A holistic approach to low-field NMR and MICP data integration for the accurate determination of the absolute pore size distribution in siliciclastic rocks

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Introduction

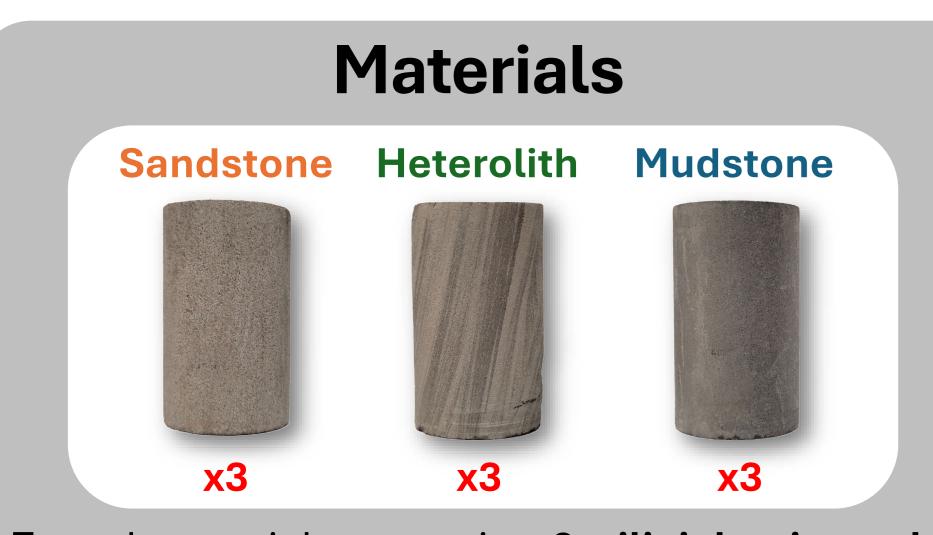
Determining the exact pore size distribution (PSD) in a rock is crucial for predicting its reservoir properties. In the standard low-field NMR approach, the PSD estimation relies only on a linear fit of the surface relaxivity (ρ_2) , assuming free diffusion and neglecting internal gradients (G). However, in nano- and micropore spaces, diffusion is restricted. Additionally, the magnitude of **G** is inversely proportional to the pore size and can cause substantial distortions in the registered, short **T**₂ relaxation times.

