

Analysis of the decoupling between surface heat flux and temperature gradient during afternoon transition

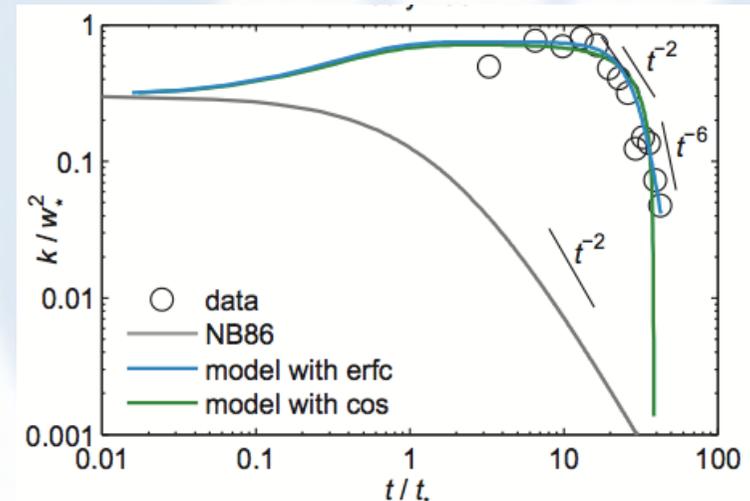
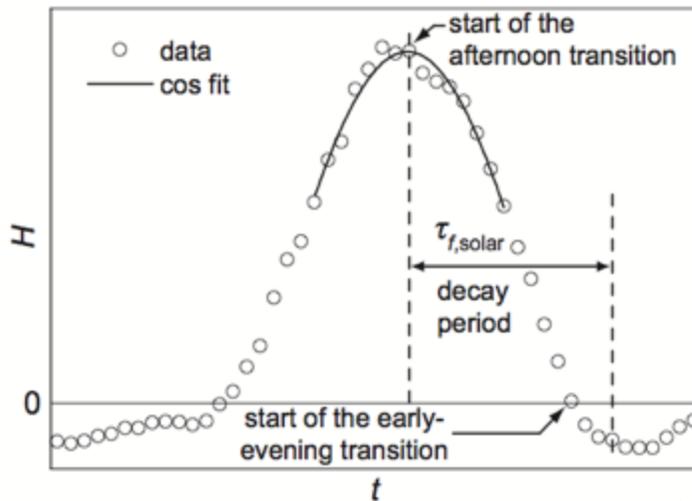
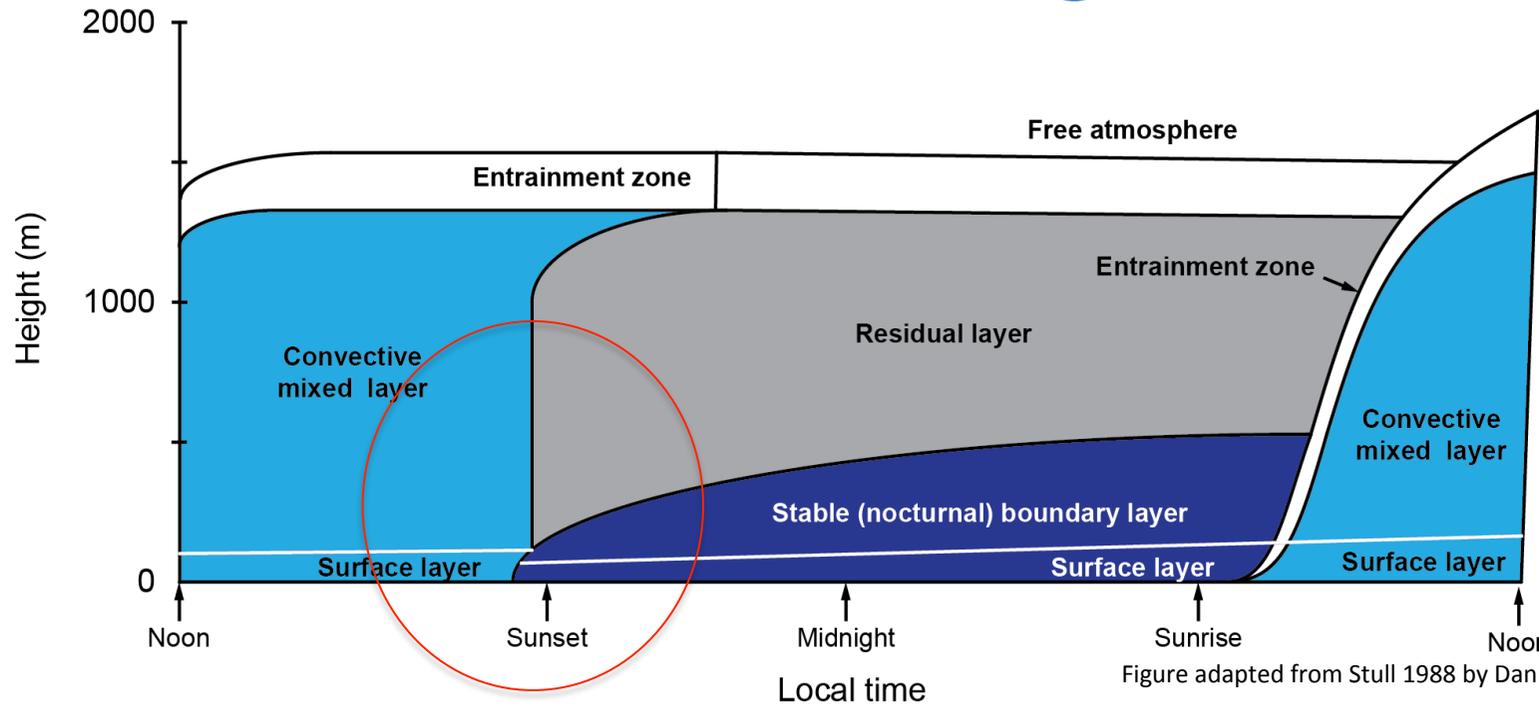
Estel Blay Carreras¹, E. Pardyjak², D. Pino¹, M. Lothon³, F. Lohou³

