

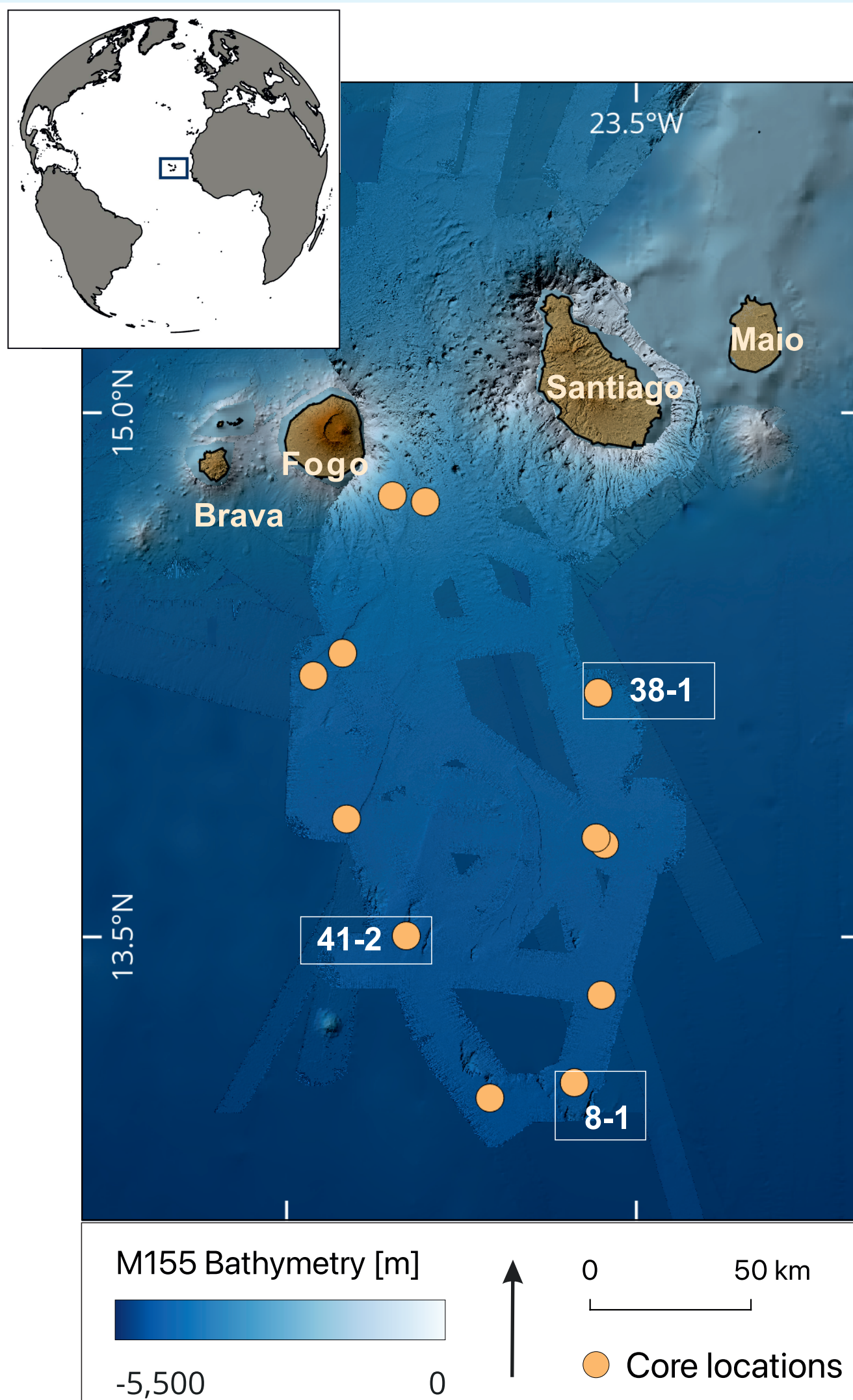
Linking composition and emplacement mechanisms of volcanoclastic mass transport deposits offshore Fogo, Cabo Verde

