

Future Trends in Upper-Atmospheric Shear Instability from Climate Change

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Why This Matters

Aviation Impact

- >70% of weather-related accidents
- Injuries & financial costs
- Flight routing & efficiency
- Certification standards (**outdated**)

Climate Connection

- Changes in jet stream structure
- Changes in meridional temperature gradient
- Increased vertical wind shear

Clear-air turbulence (CAT) is invisible, dangerous, and projected to worsen



Background

Earth and Space Science

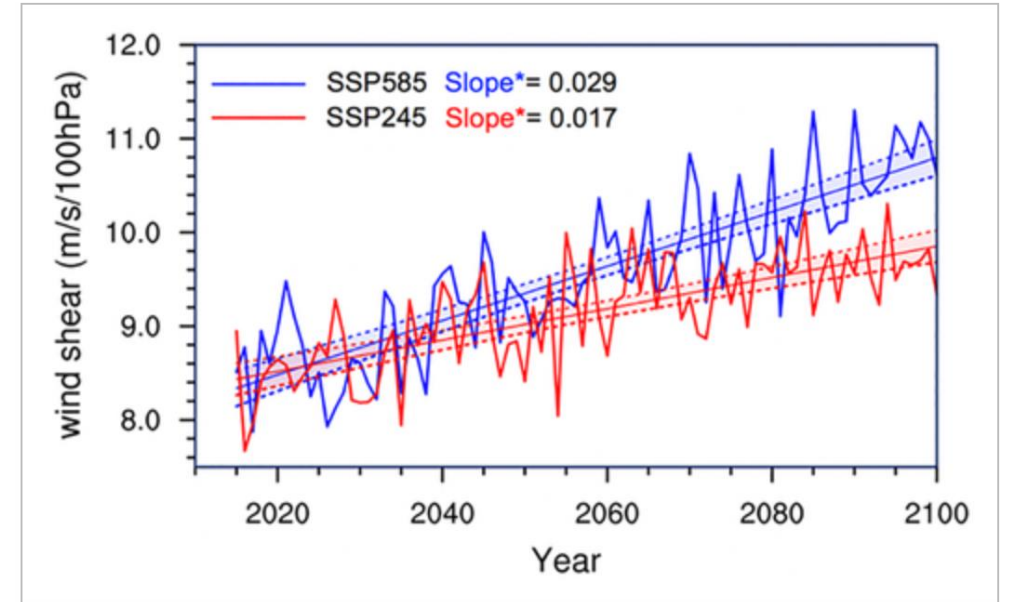
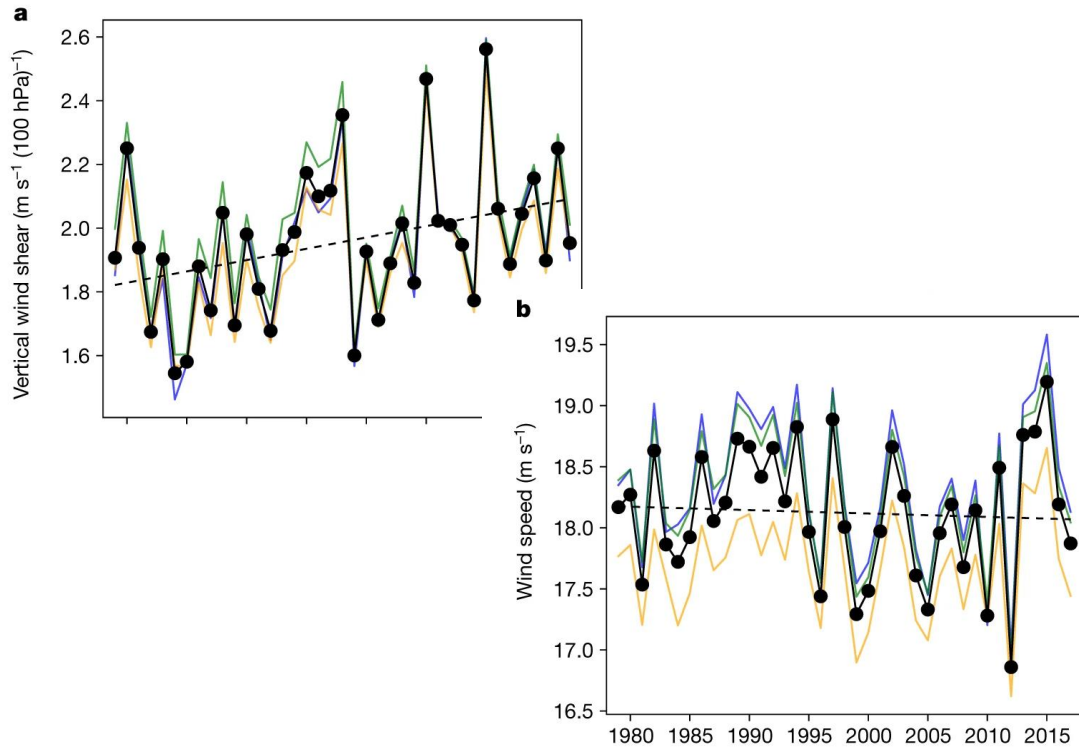
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Increased Turbulence in the Eurasian Upper-Level Jet Stream in Winter: Past and Future

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— ERA-Interim — NCEP/NCAR — JRA-55 ● Mean -- Mean t



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Letter | Published: 07 August 2019

Increased shear in the North Atlantic upper-level jet stream over the past four decades

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[Nature](#) 572, 639–642 (2019) | [Cite this article](#)

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Future Trends in Upper-Atmospheric Shear Instability from Climate Change



Data: 26 CMIP6 models, SSP2-4.5 & SSP5-8.5, 2015 – 2100, Global & Regional Analyses