(1)Universitat Politècnica de Catalunya (UPC), SPAIN

(2)University of Utah, USA

(3)Laboratoire d'Aérodynamique, CNRS, FRANCE

Afternoon and Evening Transition



From Nadeau et al. 2011, BLM, **26** (2), 301-324

Motivation

- 1) Does the buoyancy flux cease at the same time the local gradient of the virtual potential temperature becomes positive (as predicted by flux gradient theory)?
- 2) If a delay exists, can it be parameterized? What are the physics that govern the delay?
- 3) Can these shortcomings be explained with a simple counter gradient formulation?
- 4) Do turbulent viscosity and thermal diffusivity play an important role during afternoon transition at the surface layer?



BLLAST - Boundary Layer Late Afternoon and Sunset Transition Experiment

- Lannemezan, France 14 June – 8 July 2011
- Lead Scientist - M. Lothon, Laboratoire d'Aérodologie, CNRS, FRANCE

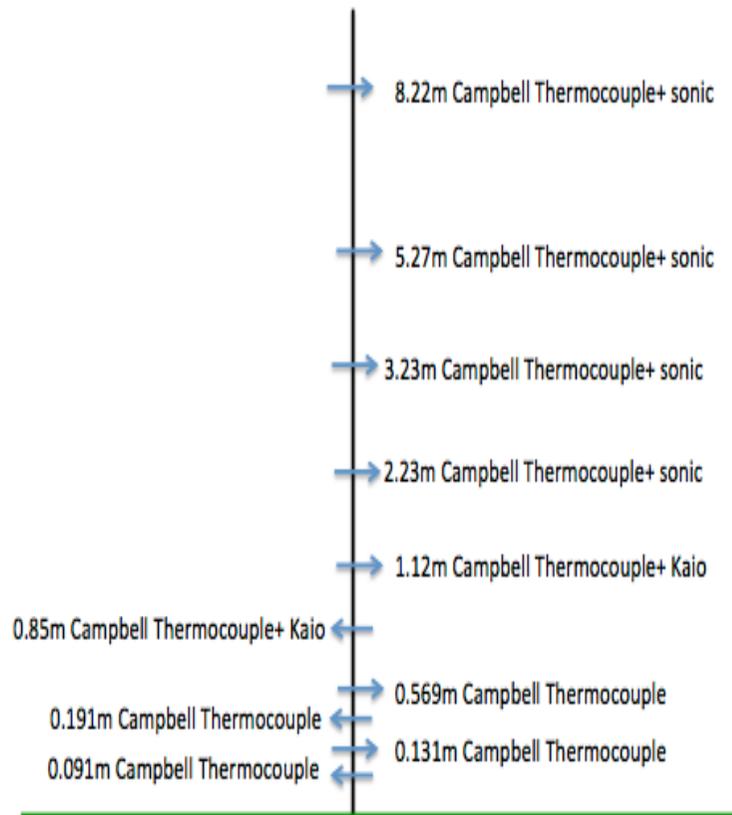
Objectives of the BLLAST Campaign

- To understand the importance of surface heterogeneity during the Late Afternoon Transition (LAT).
- To study the vertical structure and evolution of the boundary layer itself during LAT.

