

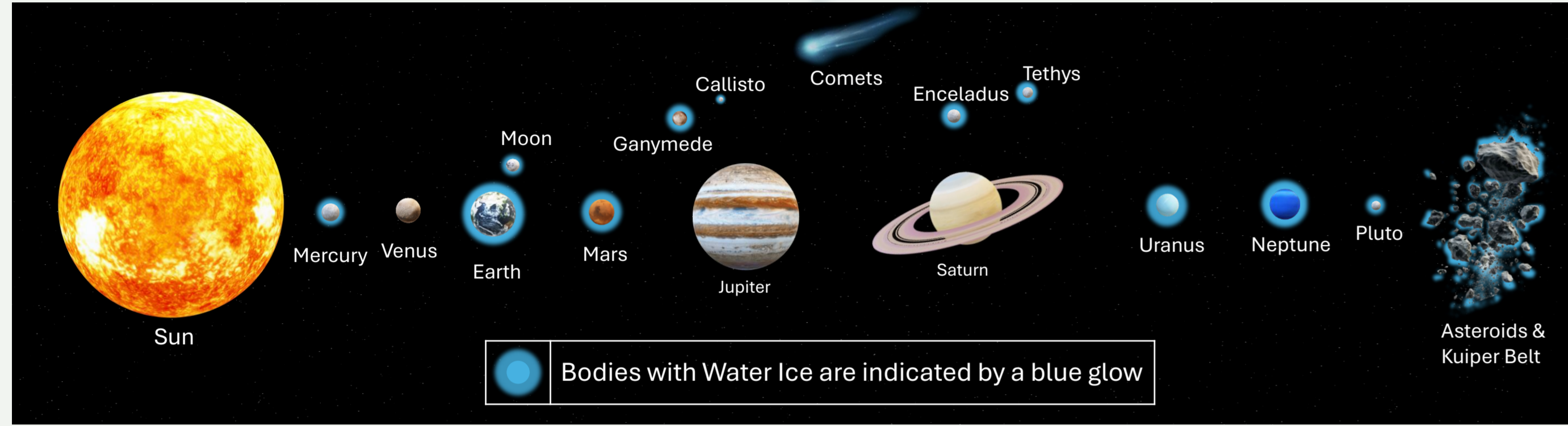
Water Ice Sublimation and O-H Isotopic Fractionation in Terrestrial and Extraterrestrial Environments: New insights gained from numerical modelling and laboratory experiments

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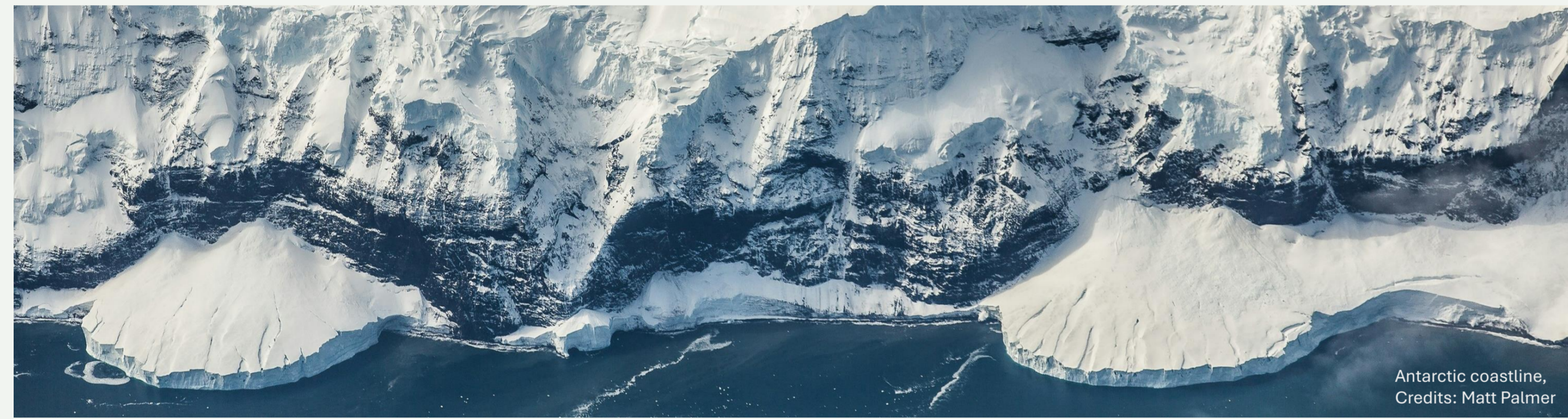
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1. Introduction