Challenges

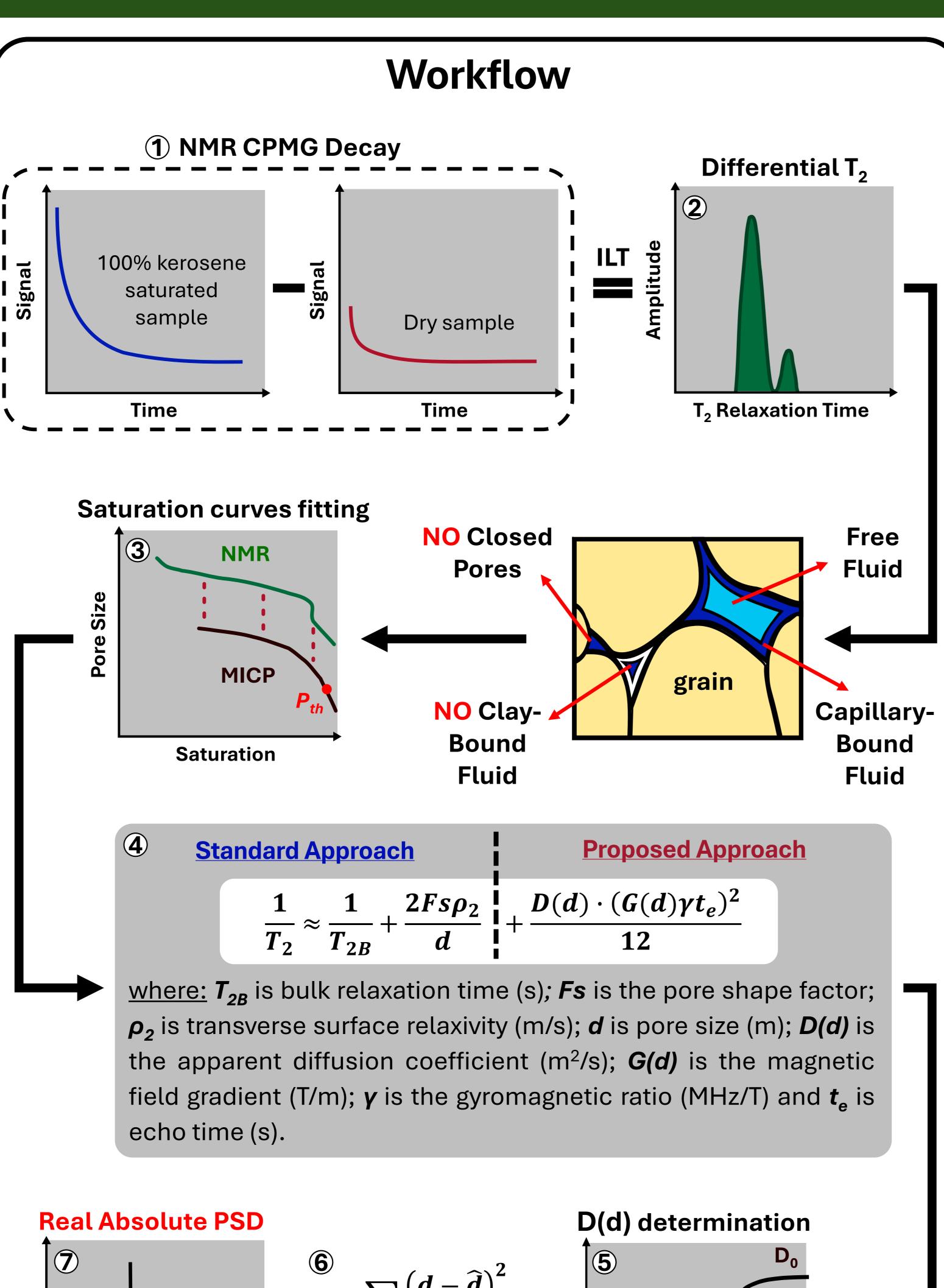
- In fine-grained siliciclastic rocks magnetic susceptibility (X) significantly increases the magnitude of internal gradients.
- **Diffusion (D)** in nano- and micropores for attainable experimental time is in the motional averaging regime (MAV) thus, the observed diffusion coefficient decreases significantly.