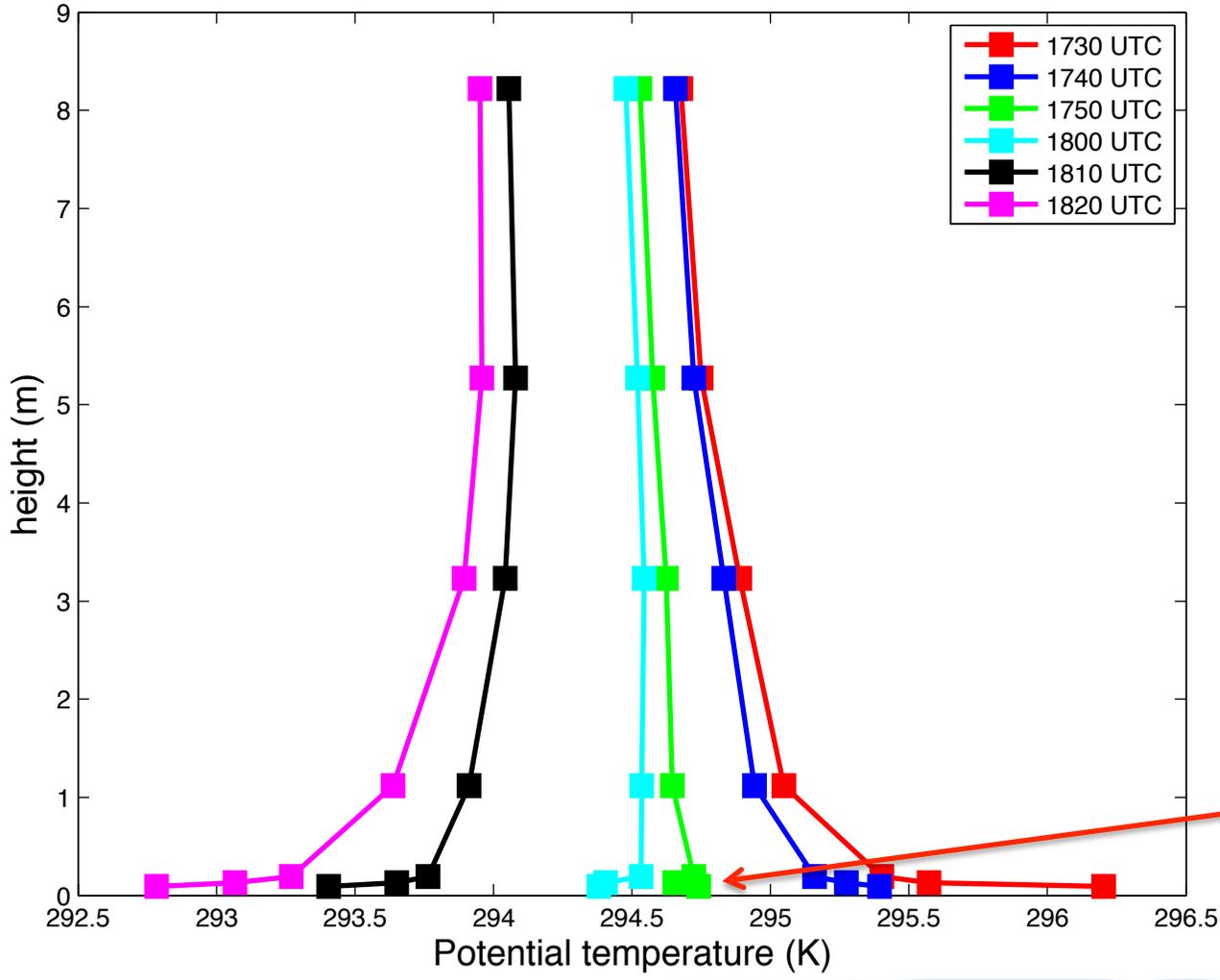


Skin Flower Tower – Site 1

- 3D Sonics at 6 Levels
- FW TCs at 9 levels



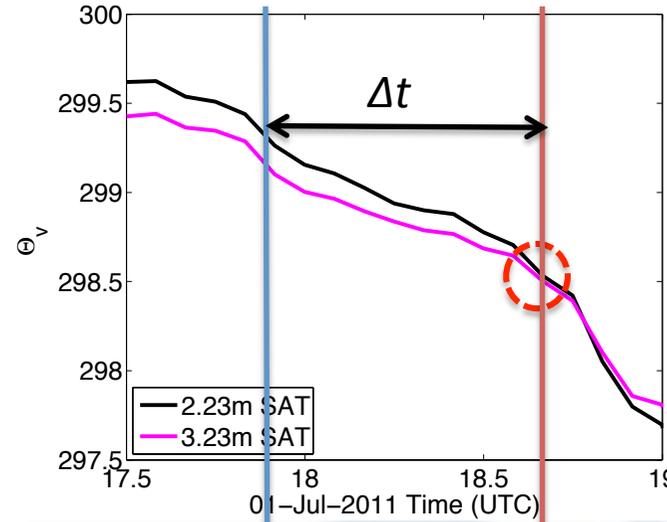
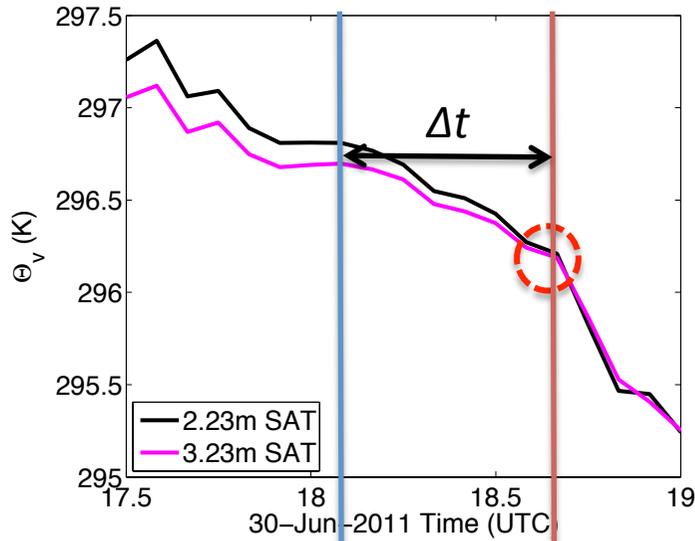