Janne Scheffler¹, Steffen Kutterolf¹, Emma Hadré², Ricardo Ramalho³, Andreas Klügel⁴, Josephin Wolf², Johanna Schenk¹, Sebastian Krastel²

¹GEOMAR Helmholtz Centre for Ocean Research, Kiel, Germany; ²Kiel University, Kiel, Germany; ³Cardiff University, Cardiff, Wales, United Kingdom; ⁴University of Bremen, Bremen, Germany



1. Introduction and Motivation



- Fogo (Cabo Verde) is one of the most active volcanoes in the Atlantic^[1].
- The island is shaped by an eastward facing scarp, which is the result of the Monte Amarelo flank collapse (68 ka)^[2]. The collapse is interpreted to have triggered a tsunami^[3].
- It is the source of several large sector collapses and mass-wasting events, whose debris fields extent far out into the surrounding abyssal plain^[4].
- These events pose threats to communities in coastal areas and across ocean basins.
- Marine tephra layers are used as stratigraphic markers as they provide the most complete archives of volcanic eruptions and mass-wasting events^[5].

Objectives

- Establish diagnostic criteria for **distinguishing** different types of **volcanoclastic mass-transport deposits** (VMTDs), interpret their **origin** and assign them to **emplacement mechanisms**
- Construct a spatially resolved framework of **depositional patterns** and **flow behaviour** of volcanogenic turbidites south of Fogo

2. Methods and Approach

Sampling

- Marine sediment cores were collected during RV Meteor voyage M155 using a gravity corer.

Sedimentology

- Epoxy peels were prepared to examine sedimentary structures in detail.

Petrology

- Major element oxides: Electron microprobe
- Trace elements: Laser-ICP-MS

Petrography

- Component inventory analysis using smear-slide microscopy

- Quantification of mineral-, lithic- and glass-content

Age dating

- Radiocarbon (¹⁴C) dating of foraminifera
- Calculation of sedimentation rates and age model