Vertical uShear ↑

+16-34% at 250 hPa

0.04-0.11 m/s per 100 hPa
decade⁻¹

Stability (N^2) ↓

-10% to -20% at 250 hPa

Reduced buoyancy resistance

Richardson (Ri) ↓

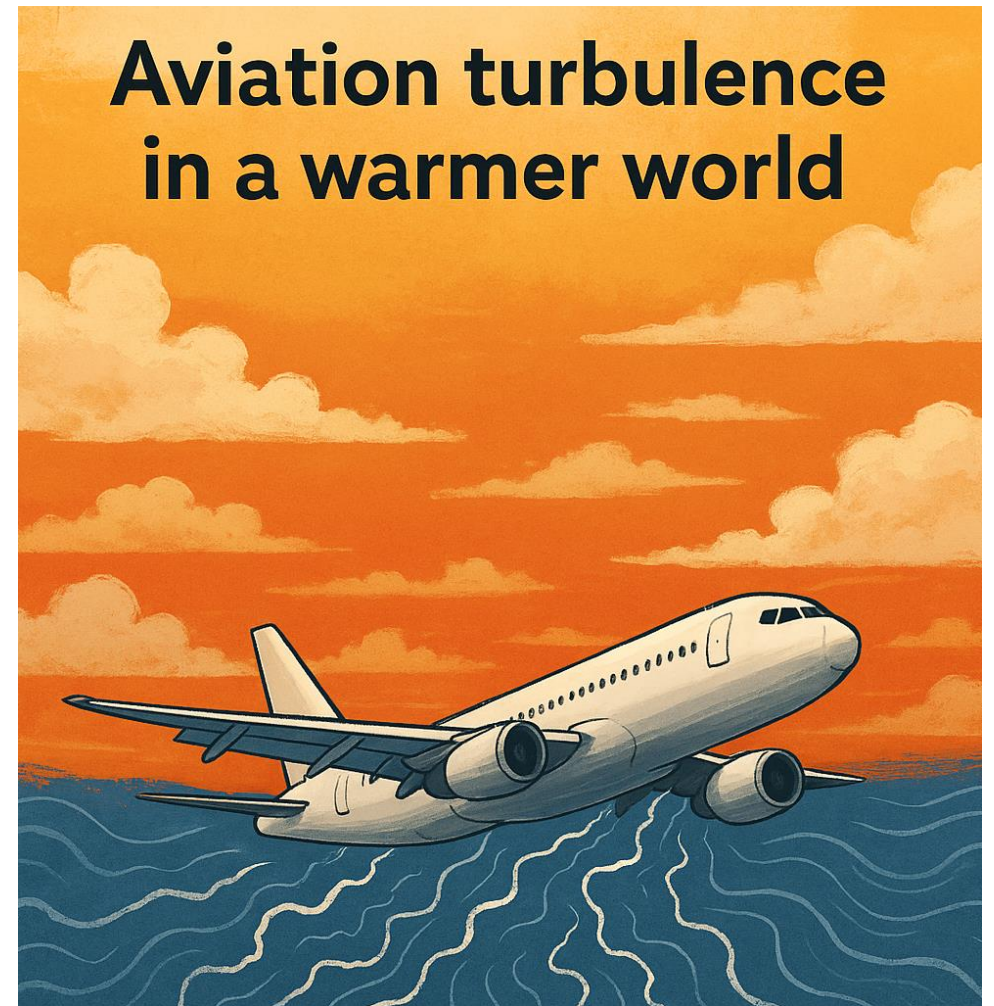
-38% to -47% at 250hPa

More turbulence-prone

$$Ri = N^2 / \text{Shear}^2$$

→ Atmosphere becoming more favorable for CAT generation

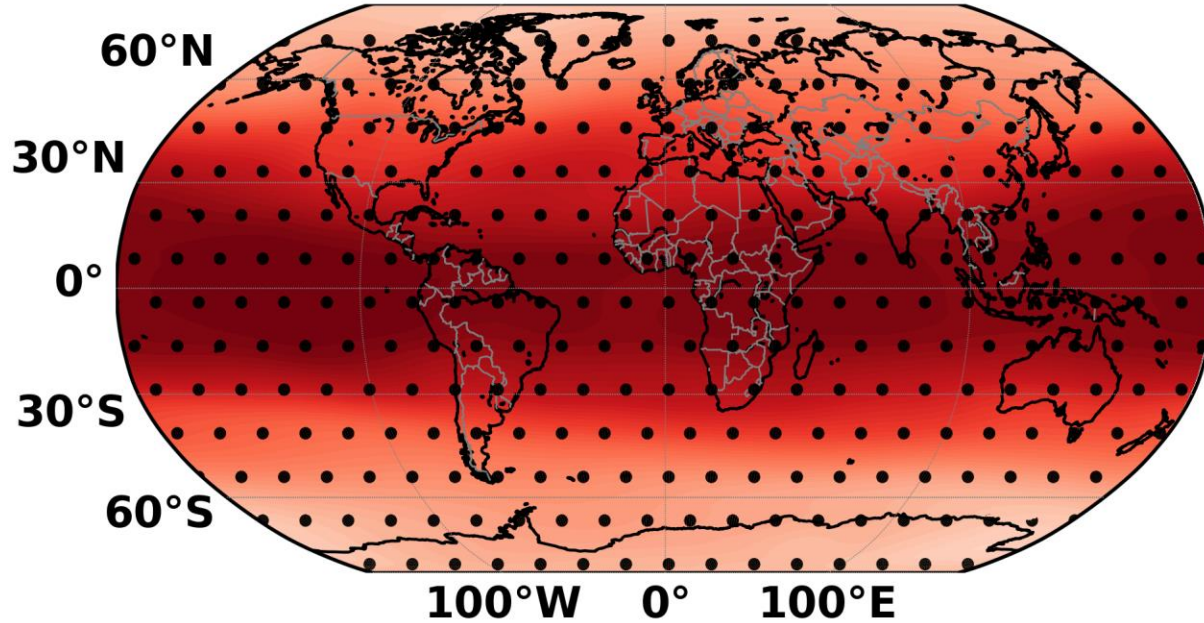
- **Results**
 - a) **Temperature @250 hPa**
 - b) **Thermal Gradient Wind Balance @250 hPa**
 - c) **Vertical Shear of Zonal Wind**
 - d) **Meridional Temperature Gradient**
 - e) **Atmospheric Stability (N^2) and Ri @250 hPa**
 - f) **Time series and Trends of uShear, N^2 and Ri @250 hPa**
- **Summary & Discussion**