Temperature Profile Observations



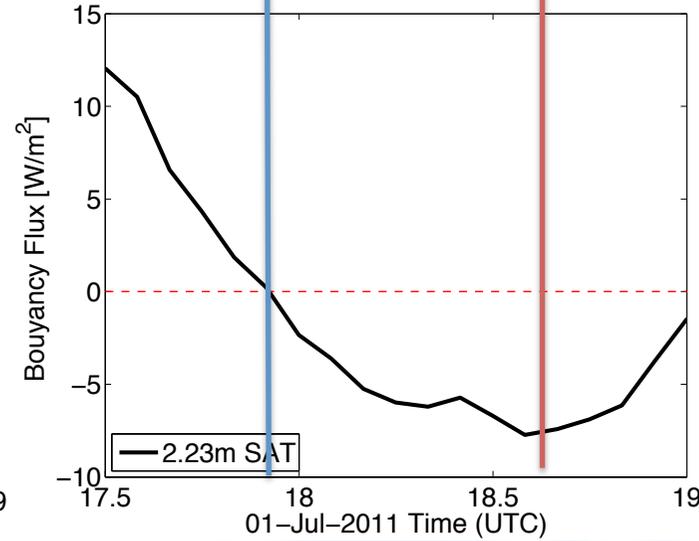
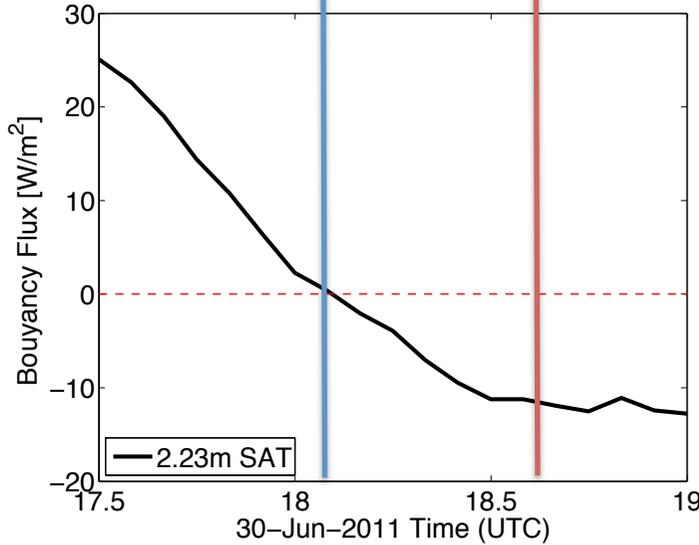
Possible presence of Lifted Temperature Minimum (Future Work)