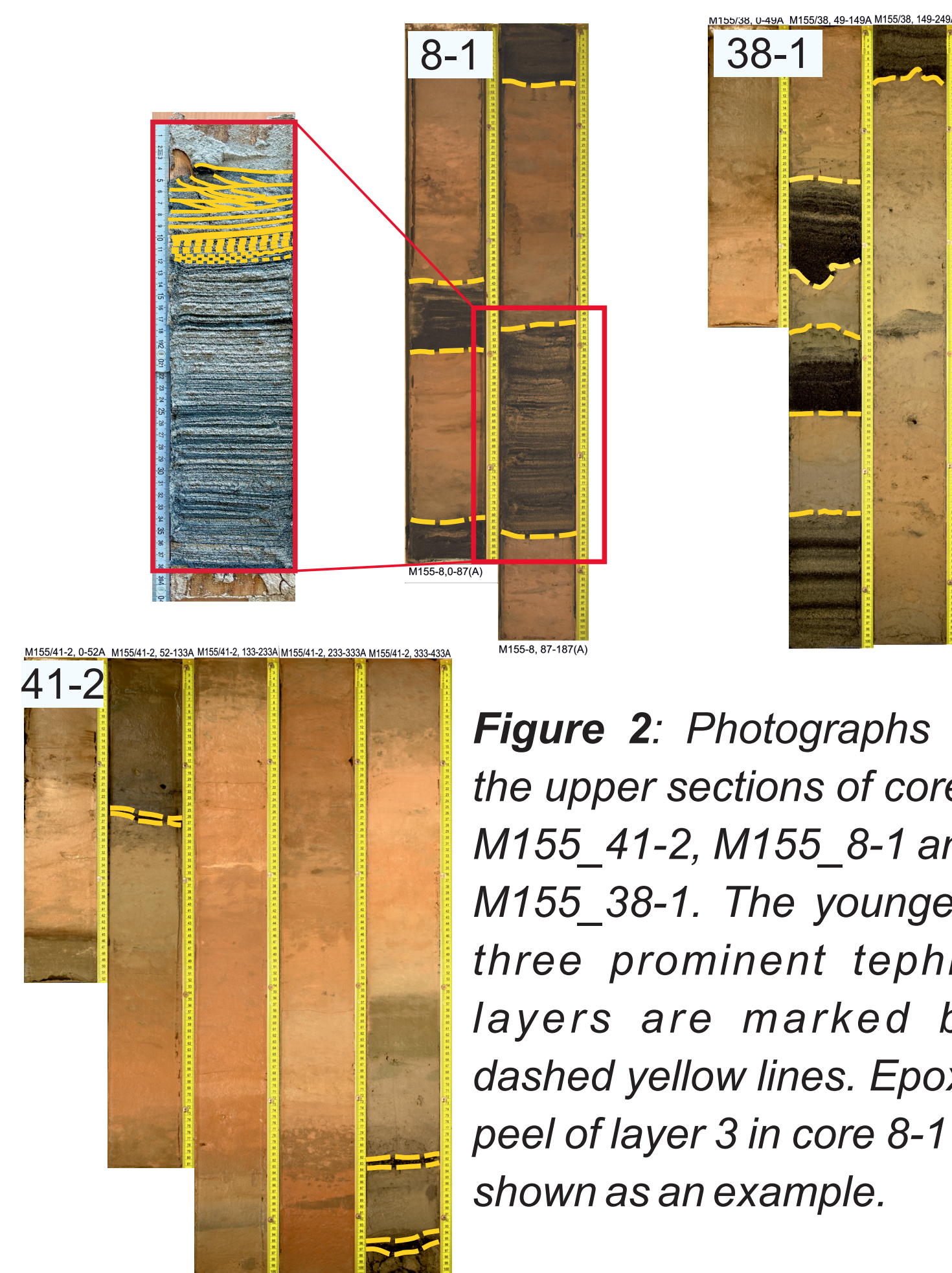


Figure 2: Photographs of the upper sections of cores M155_41-2, M155_8-1 and M155_38-1. The youngest three prominent tephra layers are marked by dashed yellow lines. Epoxy peel of layer 3 in core 8-1 is shown as an example.

