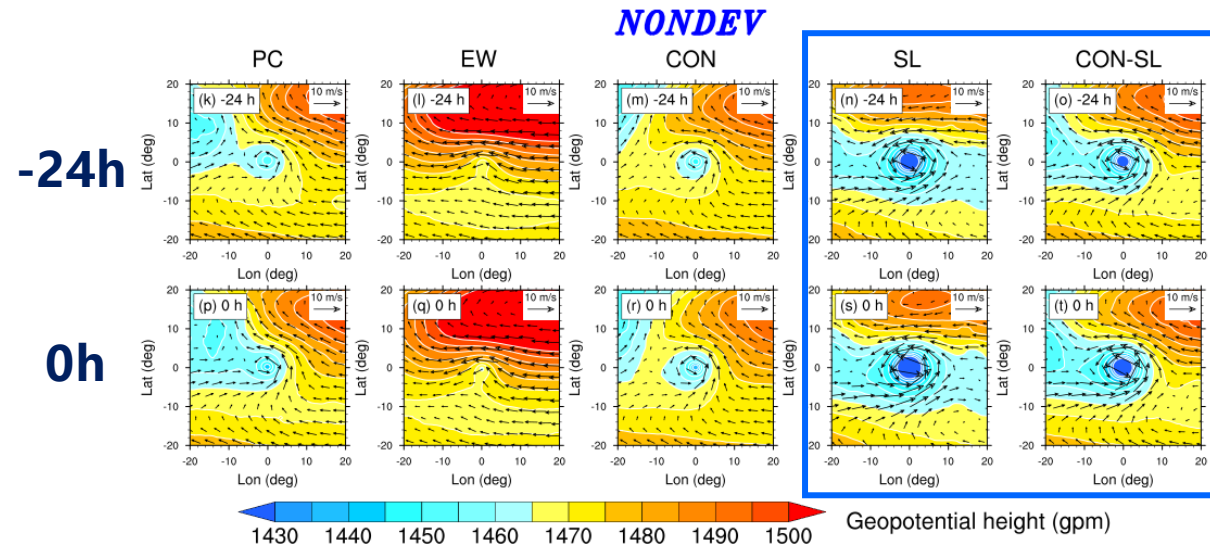
20-day low-pass filtered 850-hPa wind and height for the five patterns at 0 h

**DEV intensifies rapidly approaching genesis**

**NONDEV remains nearly unchanged approaching its maximum strength**



**Rapid development**



**Even stronger at -24 h, but develop slowly**

Unfiltered 850-hPa wind and height evolution



## A box difference index (BDI)

by Peng et al (2012)