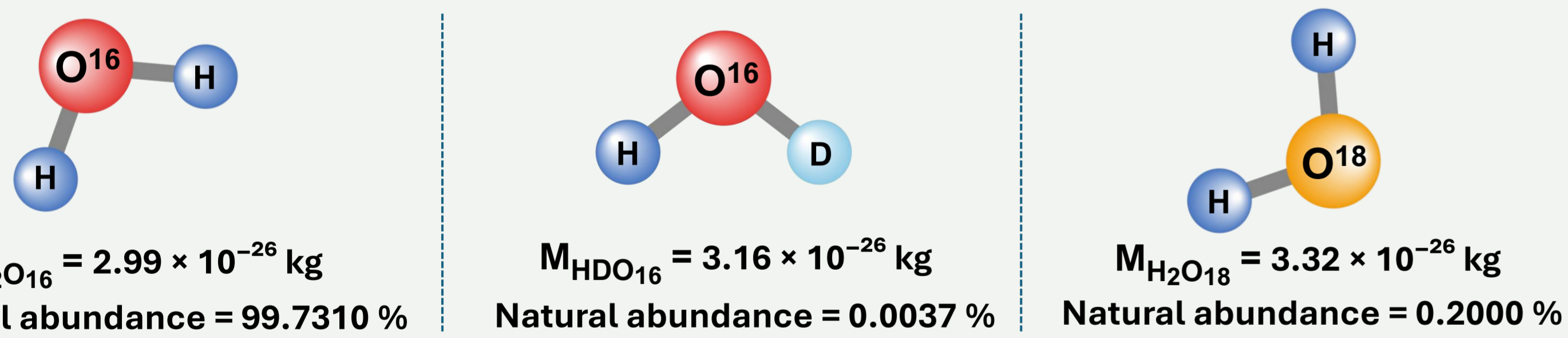
- Tremendously large reservoirs of **extraterrestrial water ice** exist on previously supposed bone dry solar and extrasolar bodies serving as quintessential **resources** for space exploration.



- Terrestrial water ice** (glaciers, permafrost, snow etc.) store valuable **climatological records** on timescales of a few days up to over hundreds of thousands of years.



- Stable isotopes** of oxygen and hydrogen (O-H) are powerful **tracers** of natural phenomena widely used across multiple disciplines in Earth and Planetary sciences.



2. Motivation

- Few experimental studies** investigated **water ice sublimation rates** and **O-H isotopic fractionation** under extreme conditions like the **lunar environment**.
- Many unknowns and uncertainties** in sublimation - fast diffusion - condensation of water ice and related O-H isotopic fractionation under extreme low temperatures and vacuum.
- Experimental and numerical modelling approaches will help **guide space technologies** for water ice extraction and **inform climate science** in past climate reconstructions.

3. Experiments

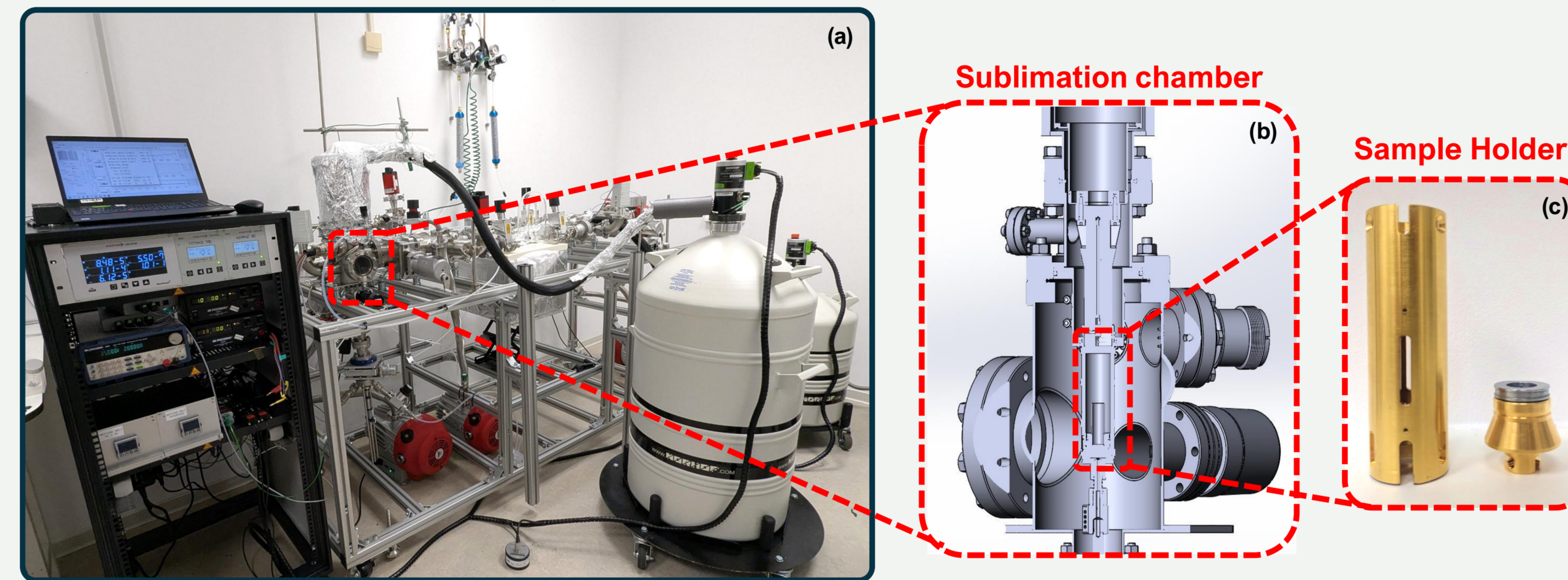
LUNAQUA is a laboratory scale prototype developed at LIST, capable of producing ultra high vacuum ($UHV \sim 10^{-7} \text{ Pa}$) and cryogenic temperatures ($\sim 110 \text{ K}$), fig. (a).