3.1 Characterisation of volcanoclastic event layers

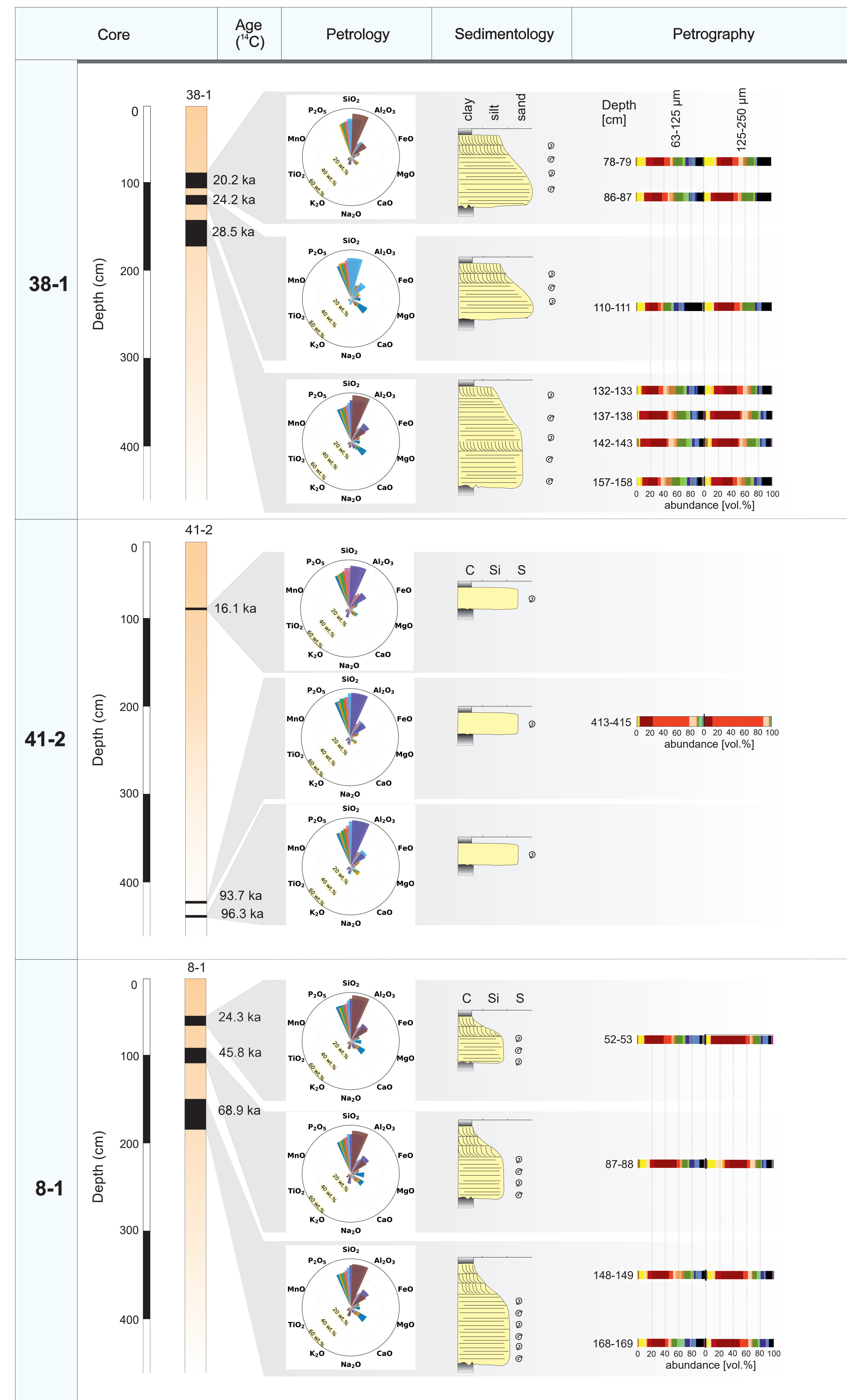


Figure 3: Integrated radiocarbon age, petrological, sedimentological and petrographic data from the youngest three ash layers in marine sediment cores 38-1, 41-2 and 8-1.

Chemical composition

- Population 1
- Population 2
- Population 3
- Population 4
- Population 5
- Population 6
- Population 7
- Population 8

Sedimentary feature

- Foraminifera
- Scour mark
- Massive
- Cross-laminae
- Planar laminae

Component

- Altered glass
- Brown glass
- Tachylitic glass
- Transparent glass
- Pumice
- Cumulate clast
- Volcanic lithic dark
- Volcanic lithic with glass
- Bioclast (Foraminifera)
- Carbonate clast
- Sedimentary clast
- Augite
- Enstatite
- Hypersthene
- Hornblende
- Plagioclase
- Potassium feldspar
- Olivine
- Rutile
- Titanite
- Fluorite
- Zeolite
- Zircon
- Opaque
- Quartz
- Nepheline
- Sodalith group