$$BDI_D = \frac{m_{DEV} - m_{NONDEV}}{s_{DEV} + s_{NONDEV}}$$

$m$ : mean

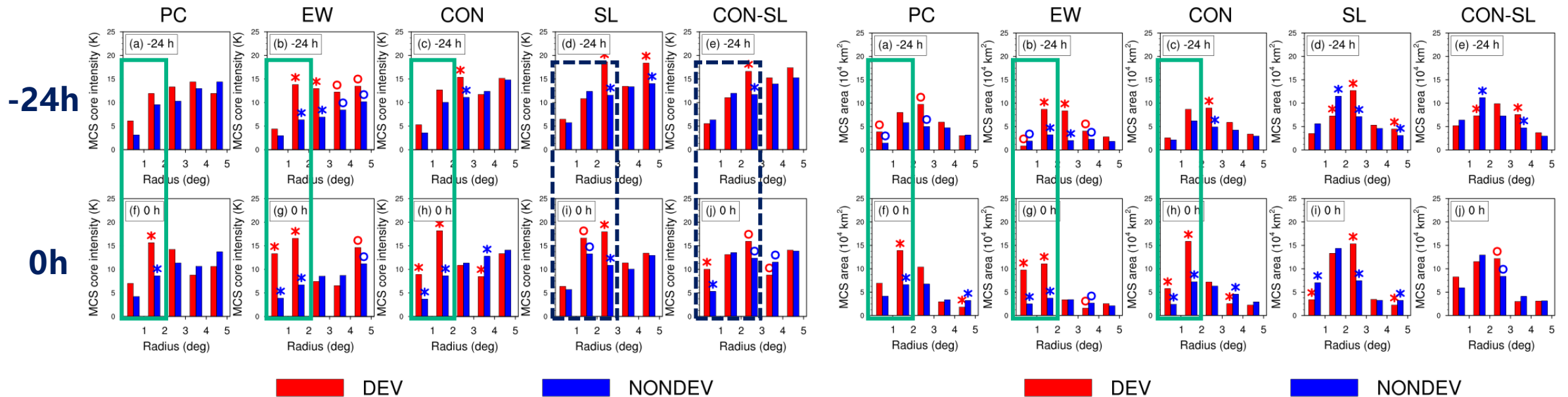
$s$ : standard deviation

Variable	$BDI_D$				
	PC	EW	CON	SL	CON-SL
<b>Dynamic variables</b>					
200-hPa divergence	<b>0.42*</b>	<b>0.65*</b>	<b>0.42*</b>	<b>0.20*</b>	<b>0.22*</b>
850-hPa convergence	0.02	<b>0.28*</b>	0.11	0.01	0.05
850-hPa vorticity	<b>0.37*</b>	<b>0.48*</b>	<b>0.30*</b>	0.05	-0.01
850-hPa vorticity growth	0.06	0.05	0.11	<b>0.17*</b>	<b>0.13*</b>
500-hPa vorticity	<b>0.33*</b>	<b>0.46*</b>	<b>0.29*</b>	0.03	-0.01
500-hPa vorticity growth	0.06	0.04	<b>0.18*</b>	<b>0.16*</b>	<b>0.14*</b>
850-200-hPa VWS	0.02	<b>-0.15*</b>	-0.04	0.03	-0.01
<b>Thermodynamic variables</b>					
500-hPa relative humidity	<b>0.26*</b>	<b>0.45*</b>	<b>0.24*</b>	<b>0.15*</b>	<b>0.16*</b>
Surface heat flux (SHF)	<b>0.26*</b>	0.06	<b>0.17*</b>	-0.04	0.01
SHF growth	<b>0.20*</b>	<b>0.17*</b>	<b>0.23*</b>	<b>0.15*</b>	<b>0.12*</b>

The values are in bold with \* when the difference between **developing and nondeveloping** disturbances is significant at a 95% confidence level.

## Roles of mesoscale convective systems (MCSs)

Criteria for MCSs within 500 km from the disturbance center,  $BT < 233 \text{ K}$ ,  $\text{area} \geq 5000 \text{ km}^2$



MCS intensity

Total area of MCSs

- PC, EW and CON: DEV has stronger MCSs close to the disturbance center
- SL and CON-SL: DEV has deeper convection approaching genesis.