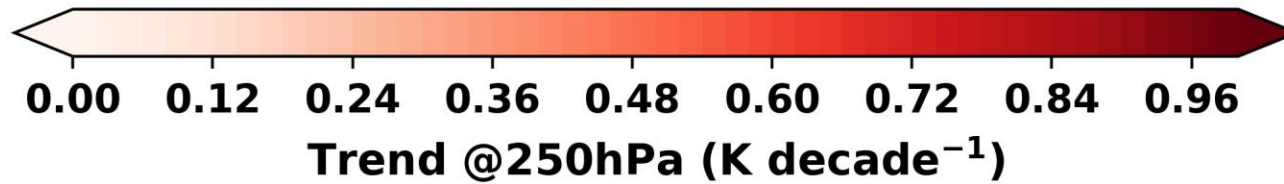
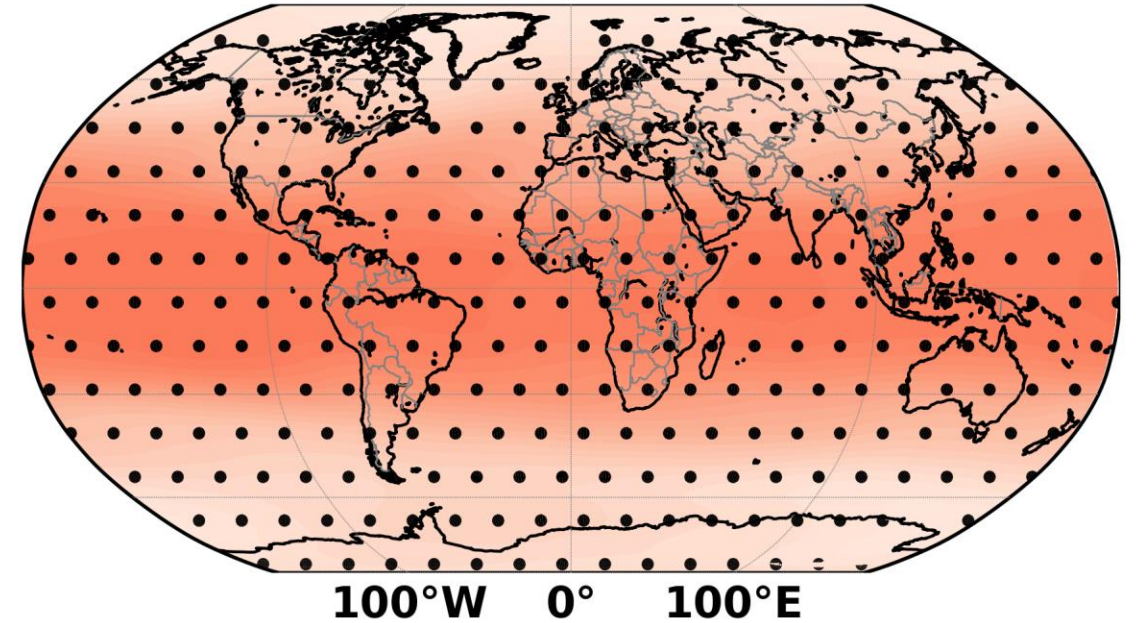
a. Temperature @250 hPa

Ensemble mean of temperature trends @250 hPa

temperature
SSP 585



temperature
SSP 245



Warmer conditions (**stronger** trends) near the equator and progressively **cooler** temperatures (**weaker** trends) toward the poles,

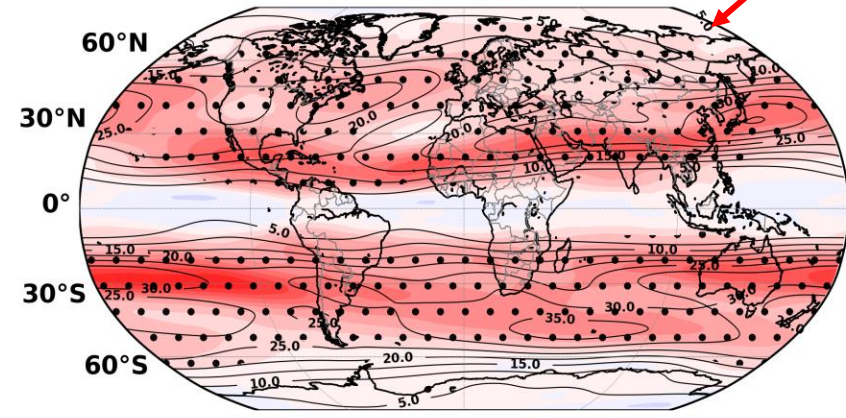
$$\frac{U_{gr}^2}{R} + fU_{gr} = -g \frac{\partial Z}{\partial Y}$$

$$\Rightarrow \frac{d}{dp} \left(\frac{U_{gr}^2}{R} + fU_{gr} \right) = \frac{R_d}{p} \frac{\partial T}{\partial Y} \iff$$