3.2 Provenance

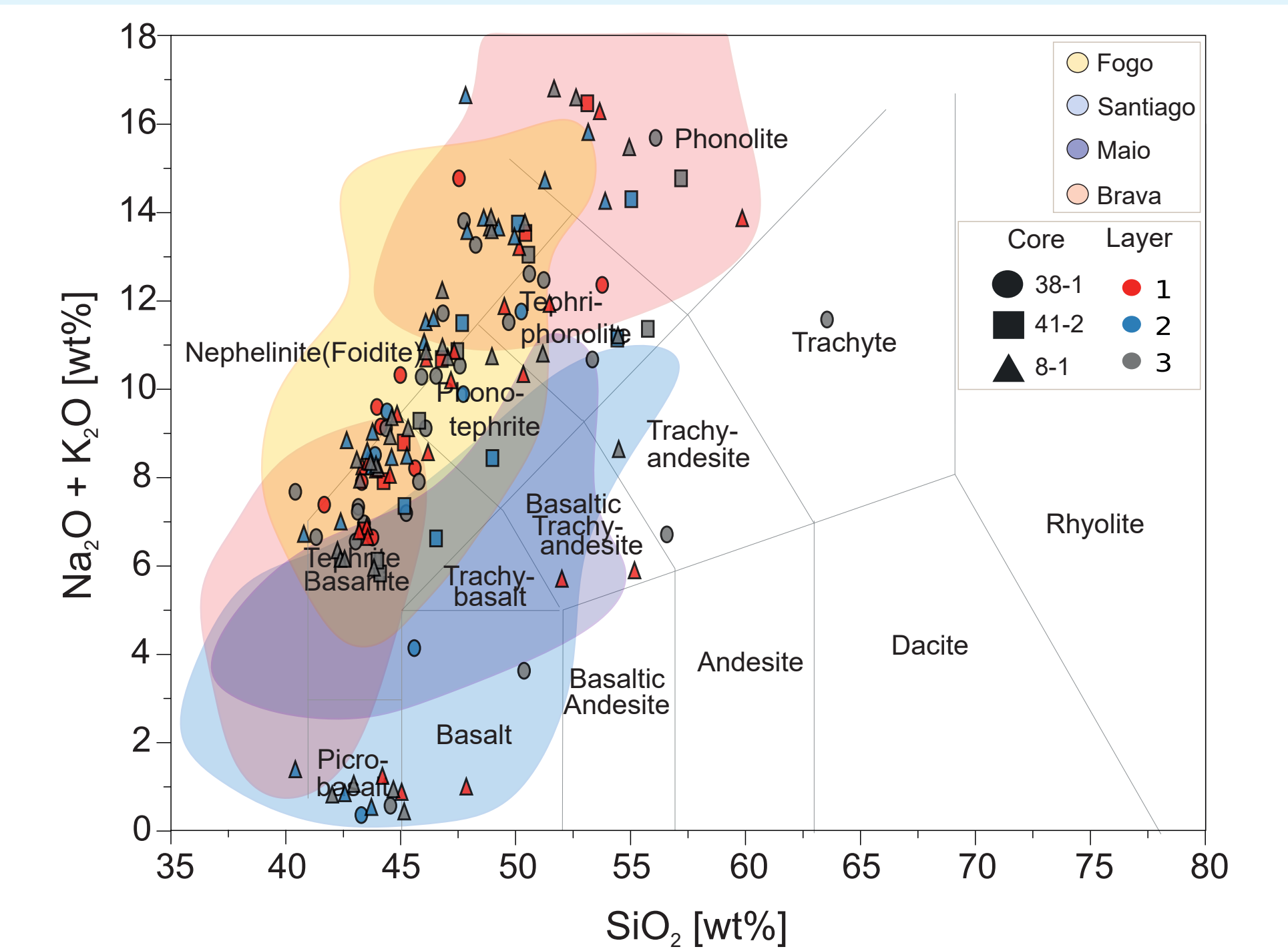


Figure 4: Total alkalis versus silica diagram with compositional fields of the southern island chain^{[6],[7],[8]}.

4. Interpretation

Core	Layer	Provenance				Emplacement	
		Reworked	Fogo	Santiago	Brava		Bouma
38-1	1	X	X		X	T _C , T _B	Alternating volcanoclastic and bioclastic laminae indicate episodic reworking of ash deposits and incorporation of biogenic material.
	2	X	X	X		T _C , T _B	
	3	X	X	X	X		
41-2	1	X	X		X	T _A	Massive layers (>80 vol.% volcanic lithics and glass) with minor foraminifera content indicate volcanoclastic density flow deposition.
	2	X	X	X	X	T _A	
	3	X	X		X		
8-1	1	X	X	X	X	T _C , T _B	Alternating planar laminae reflect density separation of volcanoclastic and bioclastic components during prolonged, pulsating flow conditions with internal fluctuations in velocity.
	2	X	X	X		T _C , T _B	
	3	X	X	X	X		