## Secondary circulation

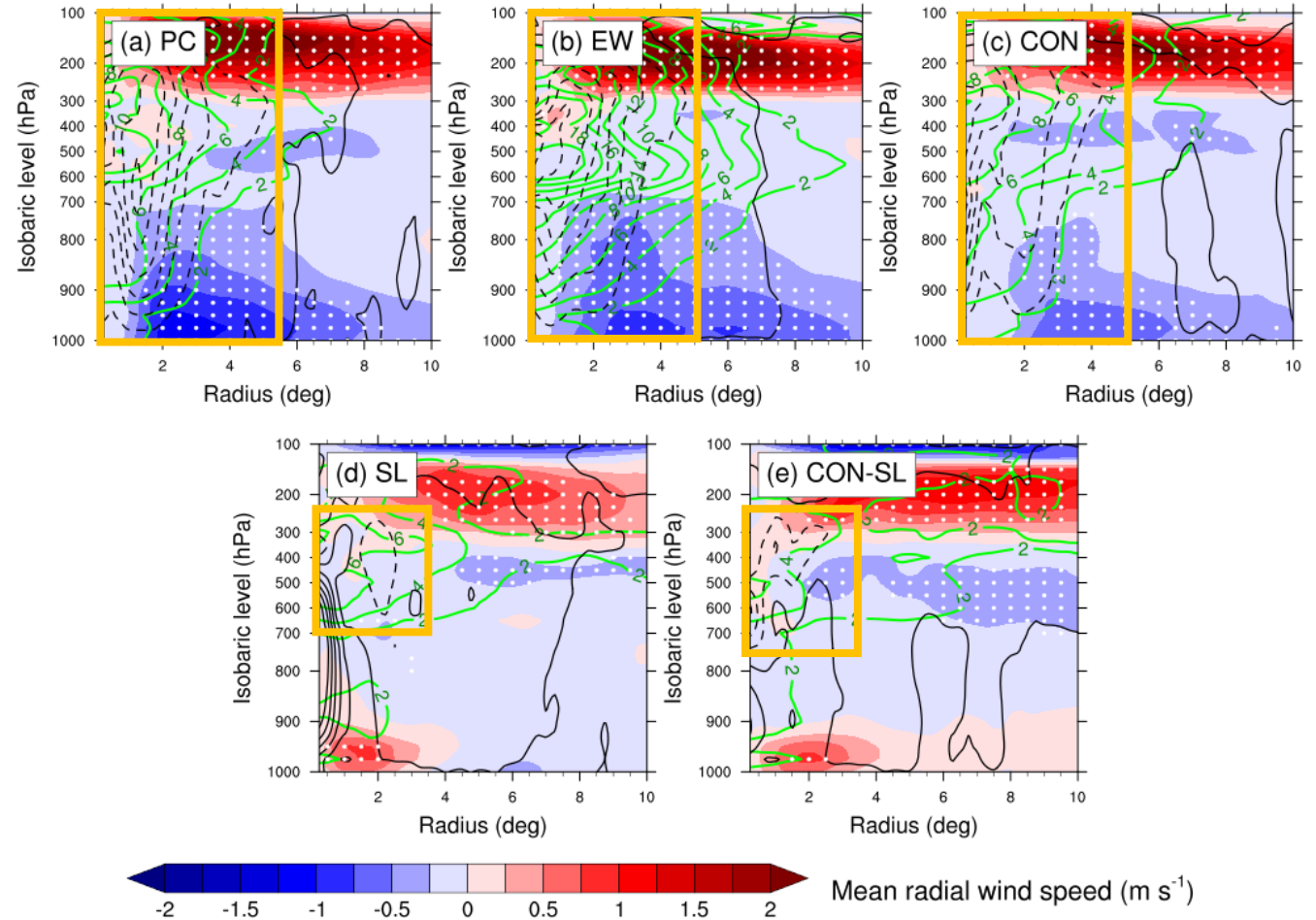
DEV has **stronger upper-level outflow**

