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Key Points:

- New framework reveals overestimated impact of ER on water storage when mining and agriculture are ignored
- GRACE satellite data in Mu Us Sandyland is affected by coal mining, necessitating careful interpretation
- Comprehensive analysis shows farming consumes more water than ER, informing sustainable practices

Supporting Information:

Supporting Information may be found in the online version of this article.

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Disentangling Ecological Restoration's Impact on Terrestrial Water Storage

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Abstract Large-scale ecological restoration (ER) in semiarid regions is often associated with substantial terrestrial water storage (TWS) depletion. This study challenged previous estimates by demonstrating the critical importance of considering other human activities when assessing ER impacts on TWS. Using a novel analytical framework integrating GRACE satellite data and ground observations, we analyzed TWS changes in China's Mu Us Sandyland under two scenarios: with and without considering mining and farming activities. Our results show that ER consumed TWS at an average rate of 11.7 ± 12.2 mm yr⁻¹ from 2003 to 2022. Neglecting the impacts of mining and farming led to a 251% overestimation of ER's effect on TWS. This study provided a more nuanced understanding of water resource dynamics in restored ecosystems, emphasizing the need for comprehensive approaches in TWS assessments and informing sustainable land management strategies globally.

Plain Language Summary Planting trees and restoring landscapes, known as ecological restoration (ER), helps mitigate climate change and improve dry environments. However, some are concerned that these efforts may delete significant amounts of water, exacerbating the strain on already limited water resources. Our study took a fresh look at this issue in the Mu Us Sandyland, a key area for restoration in northern China. We used satellite data, ground measurements, and computer models to track water changes before and after restoration initiatives. Importantly, we also considered how other human activities, including coal mining and farming, affect water resources. We discovered that previous studies may have overestimated water restoration usage by approximately 251% due to the failure to account for the impacts of other activities. This research helps us better understand how restoring landscapes affects water resources in dry areas, which is crucial for planning sustainable environmental projects worldwide.

1. Introduction

Large-scale ecological restoration (ER) of degraded lands has emerged as a critical strategy for enhancing ecosystem services and ensuring environmental security at global scales (Feng et al., 2013). Since 1998, China has initiated an unprecedented series of ER programs, aiming at combating soil erosion, desertification, and climate change while improving dryland ecosystems (Bryan et al., 2018; Lü et al., 2015; Ouyang et al., 2016). Initiatives such as the Three-North Shelterbelt Development Program and the Grain for Green Program have significantly increased vegetation greenness and land improvement (Deng et al., 2014; Mu et al., 2022). However, the success of these programs has led to an unexpected consequence: the intensification of evapotranspiration (ET) due to increased vegetation cover. This has resulted in a decrease of soil water storage and groundwater levels, ultimately reducing the total terrestrial water storage (TWS) and exacerbating water scarcity in affected regions (Deng et al., 2016; Feng et al., 2016; Jia et al., 2017; Lu et al., 2018; Zastrow, 2019; Zhao et al., 2020).

The reduction in TWS poses significant challenges for both socioeconomic development and the protection of water supplies for human and ecosystem needs (Cosgrove & Loucks, 2015). However, some studies suggest that large-scale vegetation recovery may enhance the atmospheric water cycle and increase precipitation through

ecosystem feedback mechanisms, potentially mitigating some of the impacts on water resources (Chen et al., 2023; Hoek Van Dijke et al., 2022; Tian et al., 2022; Zhang et al., 2023). Given these complex interactions, accurately quantifying changes in TWS and the impact of ER is crucial for ensuring sustainable water resource management in restoration practices, particularly in water-limited drylands.

To address this challenge, researchers have utilized gravity field measurements from the Gravity Recovery and Climate Experiment (GRACE) satellite, combined with multiple other environmental observations, to estimate changes in TWS and the impact of ER (An et al., 2021; Rodell et al., 2018; Tapley et al., 2019; Zhao et al., 2020). GRACE provides monthly anomalies in the Earth's gravitational field with unprecedented precision (Li et al., 2022; Tapley et al., 2004). While these gravitational variations are primarily driven by water redistribution at monthly scales (Tapley et al., 2004), it is essential to consider other anthropogenic impacts, such as reservoir regulation, water diversion, and bulk commodity transport, when interpreting GRACE data in regions with intensive human activities (Tang et al., 2013; Xie et al., 2018; Zhou et al., 2023). Previous studies have shown that in some regions, such as the North China Plain, the mass-loss rate of groundwater depletion observed by GRACE is significantly offset by mass gains from reservoir regulation, water diversion, and coal transport (Tang et al., 2013). This underscores the necessity of adopting a more comprehensive approach to estimating TWS changes, one that considers region-specific anthropogenic impacts on mass variations.