5. Conclusions and Outlook

- The integrated and detailed characterisation of volcanoclastic event layers is proven as useful for distinguishing deposit types and ultimately infer the materials origin.
- Ongoing work focuses on establishing a continuous composite VMTD chronostratigraphy through correlating marine tephra layers along and across arrays south of Fogo.
- Eventstratigraphy is applied to perform hazard assessment using integrated time-series analysis and probabilistic models of distribution and volume.

Get in touch

jscheffler@geomar.de



Literature

- ^[1]Masson, D. G., Le Bas, T. P., Grevenmeyer, I., & Weinrebe, W. (2008). Flank collapse and large-scale landsliding in the Cape Verde Islands, off West Africa. *Geochemistry, Geophysics, Geosystems*, 9(7). <https://doi.org/10.1029/2008GC001983>
- ^[2]Day, S. J., Heleno Da Silva, S. I. N., & Fonseca, J. F. B. D. (1999). A past giant lateral collapse and present-day flank instability of Fogo, Cape Verde Islands. *Journal of Volcanology and Geothermal Research*, 94(1-4), 191-218. [https://doi.org/10.1016/S0377-0273\(99\)00103-1](https://doi.org/10.1016/S0377-0273(99)00103-1)
- ^[3]Ramalho, R. S., Winkler, G., Madeira, J., Helfrich, G. R., Hipólito, A., Quartau, R., et al. (2015). Hazard potential of volcanic flank collapses raised by new megatsunami evidence. *Science Advances*, 1(9), e1500456. <https://doi.org/10.1126/sciadv.1500456>
- ^[4]Le Bas, T. P., Masson, D. G., Holtom, R. T., & Grevenmeyer, I. (2007). Slope Failures Of The Flanks Of The Southern Cape Verde Islands. In V. Lykousis, D. Sakellariou, & J. Locat (Eds.), *Submarine Mass Movements and Their Consequences* (pp. 337-345). Dordrecht: Springer Netherlands. https://doi.org/10.1007/978-1-4020-6512-5_35
- ^[5]Kutterolf, S., Freundt, A., Pérez, W., Mörz, T., Schacht, U., Wehrmann, H., & Schmincke, H.-U. (2008). Pacific offshore record of plinian arc volcanism in Central America: 1. Along-arc correlations: PLINIAN ARC VOLCANISM, 1. *Geochemistry, Geophysics, Geosystems*, 9(2), n/a-n/a. <https://doi.org/10.1029/2007GC001631>
- ^[6]Davies, G., Cliff, R., Norry, M., & Gerlach, D. (1989). A combined chemical and Pb-Sr-Nd isotope study of the Azores and Cape Verde hot-spots: the geodynamic implications. *Geological Society, London, Special Publications*, 42(1), 231-255. <https://doi.org/10.1144/GSL.SP.1989.042.01.15>
- ^[7]Barker, A. K., Holm, P. M., Peate, D. W., & Baker, J. A. (2009). Geochemical Stratigraphy of Submarine Lavas (3-5 Ma) from the Flamengos Valley, Santiago, Southern Cape Verde Islands. *Journal of Petrology*, 50(1), 169-193. <https://doi.org/10.1093/petrology/egn081>
- ^[8]Mourão, C., Mata, J., Doucelance, R., Madeira, J., Millet, M.-A., & Moreira, M. (2012). Geochemical temporal evolution of Brava Island magmatism: Constraints on the variability of Cape Verde mantle sources and on carbonatite-silicate magma link. *Chemical Geology*, 334, 44-61. <https://doi.org/10.1016/j.chemgeo.2012.09.031>

Acknowledgements

This work is funded by the German Research Foundation (DFG) with grant no.507049253. We acknowledge the work of the shipboard crew and scientific party of M155, who ensured successful operations at sea.

Abstract:

