





## 1) Background

- The 2016 Mw 6.1 shallow crustal left-lateral strike-slip Tottori earthquake occurred in the central part of the Tottori prefecture in the Chugoku region in western Japan.
- It caused strong ground motions with a maximum PGA of 1.4 gal.
- The mainshock was preceded by 70 foreshocks over a 12-hour period (the largest with Mw 4.2) and followed by more than 10,000 aftershocks over the following 10 days (the largest with Mw 5.0). • This earthquake occurred in the San-in shear zone, which is part of the largest right-lateral
- Northern Chugoku shear zone. • The Tottori prefecture has been struck by several large crustal earthquakes with severe damage
- during the last decade (the 1947 Mw 7.0 Tottori, the 1983 Mw 5.5 Central Tottori, the 2000 Mw 6.7 Western Tottori).

# 2) Motivation

- Published kinematic rupture models inferred either from geodetic or seismic data exhibit significant discrepancies (Ross et al., 2018; Kubo et al., 2017; Meneses-Gutierrez et al., 2019; Amey et al., 2019; Shibata et al., 2025).
- heterogeneous prestress and friction parameters

# 3) Data

- We use data from 21 seismic stations located within 50 km of the epicenter, obtained from the  $_{35^{\circ}30'N}$ NIED strong-motion seismograph network.
- The data are integrated to velocities and bandpass filtered between 0.05-1.2 Hz (TTR005), 0.1-0.6 Hz (TTRH07) and 0.05-0.6 Hz (others).
- We use horizontal and vertical coseismic static displacements from 33 GNSS Earth Observation <sup>35°1</sup> Network stations inferred by Meneses-Gutierrez et al. (2019) and Amey et al. (2019).
- 1D layered Crustal model is extrapolated from the Japan Integrated Velocity Structure Model (Koketsu et al., 2008, 2012) under each station <sup>35°00'N</sup> separately.
- Location of foreshocks and aftershocks are adopted from Ross et al. (2018).



133°45'E 134°15'E 134°30'E 134°00'F Figure 2: Map of NIED seismic stations (triangles) and geodetic GNSS observations (circles, DPRI marked with their name and GSI marked with numbers) used in this work. The black and gray lines represent the surface projections of the upper and lower along-strike sections of the fault, respectively. The red star indicates the location of the hypocenter, whereas the black dots denote the positions of foreshocks and aftershocks.

## 4) Method

We utilize fd3d\_tsn\_pt code (Gallovič et al., 2020) for dynamic simulation of rupture propagation, calculation of synthetic seismograms and static displacements, and inversion.

## • Forward problem:

- We utilize the 3D finite-difference staggered grid code FD3D\_TSN (Premus et al., 2020) with GPU acceleration to simulate the dynamic rupture propagation.
- We assume the classic linear slip-weakening friction law with spatially variable  $T_c = T$ dynamic parameters on a planar vertical fault.
- Green's functions are precalculated in 1D velocity models acquired for each station from a 3D model (seismic: Axitra code by Cotton and Coutant, 1997; geodetic: Okada, 1985).
- The result of dynamic simulation is a spatio-temporal distribution of slip rates along the fault
- Synthetic seismograms are calculated using the representation theorem (Aki & Richards, 2002).
- Parameterization :
  - For forward problem, we prescribe inhomogeneous distributions of the three dynamic parameters of the linear slip-weakening friction law on the fault
- Dynamic parameters are prestress  $T_0 = T_i T_d$  (i.e. the difference between the initial and dynamic value of the shear traction), *friction drop*  $\Delta\mu$  (i.e. the difference between the static and dynamic friction coefficient) and slip-weakening distance  $D_{c}$ .

## Inverse problem:

- The inverse problem is formulated in a Bayesian framework, employing the Parallel Tempering Markov Chains Monte Carlo approach to sample the posterior distribution of the model parameters (Sambridge, 2013) with modified Metropolis-Hastings acceptance rule.
- We assume uniform prior PDFs of the model parameters within wide ranges; see Table 1
- We explore the model space using 7 MPI processes on three Nvidia RTX 3070 GPUs, running the computation for a total of 82 days.



# INVERSE PHYSICS-BASED MODELING OF THE 2016 MW 6.1 TOTTORI EARTHQUAKE

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• In this study, we perform a so-far missing dynamic source inversion with slip-weakening friction law with spatially distribution on the fault to better understand this event

eter	Value
al	
imension	Length 16 km, width 16 km
nechanism	Strike $160^{\circ}$ , dip $88^{\circ}$ , rake $0^{\circ}$
discretization	100 m
f-domain size (along strike, normal, dip)	161, 110, 161
r of time step (FD)	5 000
tep in seconds (FD)	$2\mathrm{ms}$
	90°
tational time (GPU)	$15\mathrm{s}$
's functions	
fault discretization	$1\mathrm{km}$
um computed frequency	2.5 Hz (Nyquist)
ampling	0.2 s
ranges of parameters for inversion	
r of model parameters (discretization)	11, 11
ss prior	(0;100) MPa
Dynamic friction coefficient prior	(0; 1.0)
teristic slip-weakening distance prior	$(0.04; 2.0) \mathrm{m}$
tion position (along strike, along dip)	$11\mathrm{km},4.5\mathrm{km}$
tion radius prior	1000 m
and modeling errors	
data uncertainty	0.05 m

(6.05; 6.25)Table 1: Parameters for forward FD simulation, Green's functions and dynamic inversion model setup.