Our study focused on the Mu Us Sandyland (MUS) and its surrounding areas in northern China (Figures 1a and 1b), a semiarid region that has undergone over 20 years of large-scale ER practices. This area is regarded as one of the most successful examples of vegetation recovery and desertification reversal worldwide (Han et al., 2020; Zhang & Wu, 2020; Zheng et al., 2020). Satellite observations of the Normalized Difference Vegetation Index (NDVI) from 1981 to 2022 revealed a fluctuating yet overall upward trend, rising from 0.15 in 1981 to 0.22 in 2022 (Figures 1c and 1d). Annual land cover data indicated a rapid decrease in bare land area and an increase in grassland area from 1998 to 2002 (Yang & Huang, 2021; Figure 1e), marking the transition between pre-ER (1981–1997) and post-ER (2003–2022) periods.

Previous studies using GRACE time-series data reported that ER in these regions consumed substantial TWS (Cao et al., 2022; Zhao et al., 2020). This finding has raised concerns among the Chinese government and scientists regarding the sustainable water use in ER programs. However, these estimates are questionable for overlooking other human activities, such as coal mining and agricultural water consumption, which may significantly impact mass variations in the region. To address these limitations and provide a more accurate assessment of ER's impact on TWS, we propose a new analytical framework for examining the impact of ER on the TWS in regions with intensive restoration and human activities (Figure 2). Our approach integrates multiple GRACE satellite solutions, field observations, government reports, and ecohydrological modeling to generate a comprehensive picture of TWS changes from 2003 to 2022.

The uniqueness of this study lies in its holistic approach to disentangling the effects of ER from other anthropogenic impacts on TWS. By considering coal mining and agricultural activities, we aim to provide a more nuanced and accurate assessment of ER's impact on water resources. Our findings will have significant implications for the design and implementation of future ER programs, not only in China but also in other regions facing similar challenges in balancing ER with water resource management.

2. Methods

Figure 2 illustrates the approach used to disentangle the effects of ER from those of coal mining and farming on TWS during 2003–2022 period, by combining GRACE data with a water balance approach. The detailed steps are explained below. Note that all variables that lead to TWS loss have negative values, such as coal and water mass flow from mining, ET, and runoff.

2.1. Estimate of Original GRACE TWS Changes (ΔTWS_{ori})

A time series of changes in original GRACE TWS (Δ TWS_{ori}) from 2003 to 2022 was estimated by analyzing five GRACE and GRACE Follow-On (GRACE-FO) solutions, including three standard spherical harmonic (SSH) GRACE solutions (CSR, JPL, and GFZ) and two Mascon solutions (CSR and JPL). Comparison between observed water table data (2018–2022) from 97 wells in MUS (Figure 1b) and derived groundwater changes for various GRACE solutions revealed that CSR and JPL Mascon solutions exhibited the highest correlation with



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Figure 1. Location of MUS and observations. (a) Location of MUS (black polygon); (b) Geographical position of groundwater monitoring wells, meteorological stations, and runoff hydrological stations in the study region; (c) Trend in annual NDVI over MUS during 1981–2022; (d) Annual mean NDVI time series from GIMMS-3g and MODIS averaged over the study region; and (e) Temporal changes in the area percentage of two major land cover types (bare land and grassland) from 1990 to 2022.

measured water table data (Text S1 in Supporting Information S1; Figure S1 in Supporting Information S1). The average of these two solutions was used to calculate the annual changes in TWS of our study area, balancing the strengths of both data sets. Note that GRACE and GRACE-FO have an 11-month data gap (July 2017 to May 2018), which limits the TWS continuity. The Variational Mode Decomposition with Long Short-Term Memory (VMD-LSTM) was used to fill the 11-month data gap between GRACE and GRACE-FO missions (Text S2 in Supporting Information S1; Figures S2 and S3 in Supporting Information S1). Interpolation results closely matched GRACE data, ensuring continuity in our TWS time series (Figure S3 in Supporting Information S1).

