EGUGeneral Assembly 2024

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1) DATAAND RESULTS

Field data and chronological analyses using U-Th dating as well as amino acid racemization (AAR) methods allowed for the identification and dating of three terraced deposits (TD) as well as their related palaeoshorelines (PS) along the Ionian coast of the Apulia region, southern Italy (De Santis et al., in press). In descending order of elevation, they are (Figs 1 and 2; Table 1):

-Terraced Deposit 1 (TD1) and Palaeoshoreline 1 (PS1, located at $+40\pm5$), dated to marine isotope stages 7-8 (MIS 7–8).

-Terraced Deposit 2 (TD2; Figs 3-6), characterised by five subunits (from TD2-sul to TD2-su5; see Tab. 1) marking transgressional episode and the first MIS 5.5 highstand. The subunit which marks an earlier phase of transgression (TD2-su1), was found to have an age of ~128 ky BP using U-Th dating. AAR analyses revealed an age of 128.4±26.2 ky BP for one of the subunits marking the later stages of transgression (TD2-su3) and aminozone E (MIS 5.5) for one of the sub-units marking the highstand (TD2-su4); *Thetystrombus latus* (Gmelin) was also found within this latter unit.

-Palaeoshoreline 2 (PS2; located at $+30 \pm 2$ m), present at the inner margin of TD2 and associated with the first, higher elevation MIS 5.5 highstand.

-Terraced Deposit 3 (TD3; Figs 7, 8), which marks the occurrence of a second, lower elevation MIS 5.5 highstand, can be subdivided into two sub-units TD3-su1 and TD3-su2 (Tab.1), lying one over another in continuous segmentation. AAR analyses revealed that Argille subappennine (ASP)-Early to Middle Pleistocene Boreholes TD3-su2 corresponds to aminozones E–D (MIS 5.5-MIS 5.3). It Calcarenite di Gravina (GRA)-Pliocene-Early Pleistocene was also found to have an age of 122.49 ± 1.25 ky BP using U-Th Mesozoic limestone

details): red numbers and red dotted lines represent the sectors of the Ostone valley and -Palaeoshoreline 3 (Ps3; located at $+19 \pm 2$ m), present at the inner their boundaries, respectively. c) Schematic geological map of the study area. margin of TD3 and associated with the second, lower elevation MIS 5.5 highstand that occurred ~122 ky BP.



Fig. 7. a) photograph and b) interpretation of stop LIZ27 (see Fig. 1c for its location), where TD3 outcrops; this unit is composed of the lower sub-unit TD3-su1 and the upper TD3-su2. TD3 overlies TD2-su2. c) Overview of the outcrops the steep flank of the Ostone valley in Sector 3 (Fig. 1a), where the section of LIZ27 reported in a) and b) was measured (black square). The black stars indicate the locations from which samples LIZ27/2 and LIZ27/2 and LIZ27/5 were of *B. rugosa* that are abundant in TD3-su2. e) Specimen of *P. jacobaeus* within

3) DISCUSSION-PALAEOCLIMATIC CONSIDERATIONS In addition, we observed the presence of the marine gastropod *Bolma rugosa* (Linneo) and the absence of the marine gastropod *Thetystrombus latus* (Gmelin) in TD1, in the lowest transgressive sub-unit of TD2, and in TD3, and, conversely, the presence of T. latus and absence of B. rugosa in the upper sub-units of TD2 that mark the last phase of transgression and the first $\frac{1}{2}$ higher highstand of MIS 5.5 (Fig. 10, left). Based on the present-day distribution of these two species (Fig. 10, right) and the reconstructed RSL trend, we have tentatively identified, in a qualitative way, three climatic phases during MIS 5.5 (Fig. 11): 1) a first, warmtemperate/subtropical in the early MIS 5.5, characterised by the presence of B. rugosa; 2) a second, fully tropical, during the mid part of MIS 5.5, with a sea level at a mean value of ca. +7.5 ± 1.5 m and characterised by the presence of T. latus and disappearance of B. rugosa; 2) a third, again warm-temperate/subtropical, during the late MIS 5.5, characterised by the disappearance of *T. latus* and the reappearance of *B. rugosa*.









