

# Calibration of a diffusion model for aggrading channel after sediment overloading experiments with near-critical flow

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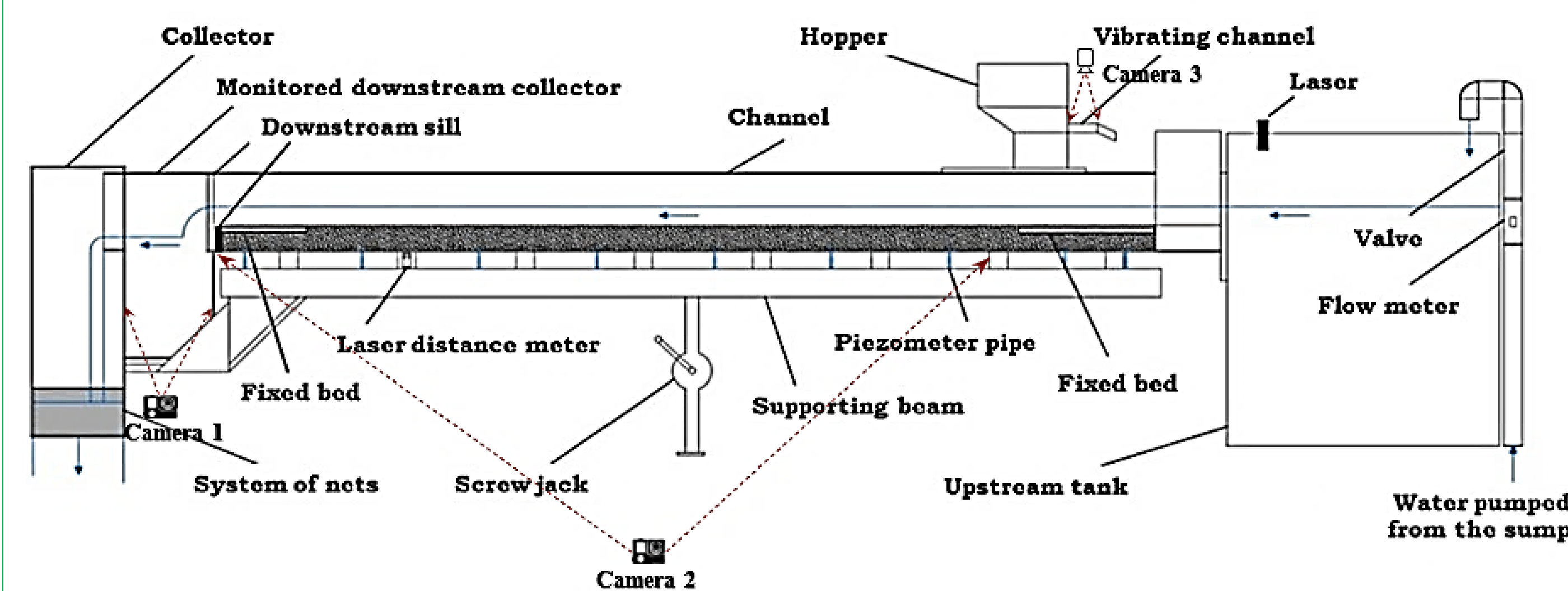


## 1. Objectives

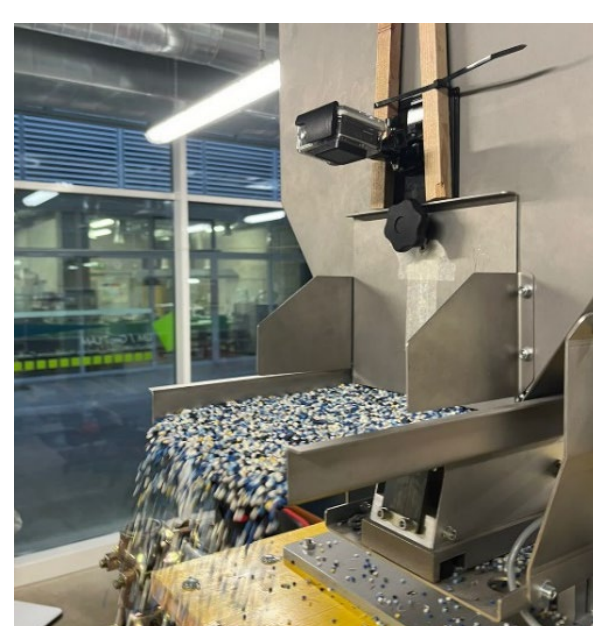
- Application of a mathematical model based on linear diffusion to eighteen aggradation experiments carried out in near-critical condition.
- Calibration of the linear diffusion model based on the experimental results.
- Investigation of the dependency of the calibrated parameter (linear diffusion coefficient,  $K$ ) on the controlling parameters e.g. overloading ratio and Froude number.