Step 1: Water ice samples of different isotopic compositions ($\delta^{18}\text{O}$ and $\delta^2\text{H}$) are introduced in the prototype's sublimation chamber [fig. (b)] within a sample holder [fig. (c)] and sublimated ($T_{\text{sub}} \sim 195 \text{ K}$).

Step 2: A cold trap on the other end at $\sim 90\text{K}$ condenses the sublimated water vapor.

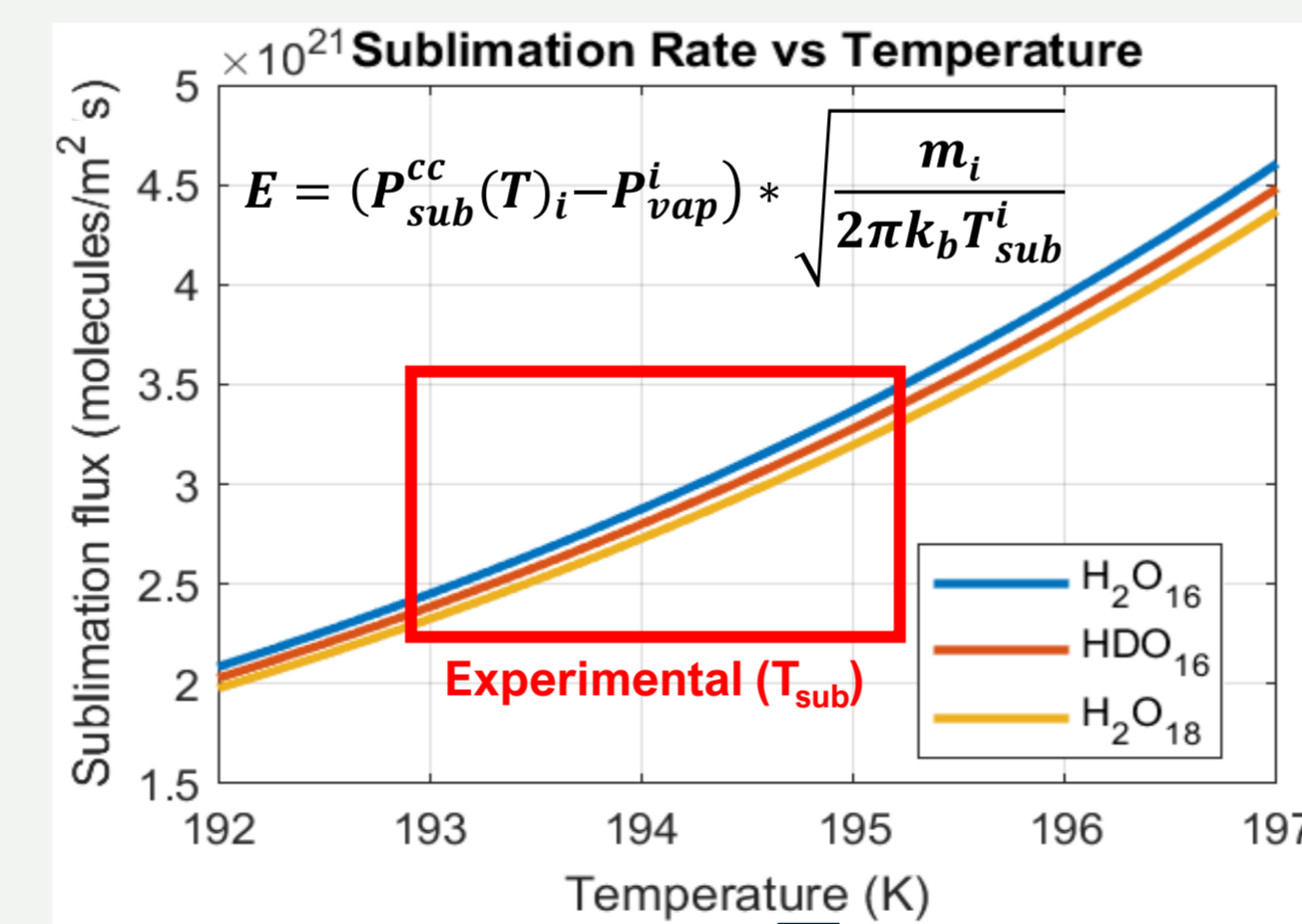
Step 3: Fully and sequentially condensed water is analysed for $\delta^{18}\text{O}$ and $\delta^2\text{H}$ values.

