

# Hydrothermal fluid circulation from soil CO<sub>2</sub> degassing, thermal flux, and self-potential

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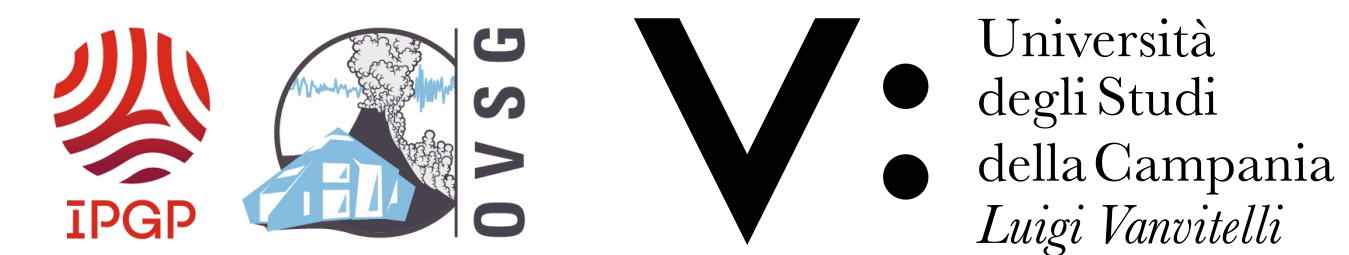
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## La Soufrière de Guadeloupe

- Andesitic dome with extensive hydrothermal system, fueled by magma reservoir 6-7 km below the summit<sup>1</sup> and tropical rainfall
- Growing unrest with peak in activity in 2018<sup>2</sup>
- Host rock alteration and sealing by acid hydrothermal fluids:
  - affects dome permeability
  - promotes pore fluid pressurisation
  - increases likelihood of flank collapses and explosive activity<sup>3</sup>.

Dome permeability reflects in ground heat and gas fluxes. Thus our **objectives** were to:

- constrain the spatial distribution of hydrothermal activity on the summit
- quantify ground heating and, for the first time, ground CO<sub>2</sub> flux.

Results provide important clues for interpreting the development of volcanic unrest and assessing hazards linked to fluid pressurisation.

