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Session: HS7.4 – Future hydroclimatic scenarios in a changing world



Stochastic Analysis of the Hydrological Cycle in the Mediterranean and its Recent Climatic Variations

Marianna Lada, Christina-Ioanna Stavropoulou, Dimitra-Myrto Tourlaki, Nikos Tepetidis, Panayiotis Dimitriadis, Theano Iliopoulou and Demetris Koutsoyiannis



School of Civil Engineering Department of Water Resources and Environmental Engineering National Technical University of Athens









Having collected data from the **Climexp** (KNMI) ERA5 Reanalysis dataset [1] for multiple climate variables across different altitudes in the Mediterranean region, for the period 1950-2024, we investigate the long-term persistent (LTP) trends/slopes and the stochastic variability [2].

Climate variables included:

- Temperature [3]
- Precipitation [4]
- Evaporation and potential evaporation [5]
- Wind speed [6]
- Zonal and meridional wind
- Column water and relative humidity [7]
- Latent, sensible heat flux and specific humidity [7]

Evaporation Stability Despite Warming: The Role of Wind Speed



0.0236

-0.0011

0.0002

3

Slope



Temperature 2-10 m





Although temperature shows a clear positive trend, evaporation does not follow, suggesting it is not primarily temperature-driven (r=0.05).

Instead, it correlates more with wind speed (r=0.44, [8]), which, like evaporation [9], remains nearly stable over time (slopes: -0.0011 and +0.0002, respectively).



To capture the **stochastic and persistent nature** of hydrological processes, we apply the **Hurst-Kolmogorov model** using observed climatic data. [10]

$$\gamma_k \coloneqq var[\underline{x}_j^{(k)}] = \gamma_1/k^{2-2H}$$

Hurst	Empirical	Theoritical (Corrected for Bias)
Temperature	0.90	0.96
Zonal Wind	0.62	0.65
Column Water	0.77	0.82
Meridional Wind	0.53	0.55
Wind Speed	0.64	0.68
Evaporation	0.79	0.85
Precipitation	0.73	0.78
Latent Heat Flux	0.79	0.85
Sensible Heat Flux	0.92	0.97





Zonal Wind (Zonal Flow)

- Refers to air movement that follows a path parallel to the equator and lines of latitude – flowing either from west to east (westerly) or from east to west (easterly).
- This pattern of circulation tends to support **consistent temperatures along those latitudinal zones**. [15]

Sectional Wind (Meridional Flow)

- Indicates to air movement that follows the lines of longitude (meridians) – flowing either from north to south or from south to north.
- Responsible for: cold air outbreaks from the poles moving south
 warm air surges from the tropics moving north [15]



Zonal Wind

Meridional Wind





Total Column Water Vapour

- Refers to the total amount of **water vapour in a vertical column** of the atmosphere, from the surface to the top of the atmosphere.
- If fully condensed and precipitated, it would form a surface layer of water of equivalent thickness. [15]

Relative Humidity

- The ratio of the current amount of water vapour in the air to the maximum amount it could hold at that temperature.
- **Highly temperature-dependent**: warmer air can hold more water vapour. [15]

Units of measurement: mm or kg/m²

Expressed as a percentage (%)

The available data provide **Total Column Water Vapour at the surface** and **Relative Humidity at** pressure levels of **850 hPa** (≈1.5 km altitude), **500 hPa** (≈5.6 km) and **200 hPa** (≈11.8 km).



S Latent Heat Flux

- Involves heat transfer associated with phase changes of water, mainly evaporation and condensation.
- Specifically, latent heat flux is described by the following equation, where E' represents evaporation.

 $\Lambda = \lambda \times E'$

where λ (kJ/kg) is calculated as λ = 3139 - 2.336 T_s, with T_s the temperature in the surface of water in K.[16,18]

Sensible Heat Flux

• Refers to the **transfer of thermal energy** between the Earth's surface and the atmosphere due to temperature differences and occurs **without phase changes**. [15]

Specific Humidity

• Is equal to the **ratio of mass water vapor to the total mass of air**, regardless of whether the air is saturated – unlike relative humidity. [15]



$$p = 1013.25 \times [1-2.256 \times 10^{-5} z]^{5.256} \Leftrightarrow z = 44332 \left[1 - \left(\frac{p}{1013.25}\right)^{0.1925} \right]$$

while for $11,000 \le z \le 20,000$ m, pressure follows the expression below:

$$p = 226.27 \text{ e}^{-0.00015769(z-11000)} \Leftrightarrow z = 11000 + 6341.6 \ln\left(\frac{226.27}{p}\right)$$

• Units of measurement: altitude (z): m, pressure (p): hPa

Based on the above equations, the pressure values were converted to the corresponding altitudes as shown below:

Pressure	Altitude				
(hPa)	(km)				
Surface - 1013.25	0				
850	1.5				
500	5.6				
200	11.8				