2.2. Effects of Coal Mining (TWS_M) on TWS and Corrected GRACE TWS (ΔTWS_{cor})

Effects of coal mining on GRACE TWS (TWS_M) were calculated by considering both the annual mass loss due to coal mining (M_{cmf}) and the associated groundwater co-production (M_{wmf}) as:





Figure 2. Analytical framework to study the potential impact of ER on total TWS. ΔTWS_{ori} is annual changes in original GRACE TWS. ΔTWS_{cor} is annual changes in corrected GRACE TWS with coal and water mass flow deductions from mining. M_{cmf} and M_{wmf} are equivalent water thickness of annual coal and water mass flow from mining, respectively. ΔTWS_{exEF} is the estimated annual changes in TWS with the exclusion of ER and farming. ΔTWS_{exEF} is the estimated TWS trend with the exclusion of ER and farming. ΔTWS_{exE} is the estimated TWS trend with the exclusion of ER only. ΔET_c is the increased ET from farming. ET_{exEF} is the simulated ET excluding ER and farming impact. TWS_M , TWS_F and TWS_E represent impact of mining, farming and ER on TWS, respectively. Bule data represent the multi-year averaged values for each variable and red data represent the effects of various activities on TWS in 2003–2022. Negative values represent the loss of TWS.

$$TWS_{M} = M_{cmf} + M_{wmf}$$
(1)

The annual mass loss due to coal mining (i.e., the amount of raw coal mined) (M_{cmf}) (ton yr⁻¹) was obtained from China Statistical Yearbook (NBS, 2003–2022) and converted to equivalent water thickness (mm yr⁻¹) by assuming water density of 1 ton m⁻³.

During coal mining, significant groundwater volumes are pumped to dewater coal seams (Doulati Ardejani et al., 2011; Mu et al., 2018). This groundwater is discharged directly into surface water runoff or evaporated into the atmosphere. Given the scale of coal production in the study region, both coal transport and the associated water co-production are major sources of mass loss. Groundwater consumption for coal mining can be estimated using the water consumption coefficient per ton of coal (Xie et al., 2018) as:

$$M_{\rm wmf} = \frac{\mu M_{\rm cmf}}{\rm S} \times 1000 \tag{2}$$

where M_{wmf} (mm yr⁻¹) is annual groundwater consumption from coal mining, μ (m³ ton⁻¹) is water consumption coefficient per ton of coal with an average value of 0.45 m³ ton⁻¹ from 25 coal mines in MUS (Zhang et al., 2013), M_{cmf} (ton yr⁻¹) is the annual mass loss due to coal mining, *S* is the area of the study region (m²), and 1000 is the coefficient to convert meters to millimeters.

After removing the effects of mining activities, the corrected annual changes in GRACE TWS (Δ TWS_{cor}) from 2003 to 2022 was calculated as:





Figure 3. Impact of coal mining, farming, and ER on TWS. (a) Annual changes in original TWS (ΔTWS_{ori}) and TWS with coal and water mass flow corrections (ΔTWS_{cor}) using GRACE solutions from 2003 to 2022. The shaded area represents ±1sd; (b) Annual changes in TWS from 2003 to 2022 excluding ER and farming effect (ΔTWS_{exEF}) and TWS excluding ER effect only (ΔTWS_{exE}). The shaded area represents ±1sd; (c) Blue bars represent annual coal mass flow from 2003 to 2022, and gray points represent annual cropland area proportion during the study period; (d) Comparison between observed and simulated multi-year averaged TWS trend during post-ER period 2003–2022.

$$\Delta TWS_{cor} = \Delta TWS_{ori} - TWS_{M}$$
(3)

2.3. Effects of Farming (TWS_F) on the TWS

Cropland area in the study region increased by 5.2% (from 6.9% to 12.1%) over the past 20 years (Figure 3c and Figure S4 in Supporting Information S1). After irrigation with groundwater, excess water is often replenished back into the aquifer, the net effect of agricultural activities on TWS is primarily from ET (Kendy et al., 2004; Zou et al., 2017). Therefore, the effect of farming on the TWS (TWS_F) during 2003–2022 period was calculated as:

$$TWS_{\rm F} = \Delta ET_{\rm c} \times \frac{S_c}{S} = (ET_{\rm c} - ET_{\rm exEF}) \times \frac{S_c}{S}$$
⁽⁴⁾

where ΔET_c (mm yr⁻¹) is the increased ET from farming, ET_c (mm yr⁻¹) is cropland ET, ET_{exEF} is the simulated ET of natural land without considering ER and farming (see Section 2.4), S_c and S is the area of cropland and the entire study region, respectively.