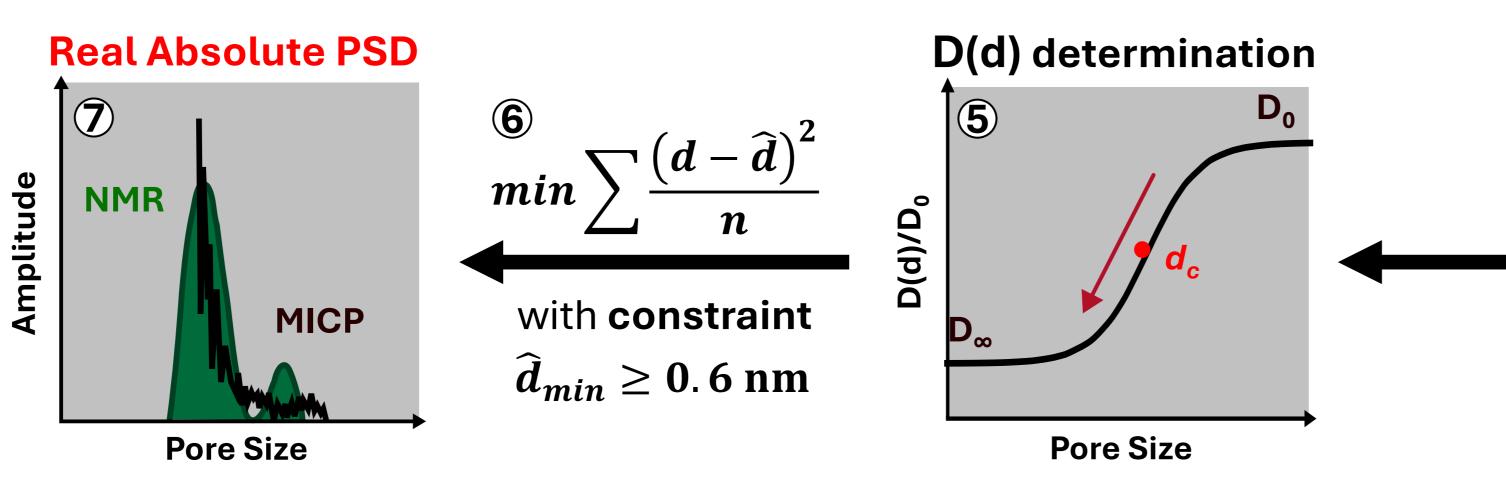
Objective

To introduce an integrated method based on differential LF-NMR and porosimetry (MICP), allowing a realistic assessment of PSD in full range, including nano- and micropore space, by considering the influence of ρ_2 , G and D.



Tested materials comprise 9 siliciclastic rock**core** samples of the Miocene age. Samples were from the wells located in South-East taken Poland, Carpathian Foredeep area.





Methods

- Low-field nuclear magnetic resonance (LF-NMR) measurements were performed using a 2 MHz Rock Core Analyser (Magritek) with field induction of 0.05 T. CPMG-T₂ sequence was applied using $t_e = 60 \ \mu s$, $RT = 5 \ 000 \ ms$, *NoE* = 50 000 and *NoS* = 512.
- **Reference measurements** were conducted using **Mercury porosimetry** (MICP) and Nitrogen adsorption (BET) methods. In addition to the PSD, low-field AC magnetic susceptibility (X) measurements were performed.

Procedure

- 1 Measurements of 100% kerosene-saturated (vacuum) and dry (110°C, vacuum) rock core samples.
- **2** Determination of differential distributions subtracting dry and 100% saturated sample signals before **Inverse Laplace Transform (ILT)**.
- **3** Selection of NMR-MICP conversion points based on the comparison of the saturation curves in the percolation thresholds framework (P_{th}).
- **4** Calculation of PSD via our **Proposed Approach**, incorporating the diffusion component and accounting for the influence of induced gradients, G(d) and apparent diffusion coefficient, D(d):

5)
$$\frac{D(d)}{D_0} = \frac{(D_\infty - 1)d_c}{d_c + d} + 1$$

where: D_0 is the self-diffusion coefficient of the bulk kerosene (0.88.10⁻⁹) m²/s), D_{∞} is the effective diffusion coefficient (m²/s) and d_{c} is the characteristic pore size (m) corresponding to the centre of the sigmoid.

- **6** The Mean Squared Error (MSE) minimization between estimated (NMR) and measured (MICP) pore size to solve for ρ_2 and D_{∞} .
- **Validation of the NMR T₂-PSD** transformation results.