Two highstands during the Last Interglacial: insights from palaeoshorelines and marine terraced deposits of ionian coast of Apulia (southern Italy)



present-day sea level (see the upper-right box in Fig. 1b). The black dotted line represents the trace of the section presented in 2b. b) Conceptual geological section showing the terraced deposits (TD1, TD2, and TD3) and the elevation above present-day sea level of their associated palaeoshorelines (PS1, PS2, and PS3). The position of boreholes from which stratigraphic samples were recovered (DEP-S1, DEP-S7, DEP-S5, DEP-S6, BON-S1, and



	Teraced deposit	Associated palaeoshoreline with its elevations in metres above the present sea level (a.p.s.l.)	Facies of terraced deposit	Depositional environment	Depositional context	Samples analysed with Aminozone (capital letter) or numerical age (ky BP)	Dated material and dating method	MIS
	TD1	PS1 (40±5)	Not recognised	high to moderate energy submerged beach	Not defined	LIZ19/1: 264.28 ±5.71	<i>Cladocora caespitosa</i> (U/Th)	7-8
						LIZ18/1:194.81±1.16	<i>Alocyathus</i> sp. (U/Th)	
	TD2	PS2 (30±2)	TD2-su5	aeolian coastal dune	highstand			early 5.5
			TD2-su4	high energy beach, from backshore to shoreface.	transgession in the lower part highstand in the upper part	LIZ17a: aminozone E	<i>Glycymeris</i> <i>pilosa.</i> (Amino acid racemization)	
						LIZ17b: aminozone E	<i>Glycymeris</i> <i>pilosa</i> (Amino acid racemization)	
G (seaward)			TD2-su3	lagoon	transgression	LIZ4a: 128.4±26.2	<i>Cyprideis</i> <i>torosa</i> (Amino acid racemization)	
x			TD2-su2	low energy environment; offshore	highstand			
-			TD2-su1	high to moderate energy submerged beach passing laterally to more protected environments and upwards to nearshore/offshore transition	transgression	T.CAS-U1/1: 127.851±1.47	<i>Cladocora caespitosa</i> (U/Th)	
						T.CAS-U1/a: aminozone E-D	<i>Glycymeris</i> <i>pilosa</i> (Amino acid racemization)	
	TD3					T.CAS-U1/b: aminozone E	<i>Bolma rugosa</i> (Amino acid racemization)	
d						T.CAS-U1/c: aminozone E	Arca noae (Amino acid racemization)	
		PS3 (19±2)	TD3-su2	infralittoral beach environment under the influence of bottom currents	highstand	LIZ 27/2: aminozone E-D	<i>Glycymeris</i> <i>pilosa</i> (Amino acid racemization)	late 5.
						LIZ 27/5: 122.49 ±1.25	<i>Alocyathus</i> sp. (U/Th)	
			TD3-su1	lower energy protected environment	highstand		· · /	
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Both scenarios give an uplift rate of 0.18 ± 0.03 mm/yr. The mean elevation of PS3 at the time of its formation (i.e., the sea level at ~122 ky BP, corresponding to the second highstand of MIS 5.5) was calculated to be -2.96 and -3.32 m for Scenarios 1 and 2, respectively (Fig. 9 a, b). After taking all uncertainties into account, the range of elevation of PS3 was found to be -2.96 ± 5.42 m for Scenario 1 and -3.32 ± 5.47 m for Scenario 2 (Fig. 9a, b). We also calculated the position of the dated layer of the lowest transgressive sub-unit of TD2 (TD2-su1; ~128 ky BP) using the uplift rates obtained from the previous analysis and consequently, assuming a palaeo-water depth of 5-10 m, the sea-level at the same age, which was found to have a value of -11.79 ± 6.29 m (fig. 9a, b). Although the mean RSL trend presented in this study were obtained from local data, they are consistent with other global and local MIS 5.5 sea-level reconstructions, which suggest the existence of two highstands, with the earlier highstand having a higher peak sea level than the second highstand (Kopp er al., 2013; Düsterhus et al., 2016; Rohling et al., 2019; Bini et al., 2020; De Santis et al., 2023).

