Global Soil Erosion Modelling, First Results V. Naipal, Ch. H. Reick and J. Pongratz Land Department, Max-Planck Institute for Meteorology, Hamburg Contact: victoria.naipal@zmaw.de



Background

In this study we try to understand the effects of soil erosion on the global carbon cycle in different climate and land-use scenarios

Aim of the study: To model large scale historical soil erosion events and their impacts on the carbon cycle

Here we discuss:

- How to model soil erosion processes
- How to make RUSLE applicable on global scale focusing on the scale dependency of topography
- Uncertainties that are related to the used methods

Tool: RUSLE Model (*Renard et al., 1997*)

First Results



- RUSLE is a semi empirical/process-based model ullet
- Parameterized for agricultural areas on the plot scale
- Does not simulate deposition and sediment transport

Global mean soil erosion rates for July 2012, on a 0.5 degree scale



- Global mean soil erosion July 6.3 ton ha⁻¹ year⁻¹ and December 11.1 ton ha⁻¹ year⁻¹ \bullet
- We found a general underestimation of erosion on agricultural areas due to the coarse \bullet scale of the datasets compared to the local scale of erosion
- We also found unrealistic high erosion rates in mountainous areas due to the absence of ۲ a "limit" to erosion in the model

Scale dependency of the steepness factor (S)

Scale dependency of the slope-length factor (L)

- The S factor can be calculated by different methods, which show the largest deviation for large slopes (figure right panel)
- The S factor is a function of slope and the slope is highly dependent on the spatial resolution of the Digital Elevation Model, DEM (figure left panel)



Fractal method for deriving high resolution slope from coarse-scale **data** (*Zhang et al., 1999*)

 $S = \alpha d^{1-D}$

S = percentage slope

 α = coefficient

d = grid resolution in meters

D = fractal dimension



- Assumption 1: Topography obeys fractal statistics
- Assumption 2: Standard deviation of elevation independent of DEM resolution



- The L factor can be estimated by different methods, among which flow accumulation is for high resolution and local scale, and Yang (2001) method for the hydrological catchment scale
- The L factor is a function of slope and slope-length and therefore depends on the spatial resolution

Method for calculation	Global	Global
of L	Mean	Standard-
		deviation
L – Flow Accumulation	7.73	26.92
L – Yang (2001)	1.48	1.31
L - slope length = 100 m	1.30	0.35

Flow accumulation vs grid resolution for a 1



Method 1: slope-length = flow accumulation * grid resolution

This does not apply for very coarse datasets (global scale) due to the large grid resolution







Slope derived from 1km DEM

Fractal slope derived from 1km DEM using grid resolution of 30m

We find a clear gain in detail in the 30m fractal slope derived from the 1km DEM compared to the original slope derived from the 1km DEM

Take-Home Messages

- This study reveals the strong dependence of soil erosion modeling on the spatial resolution
- The scale dependency, as showed for topography, is also expected \bullet to play an important role for the other RUSLE factors
- To make RUSLE applicable on global scale both temporal and spatial scaling of the RUSLE factors is needed