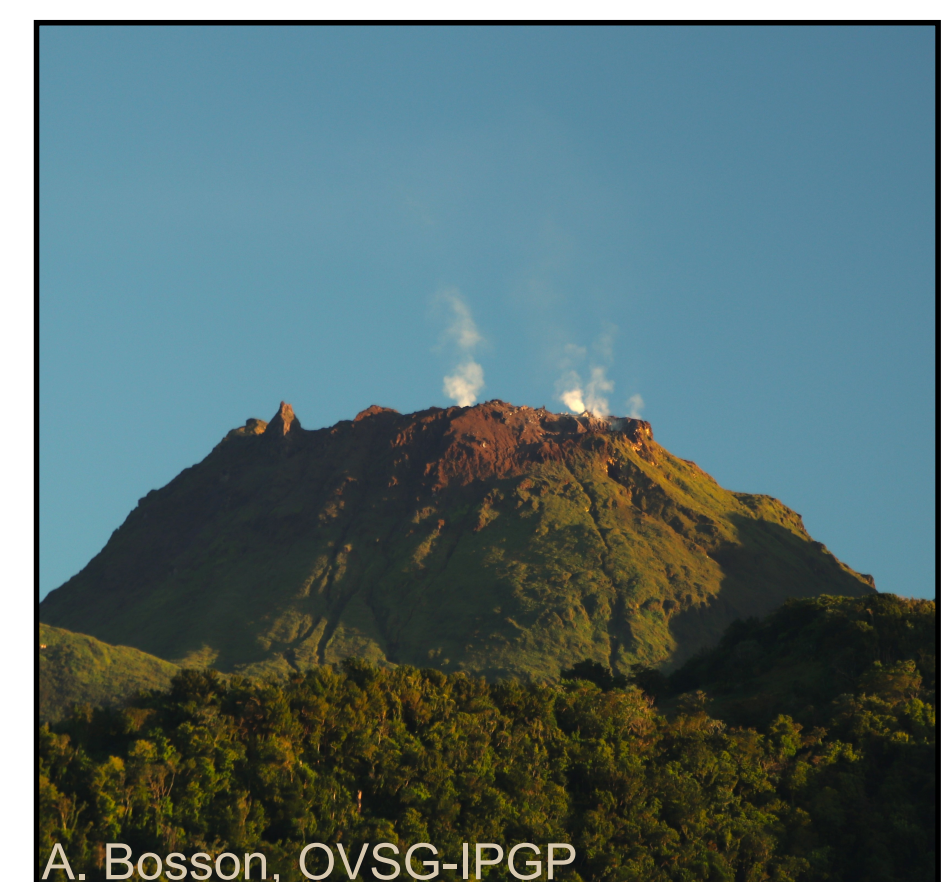
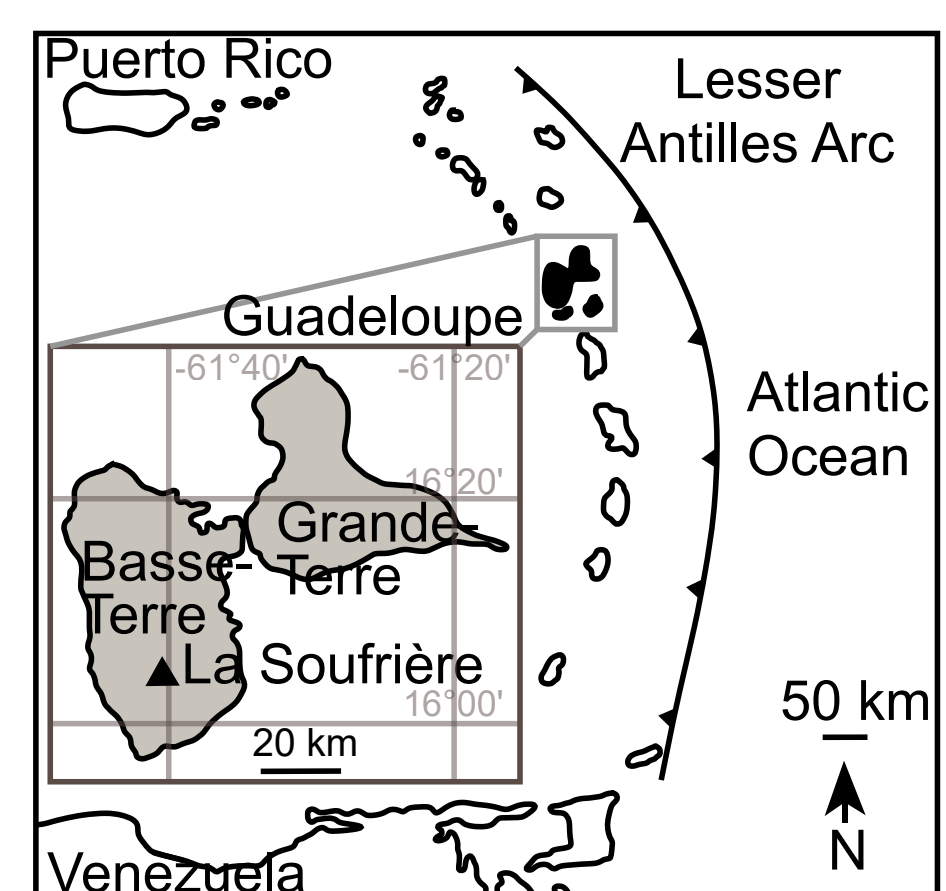


