

Aerosol Climatology on Mars as Observed by NOMAD UVIS/SO on ExoMars TGO

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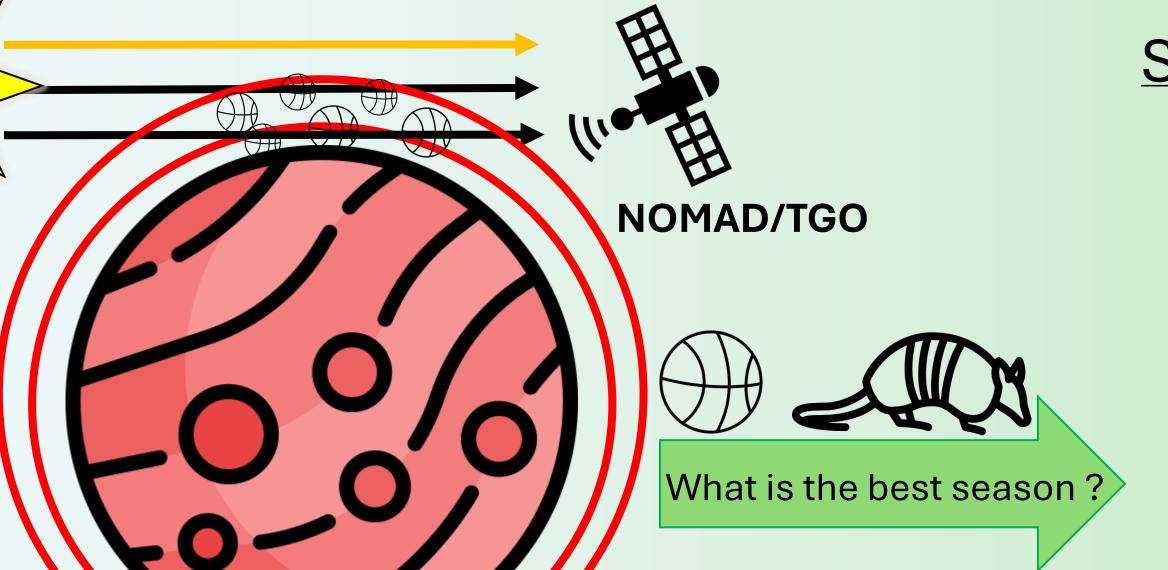


How well can we constrain aerosols with remote sensing?

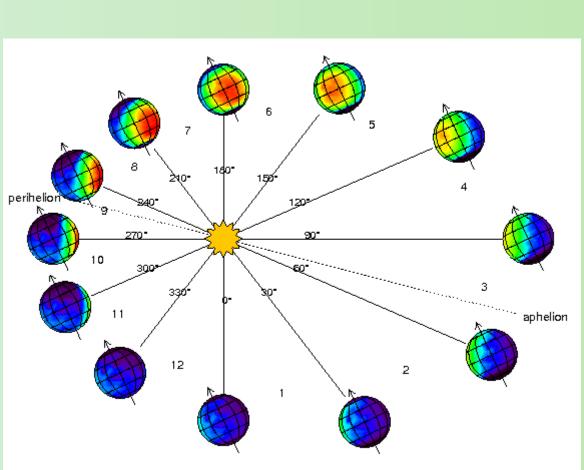
What if there were armadillos in the atmosphere?



To compute aerosols property, we are doing the following methodology



Seasonal cycle on Mars



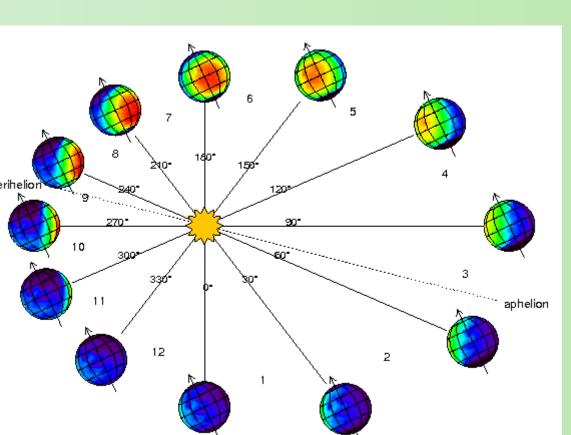
How to differentiate armadillos from aerosols?