## 5) Results

- The total number of visited models is approximately After the removal of the burn-in phase models, the er used for further analysis consists of 14,764 models.
- Here, we first present the maximum a posteriori (MAP) model,

follo	owed by the enser	nble with its statisti	CS.
TTDOOF	N-S	E-W	Z
TTR005 -		- Museum -	
TTRH07 -			
OKY015 -			
TTR004 -		M	
TTRH04 -	- Amma -		
TTR006 -		- Asper -	
OKY001 -			- AA
OKYH09 -	- Marine -		- Massim
OKY002 -			
TTRH06 -	- marine -	- My vara -	
TTR003 -		- My Man-	
TTRH03 -	- Masan -		
TTR007 -			
TTR002 -	hopen -	- Alexandre -	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~
OKY006 -			
OKYH11 -	- Andres -		
OKY005 -			
TTRH02 -	- Alar -	Am -	
TTR008 -	- And -		
OKY003 -	- Avar -	- Ara -	
TTR001 -	- AMA -		

Figure 4: Comparison between observed (black) and synthetic (red) seismograms for the MAP rupture model. The left column lists the seismic station codes, followed by the three components: north-south, east-west, and vertical. The right column displays the amplitudes in cm/s.

Figure 5: Distribution of the rupture parameters on the fault of the MAP rupture mode Black lines contour the region where the slip exceeds 10% of the maximum slip value violet line bounds the nucleation and black dots denote foreshocks and aftershocks projected on the fault plane.

1	.2	million.
n	se	mble











4 6 8 10 12 14 16 Along strike (km)



Parameter	Value
Ruptured area $A_R$	$92.76{\rm km^2}$
Average speed $\overline{v_R}$	$2059\mathrm{m/s}$
Seismic moment $M_0$	$2.34 \times 10^{18} \mathrm{Nm}$
Average stress drop $\overline{\Delta T}$	7.75 MPa
Fracture energy $E_F$	$2.55 \times 10^{14} \mathrm{J}$
Radiated energy $E_R$	$4.01 \times 10^{13} \mathrm{J}$
Radiation efficiency $\eta_{rad}$	13.6%
Seismic misfit	43.69
Seismic VR	$54.99\%$ for shift $0.2 \mathrm{s}$
GPS misfit	2.31
GPS VR	98.94%



Figure 6: Observed (black) and synthetic (red) GNSS static displacements (scaled) for the MAP rupture model. The thick black and gray lines represent the surface projection of the lower and upper along strike parts of the fault, the red star marks the epicenter location, and the gray dots indicate the positions of foreshocks and aftershocks.



Figure 7: Slip distribution and slip rate functions (violet curves) on the fault (averaged over 1km x 1km subfaults). Black star denotes the position of hypocenter and black dots are foreshocks and aftershocks projected on the fault plane.



Figure 8: Slip contours on the fault for all accepted model samples The white-and-black scale represents the normalized probability density function (PDF). Thick violet lines indicate the contours of the mean model, while the two thin violet lines represent the averaged slip model with added and subtracted one-sigma uncertainty.



Figure 9: Mean input model parameters on the fault (averaged over the ensemble models) with their uncertainties in terms of one sigma interval, as inferred by the Bayesian dynamic finite fault inversion. Black and violet curves outline the contours of the mean slip model with one sigma uncertainties. Black dots denote the position of aftershocks and foreshocks.



Figure 13: Averaged Dc values over the specified annulus of distance from nucleation center for the mean (blue), MAP (black) and initial model (yellow). The values are fitted with linear regression lines (see legend).

earthquake. Journal of Geophysical Research: Solid Earth, 124, 6970–6988. https://doi.org/10.1029/2019JB017512



![](_page_0_Picture_73.jpeg)

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![](_page_0_Figure_75.jpeg)

Distance from nucleation center (R) [m]

Amey R. M. J., Hooper A., & Morishita Y. (2019). Going to any lengths: Solving for fault size and fractal slip for the 2016, Mw 6.2 Central Tottori earthquake Hisahiko Kubo, Wataru Suzuki, Shin Aoi, and Haruko Sekiguchi. Source rupture process of the 2016 central Tottori, Japan, earthquake (M JMA 6.6) Gallovič F., Valentová Ľ., Ampuero J.-P., & Gabriel A.-A. (2019). Bayesian dynamic finite-fault inversion: 1. Method and synthetic test. Journal of Geophysi

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Koketsu, K.; Miyake, H., et al., 2012. An updated integrated velocity structure model of Japan based on recent geological and geophysical observations. Journal or

Meneses-Gutierrez A., Nishimura T., & Hashimoto M. (2019). Coseismic and postseismic deformation of the 2016 Central Tottori earthquake and its slip model

Figure 12: The relationship between selected inferred

correlation coefficient (black number in the subplot)

Disaster Research. Available also from: https://www.iitk.ac.in/nicee/ publications.

parameters, with the respective Pearson product-moment

0.53

2 4 6 8