1-Jun-2011

Methodology

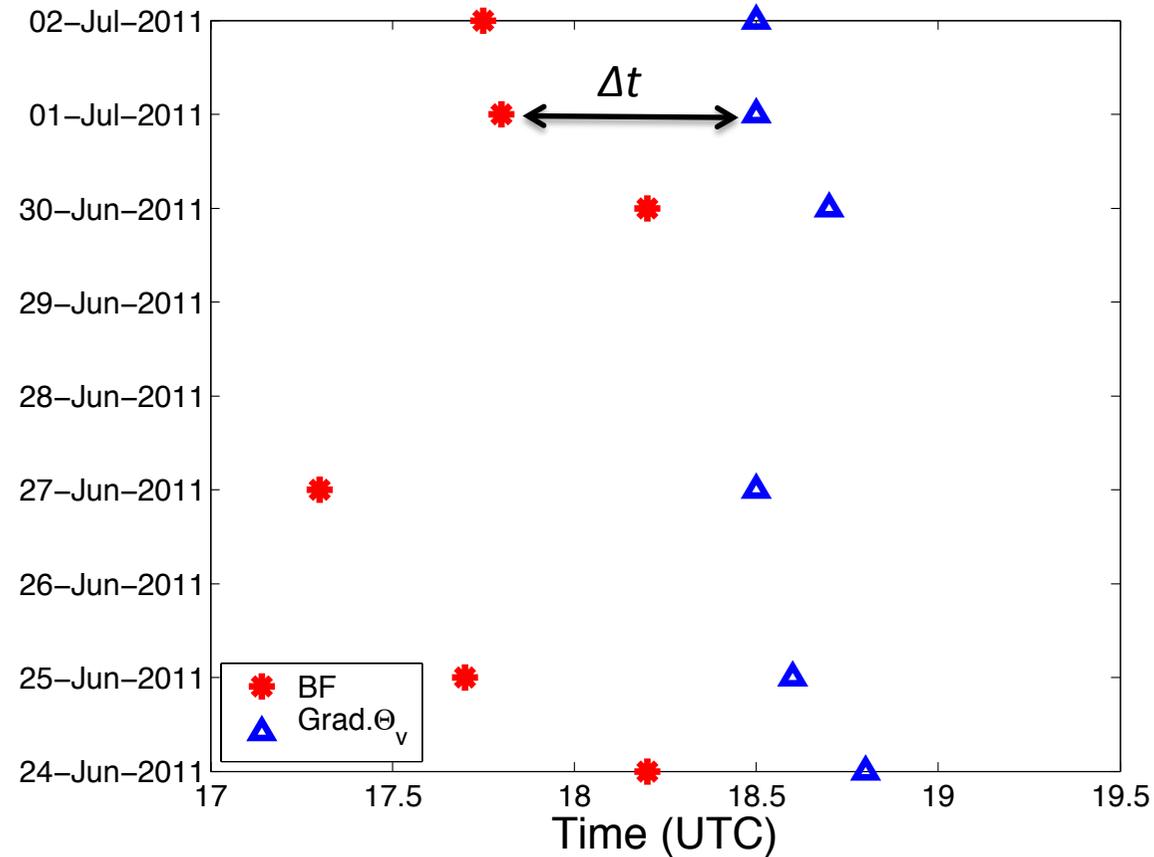


Change of the sign of the θ_v gradient



Change of sign of the buoyancy flux

Results

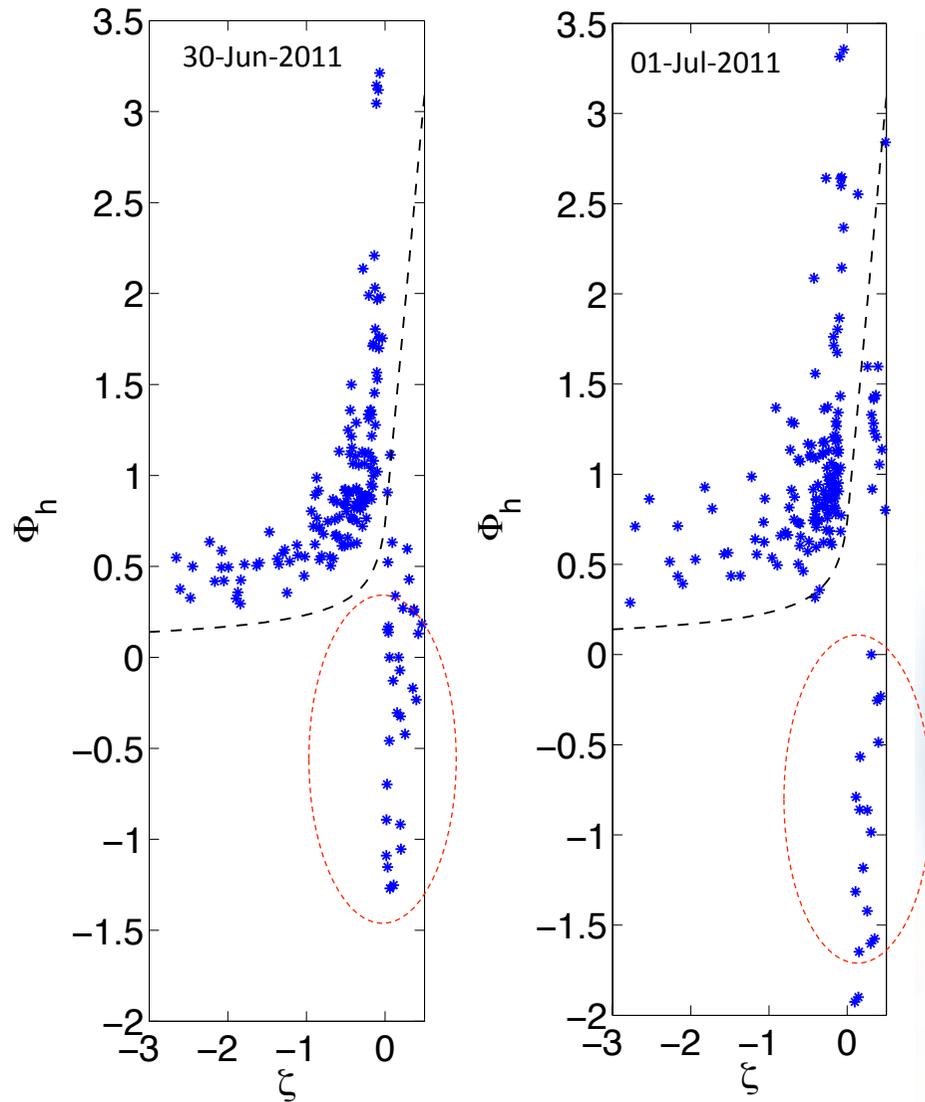


$$t\left(\overline{w'\theta'_v}\Big|_s \approx 0\right) \text{ vs } t\left(\partial\theta_v / \partial z \approx 0\right)$$

There is a delay time in all the IOP analyzed (30-70 min)

Results

MOST Temperature Gradients



Presence of Counter Gradient Fluxes