$$\delta^A X = \frac{R_{\text{Sample}}}{R_{\text{Standard}}} - 1 (\times 10^3 \text{‰}) \quad R = \frac{\text{Abundance of rare isotope}}{\text{Abundance of abundant isotope}} \quad \alpha_{\text{sub}} = \frac{R_{\text{vapor}}}{R_{\text{ice}}}$$

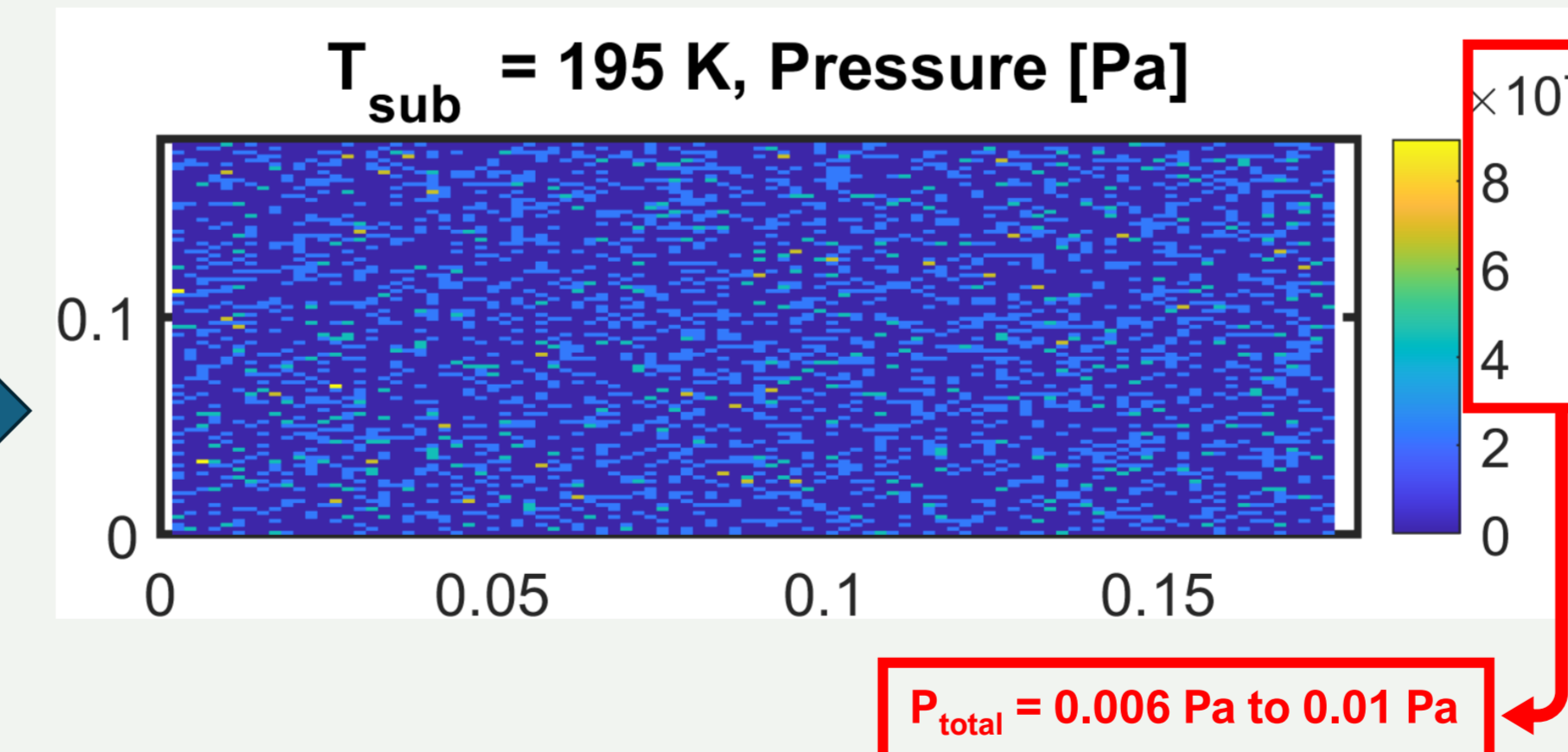