### PC, EW and CON

- **stronger boundary-layer inflow and whole-layer updrafts with DEV**

### SL and CON-SL

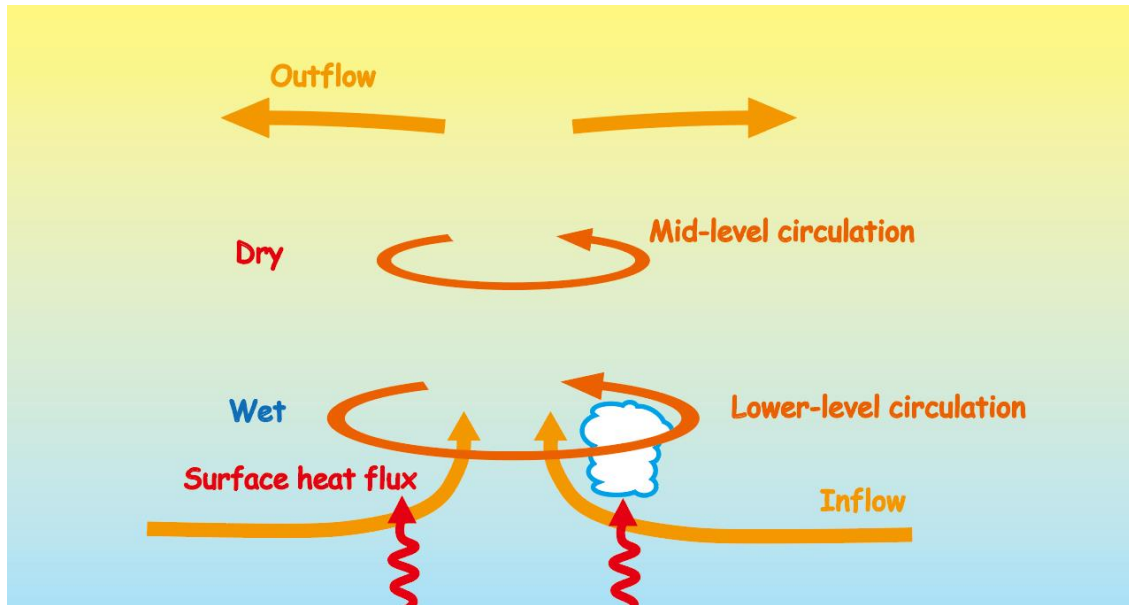
- **NONDEV** has stronger boundary-layer inflow and updrafts but **less mid to upper-level moisture** and **weaker upper-level suction**



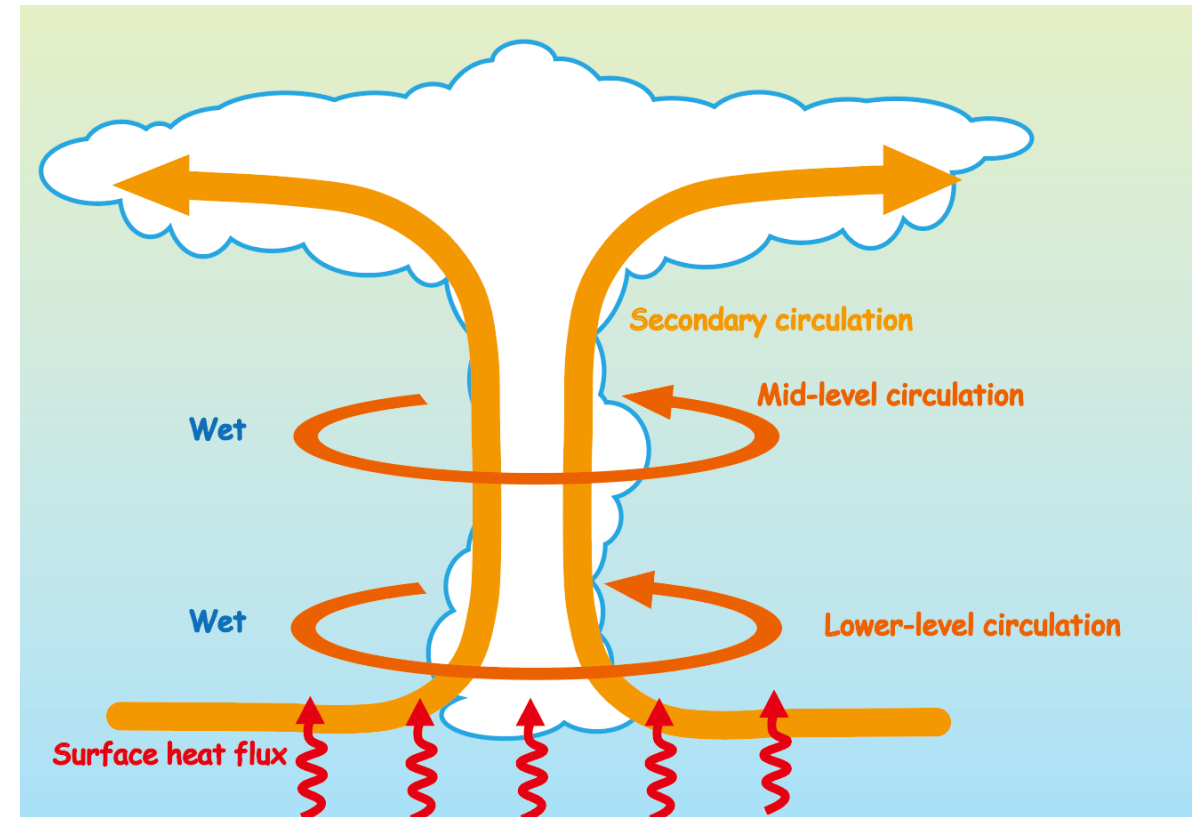
Azimuthally-averaged radial wind, vertical velocity (black) and relative humidity (green) difference between DEV and NONDEV

## Key factors for disturbance development

- Active MCS close to the center
- Mature primary circulation
- Whole-layer moisture and updrafts



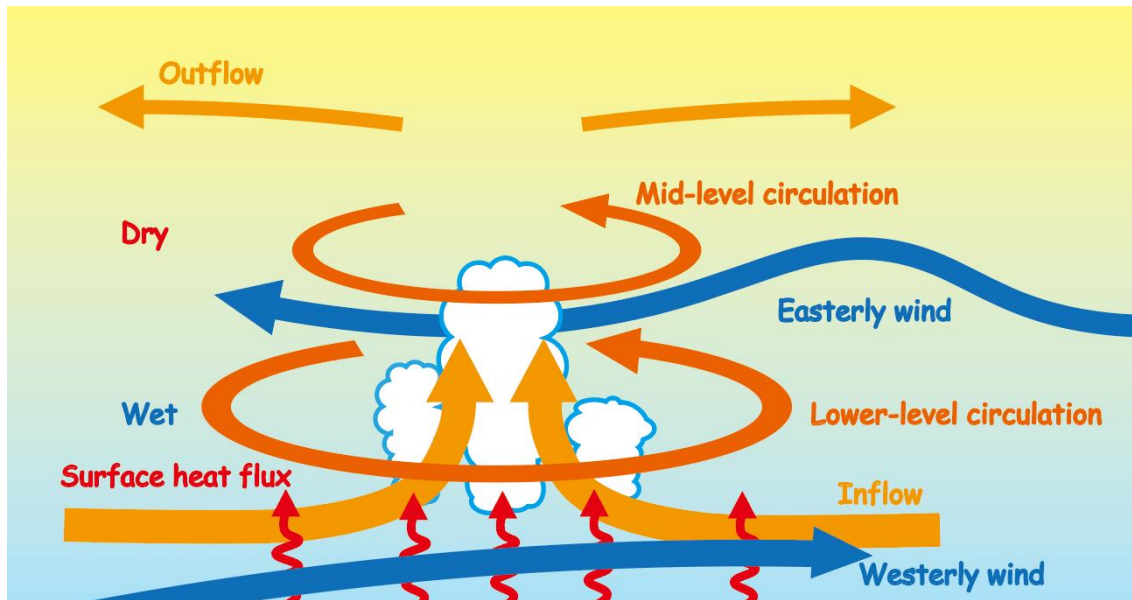
NONDEV in the PC, EW and CON pattern



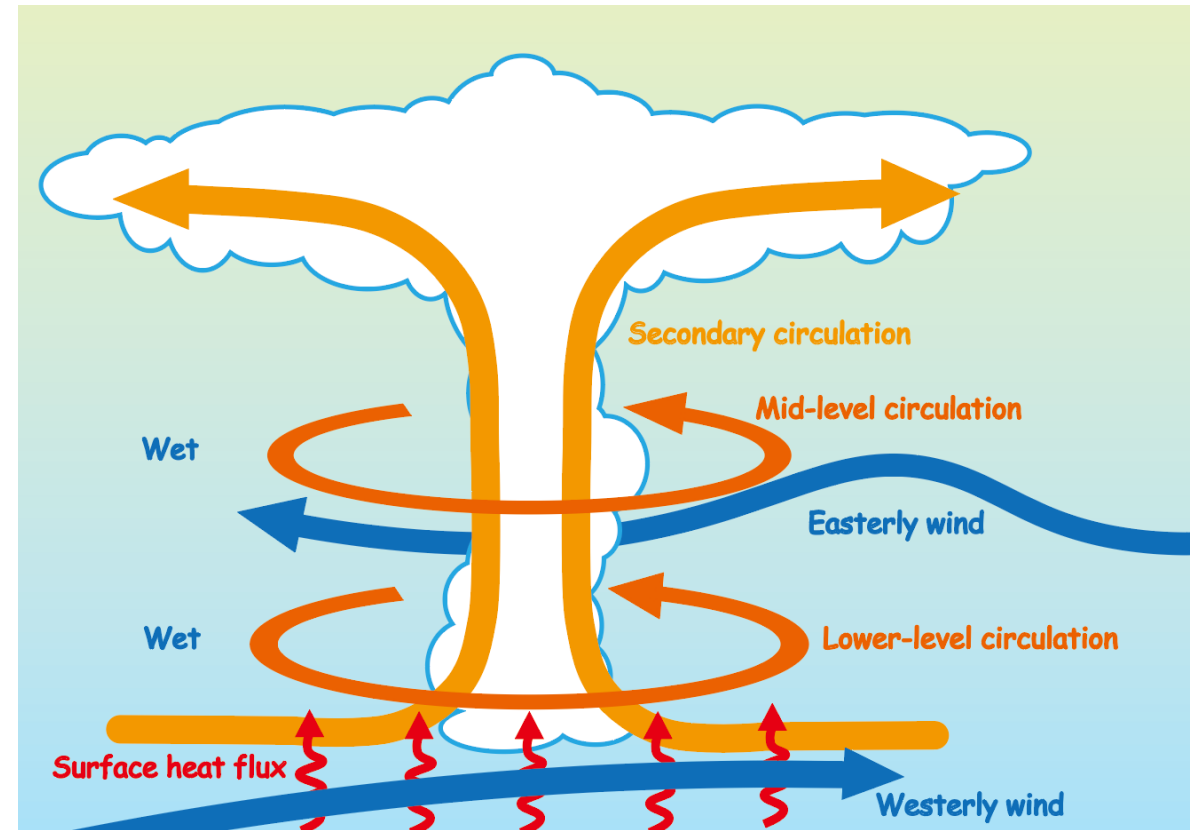
DEV in the PC, EW and CON pattern

## Key factors for disturbance development

- mid-level moisture
- sufficient upper-level suction



NONDEV in the SL and CON-SL pattern



DEV in the SL and CON-SL pattern



# Conclusions



- Over the WNP, large-scale **zonal wind shear line (SL and CON-SL)** remarkably **facilitates** TC genesis, while the **easterly wave (EW)** is the **least favorable** among canonical patterns.
- DEV benefits from **larger background westerly flows** in the south, and **intensifies more rapidly** approaching genesis time.
- In **PC, EW and CON**, **whole-layer moisture** is important, and **strong, large-area MCS close to the center** are crucial to disturbance development.
- In **SL and CON-SL**, both DEV and NONDEV usually have active MCS close to the center, with strong primary circulation, probably because of relatively **high environmental vorticity**. However, **drier mid layer in the NONDEV inhibits deep convection**, which may explain its **shallow secondary circulation** and therefore poor potential to develop further.



*Thanks for your attention!*

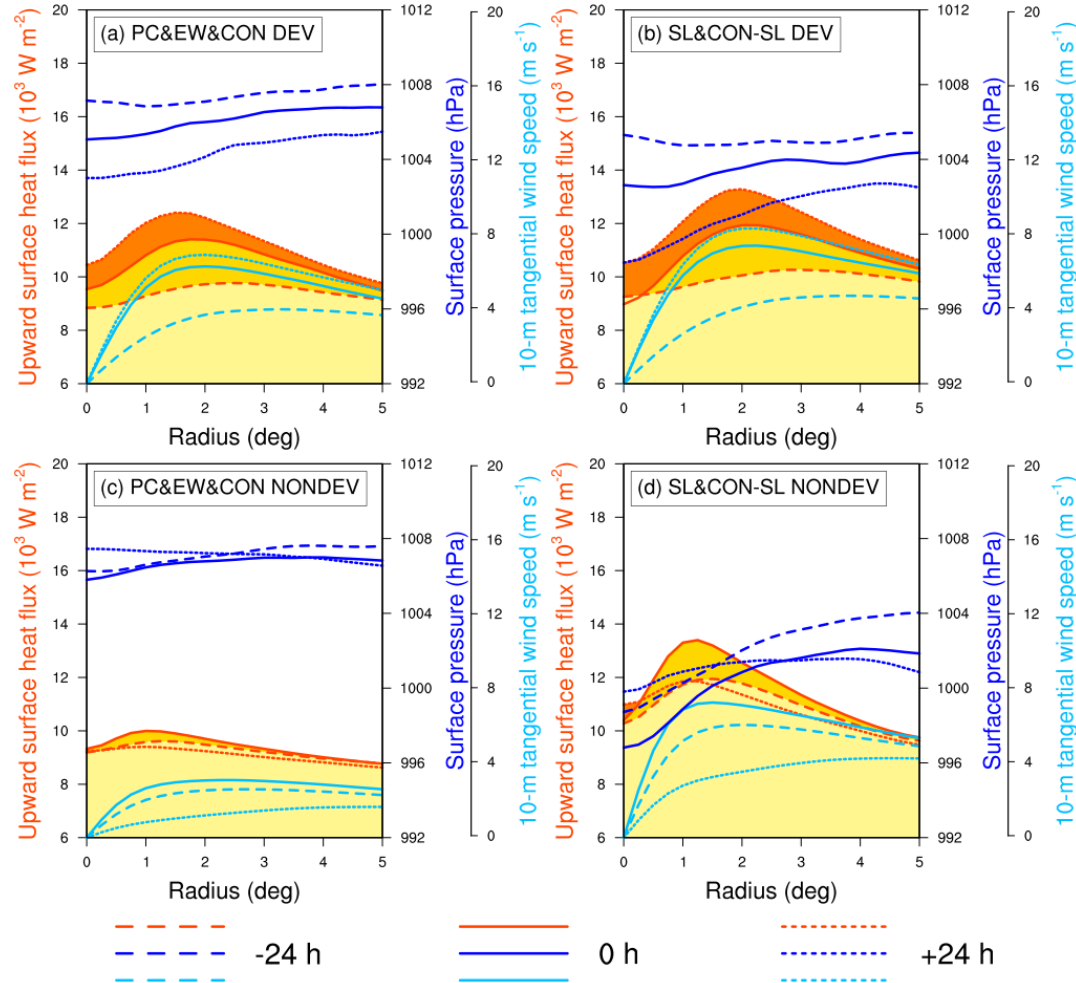
**Wang, Z. and G. H. Chen\*, 2023: Comparison between developing and nondeveloping disturbances for tropical cyclogenesis in different large-scale flow patterns over the western North Pacific. *J. Climate*, 37, 655–672, <https://doi.org/10.1175/JCLI-D-23-0401.1>.**

**PC, EW and CON  
DEV**

**PC, EW and CON  
NONDEV  
quite weak  
from -24 h to +24 h**

**SL and CON-SL  
DEV**

**SL and CON-SL  
NONDEV  
stronger at -24 h  
weakened by +24 h**



Composite azimuthally-averaged profile of 10-m tangential wind, surface pressure and SHF