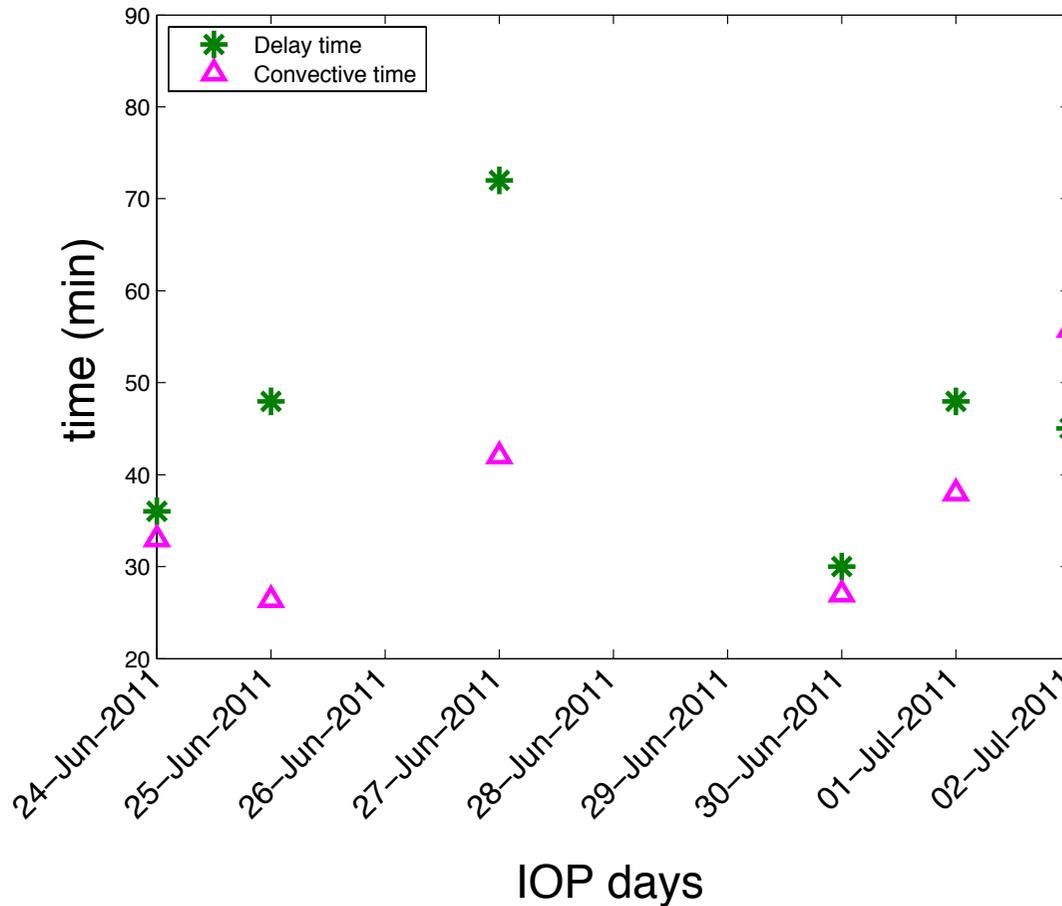
$$\phi_h = \frac{kz}{\theta_*^{SL}} \frac{\partial \theta}{\partial z} = \frac{kzu_*}{-\overline{(\omega' \theta_v')}_s} \frac{\partial \theta}{\partial z}$$

$$\zeta = \frac{z}{L} = \frac{-kzg \overline{(\omega' \theta_v')}_s}{\theta_v u_*^3}$$

Results

Convective time

$$Delay\ time\ (DT) = t(\partial\theta_v / \partial z \approx 0) - t(\overline{w'\theta'_v}|_s \approx 0)$$



Does delay time appear for the last eddy movement?