4. Numerical Modelling

Step 1: Sublimation Rate (Hertz-Knudsen Equation)



Step 3: Particle and Pressure fields

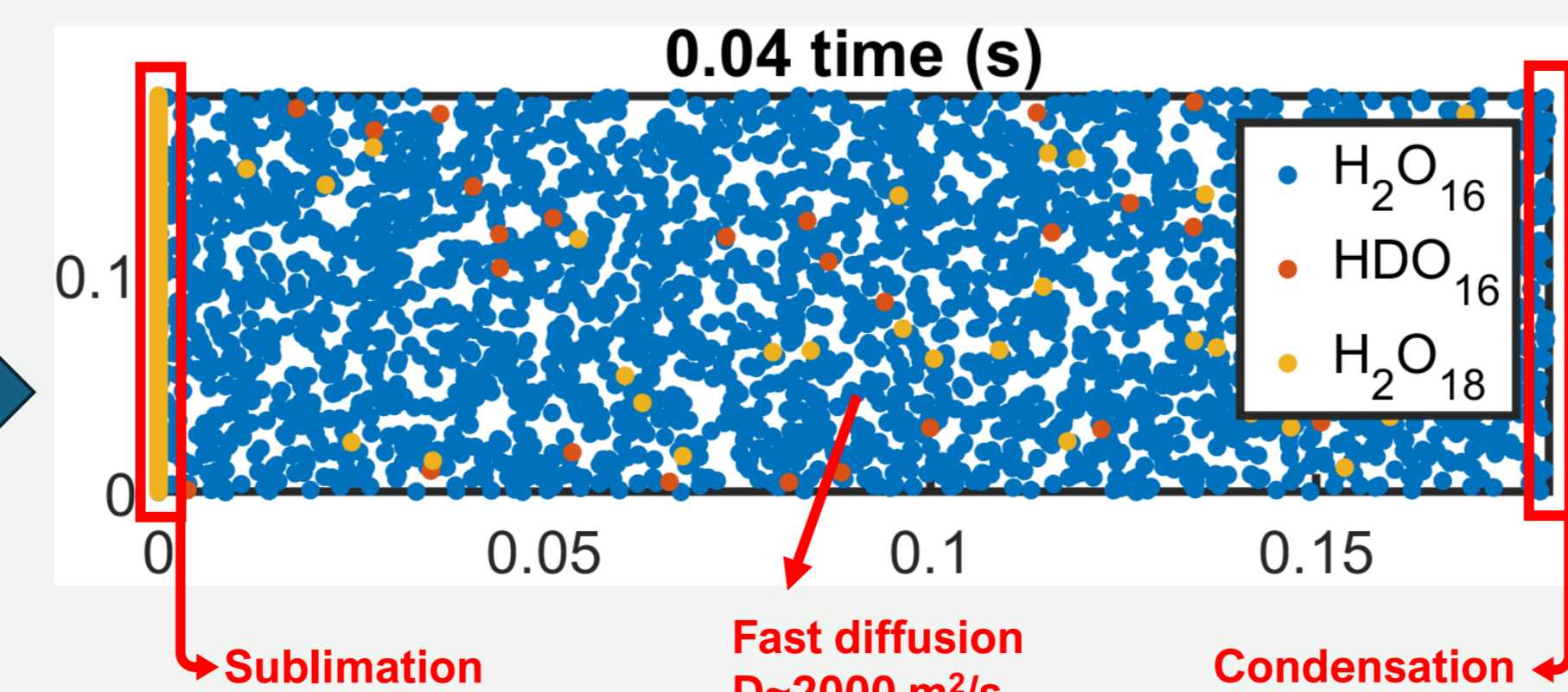


Step 2: Lagrangian Tracking: (Langevin's equation)

$$x_{\text{new}} = x_{\text{old}} + v_{\text{Boltz}} * \Delta t + v_{\text{darcy}} * \Delta t$$