Refractive indices

Dust and ice are indistinguishable in the UV

Cross sections

- Mie code (spherical shape)
- Monomodal log normal distribution
- r_{eff} : [0.1, 0.8] μm
- V_{eff}: 0.1



Fit to find the best size

- $\beta = n * C_{ext}(reff, veff)$, with β the extinction, n the number density and C_{ext} the cross section
- Fit with least square algorithm
- Best reduced chi square obtained gives us the optimal r_{eff}

Aerosols in the Martian atmosphere

Dust:

- Heats the atmosphere
- Impacts with other retrievals
- Dust storms are present, and more active during perihelion

lce:

IR (SO) channels: 2.3 – 4.3 μm

- Could be H₂O or CO₂
- Reflects sunlight and cool the atmosphere
- Forms on dust condensation nuclei
- More present during aphelion

UV (UVIS): 200-650 nm

0.100 0.125 0.150 0.175 0.200 0.225 0.250 0.275 0.300 Observation name: 20190505_070003_1p0a_UVIS_I $-r_{eff} = 0.75 \, \mu m, \, v_{eff} = 0.065$ - r_{eff} = 1.1 μm, v_{eff} = 0.065 UVIS data 1.00 1.25 1.50 1.75 2.00 0.25 0.50 0.75

Size sensitivity study on the UVIS data. The right panel shows the reduce chi squared of different r_{eff} and v_{eff} . The sensitivity limit of the dataset is given as $0.1-0.8 \mu m$.

The reduced chi square is constrained for small r_{eff}

v_{eff} doesn't change the retrieved r_{eff} by more than $0.05\,\mu m$

Large particles gives a flat spectra which make it difficult to fit accurately

Altitude (km): 53 dust : r_{eff}= 0 μm, RCS =24402 3200 Wavelength (nm)

5 orders of diffraction are used to retrieved the background

Aerosols composition can be differentiated, armadillos and H₂O ice are spectrally different

Altitude (km): 13 H_2O ice : $r_{eff} = 5 \mu m$, RCS = 681 10^{-2} CO_2 ice : r_{eff} = 2 μ m, RCS =32 8×10^{-3} 7×10^{-3} 6×10^{-3} 3000 3200 Wavelength (nm)

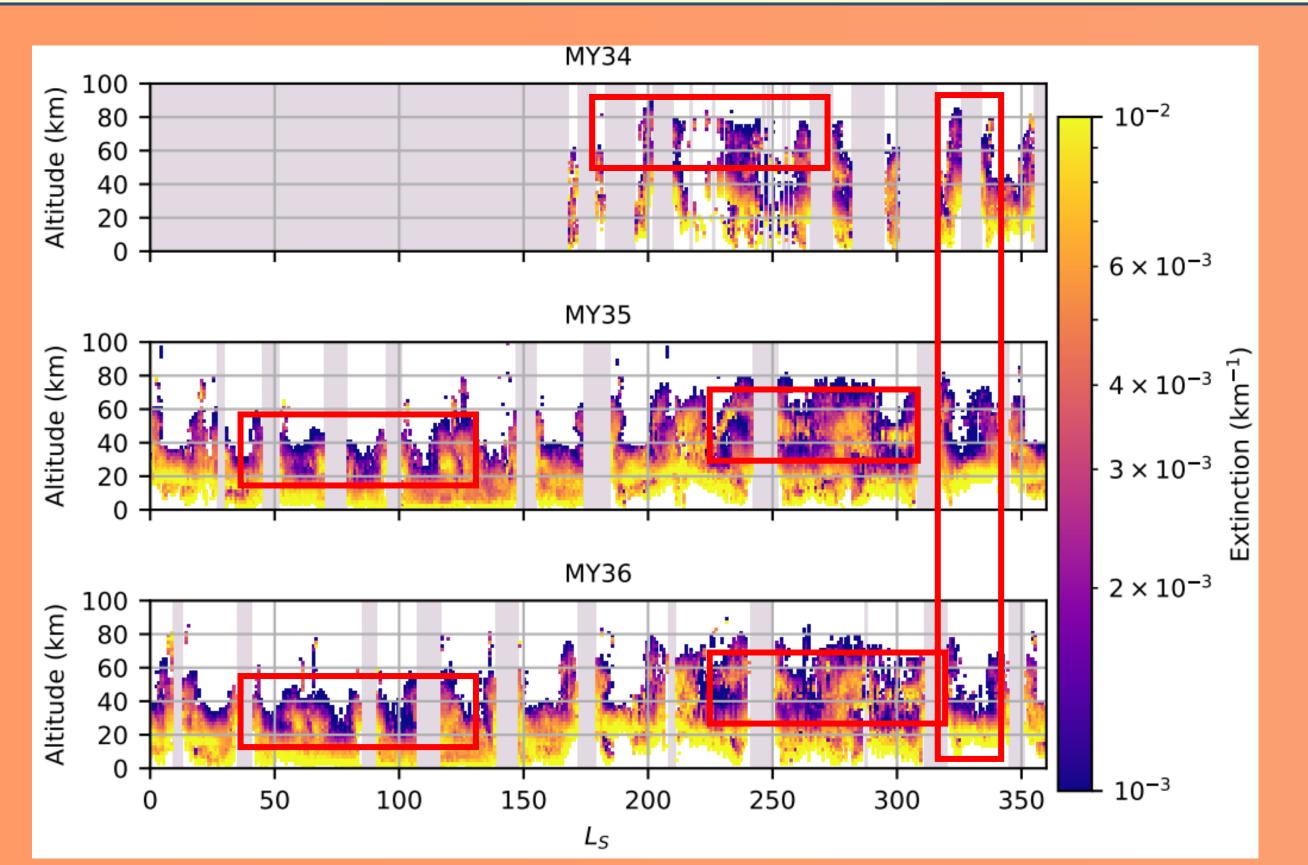
CO₂ ice signature is within 1 diffraction order of the instrument