Fig. 1: Location of La Soufrière de Guadeloupe and photo of the dome

## Soil CO<sub>2</sub> emissions

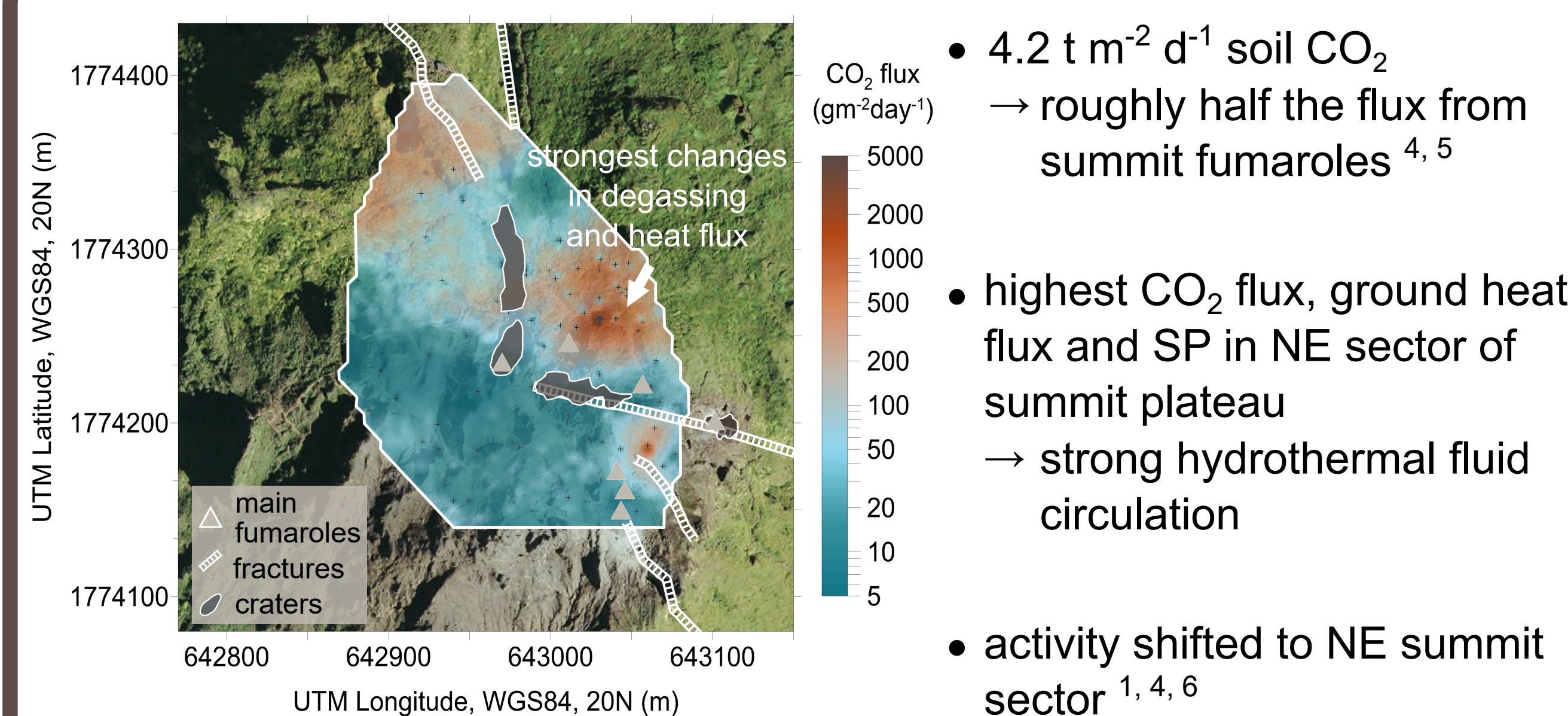


Fig. 3: Interpolated soil CO<sub>2</sub> flux on La Soufrière summit from May 2023. Background image IGN, BD ORTHO 2017

- 4.2 t m<sup>-2</sup> d<sup>-1</sup> soil CO<sub>2</sub> → roughly half the flux from summit fumaroles<sup>4,5</sup>
- highest CO<sub>2</sub> flux, ground heat flux and SP in NE sector of summit plateau → strong hydrothermal fluid circulation
- activity shifted to NE summit sector<sup>1,4,6</sup>

## Permeability as controlling factor ...

Estimated condensation depth of rising fluids from measured ground temperature profiles

Shallow condensation level in NE summit sector (few tens of cm)

Increased ground permeability may explain high ground CO<sub>2</sub> and heat flux in NE sector of summit