8



Slope									
Variable	Unit/year	Altitude							
variable		Surface	2-10 m	1.5 km	5.6 km	11.8 km			
Temperature	°C/year	-	0.023612	0.018929	0.016121	-0.000921			
Zonal Wind	m/s/year	-	-0.001470	-0.004096	-0.009096	-0.011640			
Column Water	kg/m ² /year	0.014411	-	-	-	-			
Relative Humidity	%/year	-	-	-0.000436	-0.000074	0.000076			
Meridional Wind	m/s/year	-	-0.011640	-0.001518	-0.004820	-0.010786			
Wind speed	m/s/year	-	-0.001051	0.000010	0.000021	0.000021			
Evaporation	mm/day/year	-	0.000165	-	-	-			
Potential Evaporation	mm/day/year	-	-	0.002090	-	-			
Precipitation	mm/day/year	0.000000	-	-	-	-			
Latent Heat Flux	W/m ² /year	0.000000	-	-	-	-			
Sensible Heat Flux	W/m ² /year	-	0.000000	-	-	-			
Specific Humidity	kg/kg/year	-	-	0.000002	0.000001	0.000000			

The table summarizes the **linear trends (slopes)** of key atmospheric variables across different altitudes; however, there is uncertainty in the estimation due to the nature of classical methods and the LTP [8,13].

Increase in Temperature Slope





The temperature slope is increasing, exhibiting a positive trend.

It is observed that it also follows the temperature at heights of 2-10 m, with slopes +0.0020 and +0.0236, respectively.

Temperature



Slope

0.023612

0.018929

0.016121

-0.000921



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Zonal Wind





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Surface



Relative Humidity







Slop	Slope					
1.5 km	-0.000436					
5.6 km	-0.000074					
11.8 km	0.000076					

Meridional Wind





Slope					
2-10 m	-0.011640				
1.5 km	-0.001518				
5.6 km	-0.004820				
11.8 km	-0.010786				

AT11----

2020

2020

Wind Speed





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Evaporation and Potential Evaporation









Slope				
Surface	-0.001644			





Slope				
Surface	-0.004645			







Slope				
2-10 m	-0.013819			

Specific Humidity







11	8	2m	
τ.	0	VIII	



Slope				
1.5 km	0.000002			
5.6 km	0.000001			
11.8 km	0.000000			

Pearson Correlations Between Variables at Different Altitudes (I)

1.0

- 0.8

- 0.6

- 0.4

- 0.2

- 0.0

- -0.2



- 1.0

- 0.8

- 0.6

- 0.4

- 0.2

- 0.0

- -0.2

Pearson Correlation



1.00 0.05 -0.19 -0.21 0.05 -0.40 Temperature -Zonal Wind -0.05 1.00 0.48 0.64 0.06 0.02 Pearson Correlation Meridional Wind --0.19 0.48 0.67 0.29 0.23 1.00 Wind Speed --0.21 0.64 0.67 0.44 -0.02 1.00 Evaporation -0.05 0.05 0.29 0.44 1.00 -0.23 -0.40 0.02 0.23 -0.23 Sensible Heat Flux --0.02 1.00 Wind **Femperature** Meridional Wind Wind Speed Sensible Heat Flux Evaporation Zonal

2-10 m

- -1 ≤ r ≤ 1
- Significant positive correlation: **r > 0.25**
- Significant negative correlation: **r < -0.25**



Correlation

Pearson



- 1 00				500	hPa			_	- 1 0
- 0.75	Temperature -	1.00	0.03	-0.24	-0.08	0.04	0.71		- 0.8
- 0.50	Zonal Wind -	0.03	1.00	0.30	0.61	-0.46	0.14		- 0.6
- 0.25 - Correlation	Relative Humidity -	-0.24	0.30	1.00	0.44	-0.74	0.48		- 0.4 - 0.2
- 0.00 - Dearson	Meridional Wind -	-0.08	0.61	0.44	1.00	-0.77	0.18		- 0.0
0.50	Wind Speed -	0.04	-0.46	-0.74	-0.77	1.00	-0.44		0.2 0.4
0.75	Specific Humidity -	0.71	0.14	0.48	0.18	-0.44	1.00		0.6
		Temperature -	Zonal Wind -	Relative Humidity -	Meridional Wind -	Wind Speed -	Specific Humidity -		-

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- **Rising Temperature**: A clear and statistically significant positive trend in temperature is observed across all altitude levels in the Mediterranean region, with linear slopes indicating a persistent warming since 1950.
- **Evaporation Stability Despite Warming**: Despite the ongoing temperature increase, evaporation and potential evaporation remain nearly stable. The weak correlation with temperature (r = 0.05) suggests that temperature alone does not drive evaporation.
- Wind Speed as a Key Driver: Wind speed exhibits a stronger correlation with evaporation (r = 0.44). Wind speed itself remains relatively stable, with a slight negative trend.
- Importance of Interconnected Climate Drivers: The results highlight that evaporation is influenced by a combination of factors especially wind speed and not just temperature alone. Understanding these interdependencies is crucial for accurate climate analysis.