Convective time :

$$w_* = \left[\frac{gz_i}{\theta_v} (\overline{w'\theta'_v})_s \right]^{1/3}$$

$$t_* = \frac{z_i}{w_*}$$

Why some days DT is similar to the convective time?

Results

Obukhov Length

Convective days

24/06 & 30/06: IOPs with **large** $-z/L$ averaged between 12 UTC-16:45UTC have **small** DT-CT.

Weakly convective days

25/06 & 27/06 → IOP with **small** $-z/L$ averaged between 12UTC-16:45UTC have **large** DT-CT.



Why??



Weakly convective IOPs have larger u_* → more horizontal turbulence → larger delay time

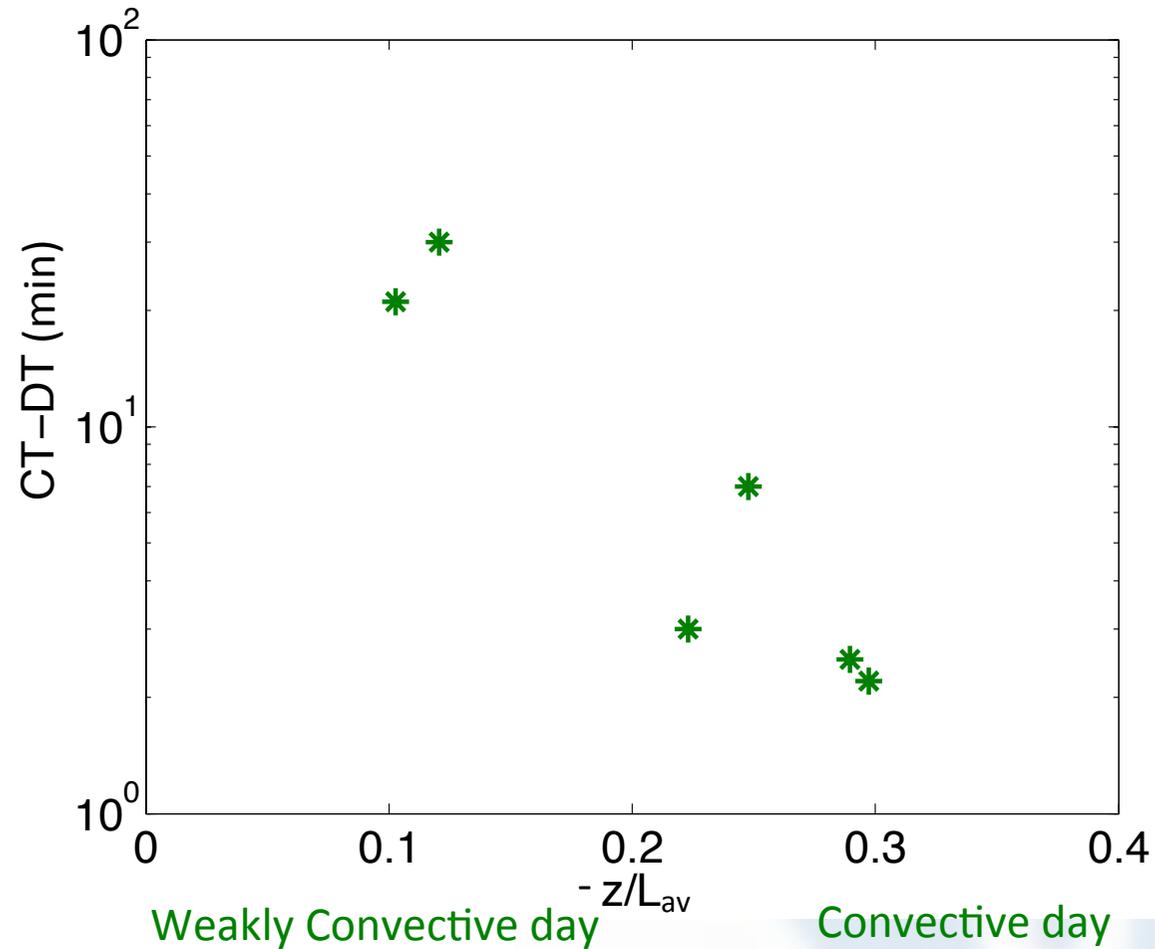
$$\zeta = -\frac{z}{L} = \frac{kzg \left(\overline{w' \theta_v'} \right)_s}{\overline{\theta_v} u_*^3}$$