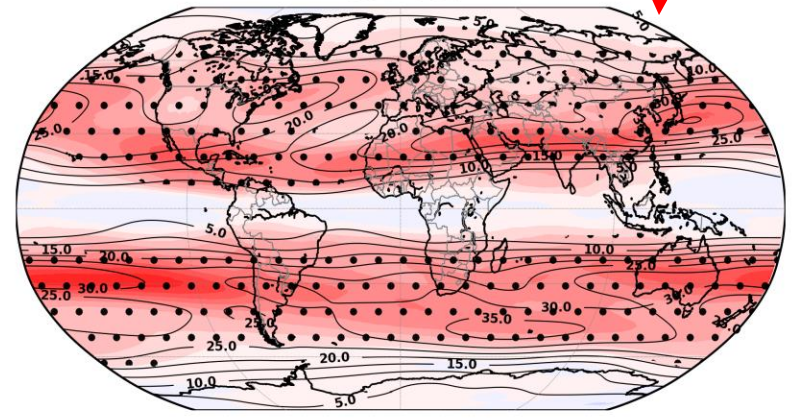
$$\frac{dU_{gr}}{dp} = \left(\frac{R_d}{fp} \frac{\partial T}{\partial Y} + \frac{U_{gr}^2}{fR^2} \frac{\partial R}{\partial p} \right) \left(1 + \frac{2U_{gr}}{R} \right)$$

(Knox, 1997)

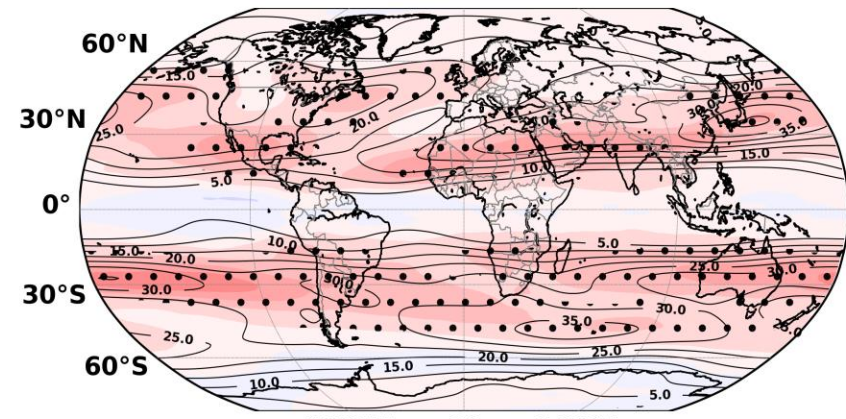
vertical grad wind shear
SSP 585



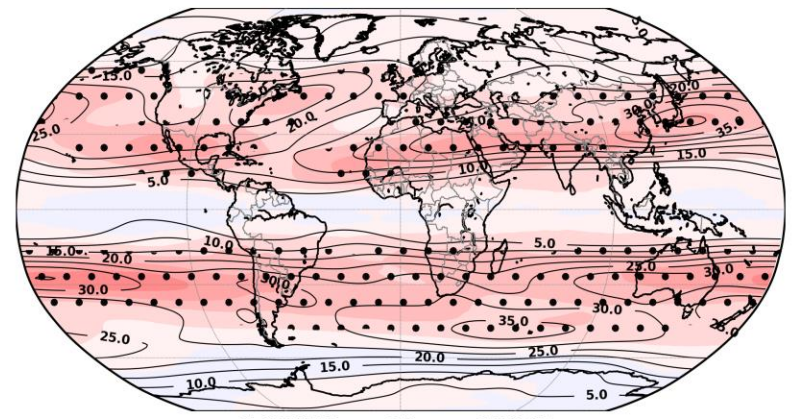
zonal thermal grad wind balance
SSP 585



vertical grad wind shear
SSP 245

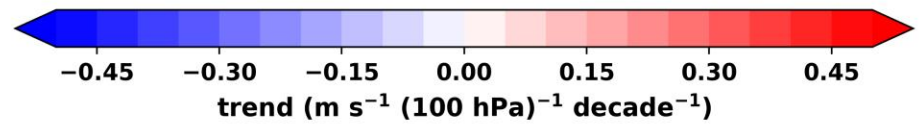


zonal thermal grad wind balance
SSP 245



Vertical Wind Gradient Shear @250 hPa

Changes driven by **intensified upper troposphere temperature gradients** that **enhance vertical wind shear** near subtropical jets; Stronger in SSP 585



d. Thermal Gradient Wind Balance @250 hPa

b. Vertical Shear of Zonal Wind

Ens mean of yearly mean trends

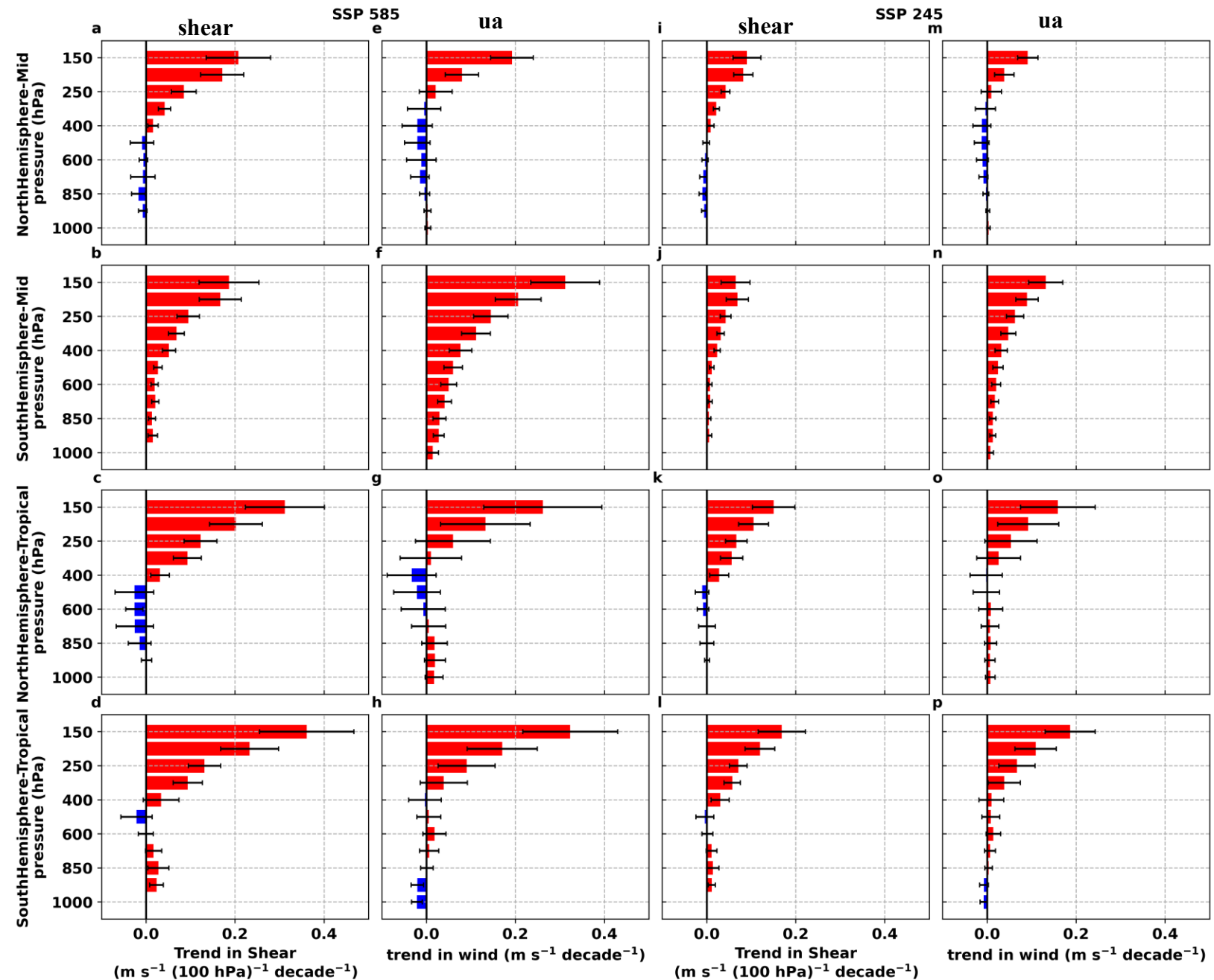
$$\text{Shear}_i = \frac{V_{i-1} - V_{i+1}}{z_{i-1} - z_{i+1}}$$