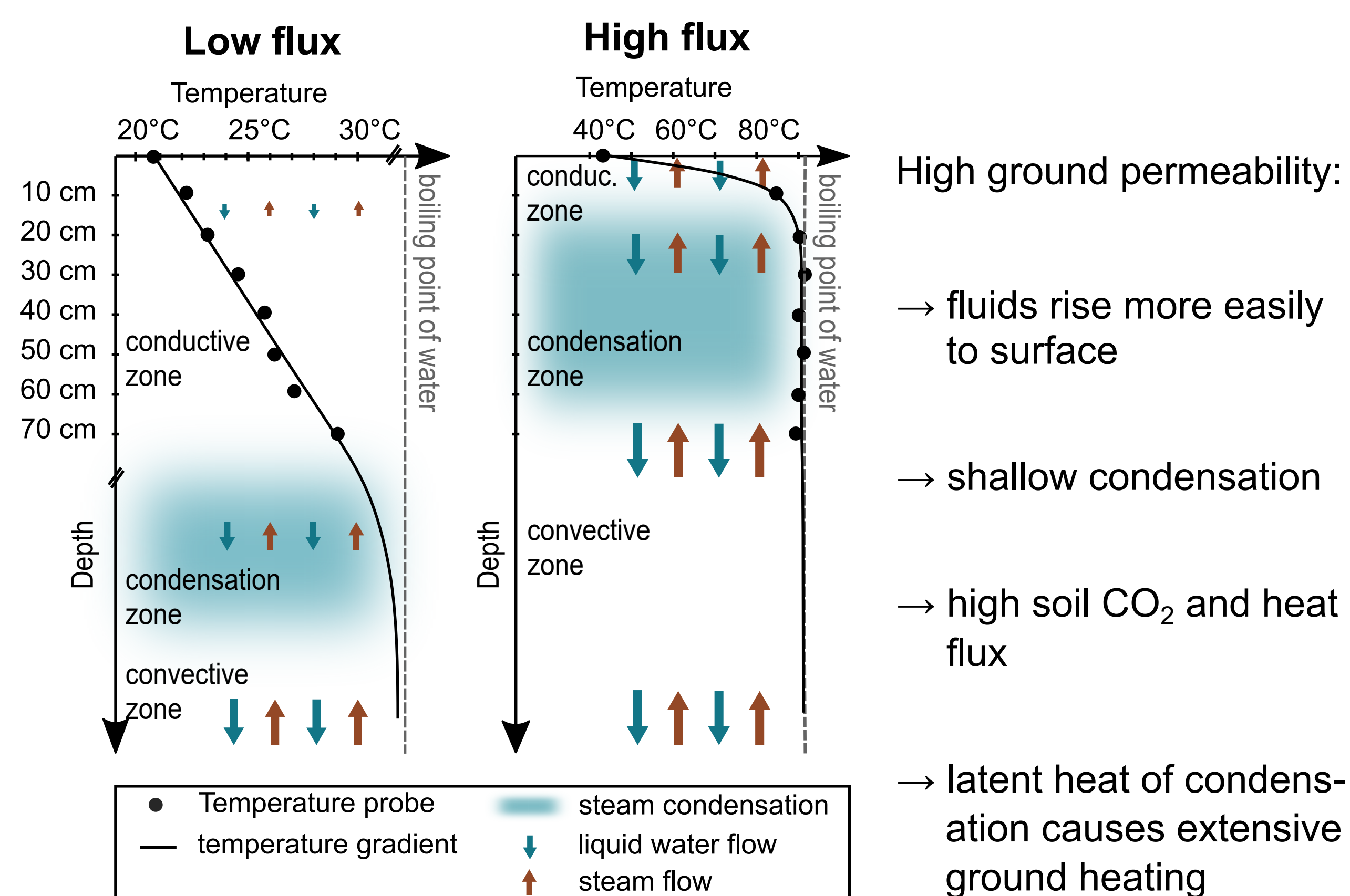


Fig. 4: Temperature profiles in diffuse degassing zones on La Soufrière summit (figure adapted from Gaudin et al., 2017)

## ... for changing patterns of unrest

- Area of ground heating has likely expanded since 2019<sup>4,6</sup>
- Increase in shallow hydrothermal circulation in NE dome sector<sup>4,6,7</sup>

Key factors:

- changing ground permeability
- opening of W-E summit fractures and sliding of SW flank<sup>2</sup>
- contribution of lower water level in hydrothermal aquifer?