## 2. Experimental setup and measurement methods

- Rectangular channel, width = 30 cm, height = 45 cm, length = 5.2 m, erodible bed with uniform sediment.
- Sediment: PVC grains,  $d = 3.8$  mm, density = 1443 kg/m<sup>3</sup>.
- During each experiment channel is fed with a sediment discharge ( $Q_{sin}$ ) via a hopper installed at the upstream part.
- Six constant water discharges ( $Q$ ) of 3, 4, 5, 6, 7 and 8 l/s.
- Five values of overloading ratio ( $Lr = \text{sediment inflow discharge/sediment transport capacity of the channel}$ ) of 1.55, 1.85, 2.15, 2.75 and 3.25.



- Different parameters including bed and water surface elevation and sediment inflow discharge are measured during each experiment by using different cameras and applying image processing methods.



## 3. Conceptual framework, literature model

- The Saint-Venant-Exner system of equations describing the evolution of bed and water surface elevation can be simplified to a linear diffusion equation by applying some assumptions (Soni et al):

$$\frac{\partial z}{\partial t} = K \frac{\partial^2 z}{\partial x^2}$$

- Gill (1983) proposed an analytical solution for this equation:

$$Z(x, t) = S_{\infty}(L - x) + \frac{8L(S_0 - S_{\infty})}{\pi^2} \sum_{n=1,3,5,\dots}^{\infty} \frac{1}{n^2} e^{-\frac{n^2 \pi^2 K t}{4L^2}} \cos \frac{n \pi x}{2L}$$

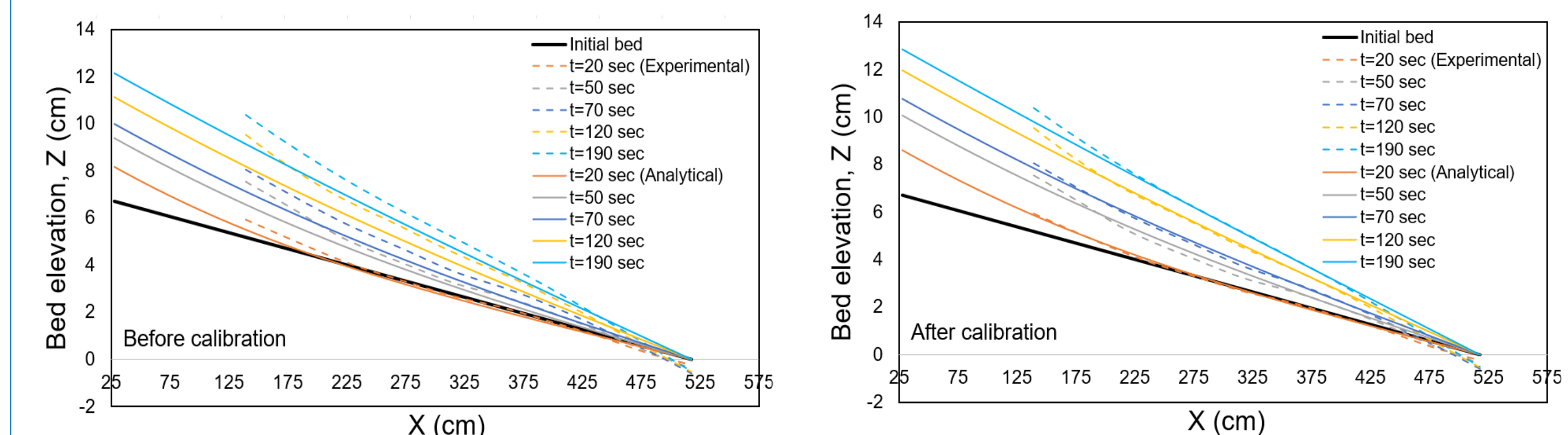
- The first-guess value for diffusion coefficient,  $K$ , is estimated by:

$$K_{Gill} = \frac{q_{\infty} - q_{s0}}{(1 - P_0)(S_{\infty} - S_0)}$$

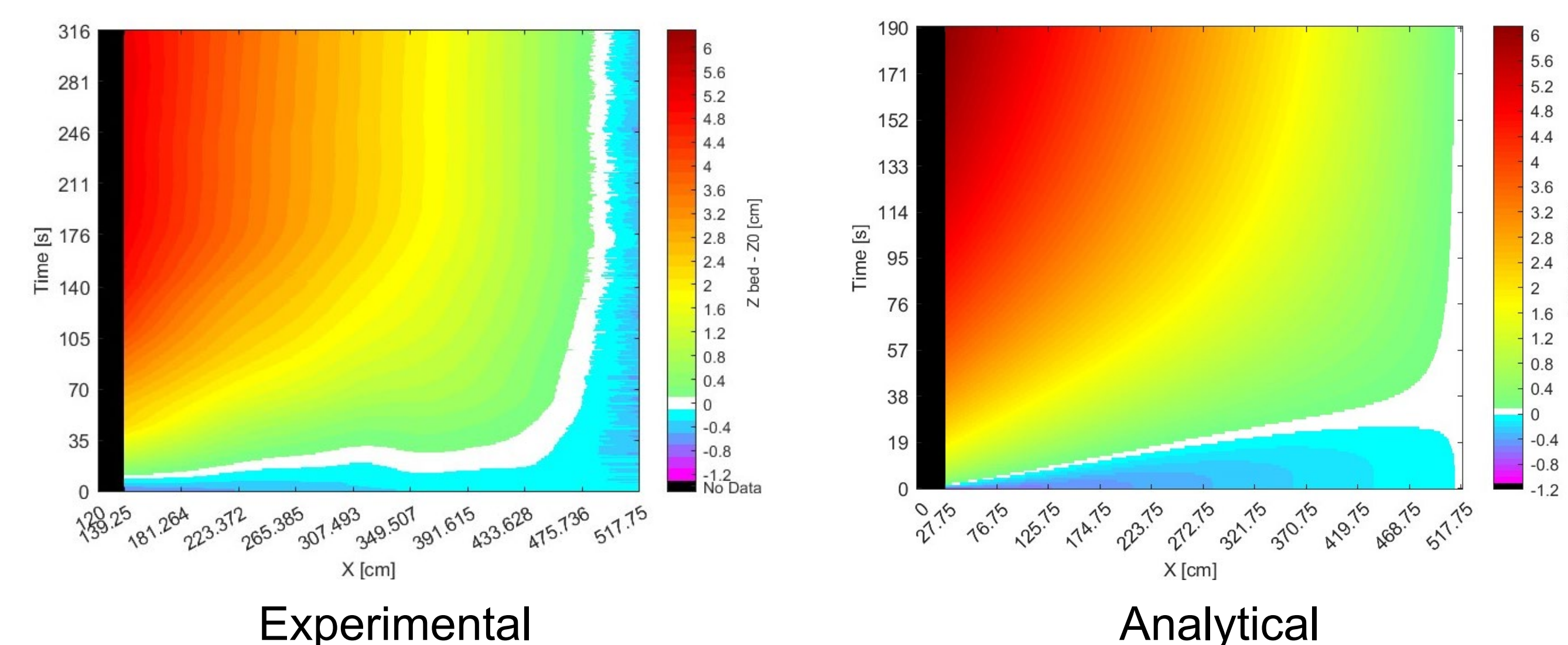
- $Z$  is bed elevation,  $S_0$  and  $S_{\infty}$  are bed slope at initial and equilibrium times,  $q_{s0}$  and  $q_{\infty}$  are sediment discharge per unit width at initial and equilibrium times,  $P_0$  is porosity.