- NH Mid-Latitudes:

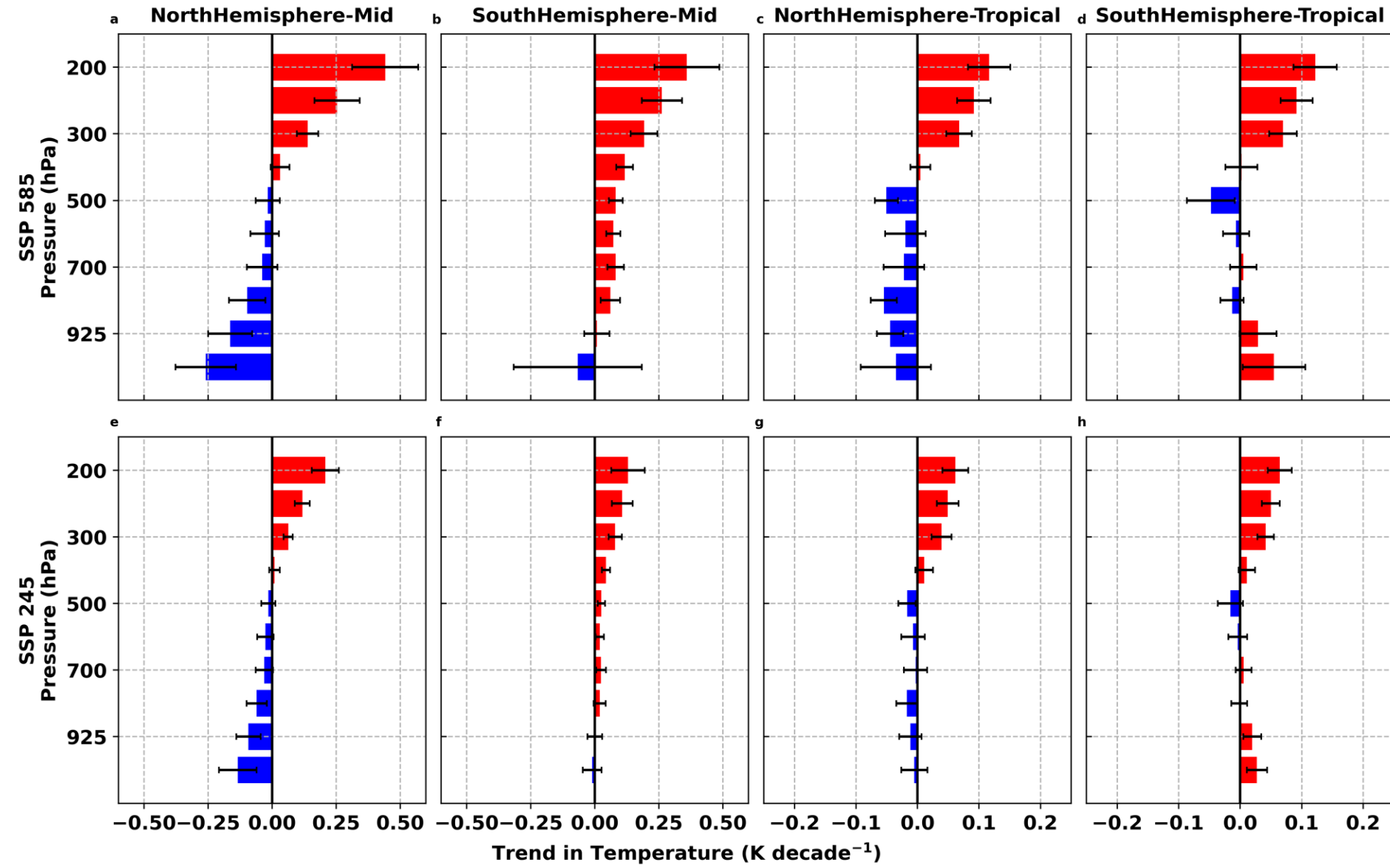
- Negative/Weak trends in wind speed up to ~350 hPa.
- Increase near 250 hPa → stronger upper-level jets (Shaw & Miyawaki, 2024).
- Vertical shear rises significantly above 250 hPa.