All croplands in the study region are irrigated. Because cropland remains relatively small and sporadically distributed, we calculated the ET_c using the Penman-Monteith equation (Wang et al., 2013; Text S3).

2.4. Calculation of 2003–2022 TWS Excluding ER Impact (ΔTWS_{exE})

We first calculated the annual changes in TWS during 2003–2022 period excluding ER and farming (ΔTWS_{exEF}). Subsequently, the associated changes in TWS excluding only ER (ΔTWS_{exE}) were determined by incorporating the farming effects (Equation 4) into the TWS values that exclude both ER and farming, as follows:

$$\Delta TWS_{exE} = \Delta TWS_{exEF} + TWS_F$$
(5)



The ΔTWS_{exEF} was calculated using the water-balance equation (Wan et al., 2015; Zeng et al., 2012; Zhao et al., 2020) as:

$$\Delta TWS_{exEF} = P + R + ET_{exEF}$$
(6)

where *P* is annual precipitation, *R* is annual runoff, and ET_{exEF} is the annual ET simulated with the modified Shuttleworth-Wallace (S-W) model (Li et al., 2023; Shuttleworth & Wallace, 1985). Runoff data were derived from the Baijiachuan and Wenjiachuan hydrological stations (Figure 1b). Meteorological data (1981–2022), including precipitation, temperature, air pressure, relative humidity, and wind speed were obtained from 28 stations across the study region (Figure 1b). For the land use and vegetation-related parameters of the model, annual average values from the pre-ER period (1981–1997) were used to exclude the impacts of ER and farming activities. The model performance was evaluated using the observed ET data from 54 flux towers across China (Figure S5 in Supporting Information S1), and the detailed calculation procedure are referred in Supporting Information S1).

2.5. Effect of ER (TWS_E) on the TWS

The impact of ER (TWS_E) on TWS was calculated as:

$$TWS_{E} = \Delta TWS_{cor} - \Delta TWS_{exE}$$
⁽⁷⁾

This can be derived by combining Equations 1-6, yielding:

$$TWS_{E} = \Delta TWS_{ori} - TWS_{M} - (P + R + ET_{exEF}) - TWS_{F}$$
(8)

Thus, the effects of ER on the TWS during 2003–2022 were disentangled from those of coal mining and farming, based on the GRACE observations and ecohydrological models (Figure 2).

3. Results

Original GRACE TWS observations (ΔTWS_{ori}) indicated an average decreasing trend of $-9.8 \pm 6.7 \text{ mm yr}^{-1}$ from 2003 to 2022 (Figures 3a and 3d). The MUS extracts coal with 341.9 million tons per year (Figure 3c). This net coal mass change in the MUS is equivalent to a TWS loss of approximately $-5.1 \pm 0.6 \text{ mm yr}^{-1}$. After deducing the coal mass flow, the adjusted GRACE-derived TWS trend was $-4.7 \pm 7.0 \text{ mm yr}^{-1}$. In addition, the average groundwater loss due to coal seam dewatering was estimated to be approximately $-2.3 \pm 0.3 \text{ mm yr}^{-1}$. Therefore, after accounting for the total coal mining impact (TWS_M), the corrected GRACE-derived TWS decreasing trend (ΔTWS_{cor}) was $-2.4 \pm 6.6 \text{ mm yr}^{-1}$ (Figures 3a and 3d).

The modified S-W model, calibrated with data from 54 flux towers across China (Figure S5 in Supporting Information S1), accurately simulated long-term ET trends (NSE = 0.49, RMSE = 1.9 cm, bias = 3.19; Figure S6a in Supporting Information S1). Despite a minor underestimation of ET, the model demonstrated greater precision in dryland ecosystems (Figure S6b in Supporting Information S1), making it suitable for our study area. During the post-ER period (2003–2022), excluding the effects of ER and farming, the estimated TWS (Δ TWS_{exEF}) based on water-balance equation showed an increasing trend at an average rate of 31.3 ± 13.2 mm yr⁻¹ (Figures 3b and 3d). The estimated multi-year average cropland ET_c, calculated using the Penman-Monteith equation, was approximately 553.3 mm yr⁻¹, and the net effect of farming on TWS (TWS_F) in the study region is approximately -22.0 ± 0.9 mm yr⁻¹. Therefore, the TWS, excluding the effects of ER only (Δ TWS_{exE}), showed an increasing trend at an average rate of 31.3 ± 12.9 mm yr⁻¹ (Figures 3b and 3d).