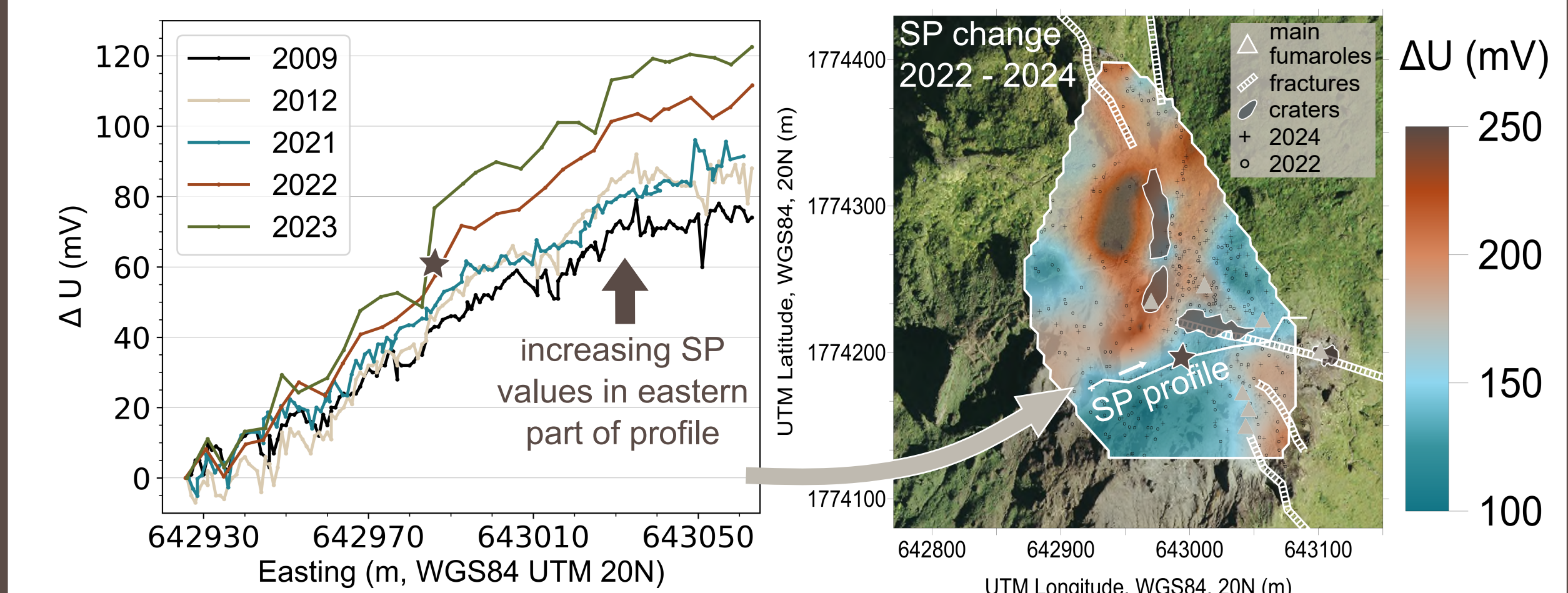


Fig. 5: Repeated self-potential (SP) measurements along W-E profile and difference in SP between January 2024 and May 2022 (interpolated). Background image IGN, BD ORTHO 2017

## Quantifying hydrothermal fluxes

Multi-parameter mapping to identify upwelling and recharge zones + estimate ground heat and CO<sub>2</sub> fluxes