References



[1] Hersbach, H.; Bell, B.; Berrisford, P.; Biavati, G.; Horányi, A.; Muñoz-Sabater, J.; Nicolas, J.; Peubey, C.; Radu, R.; Rozum, I.; et al. ERA5 Hourly Data on Single Levels from 1979 to Present; Copernicus Climate Change Service (C3S) and Climate Data Store (CDS): Bonn, Germany, 2018.

[2] Koutsoyiannis, D., Hurst-Kolmogorov dynamics and uncertainty, Workshop on Nonstationarity, Hydrologic Frequency Analysis, and Water Management, Boulder, Colorado, USA, doi:10.13140/RG.2.2.36060.39045, International Center for Integrated Water Resources Management, US Army Corps of Engineers, United States Geological Survey, US Department of the Interior - Bureau of Reclamation, National Oceanic and Atmospheric Administration, US Environmental Protection Agency, Colorado State University, 2010.
[3] Glynis, K., T. Iliopoulou, P. Dimitriadis, and D. Koutsoyiannis, Stochastic investigation of daily air temperature extremes from a global ground station network, Stochastic Environmental Research & Risk Assessment, doi:10.1007/s00477-021-02002-3, 2021.
[4] Iliopoulou, T., & Koutsoyiannis, D. (n.d.). *Have rainfall patterns changed? A global analysis of long-term rainfall records and re-analysis*

data.

[5] Dimitriadis, P., A. Tegos, and D. Koutsoyiannis, Stochastic analysis of hourly to monthly potential evapotranspiration with a focus on the long-range dependence and application with reanalysis and ground-station data, Hydrology, 8 (4), 177, doi:10.3390/hydrology8040177, 2021.

[6] Dimitriadis, P.; Koutsoyiannis, D.; Iliopoulou, T.; Papanicolaou, P. A Global-Scale Investiga-tion of Stochastic Similarities in Marginal Distribution and Dependence Structure of Key Hydrological-Cycle Processes. Hydrology 2021, 8, 59.

[7] Koutsoyiannis, D. Revisiting the global hydrological cycle: Is it intensifying? Hydrol. Earth Syst. Sci. 2020, 24, 3899–3932.

[8] Koskinas, A., E. Zacharopoulou, G. Pouliasis, I. Deligiannis, P. Dimitriadis, T. Iliopoulou, N. Mamassis, and D. Koutsoyiannis, Estimating the Statistical Significance of Cross–Correlations between Hydroclimatic Processes in the Presence of Long–Range Dependence, Earth, 3 (3), 1027-1041, doi:10.3390/earth3030059, 2022.

[9] Ma, N., Zhang, Y., & Yang, Y. (2025). Recent decline in global ocean evaporation due to wind stilling. *Geophysical Research Letters*, 52, e2024GL114256. <u>https://doi.org/10.1029/2024GL114256</u>.

References



[10] Koutsoyiannis, D. (2018). The climacogram: A tool for stochastic analysis and modeling of hydroclimatic time series. Department of Water Resources and Environmental Engineering, National Technical University of Athens.

https://www.itia.ntua.gr/el/getfile/1835/3/documents/KClimacogram_.pdf

[11] Dimitriadis, P. and D. Koutsoyiannis, Climacogram versus autocovariance and power spectrum in stochastic modelling for Markovian and Hurst–Kolmogorov processes, Stochastic Environmental Research & Risk Assessment, 29 (6), 1649–1669, doi:10.1007/s00477-015-1023-7, 2015.

[12] Koutsoyiannis, D. (2023). *Stochastics of hydroclimatic extremes: A cool look at risk* (3η έκδ.). Kallipos Open Academic Editions. <u>https://www.itia.ntua.gr/en/docinfo/2000/</u>

[13] Iliopoulou, T., and D. Koutsoyiannis, Projecting the future of rainfall extremes: better classic than trendy, Journal of Hydrology, 588, doi:10.1016/j.jhydrol.2020.125005, 2020.

[14] Lindzen, R. S. (2022). *An assessment of the conventional global warming narrative* (Technical Paper No. 5). The Global Warming Policy Foundation.

[15] American Meteorological Society (n.d.) *Welcome*. Available at: <u>https://glossary.ametsoc.org/wiki/Welcome</u> (Accessed: January 2025).

[16] NTUA (n.d.) *Μεθοδολογία Υπολογισμού Υδατικού Ισοζυγίου – Έκθεση*. Available at: <u>https://www.itia.ntua.gr/el/docinfo/116/</u> (Accessed: January 2025).

[17] Koutsoyiannis, D. et al., 2020. Climate of the past and present and its hydrological relevance. [pdf] Athens: NTUA. Available at: https://www.itia.ntua.gr/el/getfile/2065/1/documents/ClimateHydrologyMoscow.pdf (Accessed: December 2024).

[18] Koutsoyiannis, D. (2012). *Hydrometeorology: Evaporation and transpiration (2012 edition with adjustments)*. Department of Water Resources, National Technical University of Athens.