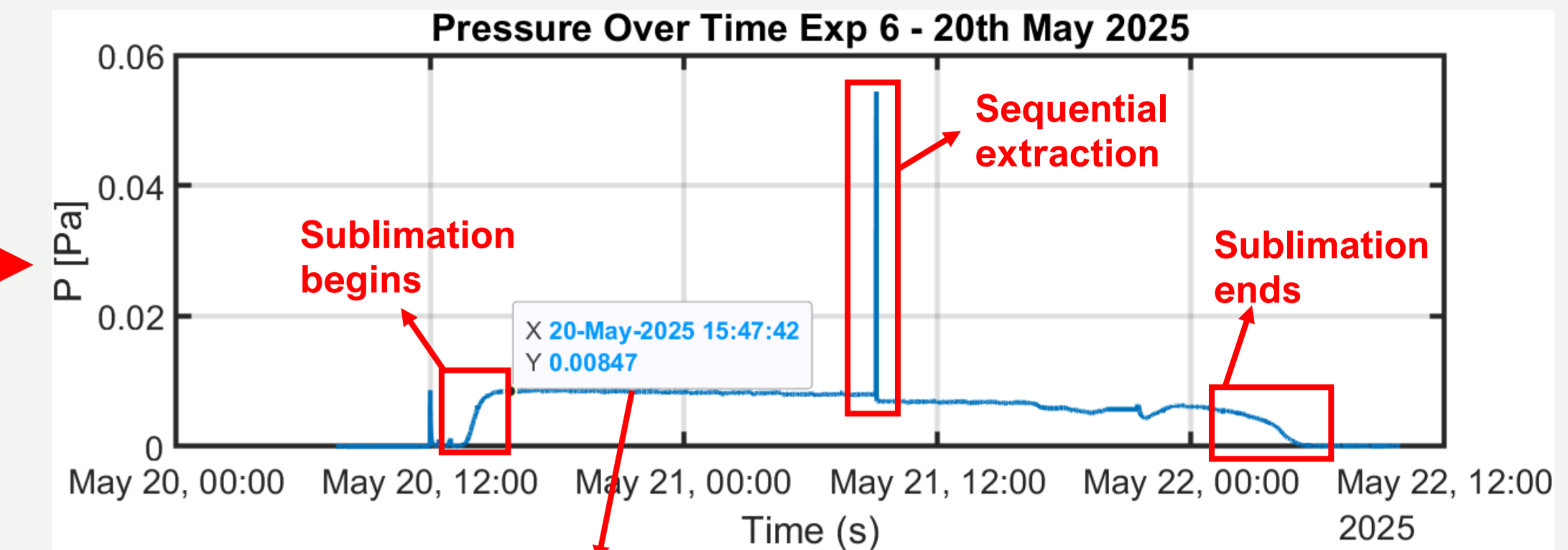
(Maxwell-Boltzmann's speed distribution)

$$f(v_{\text{Boltz}}) * dV = \frac{4}{\sqrt{\pi}} \left(\frac{m_i}{k_b T} \right)^{\frac{3}{2}} v^2 * dV * e^{-\frac{mv^2}{2k_b T}}$$