$z=2.23 \text{ m}$

Results

Obukhov Length

Exponential relation between DT-CT and $-z/L$



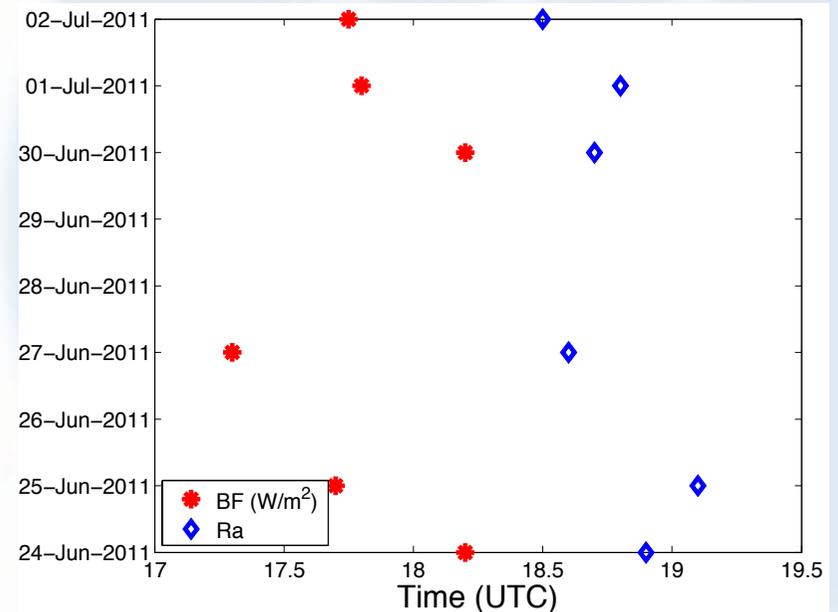
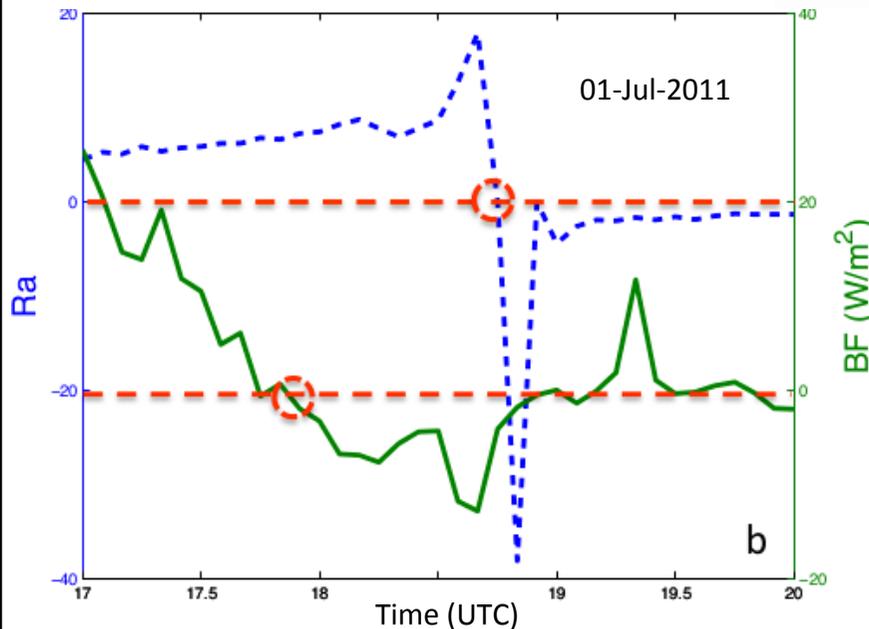
Results

Turbulent Rayleigh number physical approximation

Ra → compares the destabilizing forces (buoyancy forces) with the stabilizing forces (viscosity and thermal diffusivity).

Bénard problem → turbulent viscosity and turbulent thermal diffusivity difficult convection movements.

In all IOP buoyancy flux ceases before Ra is negative → a physical approximation → during this period viscosity and thermal diffusivity may play a role in the slow down of the cease of the convection.



Conclusions

- There is a **delay** between buoyancy flux cease and the change in the vertical gradient of θ_v .
- During moderate convective days, the delay time is small and close to the last eddy movement (**convective time**).
- When convection is lower, larger u_* , the delay time is larger due to the increase of **horizontal turbulence**.
- **Turbulent viscosity and thermal diffusivity** may help to slow down the last eddy movement increasing the convective time.

Thank you

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