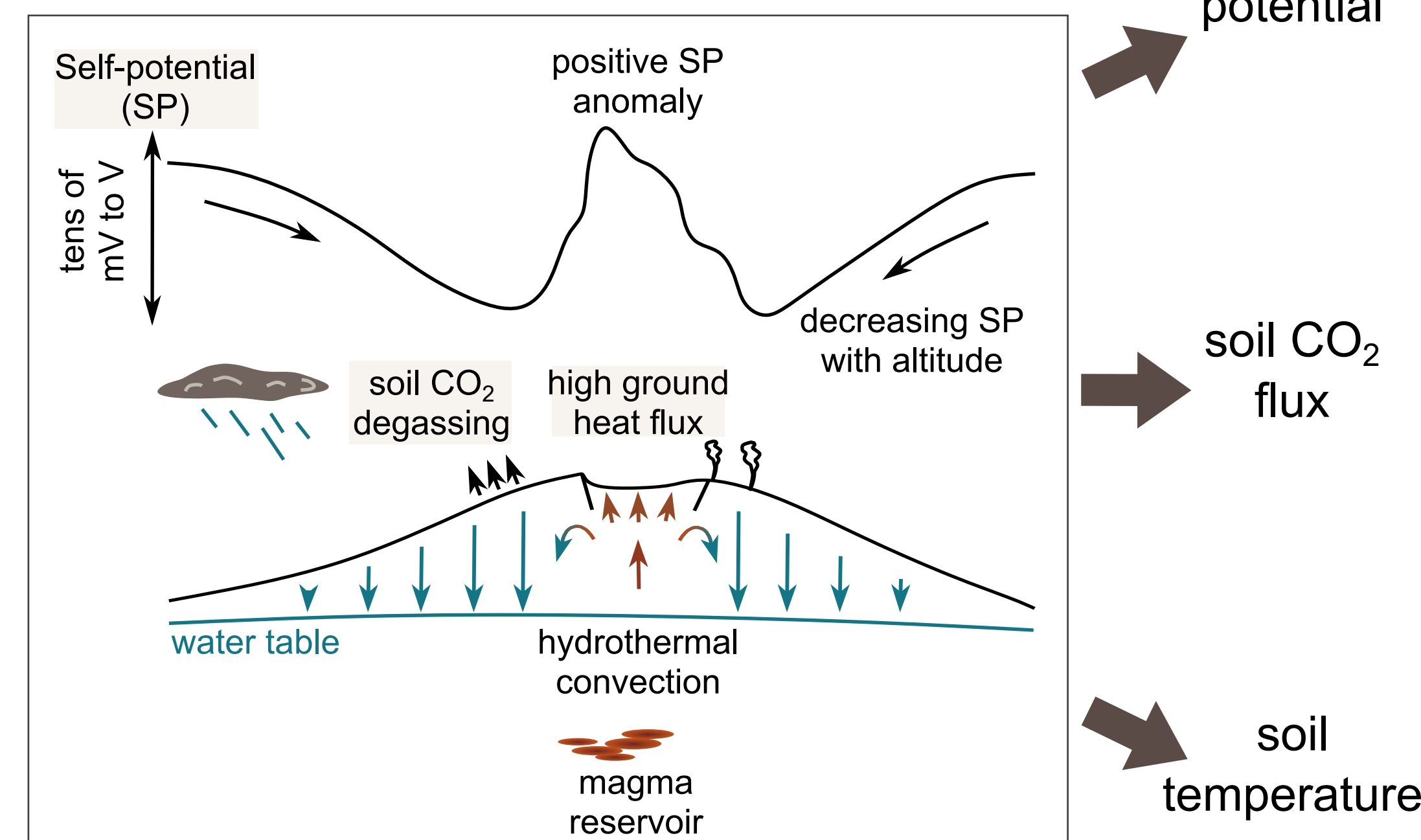


Fig. 2: Surface expressions of hydrothermal circulation and methods applied to localise and quantify subsurface fluid flow

## Outlook

How do fluid flux patterns vary and change over time?