We emphasise that the local RSL trend presented in this work (Fig. 9) is based on the interpretation of palaeoshorelines and related deposits found in the study area. Thus, this should be considered to be a broad RSL trend: we cannot rule out the possibility of minor sea-level oscillations that are not resolvable by the depositional and geomorphological record.



Fig. 11. Tentative reconstruction of the three climate phases during MIS 5.5 inferred from our results. Red vertical band: tropical climatic phase characterised by the presence of *Thetystrombus latus* (Gmelin) and absence *Bolma rugosa* (Linneo) in the study area. Orange vertical bands: termperate to warm temperate/subtropical climatic phases characterised by the presence of *B*. *rugosa* and absence of *T*. *latus* in the study area. The bands representing climatic phases are plotted together with our reconstructed RSL-trend and data point in Scenario 1 (see Figure 9a for reference). The alkenone-based SST reconstruction by Kandiano et al. (2014) from the South Balearic Islands Basin (blue line) is plotted for comparison.

10. Left: comprehensive schematic stratigraphy f the study area, showing the presence or absence of Bolma rugosa (Linneo) and Thetystrombus latus Right: present-day geographical stribution of B. rugosa (upper panel) and T. latus ower panel) from Global Biodiversity Information Facility data base (after removal of reports regarding o natural environment), which highlights the lifferent hydroclimatic conditions of the two species.

2) DISCUSSION-RSL TREND RECONSTRUCTION

We focused on the palaeoshorelines and deposits of the LIG (PS2 and TD2; PS3 and TD3).

PS2 is associated with the first highstand of MIS 5.5. The age ranging between 127.851 ± 1.47 and 128.4 ± 26.2 ky BP from sub-units of TD2 marking the transgression, suggests that the transgression towards this first highstand was still occurring at ~ 128 ky BP.

PS3 is associated with the second highstand of MIS 5.5 and has an age of ~ 122 ky BP, based on data from TD3. Also considering the above-mentioned constraints and since it was not possible to obtain an absolute age for the TD2 sub-units which mark the first MIS 5.5 highstand, we considered the following two scenarios (Kopp et al., 2013):

1) Scenario 1, which assumes that the first highstand of MIS 5.5 occurred at 125 ky Bp

2) Scenario 2, which assumes that the first highstand of MIS 5.5 occurred at 123 ky BP. In both scenarios, we assumed a mean sea level value of $+7.5 \pm 1.5$ m during the first highstand of MIS 5.5. We then calculated the uplift rate of PS2 for both scenarios, assuming a constant uplift rate to the present day.

These uplift rates were used to calculate (Fig. 9) the position of PS3 at 122.49 ± 1.25 (rounded at ~122 ky BP) as well as the position of the dated layer of the lowest transgressive sub-unit of TD2 found at an elevation of +3.80 m, dated at 127.851 \pm 1.47 ky BP (rounded at ~128 ky BP), with the latter assumed to have been deposited at a water depth of 5-10 metres.

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Fig. 9. Local highstands and RSL trends in the study area. a) Scenario 1: Elevations a.p.s.l. for PS2 (orange triangle; +7.5 m), PS3 (yellow triangle; -2.96 m), and sea level of the dated layer from Torre Castelluccia (blue triangle; -11.79 m) plotted along a resolved global mean sea level (GMSL) curve reported by Kopp et al (2013) (blue line); the 95% confidence intervals are indicated by the grey band. The dashed line represents a tentative reconstructed RSL trend for the study area based only on local geomorphological and chronological evidence. We emphasise that the local RSL trend should be considered to be a broad trend: we cannot rule out the possibility of minor sea-level oscillations that are not resolvable by the depositional and geomorphological record. The vertical bar on the yellow triangle represents the possible range of RSL values during the second highstand of MIS 5.5 (from -8.38 to +2.45 m). The vertical bar on the orange triangle represents the uncertainty of the derived sea level obtained for the dated layer from Torre Castelluccia (from -5.5 to -18.08 m). b) Scenario 2: Same colour scheme as in a) but with the yellow triangle representing a value of -3.32 m and the vertical bar on the yellow triangle representing the entire range of possible RSL values for the second highstand during MIS 5.5 (from -8.79 to 2.15 m).

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