Dust and armadillos

don't have signatures

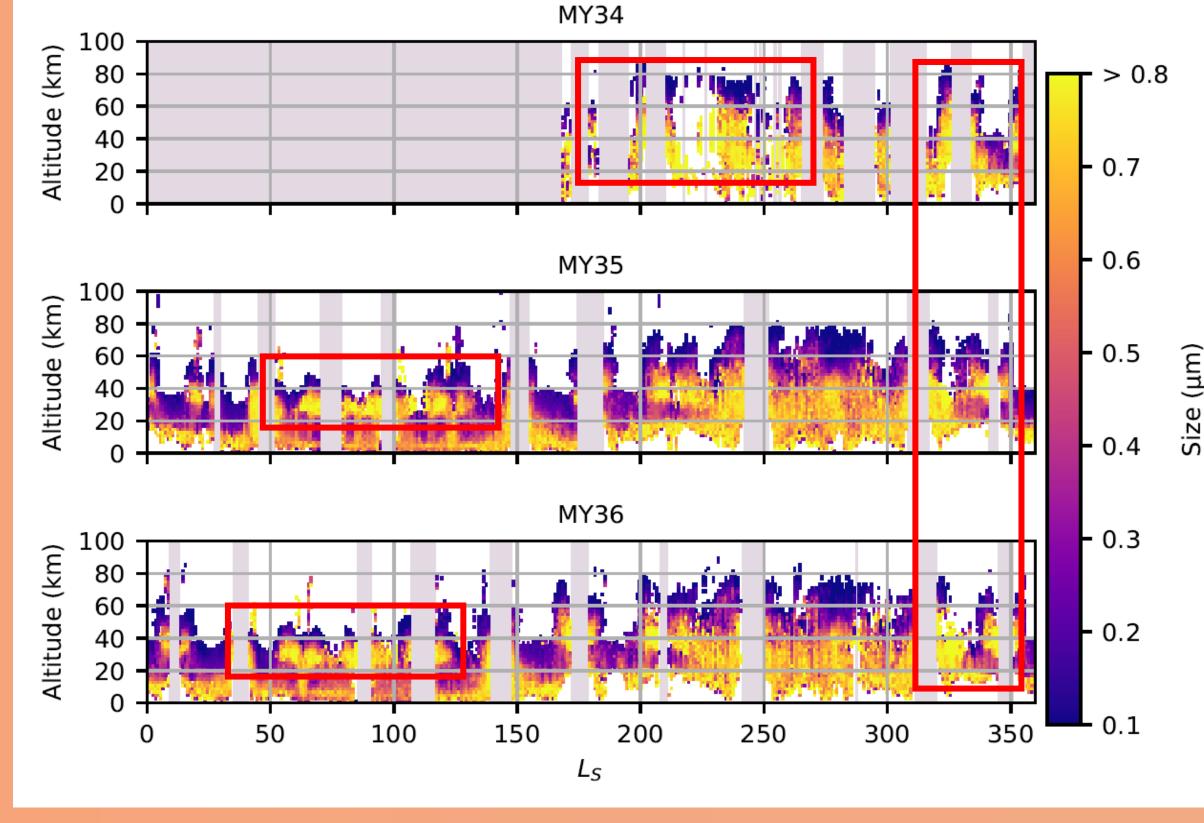
SO extinction profiles with different aerosols fit. Panel A shows water ice absorption and panel B a non water ice profile.

Aerosol climatology



Vertical extinction profiles for latitude bins by row [30° 70°], [-30° 30°], [-70° -30°], respectively from top to bottom

- Similar vertical structure are reproduced from one Martian year to another
- Detached layers are present around 50 km in the northern regions and up to 70 km in the equator
- Dust storm can be observed at L_S 330° in all MY
- The global dust storm is present in MY 34 around L_S 200°



Vertical size profiles for latitude bins by row [30° 70°], [-30° 30°], [-70° -30°], respectively from top to bottom

Conclusion

- We produced an aerosol climatology covering three different Martian years
- We derived particle sizes between **0.1** and **0.8** µm and information on the presence of larger particles
- No composition can be derived spectrally in the UV but the use of the IR channel can help to constrain it
- In future work, we will retrieve the aerosols property with both UVIS and SO to add more constraints on our retrieval and gains new insight.