- SP signals correlate with rainfall and seismicity
- Effect of aquifer water level?
- Periodicity/response times?
- Repeated mapping to track evolution of hydrothermal unrest

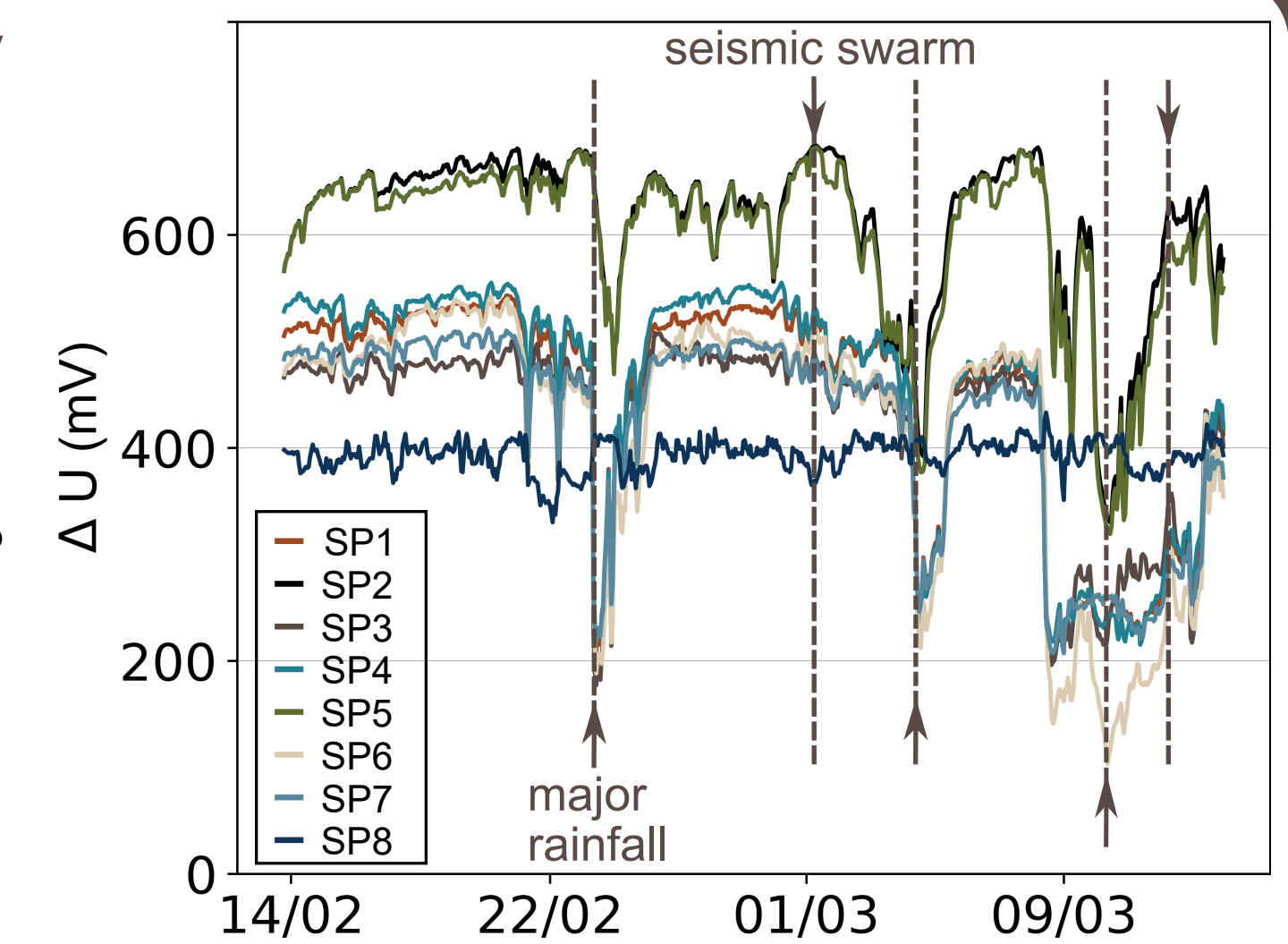


Fig. 6: SP timeseries extract of eight permanently installed electrodes on La Soufrière's summit

## References

- Allard et al., 2014. Chem Geol 384.
- Moretti et al., 2020. J Volcanol Geoth Res 393.
- Heap et al., 2021. Geophys Res. Solid Earth 126.
- Klein et al., 2024. Bull Volcanol 86.
- Moune et al., 2022. Front. Earth Sci 10.
- Jessop et al., 2021. Bull Volcanol 83.
- Brothelande et al. 2014. J Volcanol Geoth Res 288.

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