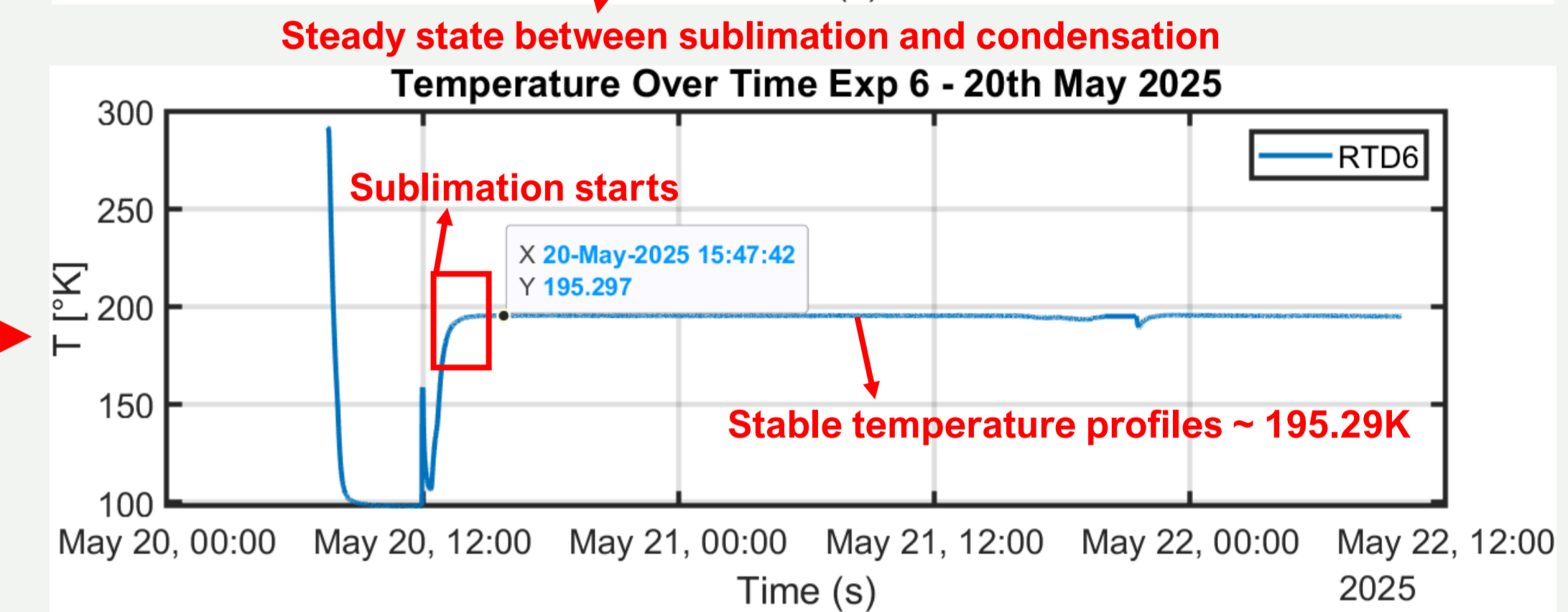


5. Results

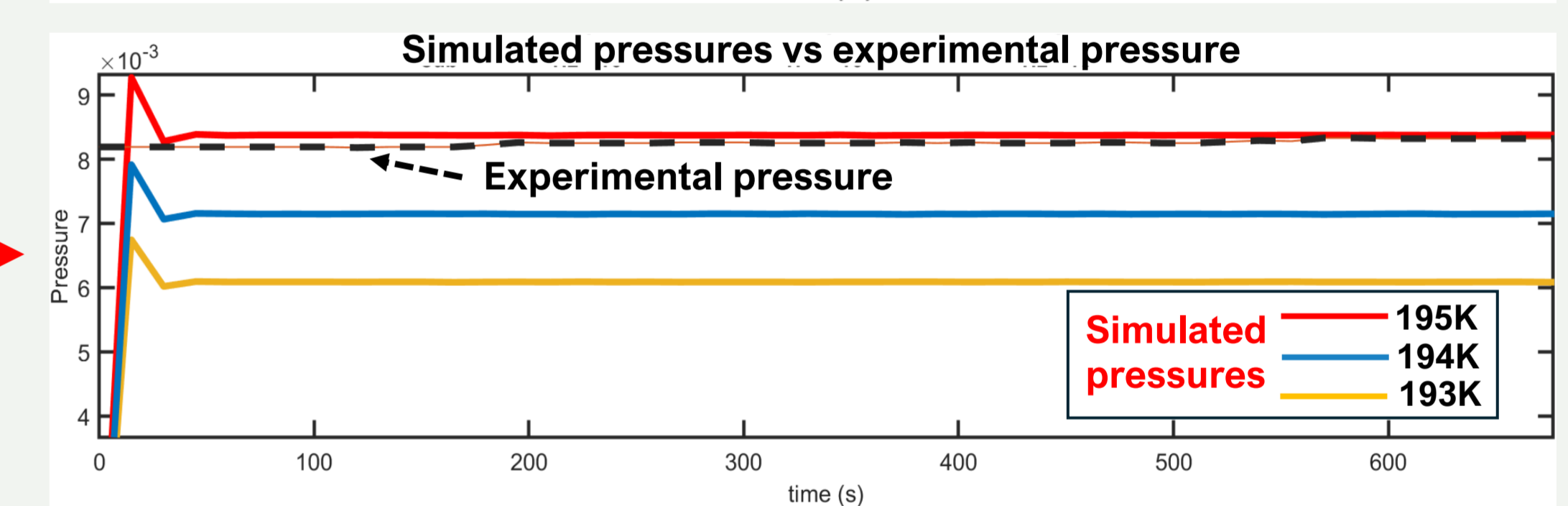
Experimental Pressure
 $\sim 0.008 \text{ Pa}$ at 195.29 K