- SH Mid-Latitudes:

- Continuous increase in wind speed & shear across all levels.



e. Meridional Temperature Gradient



Hemispheric asymmetry reflects contrasting meridional temp. gradients:

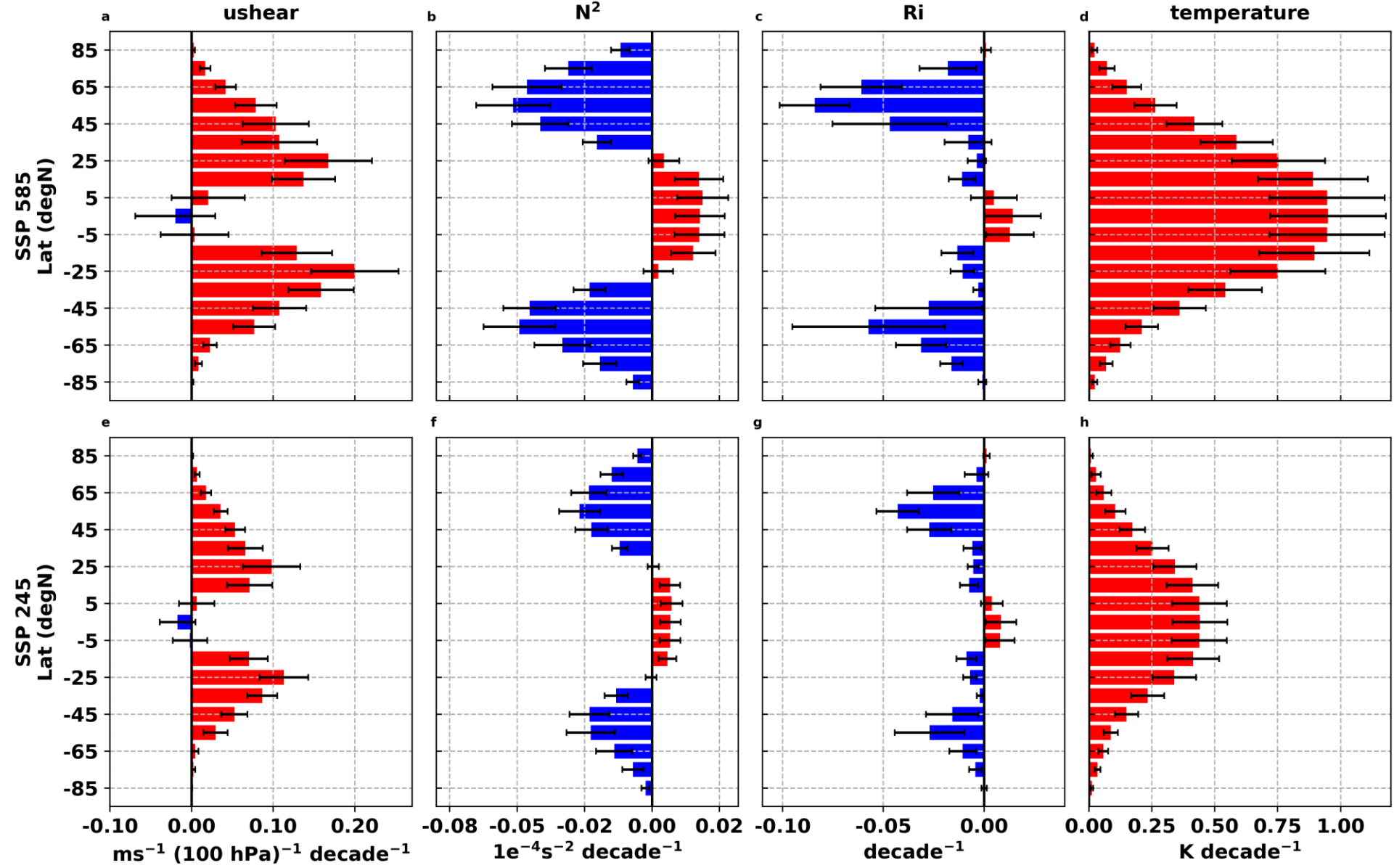
- **NH mid-lats:** Arctic amplification weakens low-level gradients but enhances upper-trop wind shear.
- **SH mid-lats:** Milder gradients, but widespread warming supports positive wind trends at both 200/300 hPa

Aligned with prior findings (Smith et al., 2019; Caporale et al., 2024).

- Mid latitudes: polar box [50 – 70 °N/S, -180 - 180 °E] – subtropical box [30 – 50 °N/S, -180 - 180 °E]
- Tropical latitudes: upper tropical box [15 – 30 °N/S, -180 - 180 °E] - lower tropical box [0 – 15 °N/S, -180 - 180 °E]