By comparing ΔTWS_{cor} and ΔTWS_{exE} , results showed that ER consumed TWS at an average rate of $-11.7 \pm 12.2 \text{ mm yr}^{-1}$ during 2003–2022 (Figure 3d). Importantly, when coal mining and agricultural water consumption are not accounted for, the apparent ER impact on TWS (i.e., $\Delta TWS_{ori} - \Delta TWS_{exEF}$) is $-41.1 \pm 13.1 \text{ mm yr}^{-1}$, an overestimation of approximately 251%.





Figure 4. Impact factors and mechanisms of TWS in semiarid sandy areas. (a) ER leads to a reduction in water resources through the use of aquifers; (b) Groundwater withdrawal for crop irrigation from aquifers leads to the reduction in water resources. The blue arrow indicates pumping for irrigation; and (c) Mining operations below aquifer cause a drop in the water table because of pumping groundwater for dewatering coal seams. The upward blue arrow represents dewatering coal seams, the downward blue arrows indicate groundwater infiltration due to cracks of the aquifer, and the upward white arrows indicate coal extraction processes. Dashed lines represent the water table before ER, irrigation, and mining, and solid lines represent the water table after various human activities.

4. Discussion and Conclusions

Our study presents a novel analytical framework for accurately evaluating the impacts of ER and other human activities on TWS in semiarid regions. By integrating multi-source data and accounting for complex anthropogenic factors, we provide a more nuanced understanding of TWS dynamics and mechanisms in the MUS (Figure 4), with implications for water resource management in similar regions worldwide.

4.1. Feasibility of the Analytical Framework

We found that failing to account for the mass-loss from coal mining and transport resulted in a substantial overestimation of water depletion. The importance of this correction is evident when comparing our adjusted GRACE-derived TWS change (mean = -2.4 mm yr^{-1}) with ground-based shallow aquifer measurements (mean = -2.7 mm yr^{-1}) in Yellow River Water Resource Bulletins (YRCC, 2003–2022). The close agreement validates our approach and highlights the potential for misinterpretation when relying solely on uncorrected GRACE data. Our findings have broad implications for TWS studies in regions with significant mining or other mass-altering activities, emphasizing the need for a more comprehensive approach to satellite-based water resource assessments.

4.2. Reassessing ER Impact on Water Resources

After accounting for coal mining and agricultural activities, we found that ER reduced TWS at a rate of $11.7 \pm 12.2 \text{ mm yr}^{-1}$ during 2003–2022, which is lower than earlier estimates (Zhao et al., 2020). This finding challenges the prevailing narrative of a looming water resource crisis caused by large-scale ER programs. The reduced impact can be attributed to the Chinese government's strategic selection of water-saving tree species and the predominant use of grassland and native shrubs in restoration efforts. For instance, the average multi-year ET of shrubs in MUS (294 mm) is significantly lower than the multi-year average precipitation (402 mm) (Han et al., 2023). This water-conservative approach allows for natural replenishment of water resources consumed by vegetation, minimizing the net impact on regional water availability. Moreover, ER may not result in continued water resource depletion over time. The observed recovery of soil water content with increasing tree age indicates that restored ecosystems and individual plants adapt to water stress conditions, potentially leading to long-term stability in water use (Jia et al., 2017, 2020; Wang et al., 2024).

Note that the potential positive feedback of ER to precipitation was not considered in this study, as the contribution of local ER to enhanced precipitation remains uncertain in the relatively small area. The precipitation trend during post-ER period (2003–2022) in MUS shows an increase of 4.2 mm yr⁻¹. Regional and global analyses indicate that ER may have contributed to the increased precipitation in the study area, with estimated increases of

 0.85 mm yr^{-1} (Tian et al., 2022) and 1.6 mm yr $^{-1}$ (Hoek Van Dijke et al., 2022), respectively. Consequently, the negative impact of ER on TWS is expected to be further mitigated if the feedback effect is taken into account.