## 4. Results and calibration of diffusion model

- Comparison between analytical bed profiles and experimental ones before and after calibration of diffusion model (as an example for one experiment), calibrated parameter:  $K$

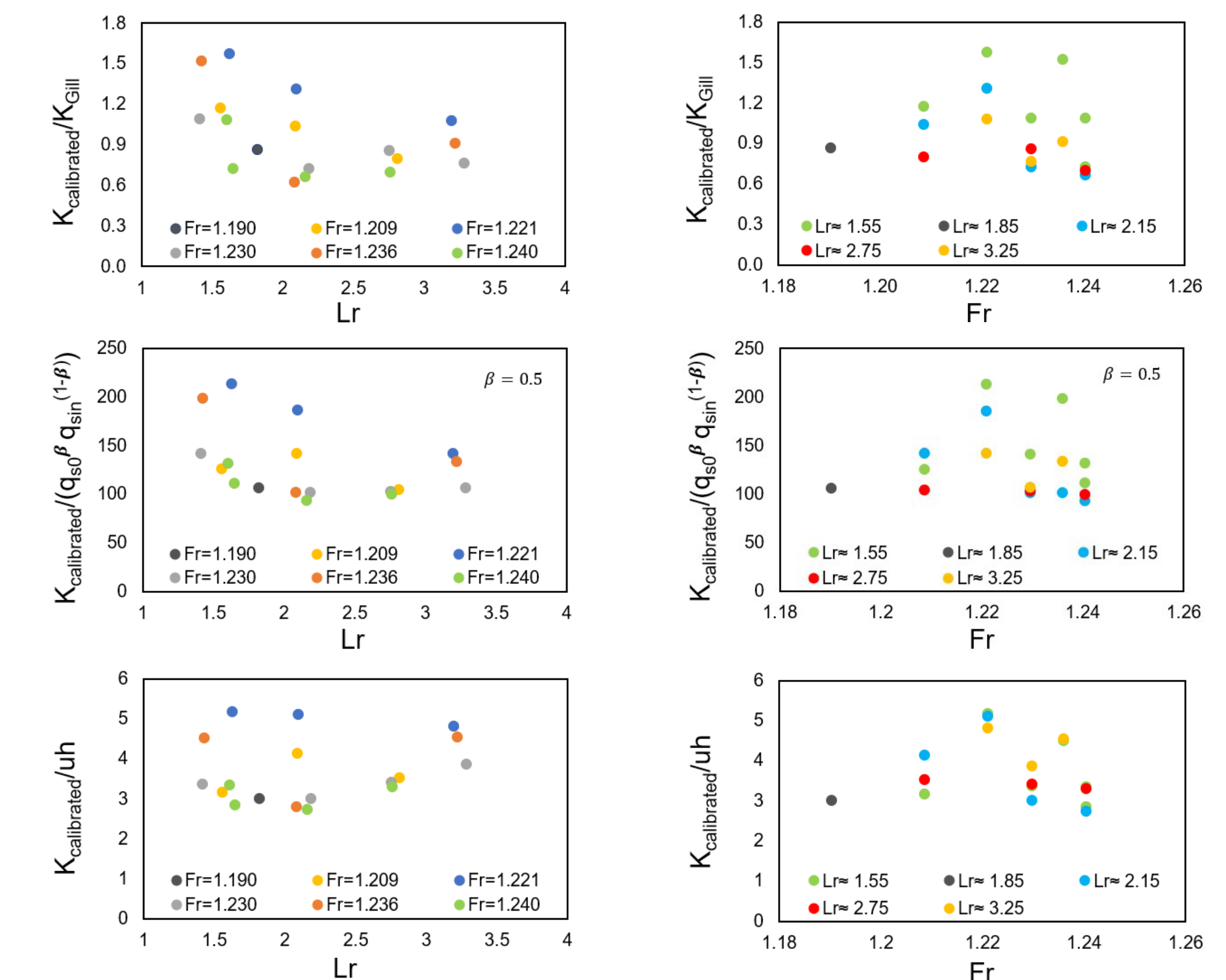


- Experimental and analytical color gradient maps of space-time evolution of bed surface after calibration:



## 5. Bulk analysis of diffusion coefficient

- Calibration was performed for each of the eighteen experiments. Moreover, the dependency of dimensionless diffusion coefficient on overloading ratio and Froude number was investigated.



## 6. Conclusions

- A constant value of the diffusion coefficient is adequate to reproduce the space-time evolution of bed surface.
- The diffusion coefficient needs to be varied from an experiment to another.
- It seems that while for  $Lr < 2$ , the dimensionless diffusion coefficient decreases with increasing overloading ratio, it remains constant for  $Lr > 2$ .
- $K_{calibrated}/K_{Gill}$  is largely different from 1 for low  $Lr$ , while it is more similar to 1 for larger  $Lr$ .
- The dimensionless diffusion coefficient appears unaffected by the Froude number, but the range of Froude number in the conducted experiments may be too small to demonstrate a clear impact.

## 7. Acknowledgements

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