f. Atmospheric Stability (N^2) and Ri @250 hPa

$$N^2 = \frac{g}{\theta} \frac{d\theta}{dz} \quad Ri = \frac{N^2}{shear^2}$$

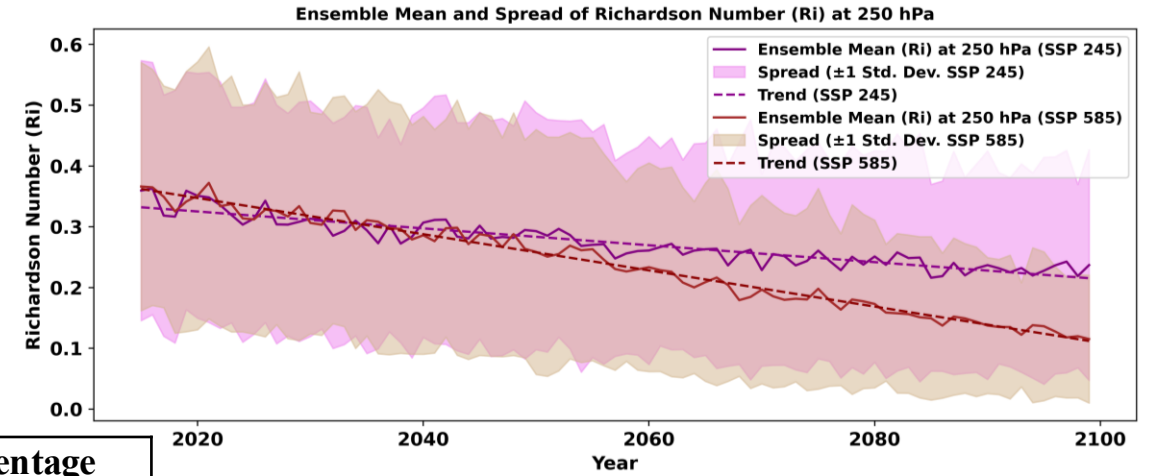
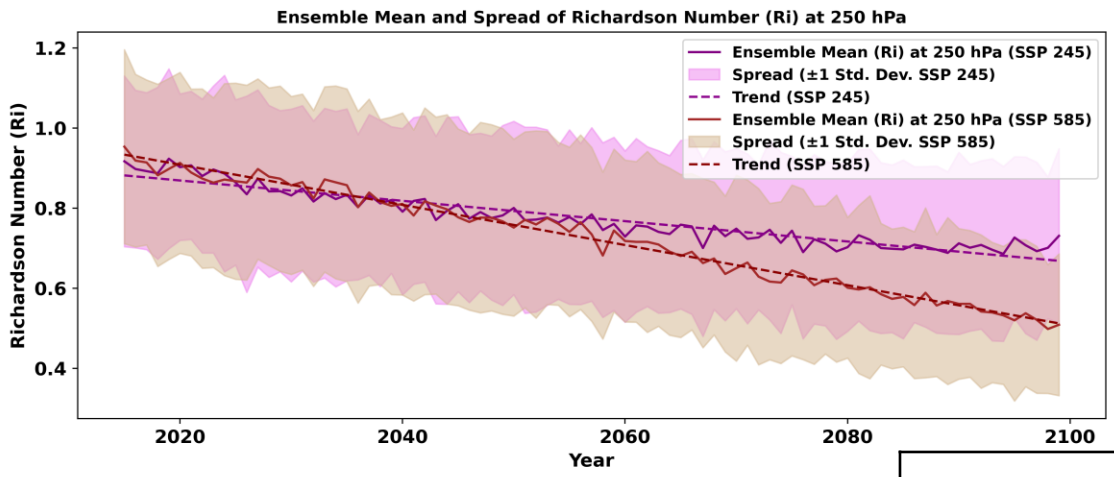
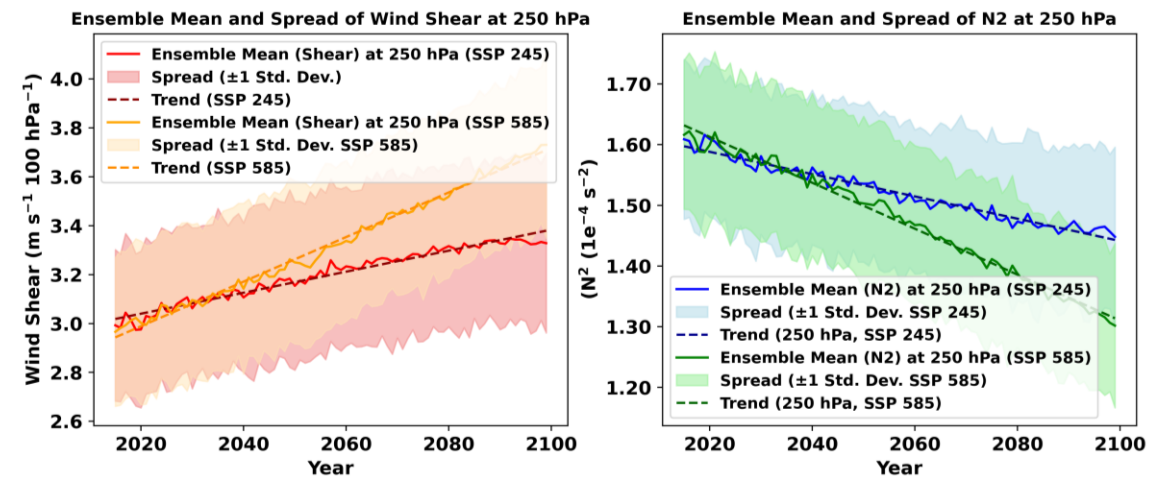
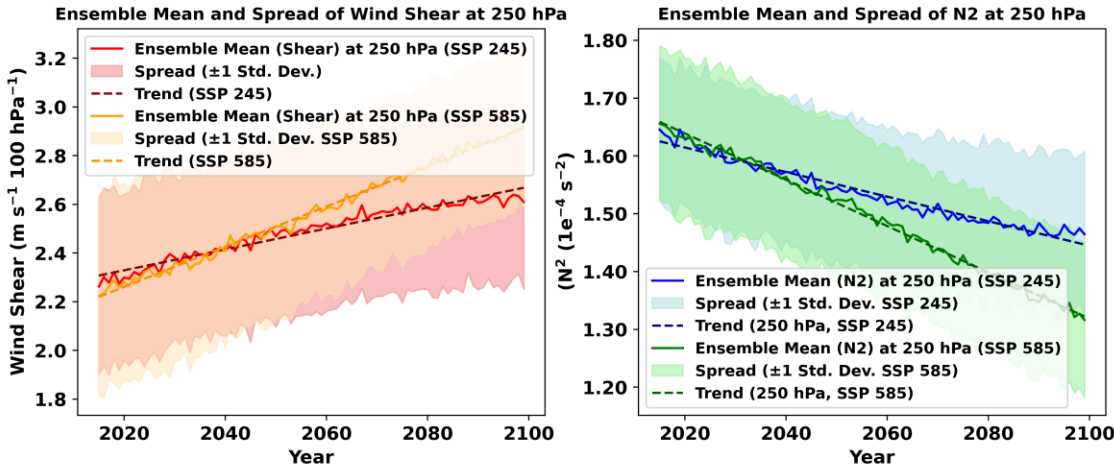