4.3. Challenges in Agricultural Water Consumption and Groundwater Management

While our study provides a more optimistic perspective on the impacts of ER, it also highlights critical challenges in agricultural water consumption and groundwater management. First, our results indicate that agricultural water consumption is a primary factor for TWS loss over MUS. This is attributed to the relatively shallow and abundant groundwater resources in the study region, which serves as the primary source for irrigation (Huo et al., 2016; Li et al., 2021b). Second, in recent years, the local government has introduced the "Requisition-Compensation Balance" initiative, a policy that has incentivized local farmers to engage in the reclamation of arable land, leading to a significant expansion of cultivated areas (Gao et al., 2024).

Additionally, the observed rapid decline in the water table (Figure S7 in Supporting Information S1), particularly near agricultural and mining areas, highlights localized issues of overexploitation and disturbance that may not be fully captured by regional-scale TWS assessments. The discrepancy between GRACE-derived TWS trends and site groundwater measurements underscores the complexity of water resource dynamics in heavily modified landscapes. Our findings suggest that water table decline in the region may be more strongly influenced by groundwater withdrawal for agricultural irrigation and mining-induced aquifer disruption than by ER activities. Mining activities conducted below the water table not only requires dewatering but also may disrupt aquifers and cause mining-induced cracks (Zhang et al., 2016), resulting in neighboring observation wells showing an exaggerated drop in water table (Figure 4c). This emphasizes the need for targeted groundwater monitoring and management strategies, particularly in areas with intensive human activity.

4.4. Global Implications and Future Directions

Our study demonstrates the importance of a holistic approach to water resource assessment that considers multiple anthropogenic factors. The analytical framework developed here can be applied to other regions with intense human activity, providing a more accurate evaluation of TWS trends globally. This approach is particularly relevant as countries worldwide implement large-scale ER projects to address climate change and land degradation. The findings challenge the notion that ER programs inevitably lead to water resource depletion, offering a more nuanced perspective on the water-vegetation relationship in semiarid ecosystems. This has important implications for global restoration efforts, suggesting that carefully planned and managed ER programs can potentially enhance both ecosystem health and water resource sustainability.

Future research should focus on.

- Refining methods to separate the impacts of various human activities on TWS.
- Investigating long-term adaptation mechanisms of restored ecosystems to water stress.
- Quantifying the feedback effects of large-scale vegetation changes on regional and global water cycles.
- Developing integrated monitoring systems that combine satellite observations with high-resolution ground measurements to provide a more comprehensive understanding of water resource dynamics.

In conclusion, our study provides a more optimistic outlook on the water resource implications of ER in semiarid regions. However, it also highlights the critical need for targeted management of groundwater resources, particularly in areas with intensive agricultural and mining activity. By adopting a more comprehensive approach to water resource assessment and management, we can better handle the complex interactions between human activities, ecosystem restoration, and water availability in a changing climate.

Conflict of Interest

The authors declare no conflicts of interest relevant to this study.

Data Availability Statement

Data sets used in this study are collected from various sources: Standard spherical harmonic (SSH) GRACE solutions are available at https://icgem.gfz-potsdam.de/home. GRACE JPL mascon solution is available at https://grace.jpl.nasa.gov/data/get-data/jpl_global_mascons/. GARCE CSR mascon solution is available at

https://www2.csr.utexas.edu/grace/RL06_mascons.html. NDVI during 1981–2015 from GIMMS-3g are available at https://climatedataguide.ucar.edu/climate-data/ndvi-normalized-difference-vegetation-index-3rd-generation-nasagfsc-gimms. NDVI during 2002–2022 from MODIS are available at Didan (2015). CO₂ concentration data are derived from Lan et al. (2025). LAI during 2002–2022 from MODIS are available at Myneni et al. (2021). LAI during 1981–2010 from AVHRR are available at http://www.glass.umd.edu/Download.html. The storage of plant canopy surface water, soil water, and snow water are available at Beaudoing and Rodell (2020). The ET data from flux tower observations are obtained at FLUXNET (https://fluxnet.org/data/fluxnet2015-dataset/), China-FLUX (http://www.nesdc.org.cn/theme/index?projectId=64e80ed07e2817429fbc7b09), and CERN (http:// www.cnern.org.cn/data/initDRsearch?classcode=STA). Soil hydraulic parameters data are derived from Zhang et al. (2018).

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