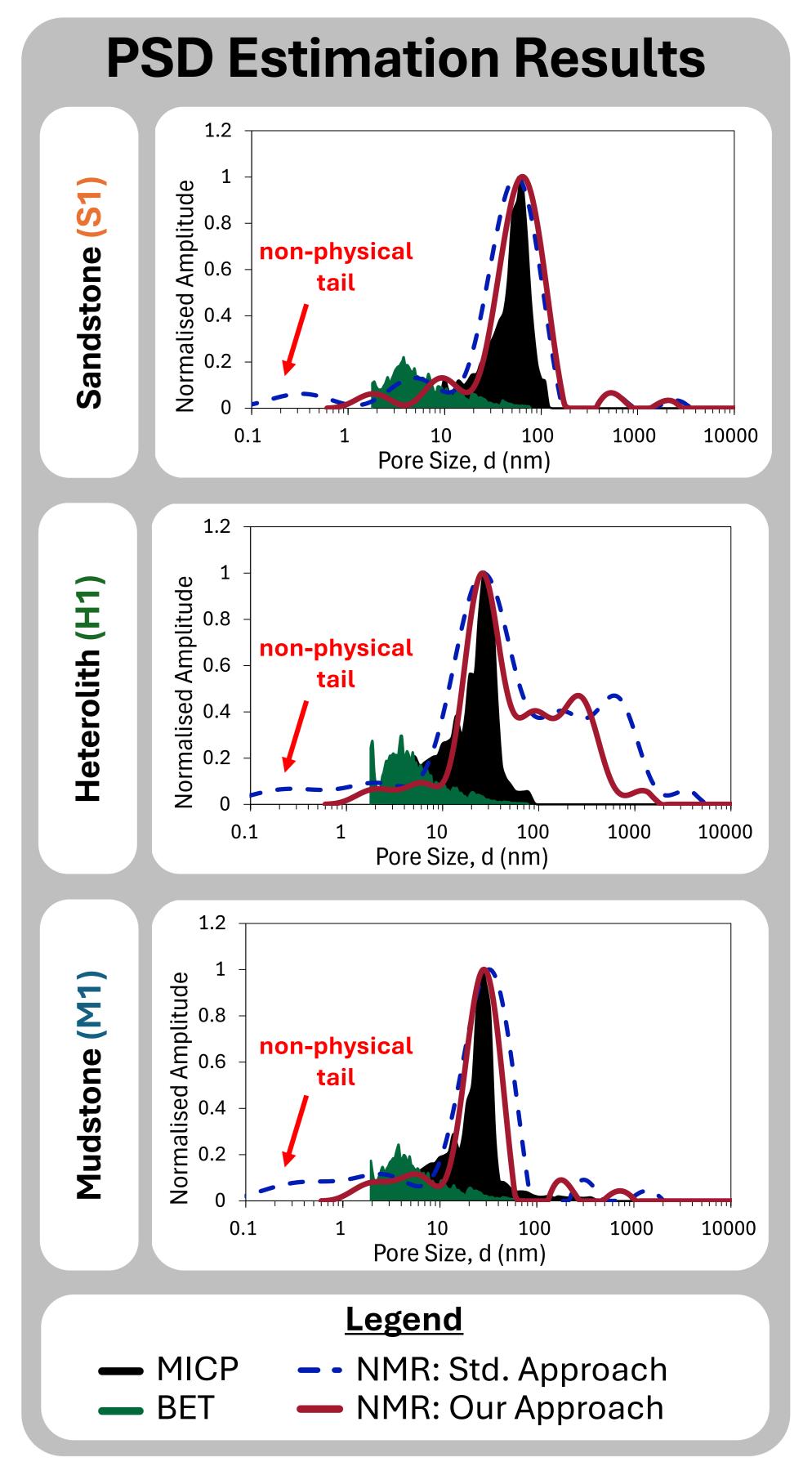
$X vs D_{\infty}$ Results **D**_∞ = (6.45·10⁻¹⁹)·X⁻² 12 **R**² = 0.99 $\begin{array}{c} \text{ed D}_{\circ} \\ 10^{11} \\ 0 \end{array}$ **H1** Estin (m **M1 S1** Measured $\chi \cdot 10^5$

Table of Param

Approach	Measured	Standard	Proposed	
Sample	X	ρ₂ (μm/s)	ρ₂ (μm/s)	D ∞ (m²/s)
Sandstone (<mark>S1</mark>)	5.6·10 ⁻⁴	1.03	0.36	2.0·10 ⁻¹²
Heterolith (H1)	1.2·10 ⁻⁴	0.61	0.15	4.5·10 ⁻¹¹
Mudstone (M1)	1.7·10 ⁻⁴	0.87	0.32	2.3·10 ⁻¹¹







Conclusions

- The results obtained with the standard **approach** exceed the physically possible pore size values in tight pore space.
- The **proposed approach** enabled precise PSD estimation in rocks with high X values.

Acknowledgements: This research was funded via AGH University of Krakow statutory funds.