Low Ri ($< 1/4$) indicates atmospheric instability, promoting turbulence such as Clear Air Turbulence (CAT)
(Sharman & Pearson, 2017)

N^2 trend :
 ↑ in tropics: upper-trop warming \gg sfc warming \Rightarrow increasing $d\theta/dz$
 ↓ at 50–60°: near sfc warming \gg aloft ($< d\theta/dz$), jet shift, tropopause rise
(Held & Soden 2006; Lu et al. 2008; Shaw et al. 2016)

NorthHemisphere-Mid

SouthHemisphere-Mid



g. Time series and Trends of ushear, N² and Ri @250 hPa

Variable	Percentage Change (%)	
	NHM	SHM
Shear (SSP 245)	15	12
N2 (SSP 245)	-10	-9
Ri (SSP 245)	-24	-35
Shear (SSP 585)	31	26
N2 (SSP 585)	-20	-19
Ri (SSP 585)	-45	-69

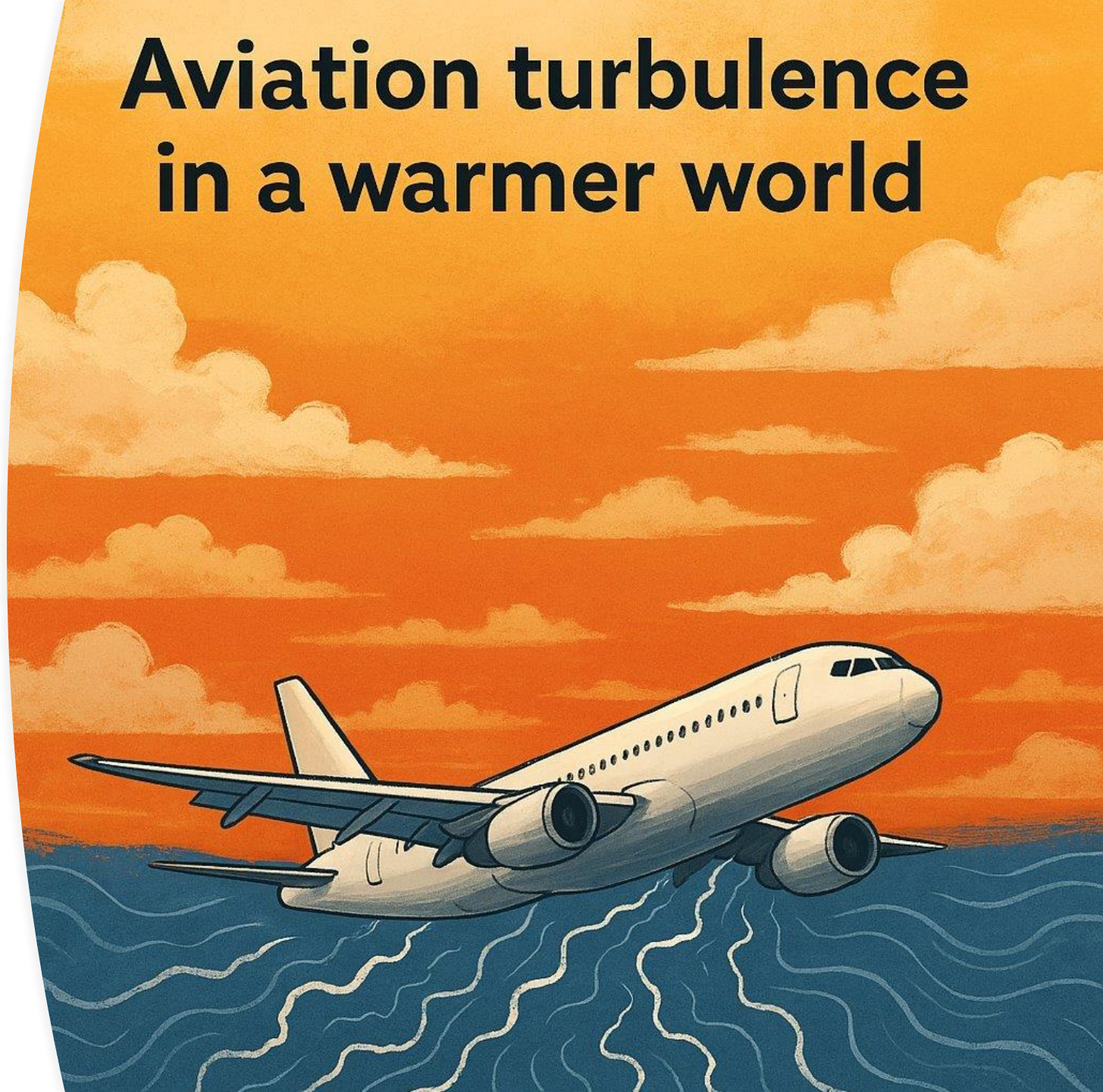
These changes are indicative of large-scale atmospheric conditions becoming increasingly favorable for turbulence.

Aviation turbulence in a warmer world

Summary & Discussion



- **Rising vertical wind shear** and **decreasing static stability (N^2)** suggest a **greater risk of clear-air turbulence (CAT)** under future climate scenarios.
- These changes, especially near **upper-level jet streams**, create more favorable conditions for turbulence (Sharman et al., 2014; Lee et al., 2023).
- **Smith et al. (2023)** confirm increasing CAT over the North Atlantic (1950–2050, multi-model study).
- While **CMIP6 resolution** limits turbulence representation, trends in **Ri , N^2 , and shear** point to real aviation and climate risks.
- **Future work** needs **high-resolution models** or **downscaling** to resolve fine-scale turbulence processes and better assess climate impacts.



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