Experimental Temperature
 $\sim 195.29 \text{ K}$



Simulation Pressure
 $\sim 0.008 \text{ Pa}$ at 195 K



Experimental Isotopic Fractionation

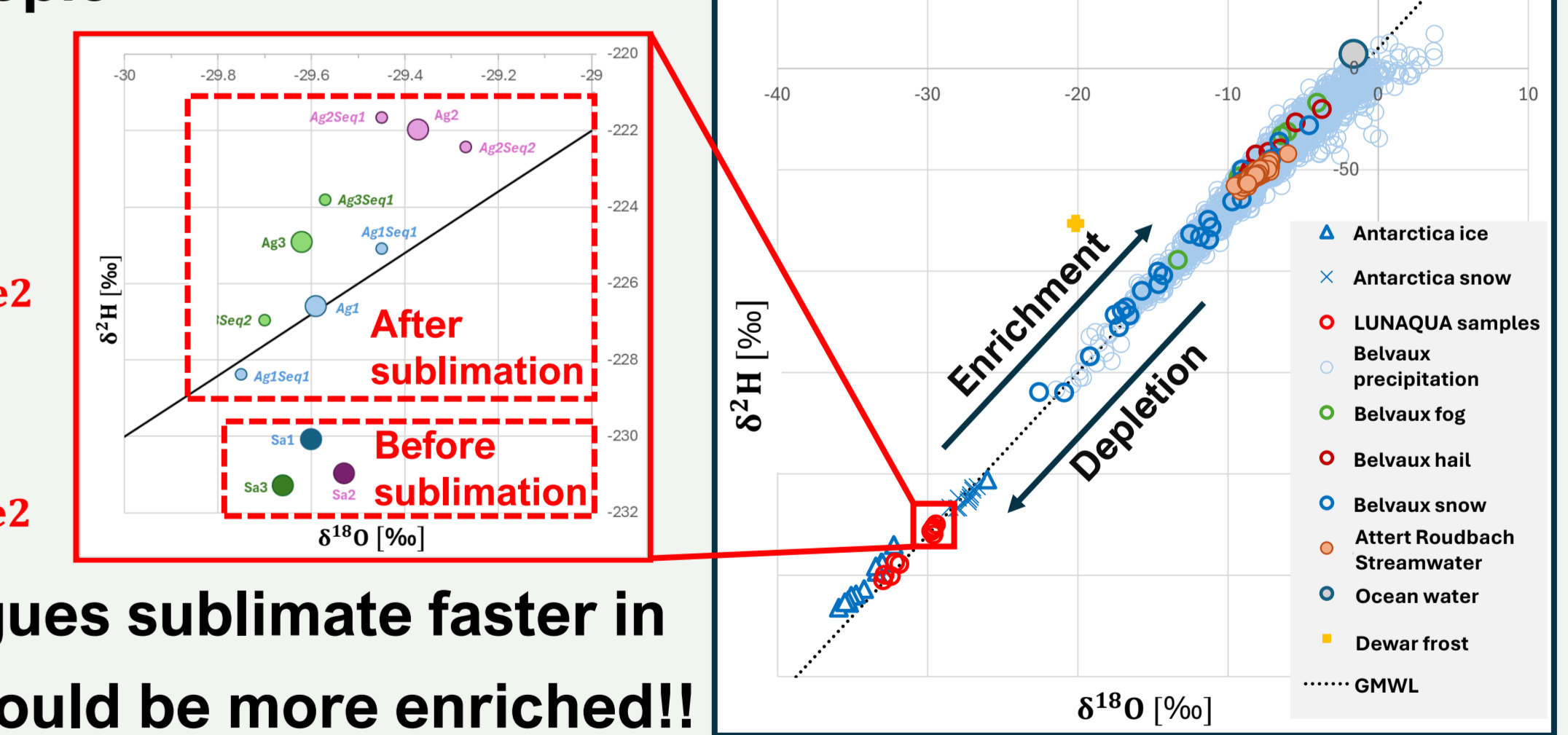
2 observations:

$$\delta_{\text{Sequence1}} > \delta_{\text{Sequence2}}$$

1 observation:

$$\delta_{\text{Sequence1}} < \delta_{\text{Sequence2}}$$

Lighter isotopologues sublimate faster in 1st half; 2nd half should be more enriched!!



6. Take Aways & Outlook

- Sublimation rates and vapor pressures are highly **sensitive** to T_{sub} .
- Simulation and experimental **pressure matching** is fundamental to finally simulate isotopic fractionation (**Kinetic Gas Theory**).
- Further experiments will collect more **isotopic fractionation evidence** to establish certainty; but fractionation can be observed.
- Further simulations will be to find the molecular origins of **isotopic fractionation factor** (α_{sub}).